

# SCIENTIFIC REPORTS

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## Quantifying Fenton reaction pathways driven by self-generated H<sub>2</sub>O<sub>2</sub> on pyrite surfaces

C. Gil-Lozano<sup>1</sup>, A. F. Davila<sup>2</sup>, E. Losa-Adams<sup>1,3</sup>, A. G. Fairén<sup>1,4</sup> & L. Gago-Duport<sup>3</sup>

Received: 19 October 2016

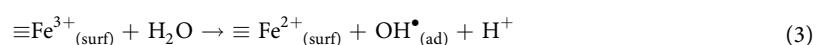
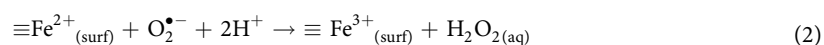
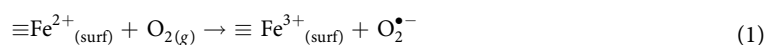
Accepted: 26 January 2017

Published: 06 March 2017

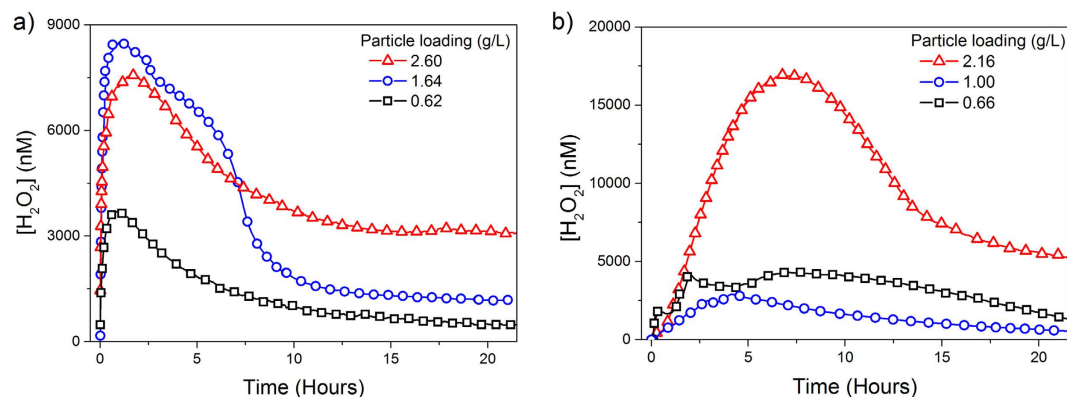
Oxidation of pyrite (FeS<sub>2</sub>) plays a significant role in the redox cycling of iron and sulfur on Earth and is the primary cause of acid mine drainage (AMD). It has been established that this process involves multi-step electron-transfer reactions between surface defects and adsorbed O<sub>2</sub> and H<sub>2</sub>O, releasing sulfoxy species (e.g., S<sub>2</sub>O<sub>3</sub><sup>2-</sup>, SO<sub>4</sub><sup>2-</sup>) and ferrous iron (Fe<sup>2+</sup>) to the solution and also producing intermediate by-products, such as hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>) and other reactive oxygen species (ROS), however, our understanding of the kinetics of these transient species is still limited. We investigated the kinetics of H<sub>2</sub>O<sub>2</sub> formation in aqueous suspensions of FeS<sub>2</sub> microparticles by monitoring, in real time, the H<sub>2</sub>O<sub>2</sub> and dissolved O<sub>2</sub> concentration under oxic and anoxic conditions using amperometric microsensors. Additional spectroscopic and structural analyses were done to track the dependencies between the process of FeS<sub>2</sub> dissolution and the degradation of H<sub>2</sub>O<sub>2</sub> through the Fenton reaction. Based on our experimental results, we built a kinetic model which explains the observed trend of H<sub>2</sub>O<sub>2</sub>, showing that FeS<sub>2</sub> dissolution can act as a natural Fenton reagent, influencing the oxidation of third-party species during the long term evolution of geochemical systems, even in oxygen-limited environments.

The chemical and electrical properties of pyrite (FeS<sub>2</sub>) – the most common iron disulfide in the Earth's crust – make it a promising material for the construction of photovoltaic panels<sup>1</sup>, as well as for wastewater remediation<sup>2–7</sup>. In the last decades, a number of experimental and field studies have addressed the oxidative dissolution of pyrite, providing an accurate characterization of the process<sup>8–17</sup>. Recent investigations have shown that, aside from the iron and sulfoxy species released during pyrite dissolution, ROS are always present as transient by-products, both under oxic and anoxic conditions<sup>18–26</sup>, and providing an important pathway to oxidize third-party species through oxygen evolution reaction (OER)<sup>18,27–30</sup>.

The rupture of S-Fe and S-S bonds over pyrite surface by mechanical fracture or during the dissolution process induces the formation of the non-stoichiometric defect sites (*i.e.*, the dangling bonds) that trigger adsorption reactions<sup>31–33</sup>. The iron surface species formed cause a reduction in the band gap of the pyrite from 0.86 electron volts (eV) in the bulk to 0.55 eV on the surface<sup>34</sup>, catalyzing the dissociation of adsorbed oxygen (O<sub>2</sub>) and water (H<sub>2</sub>O) molecules and leading to the formation of H<sub>2</sub>O<sub>2</sub> (equations 1–4)<sup>19,25,26,35</sup>:

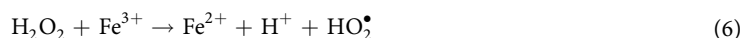
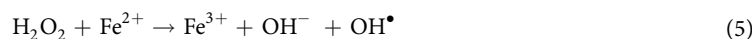


<sup>1</sup>Centro de Astrobiología (CSIC-INTA), 28850 Torrejón de Ardoz, Madrid, Spain. <sup>2</sup>Carl Sagan Center at the SETI Institute, 189 Bernardo Avenue, Suite 100, Mountain View, CA 94043, USA. <sup>3</sup>Departamento de Geociencias Marinas, Universidad de Vigo, Lagoas Marcosende, 36310-Vigo, Spain. <sup>4</sup>Department of Astronomy, Cornell University, Ithaca, 14853 NY, USA. Correspondence and requests for materials should be addressed to C.G.L. (email: cgil@inta.cab-csic.es)



**Figure 1.**  $\text{H}_2\text{O}_2$  evolution induced by pyrite slurries in unbuffered water. Inset: particle loading (g/L) (a) oxic conditions, (b) anoxic conditions.

At the same time, pyrite releases  $\text{Fe}^{2+}$  into solution than can catalyze the Fenton and Haber-Weiss reactions, leading to the generation of other ROS (equations 5–8)<sup>4,22,25,26</sup>:



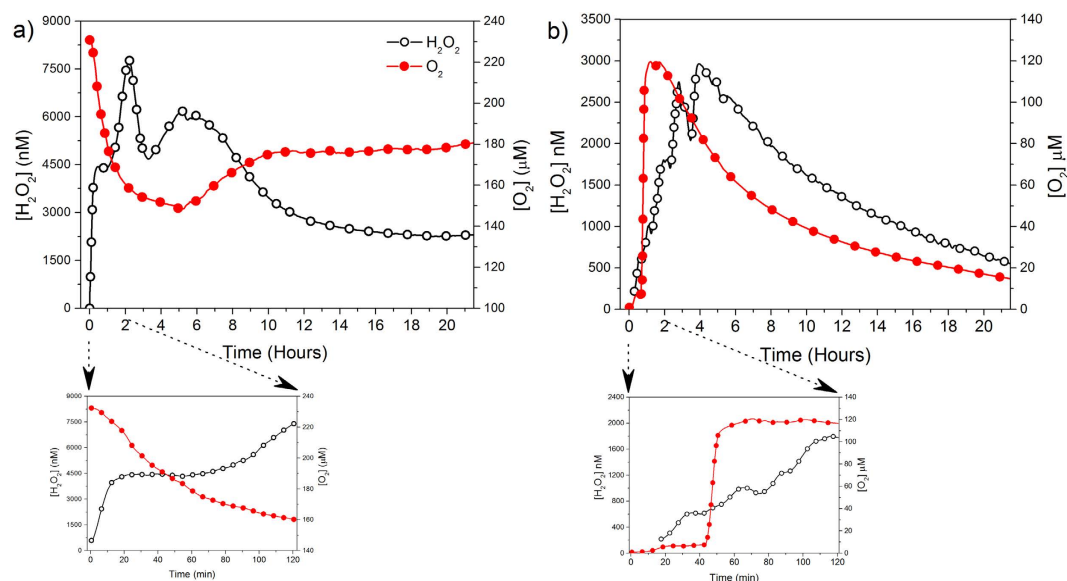
Although this sequence of reactions describes the basic features of  $\text{H}_2\text{O}_2$  formation during pyrite dissolution (equations 1–4) and its further degradation to secondary ROS in solution (equations 5–8), various important aspects remain unclear, for instance: (1) can free radicals be formed by mechanism other than the photoactivation (i.e. by mechanical bond fracture or non-stoichiometry dissolution of pyrite)? (2) can the formation of  $\text{H}_2\text{O}_2$  occur in strictly anoxic conditions overcoming the energy required to split the water molecule and the further release of  $\text{O}_2$ ? and (3) can the process of ROS generation by pyrite be sustained over long periods of time?

Real-time measurements of this process are made difficult by ROS reactivity and the subsequent redox transformations of iron and sulfur species. Spectroscopic and fluorescence methods are commonly used to measure the concentration of  $\text{H}_2\text{O}_2$ ; however, these methods usually need the presence of dyes or chelating agents that are not well suited for the kinetic analysis of transient phases. In this study, we measured the real-time generation and decomposition of  $\text{H}_2\text{O}_2$  and dissolved  $\text{O}_2$  induced by pyrite surfaces under different boundary conditions (i.e., dark/light, oxic/anoxic) to investigate the kinetic role of ROS during pyrite dissolution. In addition, we analyzed both the chemical evolution of dissolved species and surface pyrite oxidation with spectroscopy (UV-Vis, XPS), Cyclic Voltammetry (CV) and high-resolution transmission electron microscopy (HTREM) and performed specific experiments to evaluate the persistence of  $\text{H}_2\text{O}_2$  formation. Based on our results, we developed a kinetic model for the coupling between pyrite dissolution and  $\text{H}_2\text{O}_2$  generation/degradation through the Fenton reaction. When combined with the observed trends this model leads to the definition of constraints on the overall process of ROS oxidation mechanism induced by pyrite surfaces.

## Results

**$\text{H}_2\text{O}_2$  and  $\text{O}_2$  evolution: general pathway.** Figure 1 shows the concentration of  $\text{H}_2\text{O}_2$  as a function of time in aqueous suspensions of pyrite starting at circumneutral pH, under oxic and anoxic conditions.  $\text{H}_2\text{O}_2$  concentration increased at the beginning of the experiment until a maximum value was reached, and decreased thereafter asymptotically towards a nearly stationary, residual value ( $[\text{H}_2\text{O}_2] > 200$  nM) still measurable at the end of the experiment (~22 h). This coupled generation-decay response was generally observed in every experiment, although there were variations in the particular shape of the curves and, some experiments showed characteristic shoulders or secondary maxima at intermediate stages of the process. The overall process followed a sigmoidal trend, suggesting a strong interaction between the  $\text{H}_2\text{O}_2$  generation and degradation rates characteristic of an autocatalytic process.  $\text{H}_2\text{O}_2$  formation and decomposition evolved more slowly under anoxic than under oxic conditions. We did not observe any correlation between pyrite loading and the  $\text{H}_2\text{O}_2$  yield between experimental runs under our experimental conditions (i.e., 0.5–0.3 g/L, unbuffered neutral pH) because of several factors (e.g., kinks, steps, lattice anisotropy) can determine the variability of the reactive surface between samples<sup>36–39</sup>.

The evolution of  $\text{O}_2$  under oxic conditions was characterized by an asymptotic decrease followed by a slight increase and a steady stable period at the end of the experiment (Supplementary Fig. S1). The opposite trend was observed under anoxic conditions (i.e., an initial increase in  $\text{O}_2$  followed by an asymptotic decrease). To clarify the role of  $\text{O}_2$  over the formation of  $\text{H}_2\text{O}_2$ , we monitored simultaneously the evolution of  $\text{O}_2$  and  $\text{H}_2\text{O}_2$  under oxic



**Figure 2. Simultaneous trends of  $O_2$  and  $H_2O_2$  obtained from pyrite slurries in unbuffered water.** The bottom figure shows the first two hours of the reaction, under (a) oxic (particle loading = 1.75 g/l) and (b) anoxic conditions (particle loading = 1.09 g/L).

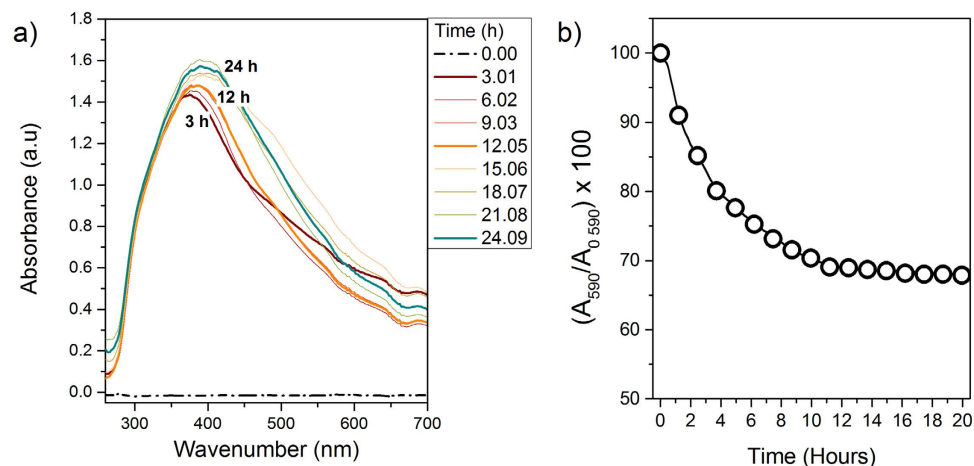
and anoxic conditions (Fig. 2). In oxic conditions (Fig. 2a), the amount of  $H_2O_2$  increased whereas  $O_2$  was rapidly consumed at the beginning of the experiment. Under anoxic conditions  $O_2$  and  $H_2O_2$  formed concomitantly (Fig. 2b).

During the experiments, pH became acidic under both oxic and anoxic conditions. Overall, pH dropped rapidly towards a nearly constant value approximately 2 to 3 pH units lower than at the initial pH value. The initial drop was more pronounced under anoxic than under oxic conditions (2 hours vs 10 hours, respectively), and in both cases the initial decline in pH values was accelerated with increasing pyrite loading (Supplementary Fig. S2).

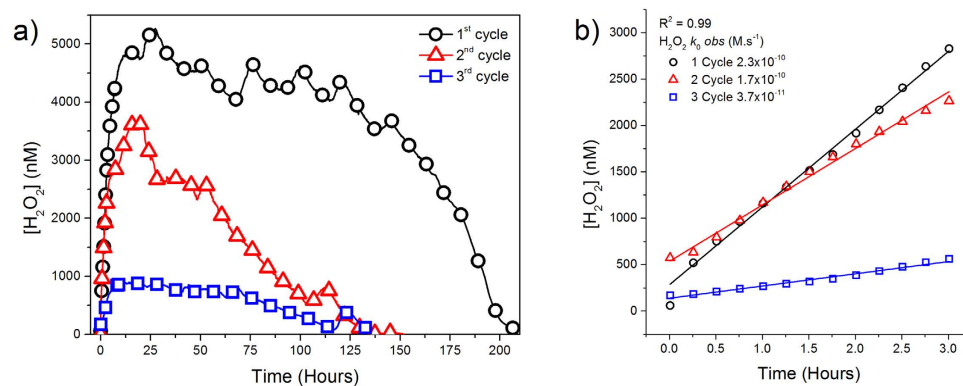
**The Fenton reaction in anoxic conditions: spectroscopic experiments.** To determine if the Fenton reaction was actually the mechanism to degrade the  $H_2O_2$  forming under anoxic conditions and in the dark (Fig. 3). First, we evaluated the intensity of the absorbance bands in the UV-Vis range of 260–700 nm to identify the formation of dissolved  $Fe^{3+}$ -complexes that coexist in solution in the range 300–450 nm<sup>40–43</sup> (Fig. 3a). The first absorbance peak ( $\lambda_{max}$  ~375 nm) evolved after three hours of reaction and was shifted to higher wavelengths (around 390–400 nm), as the reaction proceeded. Second, we monitored the formation of short-lived radicals, mainly  $OH^\cdot$  by measuring the decrease in the light absorption spectrum ( $\lambda_{max}$  = 590 nm) of Crystal Violet used as a dye probe<sup>44</sup> (Fig. 3b). The absorption spectrum of CV showed a rapid decline indicating that  $OH^\cdot$  formation occurs concomitantly from the start-up of the experiment.

**Tracking the reversibility of  $H_2O_2$  generation: cycling experiments.** We performed these experiments to better understand the long-term evolution of  $H_2O_2$  during pyrite dissolution in an open system. In order to use the same physical pyrite particles and in the same geometry, pyrite microparticles were adhered over silicone strips to form a thin film which was placed on the internal wall of the reactor as in the spectroscopic experiments. We monitored the evolution of  $H_2O_2$  until the observable amount of  $H_2O_2$  attained a constant value or a concentration of zero (Fig. 4). At that point, we replaced the solution inside the batch reactor with distilled  $H_2O$ , and monitored the evolution of  $H_2O_2$  again (we repeated this procedure twice). After a full cycle of  $H_2O_2$  generation and decay, the formation of  $H_2O_2$  was found to resume after the addition of fresh distilled  $H_2O$ , although the maximum  $H_2O_2$  yield was consistently lower than in the previous cycle (Fig. 4a). Assuming a zero-order kinetic model, the initial observed rate of the  $H_2O_2$  formation ( $k_{0,obs}$ ) showed a slight decrease between the first and the second cycle, which was more prevalent in the third cycle (Fig. 4b). This decrease could be attributed to the formation of oxide patches on pyrite surface that partially block some of the iron reactive centers.

**Surface characterization (CV, XPS and HRTEM).** Results of cyclic voltammetry under anoxic conditions revealed the anodic peaks associated with  $H_2O$  oxidation by one electron and by two electron transfer (Supplementary Fig. S3). Interestingly, in the cathodic counterpart, the peak assigned to iron reduction is split into two peaks (0.1 V NHE and 0.2V NHE, respectively), indicating that a fraction of  $\equiv Fe^{3+}$  is reduced in a nearly spontaneous manner. The analyzed XPS spectra of the (001) face of pyrite after aqueous reaction in the presence and absence of dissolved  $O_2$  (g) resulted in significant differences, showing major surface oxidation under oxic conditions with the subsequent formation of patches (Supplementary Fig. S4). In fact, XPS analysis of the (100) face of the pyrite after anoxic reaction only showed appreciable changes in the O1s orbital showing a shift to



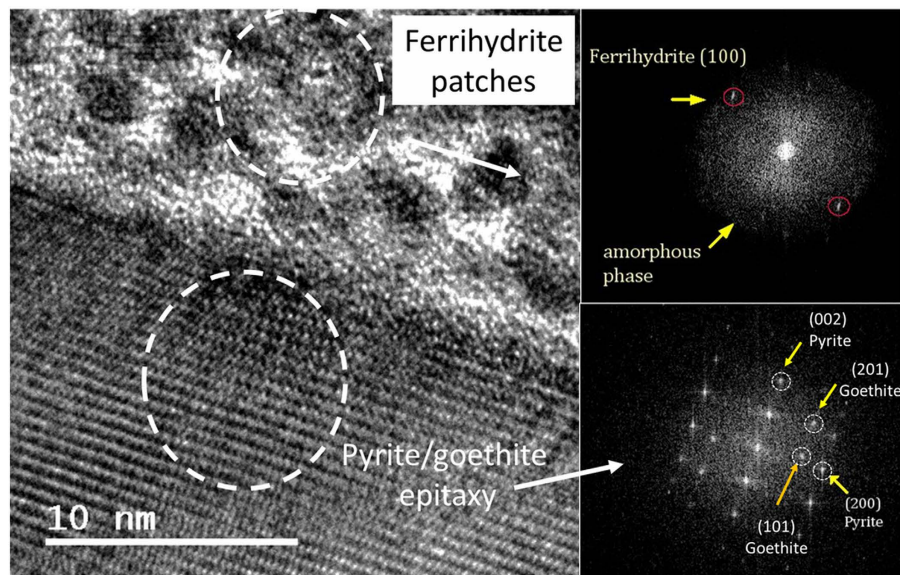
**Figure 3. Spectroscopic monitoring of  $\text{Fe}^{3+}$  and  $\text{OH}^-$  from pyrite slurries under anoxic conditions.** (a) Absorption bands detected from aqueous pyrite suspension in anoxic conditions and in the dark showing  $\text{Fe}^{3+}$ -complexes signatures (particle loading = 0.28 g/L). The numbers inserted over the absorption bands show the reaction time. Spectra were registered in real time using a liquid waveguide capillary flow cell (LWCC; path length: 250 cm; WPI), connected to the batch reactor by a peristaltic pump; (b) Degradation of CV solution upon pyrite aqueous reaction under anoxic conditions and in the dark (particle loading = 0.12 g/L,  $[\text{CV}]_0 = 225.5 \mu\text{M}$ ).



**Figure 4. Recurrence of  $\text{H}_2\text{O}_2$  formation by pyrite slurry after  $\text{H}_2\text{O}$  renewal.** (a)  $\text{H}_2\text{O}_2$  evolution from the same pyrite slurry after renewing  $\text{H}_2\text{O}$  twice, in oxic-open conditions (pyrite load particle = 0.33 g/L,  $\Delta\text{pH}_{1,\text{cycle}} = 6.8\text{--}3.8$ ,  $\Delta\text{pH}_{2,\text{cycle}} = 6.8\text{--}6.1$ ,  $\Delta\text{pH}_{3,\text{cycle}} = 6.8\text{--}6.7$ ); (b) Initial observed rate of  $\text{H}_2\text{O}_2$  formation assuming a zero kinetic order rate.

lower binding energies associated with an increase of the hydroxyl contribution ( $-\text{OH}$ ) (Supplementary Fig. S5). In order to facilitate the identification of  $\equiv\text{S}^{2-}$  and  $\equiv\text{Fe}^{3+}$  dangling bonds, one sample was ion-sputtered, which promoted the breakage of the S-S dimers (as in the grinding procedure) (Supplementary Fig. S6). The formation of iron oxidation patches was also observed with HRTEM. Figure 5 shows an image of the pyrite surface after 22 hours of reaction in a micromolar solution of  $\text{H}_2\text{O}_2$  in absence of  $\text{O}_2(\text{g})$ . The presence of discrete oxidation patches was observed in the uppermost area of the micrograph (Fig. 5a). The FFT (Fast Fourier transform) proved that the lattice fringe spacing of low contrast clusters ( $\sim 0.25 \text{ nm}$ ) were consistent with ferrihydrite nanocrystals (Fig. 5b), viewed down  $[001]$ . Additional HRTEM images (Supplementary Fig. S7) also showed the interplanar spacing characteristic of two-line ferrihydrite<sup>45,46</sup>. The FFT of the crystalline part showed the interplanar spacing of pyrite but also of goethite, suggesting that ferrihydrite can be a precursor of goethite formation (Fig. 5c)<sup>47</sup>.

**Fitting the experimental data: kinetic modeling of  $\text{H}_2\text{O}_2$  evolution.** Our experimental data suggests that the observed trend of  $\text{H}_2\text{O}_2$  is the result of the coupling between the  $\text{H}_2\text{O}_2$  generation by iron defect sites on the pyrite surface and the  $\text{H}_2\text{O}_2$  decomposition by the Fenton reaction. Based on these results, we built a kinetic model that allowed us to analyze the process in terms of elementary reactions and to determine the specific rate constant of the  $\text{H}_2\text{O}_2$  formation in oxic and anoxic conditions by fitting the experimental data. In summary, the model describes the net amount of  $\text{H}_2\text{O}_2$  according with the following expression (Supplementary, Modeling approaches):



**Figure 5. Identification of iron oxide patches in pyrite surface by HTREM.** (a) HTREM image showing the formation of secondary products over a lamella of pyrite after 22 hours immersed in a micromolar solution of  $\text{H}_2\text{O}_2$  under anoxic conditions; (b) FFT of ferrihydrite patches; (c) FFT of the crystalline part showing spacing characteristic of pyrite and goethite.

Rate constants	Oxic conditions	Anoxic conditions
$k_{\text{oxic}}$ (mol/m <sup>2</sup> s)	$10^{-3.58}$	$10^{-3.70}$
$k_{\text{anoxic}}$ (mol/m <sup>2</sup> s)	$10^{-8.25}$	$10^{-8.14}$
$^*k_{\text{pyr}}$	by $\text{O}_2$ $10^{-7.86}$ (mol/ m <sup>2</sup> s)	by $\text{H}_2\text{O}_2$ $10^{-6.03}$ (L/ m <sup>2</sup> s)

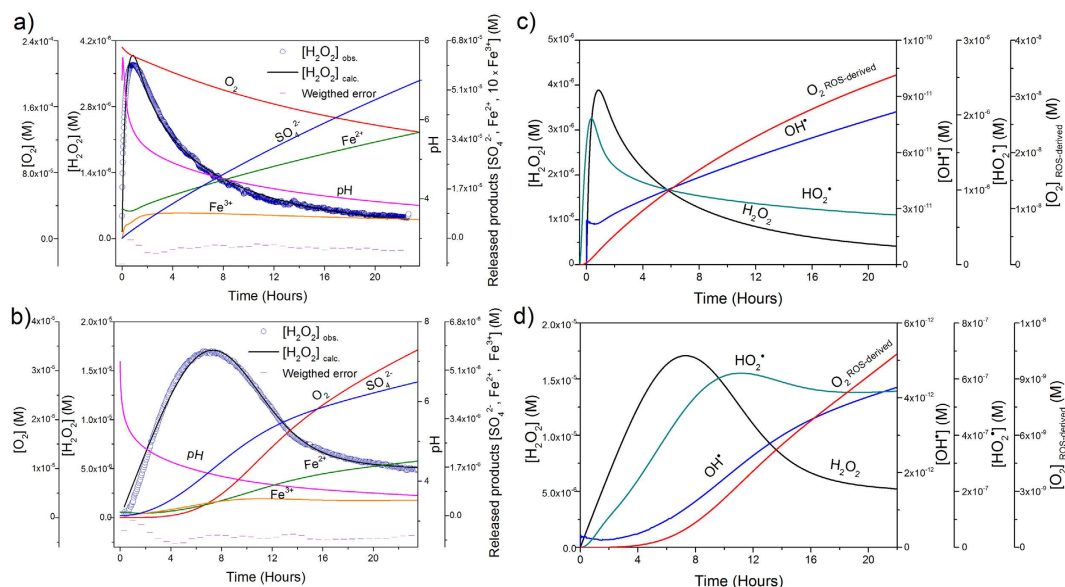
**Table 1. Rate constants obtained by fitting experimental data.** <sup>\*</sup>The rate of pyrite oxidation by  $\text{H}_2\text{O}_2$  was used as an additional adjustable parameter based on the expression given by McKibben and Barnes (1987) whereas the oxidation rate by  $\text{O}_2$  was used as a fixed parameter (Liu *et al.* 2008).

$$\frac{d[\text{H}_2\text{O}_2]}{dt} = (k_{\text{oxic}} \times \theta \times S_{\text{pyFe}^{2+}}(t) + k_{\text{anoxic}} \times S_{\text{pyFe}^{3+}}(t)) - (k_1[\text{Fe}^{2+}] + k_2[\text{Fe}^{3+}] + k_3[\text{OH}^-]) \times [\text{H}_2\text{O}_2] \quad (9)$$

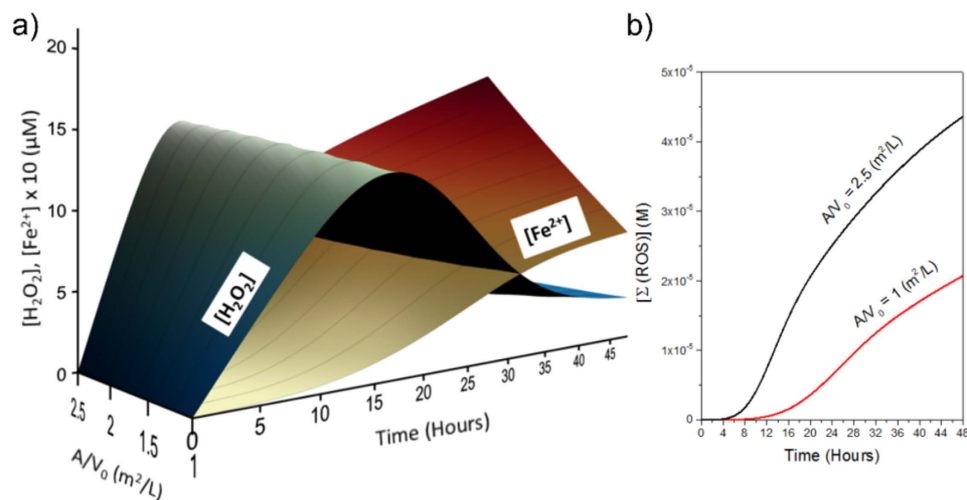
Where  $k_1$ ,  $k_2$  and  $k_3$ , represent the rate constants of each Fenton reaction step (Supplementary, Table S1). As shown in the equation, the  $\text{H}_2\text{O}_2$  formation was calculated assuming a first order dependence on the reactive surface and, both iron defect sites act simultaneously, but in anoxic conditions the  $\text{H}_2\text{O}_2$  produced by  $\text{Fe}^{2+}$  sites is limited by the  $\text{O}_2$  derived by the Fenton reaction. Minimization between experimental curves and the values calculated from equation 9 were made with a non-linear least squares approach using the Marquardt algorithm, with  $k_{\text{oxic}}$  and  $k_{\text{anoxic}}$ , the specific rate constants of  $\text{H}_2\text{O}_2$  surface generation, as adjustable parameters (Table 1).

Figure 6 presents the fitted values corresponding to the evolution of  $\text{H}_2\text{O}_2$  in oxic and anoxic conditions together with the model derived results. Under oxic conditions (Fig. 6a), when pH was higher, the ferrous surface iron was oxidized by dissolved  $\text{O}_2$  forming  $\text{H}_2\text{O}_2$  according the first order rate. Deviations of the linearity started to occur due to a change in the reaction stoichiometry by the simultaneous decrease in pH and increase in dissolved  $\text{Fe}^{2+}$  and  $\text{H}_2\text{O}_2$ . From this point, the Fenton reaction became important and the oxidation rate of  $\text{Fe}^{2+}$  by  $\text{O}_2$  slowed down leading to a minor  $\text{O}_2$  consumption. As a result, reactions forming ROS -catalyzed by  $\text{Fe}^{2+}$  and  $\text{Fe}^{3+}$ — became effective for  $\text{H}_2\text{O}_2$  degradation, which rapidly decreased, while iron species followed an opposite trend. Under anoxic conditions (Fig. 6b), the formation of  $\text{H}_2\text{O}_2$  proceeded more slowly. The decomposition of  $\text{H}_2\text{O}_2$  was also retarded because the concentration of dissolved  $[\text{Fe}^{2+}]$  supplied to solution by pyrite dissolution was lower. The Fenton reaction was initiated when pH values dropped below 4.5 and the ratio  $\text{H}_2\text{O}_2/\text{Fe}^{2+}$  increased, catalyzing the  $\text{H}_2\text{O}_2$  decomposition similar to the oxic experiments. The analysis of ROS derived from the model in both oxic and anoxic conditions (Fig. 6c and d) shows that, the first reactive species formed was  $\text{OH}^\cdot$  acting as a chain initiator, forming additional free radicals. The majority of the  $\text{OH}^\cdot$  reacted with  $\text{H}_2\text{O}_2$ , generating  $\text{HO}_2^\cdot$ ; the conjugate acid of  $\text{O}_2^{\cdot-}$ , which is the limiting reagent to assist the redox cycling of  $\text{Fe}^{3+}/\text{Fe}^{2+}$  forming  $\text{O}_2$  and helping to buffer the pH drop

Contrary to the expectation, the maximum amount of  $\text{H}_2\text{O}_2$  measured was significantly different between experimental runs, and independent of particle loading. This was likely because the effect of increasing the reactive surface was twofold: (i) an increase in the formation of  $\text{H}_2\text{O}_2$ , and (ii) an increase in the release of dissolved  $\text{Fe}^{2+}$ , accelerating the decomposition of  $\text{H}_2\text{O}_2$  by the Fenton reaction (Fig. 7). We plotted  $\text{H}_2\text{O}_2$  and  $\text{Fe}^{2+}$  evolution



**Figure 6. Experimental and fitting curves of  $\text{H}_2\text{O}_2$  evolution, together with model-derived trends of pyrite dissolution products and secondary ROS.** (a) experimental and model curves of  $\text{H}_2\text{O}_2$  together with  $\text{Fe}^{2+}/\text{Fe}^{3+}$ ,  $\text{SO}_4^{2-}$ , pH and  $\text{O}_2$  model trends, under oxidic conditions (pyrite particle loading = 0.71 g/L,  $A/V_0 = 1 \text{ m}^2/\text{L}$ ,  $\text{pH}_0 = 7$ ,  $[\text{O}_2]_0 = 232 \mu\text{M}$ ); and, (b) anoxic conditions (pyrite particle loading = 2.16 g/L,  $A/V_0 = 3.16 \text{ m}^2/\text{L}$ ,  $\text{pH}_0 = 7$ ,  $[\text{O}_2]_0 = 0 \mu\text{M}$ ). Time evolution of ROS obtained with the previous models under: (c) oxidic conditions; and (d) anoxic conditions. The evolution of  $\text{O}_2^-$  was omitted from the plot because model calculations gave negligible concentrations.



**Figure 7. Influence of pyrite reactive surface area on  $\text{H}_2\text{O}_2$  evolution.** (a) 3D plot showing the  $\text{H}_2\text{O}_2$  and  $\text{Fe}^{2+}$  simulated trends at different values of the surface area of pyrite per volume of  $\text{H}_2\text{O}$  ( $A_0/V$  from 1 to 2.5,  $\text{m}^2/\text{L}$ ). (b) Summation of ROS generated by the Fenton reaction for the lowest and the highest values of the reactive surface (Initial conditions:  $k_{\text{oxic}} = 2.33 \times 10^{-4}$ ,  $k_{\text{anoxic}} = 6.46 \times 10^{-9} \text{ mol}/\text{m}^2\text{s}$  – average values of the rate constants estimated by the model-, pyrite particle loading = 1 g/L,  $\text{pH}_0 = 7$ ,  $[\text{O}_2]_0 = 0 \mu\text{M}$ ).

as predicted by the model for a set of different values of pyrite reactive surface area (in Fig. 7a). As a result of this coupling effect, higher values of reactive surface area tend to increase the amount of secondary ROS (i.e.,  $\text{OH}^-/\text{O}_2^-$ ,  $\text{HO}_2^-/\text{O}_2^-$ ,  $\text{O}_2$ ) rather than stabilize the presence of  $\text{H}_2\text{O}_2$  in solution (in Fig. 7b)<sup>2,48</sup>.

## Discussion

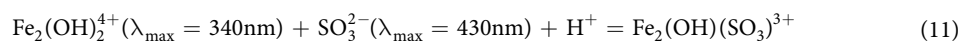
Aqueous suspensions of pyrite form  $\text{H}_2\text{O}_2$ , in the presence and absence of dissolved  $\text{O}_2$ , following a generation-decay trend (Fig. 1). The asymmetric shape observed in the experimental curves indicated that the apparent rate of  $\text{H}_2\text{O}_2$  generation was significantly faster than the apparent  $\text{H}_2\text{O}_2$  degradation. Previous studies<sup>22,25</sup>

have suggested that in the presence of dissolved  $O_2$ , pyrite slurries form  $H_2O_2$  by an electron transfer between the ferrous iron defect sites ( $Fe^{2+}-S$ ) and the adsorbed  $O_2$  through a Haber-Weiss reaction, involving the  $O_2^{\cdot-}$  radical formation (equations 1 and 2). This hypothesis is consistent with the observation that there is an inverse relationship between  $H_2O_2$  and  $O_2$  at the beginning of the oxic experiment (Fig. 2a). The mechanism of  $H_2O_2$  formation in anoxic conditions remains more controversial. Some studies have suggested that in absence of  $O_2$  the formation of  $H_2O_2$  is driven by the oxidation of adsorbed  $H_2O$  catalyzed by the pyrite surface<sup>4,19,24,26</sup>, whereas other studies have considered that this reaction is unlikely due to energetic considerations<sup>15,25</sup>. A possible reconciliation comes from considering the presence of  $\equiv Fe^{3+}$  dangling bonds generated from the cleavage of S-S bonds. Briefly, the rupture of S-S bonds generate  $S^-$  species which are highly instable and, to compensate its charge disequilibrium, donate one electron to the nearest iron atom by the autoredox reaction<sup>26,31,32,49–54</sup>:



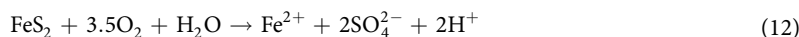
These dangling bonds could decrease the energy requirements for the chemisorption of  $H_2O$  molecules, some authors even talk about “ferry” iron dangling bonds  $\equiv Fe^{4+55}$ , leading to  $H_2O$  splitting into  $H^+$  and  $OH^-$ , with the subsequent formation of  $H_2O_2$ , as described by<sup>19</sup> equations 3 and 4. Our experiments showed that  $H_2O_2$  was generated in the absence of both dissolved  $O_2$  and light (Fig. 1b). Besides, the estimated ratio of S/Fe ( $<2$ ) in unreacted pyrite particles (Supplementary Fig. S8) indicated the presence of S vacancies.

Experiments with pyrite slurries under both oxic and anoxic conditions showed a sudden drop of pH during the first hours (Supplementary Fig. S2). This increase of  $H^+$  concentration together with the progressive accumulation of  $H_2O_2$  and  $Fe^{2+}$  in solution is expected to trigger the Fenton reaction. UV-Vis spectroscopy confirmed the presence of  $Fe^{3+}$ -complexes during anoxic pyrite dissolution (Fig. 3a) whose maximum wavelengths were compatible with the following reaction<sup>41</sup>:



The detection of both  $Fe^{3+}$ -complexes and  $OH^-$  in solution (Fig. 3) indicated that the Fenton reaction occurred even in the absence of dissolved  $O_2$ , supporting the idea that the suite of Fenton reactions conditioned the decay period of  $H_2O_2$  curves. Additionally, the formation of  $OH^-$  concomitantly with the formation of  $H_2O_2$  from the start-up of the experiment under anoxic conditions (Fig. 3b) together with the splitting of the cathodic peak associated with the nearly spontaneous iron reduction (Supplementary Fig. S3), suggest that  $\equiv Fe^{3+}$  dangling bonds, actually catalyze the  $H_2O$  oxidation by one single electron, forming  $OH^-$  as described by equation 3.

In principle, the overall process of aqueous pyrite oxidation -under oxic conditions- involves only  $O_{2(g)}$  consumption, according to:



However, considering a free radical mechanism, a simultaneous uptake and release of dissolved  $O_2$  it to be expected, because it can be both a product and a reactant. In oxic experiments,  $O_2$  concentration rapidly dropped and then reached a more constant value (Fig. 2a). Since the amount of  $Fe^{2+}$  in solution is presumably low at these early stages, the consumption of  $O_2$  during this period could be attributed to  $Fe^{2+}$  oxidation at surface sites to form the superoxide anion ( $O_2^{\cdot-}$ ) and to the subsequent production of  $H_2O_2$  (equations 1 and 2). An interesting result was the formation of  $O_{2(aq)}$  as a by-product in anoxic experiments (Fig. 2b and Supplementary Fig. S1b). Although  $O_2$  is not a direct product in the equation 3, since  $H_2O_2$  is formed, the formation of  $O_2$  can occur via several pathways such as the Fenton-like reaction, the “catalase-like reaction”<sup>25</sup> or via non-radical disproportionation of  $H_2O_2$ .

XPS and HRTEM results reported herein showed the formation of  $Fe^{3+}-O$  patches over pyrite surfaces. The implication of these  $Fe^{3+}$ -patches during aqueous pyrite oxidation is not clear. Some studies argued that the electron cycling of  $Fe^{2+}$  and  $Fe^{3+}$  between unoxidized and oxidized areas favors the electron transfer from the surface of the pyrite to molecular  $O_2$ , increasing the oxidation rate of pyrite<sup>36</sup>. However, this mechanism is based on atmospheric oxidation of pyrite and does not explain the formation of  $H_2O_2$ . In contrast, other studies have suggested that the formation of these  $Fe^{3+}$ -patches interrupts the redox cycling of  $Fe^{2+}/Fe^{3+}$ , thereby inhibiting the formation of  $H_2O_2$  and decreasing the rate of pyrite dissolution<sup>25</sup>. Our cycling experiments showed that most of the defect sites of the pyrite microparticles remained active at the end of the experiment, even when  $H_2O_2$  was no longer observable in solution (Fig. 4). Moreover, it is expected that oxidized patches will desorb at low pH during the pyrite dissolution process. HRTEM images showed an appreciable increment of the pyrite alteration layer (Supplementary Fig. S9), suggesting that these  $Fe^{3+}$ -patches failed to completely block the surface renewal during pyrite dissolution.

Based on our experimental results, we propose a model that explains the generation-decay trend of  $H_2O_2$  in terms of a kinetic competition between (1) the formation of  $H_2O_2$  by the self-oxidation of iron-sulfur cluster defect sites; and, (2) the degradation of  $H_2O_2$  by the Fenton reaction triggered by pyrite dissolution. Accordingly, the evolution of  $H_2O_2$  in solution can be summarized as follows:

$$\frac{d[H_2O_2]}{dt} = \sum \frac{d[H_2O_2]_{surf}}{dt} \pm \sum \frac{d[H_2O_2]_{Fenton}}{dt} \quad (13)$$

At the beginning of the model run ( $t \approx 0$ ), the amount of  $H_2O_2$  in solution was controlled by the surface reaction and the second term tended toward zero. However, when the Fenton reaction started,  $H_2O_2$  was progressively degraded to secondary ROS, in solution (i.e.,  $OH^-$ ,  $HO_2^-/O_2^{\cdot-}$ ,  $O_2$ ). At the end of the model run the amount of  $H_2O_2$  remained constant within a value close to zero and equation 13, became:

$$\sum \frac{d[\text{H}_2\text{O}_2]_{\text{surf}}}{dt} = \sum \frac{d[\text{H}_2\text{O}_2]_{\text{Fenton}}}{dt} \quad (14)$$

The model allowed us to estimate the rate constants of  $\text{H}_2\text{O}_2$  formation under both oxic and anoxic conditions (Table 1). Peak production of  $\text{H}_2\text{O}_2$  was shifted towards longer times in anoxic conditions (Supplementary Fig. S10), pointing to slower oxidation kinetics of pyrite in the absence of  $\text{O}_2$ , a result also supported by our XPS analysis (Supplementary Figs S4 and 5) and by the sulfate and iron released by pyrite dissolution (Supplementary Fig. S11). When compared with experimental data, the model reproduces qualitatively well the observed trends for pH,  $\text{O}_2$ ,  $\text{Fe}_{\text{total}}$  and  $\text{SO}_4^{2-}$ , as these parameters are calculated using the overall equations of pyrite dissolution<sup>17,57,58</sup> (Fig. 6a–b). An additional feature of the model is that provides a way to reconcile the classical dissolution approach with the free radical assumption, opening a new pathway to analyze the flux of secondary ROS resulting from the degradation of  $\text{H}_2\text{O}_2$  through the Fenton reaction (Fig. 6c–d).  $\text{OH}^\cdot$  radicals were the first species produced, rapidly decaying to  $\text{HO}_2^\cdot/\text{O}_2^-$ . These ROS species counterbalances the decrease in pH and promoted the so-called “Fenton like” reaction, which resulted in the formation of  $\text{O}_2(\text{g})$ .  $\text{Fe}^{2+}$  regenerated through this sequence, and also through the dissolution of pyrite, makes the Fenton reaction more efficient<sup>59</sup>. Therefore, the disappearance of  $\text{H}_2\text{O}_2$  in solution was likely due to a fast transformation into ROS, catalyzed by dissolved  $\text{Fe}^{2+}$ , rather than by the cessation of the generation mechanism itself. This result is consistent with the continuous production of  $\text{H}_2\text{O}_2$  and  $\text{OH}^\cdot$ <sup>23,26,60</sup> during pyrite dissolution. We hypothesize that as a result of the progressive acidification of the solution together with iron recycling by the Fenton reaction, the abiotic dissolution of pyrite microparticles can be considered as a natural and auto-catalytic Fenton reagent, useful to understanding long-term oxidation processes even in oxygen limited environments.

## Methods

**Sample preparation and characterization.** Natural pyrite samples (Logroño, Spain) were ground to obtain particles with average diameter of 1.4  $\mu\text{m}$ , Laser Diffraction Particle Size Analyzer (LS13320) and BET (Brunauer – Emmett – Teller) surface area of 1.46  $\text{m}^2/\text{g}$  (multi-point  $\text{N}_2$  adsorption). Prior to use, freshly ground pyrite microparticles were washed by sonication in ethanol (96%) and hydrochloric acid (HCl 0.25 M), to remove organics and oxide surface coatings. Samples were then rinsed with deoxygenated deionized water (MilliQ) and dried in a low vacuum chamber purged with nitrogen ( $\text{N}_2$ ) until used. Minor and trace elements in the acid washed samples were evaluated by scanning electron microscopy (SEM) using X-ray mapping (XRM) (Supplementary Fig. S8). X-ray diffraction (XRD) was used to assess the presence of secondary phases and to characterize the degree of structural disorder in the samples using the Rietveld method (Supplementary Fig. S12).

**Batch kinetic experiments.** The kinetics of  $\text{H}_2\text{O}_2$  formation and degradation on pyrite slurries was investigated in batch reactors utilizing amperometric sensors for  $\text{H}_2\text{O}_2$  (ISO-HPO-100, World Precision Instruments, Inc.) and dissolved  $\text{O}_2$  (Unisense DK), and a glass electrode for pH determination (Vernier FPH-BTA). Batch reactors were designed to fit with microsensors, spectroscopic probes, ports for pyrite inlet under  $\text{N}_2$  atmosphere, and valves. The valves allowed fluid circulation and solution sampling in a closed system configuration.

The production of ferric iron ( $\text{Fe}^{3+}$ ) and hydroxyl radical ( $\text{OH}^\cdot$ ) species by pyrite slurries in absence of dissolved  $\text{O}_2$  were measured using UV-vis spectroscopy by monitoring absorption bands at specific wavelengths. To prevent data masking due to particle absorption, the pyrite microparticles were deposited onto silicone strips as a thin film adhered to the inner reactor walls. Spectroscopic data were collected with a fiber optic UV-Vis spectrometer (Black-comet, Stellarnet or USB4000, Ocean Optics) and acquired with the SpectraWiz<sup>®</sup> or loggerpro3 codes. In addition, total dissolved iron and sulfate released during pyrite dissolution were measured at different time intervals using inductively coupled plasma mass spectrometry (ICP-MS) and ion chromatography (IC).

All experiments were made under continuous magnetic stirring (500 rpm), at room temperature ( $T^\circ \sim 22^\circ\text{C}$ ) and in the dark, unless other conditions are specified. Due to the autocatalytic nature of Fenton chain reactions, kinetic experiments were carried out in unbuffered distilled water. Further details of the batch reactors, electrochemical sensors, experimental procedures and test analysis are given as supplementary material (Supplementary Figs S13–18).

**Surface analysis.** Cyclic Voltammetry (CV) was employed as an additional way to assess the sequence of redox reactions involving free radicals during  $\text{H}_2\text{O}$  adsorption at pyrite interface. CV was performed using Pt/Pyrite-Np's/ Nafion<sup>®</sup>/ electrodes in  $\text{N}_2$  purged solutions vs Ag/Cl 3 M KCl. A detailed description of the experimental set-up is given as Supplementary Material. X-ray photoelectron spectroscopy (XPS) was used to analyze the surface oxidation states of (001) faces of pyrite (single-crystal,  $\sim 1 \text{ cm}^2 \times 2 \text{ mm}$ ) after aqueous reaction in oxic and anoxic conditions. Platelets parallel to (001) faces were cut ( $1 \text{ cm}^2 \times 2 \text{ mm}$ ) and cleaned following the same procedure described above and allowed to react with water in oxic and anoxic conditions. Samples were dried and stored in  $\text{N}_2$ -filled tubes until introduction into the XPS vacuum chamber. The oxidation states were analyzed at three different stages of the dissolution process: ( $t_1$ ) unreacted sample; the pyrite crystal was acid-washed to generate oxide-free surfaces by removing the normal contaminants, carbon, nitrogen, and oxygen (C, N, O) due to atmospheric oxidation; ( $t_2$ ): sample after 22 hours immersed in oxic water. ( $t_3$ ): sample after 22 hours immersed in anoxic water. Prior to this stepped analysis, the XPS spectra of a pyrite surface was analyzed after argon ion ( $\text{Ar}^+$ ) sputtering to verify the formation of  $\text{Fe}^{3+}$  and  $\text{S}^{2-}$  surface species as occurs during mechanical fracture by preferential sulfur removal. XPS Spectra were collected from a take-off angle of  $90^\circ$  relative to the sample surface in a Thermo Scientific K-Alpha ESCA analyzer using monochromatic Al K $\alpha$  (1486.6 eV) radiation and pass energies of 100 eV and 20 eV for survey spectra and narrow region spectra, respectively. Spectra were aligned by setting the C1s peak to a binding energy of 285 eV<sup>61</sup>. Deconvolution and fitting of experimental data were done



with the XPSpeak4.2 software (<http://www.phy.cuhk.edu.hk/~surface/XPSPEAK/>). The Shirley method was used for background subtraction and the binding energies of the species identified were assigned using values taken from literature (Supplementary Tables S2 and 3).

High resolution Electron Transmission Microscopy (HTREM) analysis were performed to identify nano-domains of secondary oxidation products at the pyrite interface. HTREM images were acquired on a JEOL JEM-3011 microscope with accelerating voltage of 200 kV using a Gatan Ultrascan 1000 CCD camera and Digital Micrograph software. Data processing was performed with the GADDS and image-J codes. Pyrite lamellas were prepared using a focused ion beam (FIB) with a high resolution JEOL JSM-6700 f.

**Kinetic model.** The model was run using the computer code Copasi 4.8 (COMplex PATHway SIMulator)<sup>62</sup>. We assume that the experimental trend of H<sub>2</sub>O<sub>2</sub>, in the presence and absence of O<sub>2</sub> (g), is shaped by three main processes: (i) the rate of H<sub>2</sub>O<sub>2</sub> generation by the iron defect sites on the surface of pyrite particles; (ii) the production rates of Fe<sup>2+</sup> and SO<sub>4</sub><sup>2-</sup> by pyrite dissolution; and, (iii) the kinetics of H<sub>2</sub>O<sub>2</sub> degradation by the Fenton reaction. Fits to the experimental curves that describe the evolution of H<sub>2</sub>O<sub>2</sub> were performed by using the Marquardt algorithm employing as adjustable parameters the rate constants of the H<sub>2</sub>O<sub>2</sub> formation. A detailed explanation of the modeling set-up is included in the supplementary material.

## References

- Puthussery, J., Seefeld, S., Berry, N., Gibbs, M. & Law, M. Colloidal Iron Pyrite (FeS<sub>2</sub>) Nanocrystal Inks for Thin-Film Photovoltaics. *Journal of the American Chemical Society* **133**, 716–719, doi: 10.1021/ja1096368 (2010).
- Bae, S., Kim, D. & Lee, W. Degradation of diclofenac by pyrite catalyzed Fenton oxidation. *Applied Catalysis B: Environmental* **134–135**, 93–102, doi: 10.1016/j.apcatb.2012.12.031 (2013).
- Gil-Lozano, C., Losa-Adams, E., Davila, F. A. & Gago-Duport, L. Pyrite nanoparticles as a Fenton-like reagent for *in situ* remediation of organic pollutants. *Beilstein journal of nanotechnology* **5**, 855–864, doi: 10.3762/bjnano.5.97 (2014).
- Wang, W., Qu, Y., Yang, B., Liu, X. & Su, W. Lactate oxidation in pyrite suspension: A Fenton-like process *in situ* generating H<sub>2</sub>O<sub>2</sub>. *Chemosphere* **86**, 376–382, doi: 10.1016/j.chemosphere.2011.10.026 (2012).
- Che, H. & Lee, W. Selective redox degradation of chlorinated aliphatic compounds by Fenton reaction in pyrite suspension. *Chemosphere* **82**, 1103–1108, doi: 10.1016/j.chemosphere.2010.12.002 (2011).
- Choi, K., Bae, S. & Lee, W. Degradation of pyrene in cetylpyridinium chloride-aided soil washing wastewater by pyrite Fenton reaction. *Chemical Engineering Journal* **249**, 34–41, doi: 10.1016/j.cej.2014.03.090 (2014).
- Choi, K., Bae, S. & Lee, W. Degradation of off-gas toluene in continuous pyrite Fenton system. *Journal of Hazardous Materials* **280**, 31–37, doi: 10.1016/j.jhazmat.2014.07.054 (2014).
- Chandra, A. P. & Gerson, A. R. The mechanisms of pyrite oxidation and leaching: A fundamental perspective. *Surface Science Reports* **65**, 293–315, doi: 10.1016/j.surfrep.2010.08.003 (2010).
- Descostes, M., Vitorge, P. & Beaucaire, C. Pyrite dissolution in acidic media. *Geochimica et Cosmochimica Acta* **68**, 4559–4569, doi: 10.1016/j.gca.2004.04.012 (2004).
- Druschel, G. & Borda, M. Comment on “Pyrite dissolution in acidic media” by M. Descostes, P. Vitorge, and C. Beaucaire. *Geochimica et Cosmochimica Acta* **70**, 5246–5250, doi: 10.1016/j.gca.2005.07.023 (2006).
- Heidel, C. & Tichomirowa, M. The isotopic composition of sulfate from anaerobic and low oxygen pyrite oxidation experiments with ferric iron — New insights into oxidation mechanisms. *Chemical Geology* **281**, 305–316, doi: 10.1016/j.chemgeo.2010.12.017 (2011).
- Lowson, R. T. Aqueous oxidation of pyrite by molecular oxygen. *Chemical Reviews* **82**, 461–497, doi: 10.1021/cr00051a001 (1982).
- Moses, C. O., Kirk Nordstrom, D., Herman, J. S. & Mills, A. L. Aqueous pyrite oxidation by dissolved oxygen and by ferric iron. *Geochimica et Cosmochimica Acta* **51**, 1561–1571, doi: 10.1016/0016-7037(87)90337-1 (1987).
- Rimstidt, J. D. & Vaughan, D. J. Pyrite oxidation: a state-of-the-art assessment of the reaction mechanism. *Geochimica et Cosmochimica Acta* **67**, 873–880, doi: 10.1016/s0016-7037(02)01165-1 (2003).
- Rosso, K. M., Becker, U. & Hochella Jr, M. F. The interaction of pyrite {100} surfaces with O<sub>2</sub> and H<sub>2</sub>O: Fundamental oxidation mechanisms. *American Mineralogist* **84**, 1549–1561, doi: 10.2138/am-1999-1008 (1999).
- Singer, P. C. & Stumm, W. Acidic mine drainage: The rate-determining step. *Science* **167**, 1121–1123, doi: 10.1126/science.167.3921.1121 (1970).
- Williamson, M. A. & Rimstidt, J. D. The kinetics and electrochemical rate-determining step of aqueous pyrite oxidation. *Geochimica et Cosmochimica Acta* **58**, 5443–5454, doi: 10.1016/0016-7037(94)90241-0 (1994).
- Borda, M. J., Elsetinow, A. R., Schoonen, M. A. & Strongin, D. R. Pyrite-Induced Hydrogen Peroxide Formation as a Driving Force in the Evolution of Photosynthetic Organisms on an Early Earth. *Astrobiology* **1**, 283–288, doi: 10.1089/15311070152757474 (2001).
- Borda, M. J., Elsetinow, A. R., Strongin, D. R. & Schoonen, M. A. A mechanism for the production of hydroxyl radical at surface defect sites on pyrite. *Geochimica et Cosmochimica Acta* **67**, 935–939, doi: 10.1016/s0016-7037(02)01222-x (2003).
- Cohn, C. A., Borda, M. J. & Schoonen, M. A. RNA decomposition by pyrite-induced radicals and possible role of lipids during the emergence of life. *Earth and Planetary Science Letters* **225**, 271–278, doi: 10.1016/j.epsl.2004.07.007 (2004).
- Cohn, C. A., Pak, A., Strongin, D. R. & Schoonen, M. A. Quantifying hydrogen peroxide in iron-containing solutions using leuco crystal violet. *Geochemical Transactions* **6**, 47, doi: 10.1186/1467-4866-6-47 (2005).
- Cohn, C. A. *et al.* Pyrite-induced hydroxyl radical formation and its effect on nucleic acids. *Geochemical Transactions* **7**, 11, doi: 10.1186/1467-4866-7-3 (2006).
- Cohn, C., Fisher, S., Brownawell, B. & Schoonen, M. Adenine oxidation by pyrite-generated hydroxyl radicals. *Geochemical Transactions* **11**, 8, doi: 10.1186/1467-4866-11-2 (2010).
- Javadi, N. A. & Hanumantha, R. K. Formation of hydrogen peroxide by sulphide minerals. *Hydrometallurgy* **141**, 82–88, doi: 10.1016/j.hydromet.2013.10.011 (2014).
- Schoonen, M. A. A., Harrington, A. D., Laffers, R. & Strongin, D. R. Role of hydrogen peroxide and hydroxyl radical in pyrite oxidation by molecular oxygen. *Geochimica et Cosmochimica Acta* **74**, 4971–4987, doi: 10.1016/j.gca.2010.05.028 (2010).
- Zhang, P., Yuan, S. & Liao, P. Mechanisms of hydroxyl radical production from abiotic oxidation of pyrite under acidic conditions. *Geochimica et Cosmochimica Acta* **172**, 444–457, doi: 10.1016/j.gca.2015.10.015 (2016).
- Burns, R. G. & Fisher, D. S. Iron-sulfur mineralogy of Mars: Magmatic evolution and chemical weathering products. *Journal of Geophysical Research: Solid Earth* **95**, 14415–14421, doi: 10.1029/JB095iB09p14415 (1990).
- Davila, A. F. *et al.* Subsurface formation of oxidants on Mars and implications for the preservation of organic biosignatures. *Earth and Planetary Science Letters* **272**, 456–463, doi: 10.1016/j.epsl.2008.05.015 (2008).
- Eggleston, C. M., Stern, J. R., Strellis, T. M. & Parkinson, B. A. A natural photoelectrochemical cell for water splitting: Implications for early Earth and Mars. *American Mineralogist* **97**, 1804–1807, doi: 10.2138/am.2012.4211 (2012).
- Zolotov, M. Y. & Shock, E. L. Formation of jarosite-bearing deposits through aqueous oxidation of pyrite at Meridiani Planum, Mars. *Geophys. Res. Lett.* **32**, L21203, doi: 10.1029/2005gl024253 (2005).

31. Nesbitt, H. W., Bancroft, G. M., Pratt, A. R. & Scaini, M. J. Sulfur and iron surface states on fractured pyrite surfaces. *American Mineralogist* **83**, 1067–1076, doi: 10.2138/am-1998-9-1015 (1998).
32. Schaufuß, A. G. *et al.* Reactivity of surface chemical states on fractured pyrite. *Surface Science* **411**, 321–328, doi: 10.1016/s0039-6028(98)00355-0 (1998).
33. Murphy, R. & Strongin, D. R. Surface reactivity of pyrite and related sulfides. *Surface Science Reports* **64**, 1–45, doi: 10.1016/j.surfrep.2008.09.002 (2009).
34. Krishnamoorthy, A., F. W. Herbert, S. Yip, K. J. Van Vliet, B. Yildiz, Herbert, F. W., Sidney, Y., Krystyn, J. V. V. & Bilge, Y. Electronic states of intrinsic surface and bulk vacancies in FeS<sub>2</sub>. *Journal of Physics: Condensed Matter* **25**, 045004, doi: 10.1088/0953-8984/25/4/045004 (2012).
35. Wang, J., Chen, D., Yan, D., Wei, H. & Xiang, L. Evolution from an anoxic to oxic deep ocean during the Ediacaran–Cambrian transition and implications for bioradiation. *Chemical Geology* **306–307**, 129–138, doi: 10.1016/j.chemgeo.2012.03.005 (2012).
36. Guevremont, J. M., Elseinow, A. R., Strongin, D. R., Bebie, J. & Schoonen, M. A. A. Structure sensitivity of pyrite oxidation; comparison of the (100) and (111) planes. *American Mineralogist* **83**, 1353–1356, doi: 10.2138/am-1998-11-1225 (1998).
37. Guevremont, J. M., Strongin, D. R. & Schoonen, M. A. A. *Thermal chemistry of H<sub>2</sub>S and H<sub>2</sub>O on the (100) plane of pyrite: unique reactivity of defect sites* Vol. 83 (Mineralogical Society of America, 1998).
38. de Leeuw, N. H., Parker, S. C., Sithole, H. M. & Ngoepe, P. E. Modeling the Surface Structure and Reactivity of Pyrite: Introducing a Potential Model for FeS<sub>2</sub>. *The Journal of Physical Chemistry B* **104**, 7969–7976, doi: 10.1021/jp0009498 (2000).
39. Fischer, C., Kurganskaya, I., Schäfer, T. & Lüttge, A. Variability of crystal surface reactivity: What do we know? *Applied Geochemistry* **43**, 132–157, doi: 10.1016/j.apgeochem.2014.02.002 (2014).
40. Feng, W. & Nansheng, D. Photochemistry of hydrolytic iron (III) species and photoinduced degradation of organic compounds. A minireview. *Chemosphere* **41**, 1137–1147, doi: 10.1016/s0045-6535(00)00024-2 (2000).
41. Lente, G. Reactions of the iron(III) hydroxo dimer with inorganic ligands, *University of Debrecen*, (2001).
42. Ma, J., Weng, D., Wu, X., Si, Z. & Wu, Z. Highly dispersed iron species created on alkali-treated zeolite for ammonia SCR. *Progress in Natural Science: Materials International* **23**, 493–500, doi: 10.1016/j.pnsc.2013.08.005 (2013).
43. Stefánsson, A. Iron(III) Hydrolysis and Solubility at 25 °C. *Environmental Science & Technology* **41**, 6117–6123, doi: 10.1021/es070174h (2007).
44. Fan, H.-J. *et al.* Degradation pathways of crystal violet by Fenton and Fenton-like systems: Condition optimization and intermediate separation and identification. *Journal of Hazardous Materials* **171**, 1032–1044, doi: 10.1016/j.jhazmat.2009.06.117 (2009).
45. Banfield, J. F., Welch, S. A., Zhang, H. & Ebert, T. T. & Penn, R. L. Aggregation-Based Crystal Growth and Microstructure Development in Natural Iron Oxhydroxide Biomineralization Products. *Science* **289**, 751–754, doi: 10.1126/science.289.5480.751 (2000).
46. Drits, V. A., Sakharov, B. A., Salyn, A. L. & Manceau, A. Structural model for ferrihydrite. *Clay Minerals* **28**, 185–207, doi: 10.1180/claymin.1993.028.2.02 (1993).
47. Das, S., Hendry, M. J. & Essilfie-Dughan, J. Transformation of Two-Line Ferrihydrite to Goethite and Hematite as a Function of pH and Temperature. *Environmental Science & Technology* **45**, 268–275, doi: 10.1021/es101903y (2011).
48. Che, H., Bae, S. & Lee, W. Degradation of trichloroethylene by Fenton reaction in pyrite suspension. *Journal of Hazardous Materials* **185**, 1355–1361, doi: 10.1016/j.jhazmat.2010.10.055 (2011).
49. Nesbitt, H. W. *et al.* Synchrotron XPS evidence for Fe<sup>2+</sup>–S and Fe<sup>3+</sup>–S surface species on pyrite fracture-surfaces, and their 3D electronic states. *American Mineralogist* **85**, 850–857, doi: 10.2138/am-2000-5-628 (2000).
50. Leiro, J. A., Mattila, S. S. & Laajalehto, K. XPS study of the sulphur 2p spectra of pyrite. *Surface Science* **547**, 157–161, doi: 10.1016/j.susc.2003.09.033 (2003).
51. Andersson, K. *et al.* Experimental and theoretical characterization of the structure of defects at the pyrite FeS<sub>2</sub> (100) surface. *Physical Review B* **70**, 195404, doi: 10.1103/PhysRevB.70.195404 (2004).
52. Zhang, Y. N., Hu, J., Law, M. & Wu, R. Q. Effect of surface stoichiometry on the band gap of the pyrite FeS<sub>2</sub> (100) surface. *Physical Review B* **85**, 085314, doi: 10.1103/PhysRevB.85.085314 (2012).
53. Herbert, F. W., Krishnamoorthy, A., Van Vliet, K. J. & Yildiz, B. Quantification of electronic band gap and surface states on FeS<sub>2</sub>(100). *Surface Science* **618**, 53–61, doi: 10.1016/j.susc.2013.08.014 (2013).
54. Sanchez-Arenillas, M. & Mateo-Marti, E. Pyrite surface environment drives molecular adsorption: cystine on pyrite(100) investigated by X-ray photoemission spectroscopy and low energy electron diffraction. *Physical Chemistry Chemical Physics* **18**, 27219–27225, doi: 10.1039/C6CP03760G (2016).
55. Stirling, A., Bernasconi, M. & Parrinello, M. Defective pyrite (100) surface: An ab initio study. *Physical Review B* **75**, 8, doi: 10.1103/PhysRevB.75.165406 (2007).
56. Eggleston, C. M., Ehrhardt, J.-J. & Stumm, W. Surface structural controls on pyrite oxidation kinetics; an XPS-UPS, STM, and modeling study. *American Mineralogist* **81**, 1036–1056, doi: 10.2138/am-1996-9-1002 (1996).
57. Liu, R., Wolfe, A., Dzombak, D., Stewart, B. & Caporaso, R. Comparison of dissolution under oxic acid drainage conditions for eight sedimentary and hydrothermal pyrite samples. *Environ Geol* **56**, 171–182, doi: 10.1007/s00254-007-1149-0 (2008).
58. McKibben, M. A. & Barnes, H. L. Oxidation of pyrite in low temperature acidic solutions: Rate laws and surface textures. *Geochimica et Cosmochimica Acta* **50**, 1509–1520, doi: 10.1016/0016-7037(86)90325-X (1986).
59. Duysterberg, C. K., Mylon, S. E. & Waite, T. D. pH Effects on Iron-Catalyzed Oxidation using Fenton's Reagent. *Environmental Science & Technology* **42**, 8522–8527, doi: 10.1021/es801720d (2008).
60. Fisher, S., Schoonen, M. & Brownawell, B. Phenylalanine as a hydroxyl radical-specific probe in pyrite slurries. *Geochemical Transactions* **13**, 1–18, doi: 10.1186/1467-4866-13-3 (2012).
61. Briggs, D. & Seah, M. P. *Practical Surface Analysis: Auger and X-ray photoelectron spectroscopy*. 2nd edn, Vol. 1 657 (John Wiley & Sons, 1990).
62. Hoops, S. *et al.* COPASI—a COmplex PAthway SIMulator. *Bioinformatics* **22**, 3067–3074, doi: 10.1093/bioinformatics/btl485 (2006).

## Acknowledgements

This research was funded by the Spanish Ministry of Science and Innovation in the framework of the MICINN-FEDER, Project CGL2011-30079, and from the Project “icyMARS”, funded by the European Research Council, Starting Grant no 307496. One of the authors (C.G.L.) also acknowledges the financial support by the Spanish Ministry of Science and Innovation under FPU Grant AP2008-03712. The authors would like to thank Dr. José Benito Rodríguez and Dr. Eugenio Solla for their inestimable help in the preparation of FIB samples and HREM analyses and Dra. Serra Rodríguez for the valuable help with XPS analysis.

## Author Contributions

All persons who meet authorship criteria are listed as authors, and all authors certify that they have participated sufficiently in the work to take public responsibility for the content, including participation in the concept, design, analysis, writing, or revision of the manuscript. Furthermore, each author certifies that this material or similar

material has not been and will not be submitted to or published in any other publication before its appearance in the *Scientific Reports journal*. Conception and design of study: Carolina Gil Lozano, Luis Gago Duport. Acquisition of data: Carolina Gil Lozano. Analysis and/or interpretation of data: Carolina Gil Lozano, Alfonso F. Davila, Elisabeth Losa Adams, Alberto G. Fairén, Luis Gago Duport. Drafting the manuscript: Carolina Gil Lozano, Luis Gago Duport. Revising the manuscript critically for important intellectual content: Carolina Gil Lozano, Alfonso Davila, Elisabeth Losa Adams, Alberto G. Fairén, Luis Gago Duport. Approval of the version of the manuscript to be published: Carolina Gil Lozano, Alfonso Davila, Elisabeth Losa Adams, Alberto G. Fairén, Luis Gago Duport.

### Additional Information

**Supplementary information** accompanies this paper at <http://www.nature.com/srep>

**Competing Interests:** The authors declare no competing financial interests.

**How to cite this article:** Gil-Lozano, C. *et al.* Quantifying Fenton reaction pathways driven by self-generated H<sub>2</sub>O<sub>2</sub> on pyrite surfaces. *Sci. Rep.* 7, 43703; doi: 10.1038/srep43703 (2017).

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