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RESEARCH ARTICLE

Predicting Effects of Ocean Acidification and Warming on Algae Lacking Carbon Concentrating Mechanisms

Janet E. Kübler*, Steven R. Dudgeon

Department of Biology, California State University, Northridge, California, United States of America

* janet.kubler@csun.edu

Abstract

Seaweeds that lack carbon-concentrating mechanisms are potentially inorganic carbonlimited under current air equilibrium conditions. To estimate effects of increased atmospheric carbon dioxide concentration and ocean acidification on photosynthetic rates, we modeled rates of photosynthesis in response to pCO₂, temperature, and their interaction under limiting and saturating photon flux densities. We synthesized the available data for photosynthetic responses of red seaweeds lacking carbon-concentrating mechanisms to light and temperature. The model was parameterized with published data and known carbonate system dynamics. The model predicts that direction and magnitude of response to pCO₂ and temperature, depend on photon flux density. At sub-saturating light intensities, photosynthetic rates are predicted to be low and respond positively to increasing pCO_2 , and negatively to increasing temperature. Consequently, pCO₂ and temperature are predicted to interact antagonistically to influence photosynthetic rates at low PFD. The model predicts that ρCO_2 will have a much larger effect than temperature at sub-saturating light intensities. However, photosynthetic rates under low light will not increase proportionately as pCO₂ in seawater continues to rise. In the range of light saturation (I_k) , both CO₂ and temperature have positive effects on photosynthetic rate and correspondingly strong predicted synergistic effects. At saturating light intensities, the response of photosynthetic rates to increasing pCO₂ approaches linearity, but the model also predicts increased importance of thermal over pCO₂ effects, with effects acting additively. Increasing boundary layer thickness decreased the effect of added pCO_2 and, for very thick boundary layers, overwhelmed the effect of temperature on photosynthetic rates. The maximum photosynthetic rates of strictly CO₂-using algae are low, so even large percentage increases in rates with climate change will not contribute much to changing primary production in the habitats where they commonly live.

Introduction

Continued absorption of anthropogenic emissions of CO_2 from burning fossil fuels, into seawater will inevitably lead to further declines in oceanic pH with predictable consequences for oceanic chemistry [1,2]. Changing ocean chemistry is expected to shift ratios of resources as well as environmental conditions, thereby favoring some groups of organisms at the expense of others [3–6]. Much attention has focused on calcifying organisms that are predicted to be vulnerable to ocean acidification as a result of lower saturation states of calcium carbonate species as reviewed in [7–9]. Non-calcifying phototrophs (e.g., macroalgae and seagrasses) are predicted to benefit in terms of growth, if not always in terms of photosynthetic rate, from ocean acidification (OA) due to the enhanced availability of dissolved CO_2 in the ocean.

In contrast to studies of OA on putatively vulnerable taxa, there has been less effort directed towards questions about the magnitudes of the effect on putative beneficiaries of OA. There are many important questions about these organisms in need of resolution to better predict the consequences of climate change to oceanic ecosystems, such as, "will, and if so, by how much will algal productivity be enhanced in acidified coastal water?", "will productivity be determined by changing pCO_2 or temperature?", or "how will the mode of inorganic carbon acquisition interact with other determinants of productivity as pCO_2 increases?".

Recent comprehensive meta-analyses [6,7,10] reviewing the effects of OA on marine organisms, highlight, the paucity of investigation relative to other groups, and interesting patterns of response of non-calcifying algae. Studies of fleshy macroalgae thus far, generally show significantly enhanced growth under OA, yet without significant enhancement of photosynthesis in response to moderately elevated pCO_2 . The effect of sampling few fleshy macroalgae coupled with inattention to the species' mechanisms of inorganic carbon uptake may mask variation present in the magnitude of effects OA has on macroalgal physiology, productivity and growth.

Whereas all green and brown macroalgae use both HCO_3^- and CO_2 sources of inorganic carbon for photosynthesis, not all red algae can take up HCO_3^- directly. Approximately 35% of all red algal species tested use only dissolved CO_2 for photosynthesis [4,11]. These species are characteristic of light-limited environments such as beneath overhanging rock walls in intertidal or subtidal habitats, especially in temperate and sub-arctic environments [12,13]. Non bicarbonate using red algae are found in very high abundance (> 60% of species) in low light environments around Tasmania, Australia [14]. Demand for dissolved inorganic carbon is less when PFD is low.

Photosynthesis of CO₂-using red macroalgae is potentially carbon-limited at present oceanic dissolved pCO_2 and is a hypothesized link to their occurrence in low light habitats [15,16]. Therefore, a straightforward prediction of the effects of ocean acidification is an increase in productivity and growth of strict CO₂-using algae. Greater inorganic carbon availability could lead to ranges extending into new, higher light intensity, habitats. A prior study of the CO₂using red alga, Lomentaria articulata, showed enhanced growth with increased pCO_2 [16] supporting that prediction, but there are few other well studied CO₂-using marine red algae. Critical gaps remain with respect to predicting the effects of climate change on macroalgal production. Despite extensive knowledge of the effects of temperature on macroalgal photosynthesis [17-20] the combined effects of increasing pCO₂ and temperature are less well known. Macroalgae with CCMs tend to have positive additive or synergistic effects of temperature and OA [21, 22]. A key question arising from projections of climate change is whether the known responses to temperature change and to increased inorganic carbon supply are additive or synergistic. Light supply also interacts with both temperature [20] and inorganic carbon uptake [23, 24]. We expected light supply to interact with climate change in algal productivity, possibly modulating the interaction between OA and temperature effects.

Our ability to predict the effects of climate change on algal productivity and growth are contingent upon our understanding of how changing environments affect algal physiological responses. In this context, we have developed a model specific to CO_2 -using red macroalgae. The model predicts effects of increased pCO_2 (OA) and temperature on photosynthetic production at different light intensities. Comparing theoretical predictions with observations, as more data becomes available, will provide a test of synergies in the combined effects of light intensity with increasing acidification and temperature in the oceans on productivity of CO_2 using algae. The model is parameterized using known temperature-dependent diffusivities of the different inorganic carbon species in seawater of given characteristics (salinity, pressure, viscosity, pH, total alkalinity), and photosynthetic responses of CO_2 -using marine red macroalgae to variations in light and temperature.

This model represents a starting point for choosing the combinations of multiple environmental factors and corresponding habitats that may be of special interest as ocean acidification continues. The parameterizing studies were primarily done as single stressor dose response studies rather than multi-stressor factorial manipulations and we are able to leverage areas of overlap. Our synthesis of the available data allows for the detection of habitats where physiological and ecological studies can be focused to maximize efficiency of further work. For example, if ocean acidification will alleviate other limitations to the range of the macrophytes in question, where can we predict their ranges to expand?

Materials and Methods

The Model

Our model of photosynthetic production uses the Hill-Whittingham (1955) equation, which is the appropriate model for aquatic systems where diffusive uptake of inorganic carbon can be limiting [25-27]. For red macroalgae that take up only CO₂, this is the case regardless of the presence of external carbonic anhydrase (CA) in the cell wall. CA only affects the rate of equilibration between CO₂ and HCO₃⁻ [28] and ultimately CO₂ crosses the plasmalemma by diffusion.

Photosynthetic rate is modeled as a hyperbolic function of CO₂ concentration and permeability in the following equation:

$$V_{o} = 0.5* \left\{ \left(P_{u} * K_{\frac{1}{2}} + P_{u} * S + V_{\max} \right) - \left[\left(P_{u} * K_{\frac{1}{2}} + P_{u} + V_{\max} \right)^{2} - \left(4 * P_{u} * S * V_{\max} \right) \right]^{\frac{1}{2}} \right\}$$
(1)

where;

 V_0 = Net photosynthetic rate in µmol C-fixed \cdot m⁻² \cdot s⁻¹

 P_u = permeability coefficient in m · s⁻¹; the ratio of the temperature-dependent diffusion of CO₂ and pathlength from outside the cell to the plastid

 $K_{1/2}$ = half-saturation constant of photosynthesis at specified conditions of light and temperature; in µmol · m⁻³

S = Substrate concentration in µmol · m⁻³; the dynamic flux expressed as the difference in dissolved CO₂ outside (C_{ext}) and inside (C_{int}) the cell

 V_{max} = Maximum net photosynthetic rate at specified conditions of light and temperature in µmol C-fixed · m⁻² · s⁻¹

Data obtained from the model reflect acclimation to steady state conditions of light, temperature and pH/carbon balance. We assume no adaptive evolution on the timescale of the model.

Model Parameterization

Parameters of this equation are influenced by light availability, temperature, hydrodynamics and/or nutrient concentration. Some relationships are the direct effect of water chemistry, for example, the positive effect of temperature on diffusion rate. Other parameters, such as the CO_2 concentration gradient between the seawater medium and the location of carbon fixation, depend on boundary layer characteristics and P_{max} under the prevailing conditions. Thus we have a basis to explore the potentially interactive effects of light, CO_2 availability and temperature on algal productivity, through the Hill-Whittingham equation.

The half-saturation constant ($K_{1/2}$) of photosynthesis of strict CO₂-using red algae is lightdependent [12]. We assumed a linear dependence of $K_{1/2}$ on light intensity because the functional relationship of the light-dependence of $K_{1/2}$ between sub-saturating and saturating light intensities has not been resolved.

It is well established that photosynthetic rates of macroalgae are affected by temperature and the availabilities of light and dissolved inorganic carbon (DIC) [18,23,29,30]. We used the averages of photosynthetic rate parameters (per unit thallus area) at both sub-saturating and saturating light intensities and temperatures ranging from 5 to 30°C from published studies of known CO₂-using red macroalgal species (e.g., genera including *Lomentaria*, *Delesseria*, *Plocamium*) to parameterize photosynthetic performance (i.e., V_{max}) of the model [11,12,18], [see Supporting Information, <u>S1 Table</u> for details]. The mass-specific rates of photosynthesis of *Lomentaria* species reported in [18] were converted to area-specific rates using the biomass to area conversion factor for *Lomentaria* provided in [12].

We modeled a linear increase of photosynthesis (due to presumed declining photorespiration) of 10% between the lowest modeled pCO_2 and the highest modeled pCO_2 (reduced photorespiration). At normoxia and pCO_2 of approximately 380 µatm in the gas phase, photosynthesis of CO₂-using red macroalgae is reduced by 5–6% by photorespiration [<u>31–35</u>]. As increasing CO₂ at constant pO_2 will reduce the mole fraction ratio of O₂:CO₂ at the site of carbon fixation by Ribulose-1,5-bisphosphate-carboxylase (RUBISCO), OA is expected to reduce photorespiratory losses of CO₂-using red algae.

The flux of dissolved CO₂ supplied to the site of RUBISCO is itself influenced by temperature both directly and indirectly. The direct effects of temperature on CO₂ flux include its solubility and diffusivity in seawater [36,37]. At a given seawater pH, *p*CO₂ decreases, but diffusivity of CO₂ increases, with increasing temperature. These two parameters govern the external concentration of dissolved CO₂ supplied to a cell on the thallus surface. Values for the concentrations and diffusivities of DIC species at 5°C intervals from 5 to 30°C were provided by CO₂Calc [36] and [37], respectively. Constants for use in calculating carbonate parameters were those of [38] based on the total pH scale. The formulation by [39] was used for the dissociation of KHSO₄ to estimate carbonate parameters.

An indirect effect of temperature on CO_2 flux is internal CO_2 dynamics caused by temperature-dependent rates of metabolism [12]. The boundary conditions for internal CO_2 concentration were established in the following way. Maximum internal CO_2 was assumed to occur in darkness and was taken as the value in equilibrium with CO_2 in external seawater. Minimum internal CO_2 was assumed to occur under light saturation and was taken as the steady-state equilibrium corresponding to the respiration rate at a given temperature. At sub-saturating light intensities, a linear relationship of C_{int} -depletion with increasing light intensity was modeled between these boundary conditions corresponding to the linear rate of carbon-fixation by photosynthesis in the light-limited portion of the curve. Together, these direct and indirect temperature-dependent effects determine the flux to, and concentration of, CO_2 substrate for RUBISCO. In addition to modeling photosynthetic rate, we estimated values for isotopic composition of carbon (δ^{13} C) and discrimination against 13 C (Δ) as an indicator of the physiological state of photosynthetic metabolism in response to OA and warming. At each seawater pH simulated, the mole fractions of dissolved CO₂ and HCO₃⁻⁷/CO₃²⁻ of total inorganic carbon were calculated. These values were multiplied by the carbon isotopic fractionation of dissolved CO₂ with respect to dissolved HCO₃⁻ in seawater at each temperature calculated using the temperaturedependent fractionation parameters for the carbonate system provided in [40]. These calculated values served as the isotopic composition value of the source seawater for calculating predicted values of δ^{13} C of the plant using the formula [41]:

$$\delta^{13}C_{org} = \frac{\delta^{13}C_{CO_2} - \alpha_{org} + 1}{\alpha_{org}}$$
(2)

where, $\delta^{13}C_{\text{org}}$ = carbon isotopic composition of the plant organic biomass

 $\delta^{13}CO_2$ = carbon isotopic composition of dissolved CO₂ in seawater of a given temperature α_{org} = overall fractionation of ¹³C in dissolved CO₂

We modeled the overall fractionation of ¹³C in dissolved CO₂ (α_{org}) due to diffusion in seawater ($\alpha_d = 1.0007$) and carboxylation by RUBISCO ($\alpha_c = 1.029$) following the analysis by [<u>41</u>]:

$$\alpha_{org} = \alpha_d * \left(\frac{(C_b - C_c)}{C_b}\right) + \alpha_c * \left(\frac{C_c}{C_b}\right)$$
(3)

where, $C_b = [CO_2]$ in the bulk medium,

 $C_c = [CO_2]$ internally at the site of RUBISCO

These carbon availability parameters reflect the weighting of fractionation due to diffusion in seawater and the activity of RUBISCO in the overall effect of the fractionation parameter. Possible values of α_{org} range from 1.0007 (diffusion limits any internal accumulation of CO₂ precluding discrimination by RUBISCO) to 1.029 (fractionation solely due to discrimination by RUBISCO). We model these relative weights as a function of light availability (<u>Table 1</u>). Support for the chosen values is provided by data of [<u>15</u>] showing that in low-light CO₂-using red algae, 80–85% of the limitation of the rate of photosynthesis can be attributed to carboxylation. Under light-limitation, ample CO₂ is available and, thus, RUBISCO carboxylation primarily determines α_{org} . Under high light, diffusion of CO₂ may limit photosynthesis and fractionation is increasingly influenced by the component due to diffusion.

The discrimination against ${}^{13}C(\Delta)$ was estimated for each CO₂, temperature and light level simulated from the formula in [<u>41</u>] using the light-dependent weightings of fractionation

Table 1. Weighting coefficients of fractionation of ¹³C dissolved in CO₂ due to diffusion ($w_{\alpha d}$) in seawater and carboxylation by RUBISCO ($w_{\alpha c}$) as a function of light intensity (PPFD).

PPFD	Wad	W _{αc}
10	0.02	0.98
35	0.07	0.93
50	0.01	0.90
75	0.15	0.85
100	0.20	0.80
400	0.30	0.70

PPFD measured in µmol photons m⁻² s⁻¹

above to estimate α_{org} :

$$\Delta = \frac{\delta^{13} C_{cO_2} - \delta^{13} C_{org}}{1 + \delta^{13} C_{org}} = \alpha_{org} - 1$$
(4)

Comparisons among Δ values obtained from future experiments with those of the modeled curve will elucidate whether those differences could be attributable to carbon-limitation at a given *p*CO₂, temperature and light intensity.

Model Evaluation

The model varies parameters related to photosynthetic characteristics (e.g., $K_{1/2}$ and V_{max}) and environmental conditions that determine dissolved CO₂ concentration and photosynthetic capacity. Our focus was on modeling the individual and combined effects of OA and increased temperature on photosynthetic rates and how this relationship changes with light intensity. Photosynthetic performance was modeled for nine seawater pH values (7.73, 7.76, 7.80, 7.85, 7.88, 7.91, 7.95, 8.07 and 8.10) at each of six different seawater temperatures (5, 10, 15, 20, 25, 30°C). Photosynthetic rate in each of these 54 combinations of seawater pH and temperature was estimated in simulations at six light intensities (10, 35, 50, 75, 100, 400 µmol photons \cdot m⁻² \cdot s⁻¹).

Other environmental and photosynthesis model parameters were held constant across all simulations. The physical and chemical seawater properties input as constants in CO₂Calc assumed a shallow near surface habitat characteristic of CO₂-using red macroalgae: salinity = 32‰; pressure = 20 dbars; and Total Alkalinity (TA) = 2200 μ mol/kg seawater. We also assumed a well-mixed flow environment with the minimum expected pathlength of 20 μ m for diffusive uptake of dissolved CO₂ [15,42–44]. A minimal diffusive boundary layer for model results is consistent with all results taken from the literature, which were measured under well-mixed conditions. We additionally explored the effect of increasing boundary layer thickness over a order of magnitude, from 20 through 200 μ m, on the principal model outcomes (described in <u>S1 Appendix</u>).

Analysis of Model Simulations

Modeled rates of photosynthesis at each pCO_2 and temperature combination were analyzed using polynomial regression, which was most appropriate because of the well-known curvilinear response of the photosynthesis-temperature relationship over the range of simulated temperatures [17]. A separate regression analysis was done at each simulated light intensity. The aim of the regressions was to estimate standardized regression coefficients for linear and quadratic terms for CO_2 and temperature and their interaction. These standardized coefficients were used to infer the relative magnitudes of effect on photosynthetic production caused by variation in CO_2 and temperature. A total of 13 models that varied in the number and combination of predictor variables (i.e., presence or absence of quadratic and/or linear combinations of predictor variables) were evaluated at each light intensity. The best fit models at each light level were determined from corrected AIC values and their respective weights and relative likelihoods computed. Standardized regression coefficients were estimated by model averaging with unconditional standard errors obtained from the weights of all 13 models using the AICcmodavg (version 1.35, [45]) package in R [46]. The model averaging and multi-model inference approach of [47] employed here is described in <u>S2 Appendix</u>.

Results

Modeled rates of photosynthesis varied in response to changing pCO_2 , changing temperature, and with the interaction of the two. The direction and magnitude of the predicted response to





Fig 1. Modeled rates of net photosynthesis as functions of predicted response to pCO_2 and known response to temperature at sub-saturating (35µmol photons \cdot m⁻² \cdot s⁻¹), approximately Ik (75 µmol photons \cdot m⁻² \cdot s⁻¹), and saturating (400 µmol photons \cdot m⁻² \cdot s⁻¹) saturating photon flux densities. Response surface fit using polynomial regression.

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 CO_2 and temperature, however, changed with light intensity (Fig 1). At sub-saturating light intensities, photosynthetic rates are predicted to be low and respond positively to increasing pCO_2 , and negatively to increasing temperature (Fig 1A). Consequently, CO_2 and temperature interacted antagonistically to influence photosynthetic rates. Standardized regression coefficients at low light intensities predict that CO_2 should have a much larger effect than temperature in both linear and quadratic components (Fig 2). Photosynthetic rates should be most responsive to pCO_2 under sub-saturating light among all of the combinations of light, temperature and CO_2 simulated. The large (in absolute value) predicted negative quadratic effect of CO_2 under low light reflects the expectation that photosynthetic rates will not increase proportionately as pCO_2 in seawater continues to rise (Fig 2B). This negative quadratic term approaches 0, and the linear coefficient increases, for CO_2 as light intensity increases to saturating values indicating a reduced curvature and greater linearity of the response of photosynthetic rates to increasing pCO_2 (cf. CO_2 in Fig 2A and 2B).

In the vicinity of I_k (here, ~75 µmol photons \cdot m⁻² \cdot s⁻¹), the model predicts a switch in the relative importance of CO₂ and temperature on photosynthetic rates. Linear and quadratic standardized coefficients for CO₂ decline to absolute values <0.5, whereas those for temperature increase in absolute magnitude to ≥ 1.25 (Fig 2). CO₂ and temperature coincide in their effect on photosynthetic rates (positive linear and negative quadratic components) and this interaction results in strong predicted synergistic effects on rates especially at high *p*CO₂ and moderate temperatures (Fig 1B, i.e., 15–25°C). At *p*CO₂ >700 uatm, photosynthetic rates are expected to double and be markedly responsive to temperature showing the characteristic, asymmetrical, hump-shaped curve with a maximum at 15–20°C. At historical and current *p*CO₂, the sensitivity to temperature of photosynthetic rates of CO₂-using macroalgae is less.





Fig 2. Standardized partial regression coefficients \pm standard errors for (A) linear and (B) quadratic components of the best-fit regression model using AICc. Coefficients for temperature (open columns), CO₂ (filled columns), and the temperature *CO₂ interaction (stippled columns) are plotted for each of six light intensities ranging from 10 to 400 µmol photons · m⁻² · s⁻¹.

A similar magnitude of increased photosynthetic rate with increased pCO_2 and temperature is expected under light saturation with two important differences (Fig 1C). Under light saturation, the effects of CO_2 and temperature on photosynthetic rate are expected to be additive (i.e., no interaction between CO_2 and temperature; Fig 2A). Also, modeled rates of photosynthesis of CO_2 -using macroalgae at high light intensities are less sensitive to temperature at high pCO_2 (i.e., negative quadratic coefficient of temperature approaches 0) than are algae near I_k (cf. Fig 1B and 1C). Overall under light saturation, the positive responses of photosynthetic rate to increasing CO_2 or temperature are of equivalent magnitudes with respect to the linear component. The negative quadratic component for CO_2 is small rendering a largely linear increase of photosynthetic rate to increasing CO_2 , whereas the more negative quadratic coefficient of temperature causes a more curvilinear response (Figs <u>1C</u> and <u>2B</u>).

Fig 3 illustrates how different combinations of temperature and pCO_2 in seawater at representative sub-saturating and saturating light intensities in which macroalgae are grown are predicted to affect Q_{10} responses. Estimated values of Q_{10} at sub-saturating light intensities vary little from ~0.9 at all pCO_2 and temperature levels simulated. In contrast, at light intensities approximating I_k or clearly saturating PFD, Q_{10} values exceed 1 at simulated temperatures $\leq 20^{\circ}$ C, increase with increasing pCO_2 , especially at 10°C in which Q_{10} estimates exceed 2. At 25 and 30°C, Q_{10} values are approximately 0.9–1.0.

Rates simulated under different light, pCO_2 and temperature regimes afford estimates of the percentage increase in photosynthesis relative to the recent past, pCO_2 standard of 380 ppm. At all temperature and light intensities simulated, the greatest percentage increase of photosynthetic rates of up to 50% in response to increasing pCO_2 is predicted to occur between 380 and 460 ppm CO₂ (Fig 4). Thereafter, further increases in rates to 100% (i.e., doubling photosynthetic rate relative to 380 ppm) are predicted to occur at ~700–800 ppm pCO_2 with algae at warmer temperatures and higher light intensities reaching this level of production at lower pCO_2 .

All of the results presented in Figs 1-4 assume a minimal thickness of the diffusive pathlength through the boundary layer. Fig 5 shows the predicted effect on net photosynthetic rate of increased diffusive pathlength through the boundary layer for plants at different combinations of pCO_2 and temperature at either saturating (400 µmol photons $\cdot m^{-2} \cdot s^{-1}$), or subsaturating (35 µmol photons $\cdot m^{-2} \cdot s^{-1}$) light intensity. Net photosynthetic rates are expected to decline sharply with a doubling of boundary layer from the thinnest pathlength (20 to 40 µm length). Thereafter, declines in photosynthetic rate and boundary layer thickness across light intensities is described by a negative power law (see <u>S1 Appendix</u>). Two additional predictions are noteworthy with respect to ocean acidification and warming. First, as expected, the higher pCO_2 concentrations in lower pH waters under ocean acidification partially compensates for greater DIC diffusion limitation associated with thicker boundary layers (<u>S1 Appendix</u>). Second, the effect of temperature, especially at, or above, saturating light intensities on photosynthetic rates predicted for minimal boundary layer thickness is progressively eliminated with increasingly thick boundary layers (Fig <u>5</u>).

Using discrimination against $^{13}\mathrm{C}$ to indicate physiological response to added $p\mathrm{CO}_2$

Discrimination against ¹³C by an alga (Δ) occurs with respect to the processes of diffusion of CO₂ and carboxylation by RUBISCO differently fractionating ¹³C. The line in Fig <u>6</u> represents expected values of Δ under well-mixed, minimal boundary layer conditions. Values of Δ are predicted to be inversely related to light intensity (i.e., less discrimination with increasing light



Fig 3. Modeled Q_{10} response of net photosynthesis as a function pCO_2 (µatm) for seaweeds in seawater at 10° (circles), 15° (squares), 20° (diamonds), 25° (triangles), or 30°C (pluses) and representative sub-saturating (35 µmol photons $\cdot m^{-2} \cdot s^{-1}$, top plot), or saturating (100 µmol photons $\cdot m^{-2} \cdot s^{-1}$, bottom plot) photon flux densities.



Fig 4. Percentage increase in modeled rates of net photosynthesis, relative to rates at 380 μ atm, as a function pCO_2 (μ atm) for seaweeds in 5° (filled circles), 10° (open circles), 15° (squares), 20° (diamonds), 25° (triangles), or 30°C (pluses) seawater at representative sub-saturating (35 μ mol photons \cdot m⁻² \cdot s⁻¹, top plot), or saturating (100 μ mol photons \cdot m⁻² \cdot s⁻¹, bottom plot) light intensities.





Fig 5. Predicted rates of net photosynthesis as a function of pathlength through the diffusive boundary layer at the thallus surface for (a.) subsaturating light (35 µmol photons \cdot m⁻² · s⁻¹), or (b.) saturating light (400 µmol photons \cdot m⁻² · s⁻¹) intensities. Lines represent power curves fit to predicted points at each combination of temperature (5 to 30°C at 5°C intervals) and *p*CO₂; purple, 380; blue, 460; green, 620; red, 780; black, 940 µatm)

intensity; Fig 6). Additionally, as the thickness of the boundary layer increases, so too does the weight of fractionation against ¹³C due to diffusion. This causes a downward shift of the curve of Δ versus light intensity (indicated by gray arrow in Fig 6). Thus, the line expressing the relationship between Δ and light intensity represents the upper boundary (maximal) discrimination expected. Isotopic composition in tissues (δ^{13} C) is a function of both isotopic composition of the source seawater and overall fractionation of ¹³C by the alga. Corresponding expected δ^{13} C values as a function of seawater temperature and CO₂ as a fraction of total Carbon are shown in <u>S1 Data</u>.





Discussion

Concern about effects of climate change on marine life has spawned a rapidly increasing number of studies characterizing responses of algae to increasing CO_2 and temperature in the oceans. Few of those studies include evaluation of the mechanism of inorganic carbon acquisition operating as interface between the organism and the external carbonate chemistry. Haphazard choices of fleshy algal species as experimental subjects in some cases, or the special concern about calcareous algae in other cases, and the experimental conditions to which they are subjected make an emergent synthesis difficult. Our motivation was to develop a quantitative model to predict the effects of increased CO₂ on photosynthesis of strictly CO₂ using (i.e., non-bicarbonate using) red macroalgae, which are estimated to comprise ~35% of all rhodophytes [11] and are over represented among macroalgae communities in deepwater, low transparency water and caves [14]. A second model representing some bicarbonate-using species is in a forthcoming manuscript. We restricted our model to seawater salinities but it could be extended to include the strictly CO₂ using red algae in fresh water (e.g., Lemanea spp.; [42-44, (48)). The CO₂-user model is based on the known temperature and pH-dependent fractionation of carbon species in seawater [35], dynamics of diffusive uptake of CO_2 in aquatic plants [25-27], and prior data on photosynthetic responses of CO₂-using red seaweeds to temperature and light [12,15,16,18]. Such a model provides a tool to assess understanding of the relationships of photosynthesis to key environmental parameters, a valid framework upon which to base predictions of changes in primary production, distribution and abundance, and a basis from the bottom-up to inform hypotheses about community-level implications to climate change.

Analysis of these model outcomes generates predictions that address four fundamental, unresolved questions about productivity of CO_2 -using algae in a changing climate: (1) How much (if any) will productivity be enhanced in acidified coastal waters?, (2) Will increased CO_2 or temperature be the principal driver of changing algal productivity?, (3) Will changing CO_2 and temperature act synergistically to influence algal productivity?, (4) How will light supply affect (if at all) photosynthetic responses to increasing CO_2 and temperature?

*p*CO₂ has a stronger effect on photosynthetic rates at subsaturating photon flux densities while temperature had a larger effect on photosynthetic rates under higher light.

The most striking result is the expectation that the individual and interactive effects of CO_2 and temperature on photosynthetic rates vary with light intensity. CO_2 has a stronger effect on photosynthetic rate at sub-saturating light intensities and temperature has a stronger effect at, or above, saturating light intensities. Standardized regression coefficients show that the magnitude of the CO_2 effect on photosynthetic rates is \geq 5-fold the effect of temperature when photon flux density is limiting. In addition, under limiting PFD the effects of changing pCO_2 and temperature on photosynthetic rate are predicted to be antagonistic, which leads to a moderate negative interactive effect. At, and above, light saturation, temperature has a positive effect on net photosynthetic rates that is stronger than the positive effect of CO_2 . Whereas they act synergistically in the vicinity of I_k, their effects on photosynthesis are additive well above I_k. Further evidence of the light-dependent effects of temperature are suggested by Q₁₀ values at each pCO_2 under sub-saturating or saturating light intensities. The Q₁₀ of photosynthetic rate is approximately 1 at all pCO_2 levels and temperatures under sub-saturating light, but > 1 and increasing at higher pCO_2 levels from 10–20°C under saturating light.

OA is predicted to increase photosynthetic rate in algae using dissolved CO_2 as their inorganic carbon source. The magnitude of the benefit will be largest at warmer temperatures, greater photon flux densities, and high flow (thin boundary layer) conditions.

It is frequently suggested that non-calcifying algae will benefit from ocean acidification, but these suggestions are made without regard to the source of carbon a macroalga uses for photosynthesis, nor the magnitude of the presumed effect. In regard to the former, there is likely to be little direct benefit to bicarbonate-using seaweeds, which are typically saturated for DIC at present [28] and several empirical studies corroborate this (e.g., [49,50]; see also reviews by [6,7]), though there may be indirect benefits. As all algae use dissolved CO₂ for some fraction of their inorganic carbon supply, the predicted benefits may apply to that fraction of the net primary productivity in those species, as well. Based on our analysis, OA should benefit macroalgae that lack carbon-concentrating mechanisms. The magnitude of this benefit to enhanced productivity in situ appears to depend on both light supply and, especially, temperature in the environments inhabited by these algae. The predicted increases in photosynthetic rates are similar across all light environments simulated up to $460-500 \mu$ atm pCO₂. The environmental range over which the greatest percentage (50%) benefit of OA is expected is from the current, approximate pCO_2 of 400 to 460 µatm. At higher pCO_2 , smaller proportionate increases in photosynthetic rate are predicted especially at lower PFD and temperature. At sub-saturating light intensities, a doubling of photosynthetic rate (relative to that measured at 380 ppm CO₂) is predicted at \geq 800 µatm pCO₂. By contrast, algae living under saturating light should show a doubling of current photosynthetic rates at pCO_2 's ranging from ~650, at temperatures $\geq 20^{\circ}C$, to >800 µatm at 5°C.

Stimulation of photosynthesis by OA effects on diffusive CO_2 uptake are not predicted to be a large part of total coastal primary productivity.

These large percentage increases in net photosynthetic rate, are not likely to cause large absolute changes in primary productivity. Maximum rates of net photosynthesis of red algae lacking CCMs are on the order of 1 µmol C-fixed \cdot m⁻² \cdot s⁻¹ at 380 µatm *p*CO₂. For seaweeds living in habitats typically saturated with light, photosynthetic rates may be expected to reach 2 µmol C-fixed \cdot m⁻² \cdot s⁻¹ at *p*CO₂'s greater than 700 µatm. For seaweeds living in light-limited habitats, photosynthetic rates may be expected to reach 0.75–1.0 µmol C-fixed \cdot m⁻² \cdot s⁻¹. All marine photorophs supply some part of their inorganic carbon demand through the diffusive uptake of dissolved CO₂. For most macroalgae, that is small fraction of their primary production. To the extend that additional dissolved CO₂ is increased there may be marginally great net photosynthetic rate if the CCM does not already completely saturate inorganic carbon supply.

OA is unlikely to increase total productivity of algae lacking CCMs in their current habitats but may allow for range expansions into brighter habitats

These predicted effects on photosynthesis in response to climate change have implications for the production, distribution and abundance of CO_2 -using rhodophytes and the communities in which they live. Rhodophytes lacking CCMs tend to live in temperate, low-light, cold water environments [13–16], precisely the kinds of conditions predicted to have little enhancement of productivity. Thus, the majority of CO_2 -using, non-calcifying macroalgae are expected have little potential to increase ecosystem-level production in their current, subtidal habitats.

We hypothesize that increased CO_2 may facilitate range expansion of CO_2 -using algae into new habitats. The potential increase in photosynthetic rate in bright, warmer environments due to alleviation of CO_2 limitation, predicted by the model, is consistent with this hypothesis. Nevertheless, increased production of CO_2 -using algae with continued rise of pCO_2 in seawater is not without constraint. Our estimates of Δ predict an inverse relationship with light intensity, which is supported by empirical data from three rhodophyte species [30], but even under saturating light (e.g., 400 μ mol C-fixed \cdot m⁻² \cdot s⁻¹) discrimination values >20 are predicted. Discrimination against ¹³C and isotopic composition values of CO₂-using rhodophytes are greatly influenced by the relative weights of fractionation of ${}^{13}C$ due to diffusion of CO₂ through the boundary layer and carboxylation by RUBISCO [41]. Our model can accommodate a variable boundary layer thickness parameter, but model outputs of carbon stable isotope discrimination, presented here assume a minimal diffusive boundary layer thickness. Maberly et al. [15] showed that Δ values of this magnitude (>20) from field-collected rhodophytes imply that discrimination by the organism is dominated by RUBISCO carboxylation, not CO₂ diffusion. Furthermore, [15] hypothesized that the inherently low maximum photosynthetic rates of CO_2 using red algae may represent an adaptation to low light environments that ensures the capacity to fix carbon is commensurate with low rates of light absorption as a means to avoid photodamage (see also [13]). Similar results showing more negative values of δ^{13} C in other macroalgae grown at lower PFD indicate increased isotopic discrimination by RUBISCO [51]. These combined results suggest that geographic range expansions and increased production of CO₂-using algae under OA may be constrained by a lack of adaptive capacity to exploit the added CO_2 in seawater. This interpretation is further supported by the model result that CO_2 had a stronger effect on net photosynthetic rate under light-limited, compared to light saturated, conditions.

Model outcomes predict the largest climate change effects on macroalgae will be during warming summers and in shallow, well mixed waters

It has been observed that increasing temperature and increasing pCO_2 can have competing effects on the growth and photosynthetic rates of seagrasses (R. Zimmerman, pers. comm.). In a red macroalga with a CCM, high pCO_2 increased the photosynthetic rate at low temperature [22]. Our model predicts that effect for seaweeds without CCMs, but the effect at saturating PFD is much stronger than at limiting PFD. There is a complex interrelationship between the multiple stressors of photosynthetic rate. Predicting the consequences of climate change on productivity requires knowledge of the light environment as well as the pH and temperature, with the greatest effects at high temperatures, and saturating light intensities. While some terrestrial organisms are strongly affected by warming winters (e.g., [52]), our results suggest that the largest effects on macroalgae will be during warming summers and in shallow habitats.

The stressors of photosynthetic rate, examined here, are likely to affect other aspects of algal physiology that contribute to growth and survival under climate change. Thermal effects on rates of photorespiration, respiration in the light were not modeled here. The model includes thermal effects on respiration inherent in the net photosynthetic rate versus temperature relationships used to parameterize the model. However, extrapolation from this model to growth rates would require greater knowledge of the temperature and boundary layer sensitivity of many additional uptake and metabolic processes.

The model presented here serves as a framework upon which principal variables associated with climate change (pCO_2 in, and temperature of, seawater), can be evaluated with respect to their effects on the productivity of phototrophs that depend on diffusive uptake, only, of CO_2

for photosynthesis. As data accumulate from studies targeting CO₂-users, comparisons can be made with predictions from the model about the relative importance and magnitudes of effect of these variables on primary production. Based on our understanding of the physiological responses of this group of phototrophs to multiple environmental stressors, we predict that climate change will increase their productivity slightly in the near future and may facilitate range expansions. An accurate understanding is crucial to inform community and ecosystem-level models that seek to predict the extent to which bottom-up effects on production due to changed conditions and resource availabilities will matter ecologically.

Supporting Information

S1 Appendix. Analysis of Diffusion Pathlength Effects. (DOCX)

S2 Appendix. Multimodel Inference Method. (DOCX)

S1 Data. Predicted values of isotopic composition of plants as a function of seawater temperature and CO₂ as a fraction of total Carbon. (XLSX)

S1 Table. Summary of data sources used for model parameterization. (DOCX)

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This work was inspired by and is dedicated to J.A. Raven.

Author Contributions

Conceived and designed the experiments: JEK SRD. Performed the experiments: JEK SRD. Analyzed the data: JEK SRD. Contributed reagents/materials/analysis tools: JEK SRD. Wrote the paper: JEK SRD.

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