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Negative correlation between altitudes and oxygen isotope ratios of seeds: exploring its applicability to assess vertical seed dispersal

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Keywords

Altitude, fruiting phenology, global warming, long-distance seed dispersal, oxygen stable isotope, plant distribution, vertical seed dispersal.

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Funding Information

Japan Society for the Promotion of Science (Grant/Award Number: "15K18718", "25241026", "25291101") Environment Research and Technology Development Fund (Grant/Award Number: "1-1401") Research Grant of Center for Ecological Research, Kyoto University (Grant/Award Number: "2013").

Received: 31 March 2016; Revised: 11 July 2016; Accepted: 18 July 2016

Ecology and Evolution 2016; 6(19): 6817– 6823

doi: 10.1002/ece3.2380

Introduction

Seed dispersal is one of the few methods for plant movement and thus has several essential ecological functions. Local seed dispersal strongly influences population dynamics and thereby community dynamics: It provides opportunity for avoiding disproportionate seed and seedling mortality near the parent, colonizing after disturbance and locating microhabitats suitable for establishment and growth (Howe and Smallwood 1982).

Abstract

Vertical seed dispersal, which plays a key role in plant escape and/or expansion under climate change, was recently evaluated for the first time using negative correlation between altitudes and oxygen isotope ratio of seeds. Although this method is innovative, its applicability to other plants is unknown. To explore the applicability of the method, we regressed altitudes on δ^{18} O of seeds of five woody species constituting three families in temperate forests in central Japan. Because climatic factors, including temperature and precipitation that influence δ^{18} O of plant materials, demonstrate intensive seasonal fluctuation in the temperate zone, we also evaluated the effect of fruiting season of each species on δ^{18} O of seeds using generalized linear mixed models (GLMM). Negative correlation between altitudes and δ^{18} O of seeds was found in four of five species tested. The slope of regression lines tended to be lower in late-fruiting species. The GLMM analysis revealed that altitudes and date of fruiting peak negatively affected δ^{18} O of seeds. These results indicate that the estimation of vertical seed dispersal using δ^{18} O of seeds can be applicable for various species, not just confined to specific taxa, by identifying the altitudes of plants that produced seeds. The results also suggest that the regression line between altitudes and δ^{18} O of seeds is rather species specific and that vertical seed dispersal in late-fruiting species is estimated at a low resolution due to their small regression slopes. A future study on the identification of environmental factors and plant traits that cause a difference in δ^{18} O of seeds, combined with an improvement of analysis, will lead to effective evaluation of vertical seed dispersal in various species and thereby promote our understanding about the mechanism and ecological functions of vertical seed dispersal.

> Long-distance seed dispersal is much less studied as compared to the local one, but it influences population dynamics, evolution of populations, metapopulation dynamics, biological invasions, and the dynamics and diversity of ecological communities (Cain et al. 2000). The movement of dispersed seeds is expressed in three dimensions: horizontal dispersal toward latitudinal and/or longitudinal directions (x- and y-axes), and vertical dispersal (i.e., dispersal toward lower/higher altitudes, z-axis).

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Vertical seed dispersal is considered to have important ecological functions on a large geographical scale. It shapes vertical distributions of plant communities in the mountains and facilitates gene flow among populations at lower and higher altitudes. In particular, it plays a key role in plant escape and/or expansion under climate change (Corlett and Westcott 2013; Naoe et al. 2016). This role becomes more important during intensive climate change including present global warming (ref. Briceño et al. 2015). Under present global warming, plants must disperse seeds several hundred kilometers along the latitudinal gradient until the year 2100 to keep up with the rapid rate of global warming (IPCC 2013): On the other hand, plants need to disperse seeds only several hundred meters along the vertical gradient due to the drastic decrease in temperature with increasing altitude (i.e., 100 m upslope roughly corresponds to -0.65°C) (Barry and Chorley 2009; also see Körner and Spehn 2002). Considering that long-distance seed dispersal occurs at a very low frequency (Nathan 2006), it is very likely that plants rely on vertical seed dispersal to escape global warming. The finding that anthropogenic habitat fragmentation, which prevents seed dispersal [e.g., dispersal by animals (Cordeiro and Howe 2003; McEuen 2004; Naoe et al. 2011) and wind (Damschen et al. 2014)], is less severe in mountainous areas (Corlett and Westcott 2013; Naoe et al. 2015) highlights the reliability of vertical dispersal. Actually, plant distributions are tracking global warming altitudinally rather than latitudinally, and the extent of tracking is considered to be large in plants with better dispersal traits, such as lighter seed in wind-dispersed plants (Parolo and Rossi 2008; Corlett and Westcott 2013). However, vertical seed dispersal itself has not been evaluated until recently (Naoe et al. 2016).

Two primary obstacles have made the evaluation of vertical seed dispersal difficult. First, drastic changes in many environmental factors occur with altitude, such as the climate and biome (Körner 2003). This prevents models from predicting vertical seed dispersal distance, because several essential modeling assumptions are not true with the altitudinal gradient. For example, in seed dispersal by wind, wind direction that determines the direction of seed dispersal is not spatially constant (Cousens et al. 2008); in seed dispersal by animal, animal movements that determine seed dispersal distance are limited by the animal's altitudinal range. In fact, vertical seed dispersal has never been modeled even in seed dispersal by wind, which is physically determined and thus, the causal factors have already been explored in detail (Nathan and Katul 2005). Second, observing vertical seed dispersal inevitably leads to observing long-distance seed dispersal. Slope distance, corresponding to a particular vertical distance, is much longer than the vertical distance

itself – alpinists must walk several kilometers to ascend several hundred meters vertically. It is very difficult to quantitatively measure such long-distance seed dispersal at a high resolution from a practical point of view (intensive sampling of dispersed seeds and identifying their mother plants in a landscape or regional scale are needed) (Cain et al. 2000). Thus, a new and highly cost-efficient method must be developed to evaluate vertical seed dispersal.

Naoe et al. (2016) have recently evaluated vertical seed dispersal for the first time, using oxygen stable isotope ratios (¹⁸O/¹⁶O, here after δ^{18} O) of seeds. δ^{18} O of plants is originally based on δ^{18} O of precipitation, which decreases with altitude largely because of temperature-driven distillation of the heavy isotope ¹⁸O from air masses as they move over orographic barriers (Bowen and Wilkinson 2002; Barbour 2007). Although plant internal factors, including evaporative enrichment of leaf water, isotopic exchange between water and organic molecules, also affect δ^{18} O of plants (Barbour 2007; Sternberg 2009), a decrease in the oxygen isotope ratios with altitude has been reported in plant materials such as leaves and tree rings (e.g., Burk and Stuiver 1981; Terwilliger et al. 2002). Naoe et al. (2016) found a negative correlation between altitudes and δ^{18} O of wild cherry seeds and then determined the altitude of the mother plant of a dispersed seed using δ^{18} O of the dispersed seed. As a result, they determined vertical seed dispersal distance, that is, the vertical distance between a dispersed seed and its mother plant (see Fig. 1 for the study concept). Their method is innovative in terms of simplicity and very low cost of evaluating vertical seed dispersal: It only needs the negative correlation between altitudes and δ^{18} O of nondispersed



Figure 1. Hypothetical regression line between altitude and oxygen isotope ratio of seeds (solid line) and its use to estimate vertical seed dispersal. We can locate the altitude of a mother plant using the regression line and oxygen isotope ratio of a dispersed seed (square as an example) and thus can estimate vertical seed dispersal distance by subtracting the altitude of a mother plant from that of a dispersed seed.

reference seeds, and dispersed seeds for estimating vertical seed dispersal. However, negative correlation between altitudes and δ^{18} O of seeds has been reported only in wild cherry and cultivated coffee (Rodrigues et al. 2011; Naoe et al. 2016), and thus, its applicability to other plants is unclear.

To explore the applicability of the method, we tested the relationship between altitudes and δ^{18} O of seeds of five woody species constituting of three families in temperate forests. Because climatic factors (e.g., temperature and precipitation), which affect δ^{18} O of plant materials, show intensive seasonal fluctuation in the temperate zone (Barry and Chorley 2009), we also evaluated the effect of the fruiting season of each species on δ^{18} O of seeds.

Material and Methods

This study was conducted at Okutama in the Kanto region, approximately 100 km west of Tokyo, Japan. The climate in the study area is the Japanese Pacific Ocean type, with heavy rainfall in summer and little snow in winter. The mean annual precipitation during 1981-2010 at 530 m a.s.l. was 1624 mm, and the mean annual temperature was 11.9°C (range: 1.3°C in January to 23.2°C in August) (Fig. 2) (Japan Meteorological Agency 2016). The study area is mountainous and is covered mostly with forest vegetation. Natural forests and conifer plantations (Cryptomeria japonica or Chamaecyparis obtusa) cover 41.3 and 50.3% of the area, respectively. The natural forests are dominated by Castanea crenata and Quercus serrata in the lower mountain zone (400-500 m a.s.l.); by Quercus crispula, C. crenata, and Fagus crenata in the middle zone (500-1500 m a.s.l.); and by Abies homolepis and Tsuga diversifolia in the upper zone (1500-1800 m a.s.l.) (Koike et al. 2008).



Figure 2. Monthly mean precipitation and temperature from 1981 to 2010 at a nearest meteorological station located 2 km from the study site, 530 m a.s.l. (Japan Meteorological Agency 2016). Bars and solid line indicate precipitation and temperature, respectively.

We targeted five fleshy fruited deciduous woody species of thee family, Prunus jamasakura, P. verecunda, P. grayana, Swida controversa, and Actinidia arguta, in the descending order of flowering and fruiting seasons (Table 1). These are widespread and rather abundant mountain species and thus are suitable for obtaining a wide vertical sampling range. We collected seeds directly from fruiting trees in various altitudes in their fruiting season during 2012-2014 in the radius of 10 km. Some of the samples of P. verecunda in 2012-2014 had also been used previously (Naoe et al. 2016). We analyzed three seeds per fruiting tree. We used the most external dead and hard tissue of a diaspore as the analysis subject, considering the possibility that δ^{18} O of whole seed demonstrates diurnal variation likely because of live tissues in the seeds (ref. Cernusak et al. 2002). Consequently, we analyzed the endocarp of all species except A. arguta, for which we analyzed its seed coat. In the study of δ^{18} O of plant materials, it is now rather common to extract α -cellulose for purification and analyze it (ref. Barbour 2007). However, we used raw materials for efficiency, referring to Barbour et al. (2001), who indicated that when a wide range in temperature and δ^{18} O of precipitation exists (expected situation in our study), much information may be gained from δ^{18} O of raw materials as compared to that from α -cellulose. By skipping this process, we could analyze A. arguta seeds, which otherwise were too small to analyze. The endocarp or seed coat was extracted from the seed and grinded for subsequent oxygen isotope analysis (Iacumin et al. 2009). Each endocarp or seed coat was weighed (approximately 0.15 mg) into silver capsules (Säntis Analytical, Milan, Italy) and rolled into balls for continuous flow (CF) combustion and isotope ratio mass spectrometry (IRMS) analysis using a high-temperature elemental analyzer (TC/EA, Thermo Fisher Scientific, Waltham, Massachusetts, USA) coupled online with a mass spectrometer (Delta V plus; Thermo Fisher Scientific) using a ConFlo IV interface (Thermo Fisher Scientific). The molecular water absorbed from the atmosphere was removed in advance by vacuum-drying the samples for at least one night before measurements to avoid water contamination in the samples. The oxygen isotope analyses were performed by measuring the CO obtained by high-temperature carbothermic reduction of the endocarp (1375°C) in the presence of excess of carbon by means of a TC/EA. Before entering into the mass spectrometer, the helium stream containing CO was separated from H₂ and N₂ by a 1.4-m molecular sieve chromatographic column held at 90°C. The sample gases were calibrated by measuring the reference substances (IAEA-601 and IAEA-602 benzoic acid; 23.14% and 71.28%, respectively) of known isotope composition (Brand et al. 2009). The isotopic composition of a sample is conventionally expressed as

Table 1. Family, fruiting season, life form, analyzed tissue in this study, sampling year, and altitude of target woody species.

Species	Family	Flowering season	Fruiting peak ¹	Life form	Analyzed tissue	Sampling year	Sampling altitude
Prunus jamasakura	Rosaceae	May	Early July (191)	Canopy tree	Endocarp	2012	280, 570, 1100 m
Prunus verecunda	Rosaceae	May	Mid July (198)	Canopy tree	Endocarp	2012	550, 800, 1100, 1290 m
						2013	550, 800, 1000, 1100, 1180, 1290 m
						2014	550, 1000, 1110, 1180 m
Prunus grayana	Rosaceae	May	Early Sep. (251)	Subcanopy tree	Endocarp	2014	260, 590, 940, 1110, 1180, 1290 m
Swida controversa Actinidia arguta	Swidaceae Actinidiaceae	May June	Late Sep. (265) Late Oct. (295)	Canopy tree Woody vine	Endocarp Seed coat	2012 2013	350, 680, 900 m 600, 950, 1100, 1200, 1280 m

¹Data from Masaki et al. (2012) and Naoe personal observation at ca. 650 m a.l.s. in a deciduous forest in the Kanto region. The number in the parenthesis indicates elapsed days from January 1.

" δ " value in per mill by comparison with international primary reference material (Vienna Standard Mean Ocean Water, V-SMOW) as follows:

$$\delta^{18} O = R_{\text{sample}}/R_{\text{reference}} - 1$$

where R denotes the ratio of numbers (N) of each isotope

 $R = N(^{18}\text{O})/N(^{16}\text{O})$

The standard deviation of the replicates was approximately 0.2% (1 σ) for measurements.

To estimate the calibration line for each species, we regressed δ^{18} O of seeds produced in the same year and altitudes (Table 1), considering the possibility of δ^{18} O variation among years (Naoe et al. 2016). To evaluate the effect of the fruiting season on δ^{18} O of seeds, we made generalized linear mixed model (GLMM), for which we used δ^{18} O of each seed as a response variable, altitude and fruiting peak (elapsed days from January 1st, maximum of 365) as explanatory variables, and species identity and year as random factors. We used the day of fruiting peak, assuming that the timing of fruit maturation is the same with that of seed maturation. Because we had no information regarding the exact day of fruiting peak at the study site and because fruiting peak of the same species was different among altitudes, we used the fruiting peak measured in a temperate forest in the same Kanto region at 650 m a.s.l. (Masaki et al. 2012). We normalized the explanatory variables for analysis. Correlation between both explanatory variables was <0.7, which indicates that multicollinearity is not severe (Dormann et al. 2013). In the analysis, we used Gaussian distribution for error and the function lemr in the lme4 package (Bates et al. 2014) in R (R Core Team 2014).

Results

A negative correlation between altitudes and δ^{18} O of seeds was found in four of the five species tested (Fig. 3).

The slope of regression lines tended to be lower in latefruiting species. The GLMM analysis revealed that altitudes and date of fruiting peak negatively affected δ^{18} O of seeds (Table 2).

Discussion

Significant regression lines between altitudes and δ^{18} O of seeds among multiple taxa indicate that the estimation of vertical seed dispersal using δ^{18} O of seed can be applicable for various species, not just confined to specific taxa. Although we targeted only temperate woody species, our method will not be confined to temperate plants because this is based on the negative correlation between altitudes and δ^{18} O of precipitation that is observed worldwide (Bowen and Wilkinson 2002). In fact, in the study on the identification of the geographical origin of food stuffs, a negative correlation between altitudes and δ^{18} O of cultivated coffee bean was reported in tropical Hawaii (Rodrigues et al. 2011). The fact that many plant species distribute in wide vertical range - over 500 m range is quite common, and over 1500 m range is not rare (e.g., in temperate and subarctic zone, Horikawa 1972; Flora of North America Editorial Committee 1997; in tropical zone, Soepadmo et al. 1996) - is also favorable for making a regression line between altitude and δ^{18} O of seeds. There are not negligible exceptional species which have short vertical ranges because of preferring specific habitat such as the coast. However, to begin with, these species have almost no chance to establish their populations at higher altitudes where their suitable habitats are probably not available. By applying our method for representative plants in each climatic zone, it may be possible to promote elucidating the process that determines the distribution of each species and plant community, and to upgrade the prediction of their future distributions under global warming. Significant regression lines between altitudes and δ^{18} O of seeds were not observed in Swida controversa. There are two potential reasons for this failure.



Figure 3. Relationship between altitude and oxygen isotope ratio of seeds of (A) *Prunus jamasakura*, (B) *Prunus verecunda*, (C) *Prunus grayana*, (D) *Swida controversa*, and (E) *Actinidia arguta* in ascending order of the date of fruiting peak. Squares, circles, and triangles indicate seeds sampled in 2012, 2013, and 2014, respectively. Note that regression lines were estimated for each year (2012, dotted line; 2013, solid line; 2014, dashed line) and that y-axes are different in the upper (A–C) and lower graphs (D, E).

 Table 2. Result of the GLMM, showing coefficient, standard error (SE), and t- and P-values for each explanatory variable.

Explanatory variables	Coefficient	SE	t	Р
Altitude	-2.283	0.126	-18.190	<0.001
Date of fruiting peak	-4.212	0.279	-15.100	<0.001

The first is poor sampling design for *S. controversa*: Its sampling vertical range and locations were smallest among the five species (i.e., 550 m range and three locations). This might lower the statistical power. The second is species-specific traits such as the fruiting season. The regression slopes of *S. controversa* may become very small due to intense variation in climate during its seed maturation (see the detailed discussion below). In such cases, a combination of other stable isotopes and elements for locating the mother plants (Rodrigues et al. 2011; Webb-Robertson et al. 2012) may be useful for evaluating vertical seed dispersal.

While we could gain effective regression lines from four of five species, there were remarkable differences among the species in δ^{18} O of seeds and in the regression slope: δ^{18} O of seeds and regression slope were smaller in latefruiting species. This suggests that it is not easy to apply regression lines of other species for estimating vertical seed dispersal, even if they belong to the same genus. As for the difference between δ^{18} O of seeds among species,

the timing of seed maturation, which is closely connected to climatic factors, would be the reason. Low temperature, high precipitation, and relative humidity are known to lower δ^{18} O of plant tissues (Barbour 2007). At the study site, late-fruiting species experienced higher temperature and precipitation during their seed maturation as compared with early-fruiting species (i.e., July to September, Fig. 2). Thus, although we lack information regarding relative humidity, seasonal variation in precipitation might cause the difference between δ^{18} O of seeds among species. As for the difference between the regression slopes among species, the length of seed maturation may be the reason. In the present study, the length of seed maturation of late-fruiting species is longer than that of early-fruiting species because all species flower in spring (Table 1). Therefore, late-fruiting species experienced drastic climate variation during their seed maturation (Fig. 2). For instance, temperature change from July to August in a nearest meteorological station is 4.7°C which corresponds to 723 m in altitudinal distance following the lapse late (i.e., 100 m upslope roughly corresponds to -0.65°C) (Barry and Chorley 2009). The slopes of latefruiting species may become smaller because climate variation during their seed maturation surpasses the climate variation with altitude. This suggests the low resolution of vertical seed dispersal in late-fruiting species due to small regression slopes. But it is notable that we used the date of fruit maturation instead of seed maturation in the

analysis. There were mature seeds in mature fruits in all species, but the date of seed maturation might be faster than that of fruit maturation. It will be desirable to clarify the exact timing of seed maturation for a finer explanation for δ^{18} O of seeds.

While our results indicate the potential applicability for various plants and reveal the effect of fruiting season on the altitude- δ^{18} O relationship of seeds, it is notable that our results have several caveats stemming from too little information about δ^{18} O of seeds. First, the mechanism of how altitude affects δ^{18} O of seeds remains ambiguous. In a future study, we need to test what climatic and other environmental factors affect δ^{18} O of seeds by investigating the relationships between these factors and δ^{18} O of seeds. To this end, seed sampling across a long chronosequence or various locations that include multidimensional environmental gradient will be effective. Second, it may be problematic that we focused on only the fruiting season, which is closely linked to seasonal variation in climate, as the causal factor for the difference in δ^{18} O of seeds among species. Other plant traits may also cause differences among species. For example, species-specific root structure and habitat may affect δ^{18} O of seeds through the differences in δ^{18} O of source water that the plants take up (ref. Flanagan et al. 1992; Roden and Ehleringer 2000; Yamanaka et al. 2006). In addition, species-specific differences in seed chemical composition (ref. Gray and Thompson 1977) and in physical distance from vein to apoplast and tortuosity of the diffusive pathway that drive evaporative enrichment of δ^{18} O of leaf water (Barbour 2007; Sternberg 2009) are likely to affect δ^{18} O of seeds. Although we believe that the effect of the fruiting season on δ^{18} O of seeds is robust to some extent considering that the effect was detected among multiple taxa, we need to test whether this effect can be detected among taxa that vary greatly in other plant traits. Annual δ^{18} O of precipitation at study sites estimated from global database OIPC (Online Isotopes in Precipitation Calculator, Bowen and Revenaugh 2003) were ranging from -7.7% at 200 m a.s.l. to $-9.9\%_{o}$ at 1300 m a.s.l., decreasing $-0.2\%_{o}$ with altitude increasing 100 m. Their much lower values compared to that of seeds (range from 15.1 to 35.7%), decreasing -1.3 to -0.2% with altitude increasing 100 m (Fig. 3), suggests strong but unidentified isotopic effects on δ^{18} O of seeds before seed maturation. The identification of environmental factors and plant traits that cause differences in δ^{18} O of seeds among altitudes and among species, combined with improvement of analysis, will lead to evaluation of vertical seed dispersal at a high resolution in various species and thereby promote our understanding about the mechanism and ecological functions of vertical seed dispersal.

Acknowledgments

We thank A. Kagawa, S. Ohashi, and T. Miyashita for their support and comments. We also thank T. F. Haraguchi and R. Hirasawa for assistance in the laboratory. This work was funded by JSPS KAKENHI Grant number 25241026, 25291101, 15K18718; the Ministry of the Environment (Environment Research and Technology Development Fund 1-1401). This study was conducted using Cooperative Research Facilities (Isotope Ratio Mass Spectrometer) and Joint Usage/Research Grant of Center for Ecological Research (2012, 2013, 2014 jure-cer), Kyoto University.

Conflict of Interest

None declared.

References

- Barbour, M. M. 2007. Stable oxygen isotope composition of plant tissue: a review. Funct. Plant Biol. 34:83–94.
- Barbour, M. M., T. J. Andrews, and G. D. Farquhar. 2001. Correlations between oxygen isotope ratios of wood constituents of *Quercus* and *Pinus* samples from around the world. Aust. J. Plant Physiol. 28:335–348.
- Barry, R. G., and R. J. Chorley. 2009. Atmosphere, weather and climate, 9th ed. Routledge, London.
- Bates, D., M. Maechler, B. Bolker, S. Walker, and R. H. Bojesen. 2014. lme4: Linear Mixed-Effects Models using 'Eigen' and S4, R Package version 1.1-7.
- Bowen, G. J., and J. Revenaugh. 2003. Interpolating the isotopic composition of modern meteoric precipitation. Water Resour. Res. 39:1299–1311.
- Bowen, G. J., and B. Wilkinson. 2002. Spatial distribution of δ^{18} O in meteoric precipitation. Geology 30:315–318.
- Brand, W. A., T. B. Coplen, A. T. Aerts-Bijma, J. K. Böhlke, M. Gehre, H. Geilmann, et al. 2009. Comprehensive interlaboratory calibration of reference materials for d18O versus VSMOW using various on-line high-temperature conversion techniques. Rapid Commun. Mass Spectrom. 23:999–1019.
- Briceño, V. F., G. L. Hoyle, and A. B. Nicotra. 2015. Seeds at risk: How will a changing alpine climate affect regeneration from seeds in alpine areas? Alp. Bot. 125:59–68.
- Burk, R. L., and M. Stuiver. 1981. Oxygen isotope ratios in trees reflect mean annual temperature and humidity. Science 211:1417–1419.
- Cain, M. L., B. G. Milligan, and A. E. Strand. 2000. Longdistance seed dispersal in plant populations. Am. J. Bot. 87:1217–1227.
- Cernusak, L. A., J. S. Pate, and G. D. Farquhar. 2002. Diurnal variation in the stable isotope composition of water and dry matter in fruiting *Lupinus angustifolius* under field. Plant Cell Environ. 25:893–907.

Cordeiro, N. J., and H. F. Howe. 2003. Forest fragmentation severs mutualism between seed dispersers and an endemic African tree. Proc. Natl Acad. Sci. USA 100:14052–14056.

Corlett, R. T., and D. A. Westcott. 2013. Will plant movements keep up with climate change? Trends Ecol. Evol. 28:482–488.

Cousens, R., C. Dytham, and R. Law. 2008. Dispersal in plants: a population perspective. Oxford University Press, Oxford.

Damschen, E. I., D. V. Baker, G. Bohrer, R. Nathan, J. L. Orrock, J. R. Turner, et al. 2014. How fragmentation and corridors affect wind dynamics and seed dispersal in open habitats. Proc. Natl Acad. Sci. USA 111:3484–3489.

Dormann, C. F., J. Elith, S. Bacher, C. Buchmann, G. Carl, G. Carré, et al. 2013. Collinearity: a review of methods to deal with it and a simulation study evaluating their performance. Ecography 36:27–46.

Flanagan, L., J. R. Ehleringer, and J. Marshall. 1992. Differential uptake of summer precipitation among cooccurring trees and shrubs in a pinyon-juniper woodland. Plant Cell Environ. 15:831–836.

Flora of North America Editorial Committee. 1997. Flora of North America: volume three, Magnoliophyta: Magnoliidae and Hamamelidae. Oxford University Press, New York.

Gray, J., and P. Thompson. 1977. Climatic information from ¹⁸O/¹⁶O analysis of cellulose, lignin and whole wood from tree rings. Nature 270:708–709.

Horikawa, Y. 1972. Atlas of the Japanese flora: an introduction to plant sociology of East Asia. Gakken, Tokyo.

Howe, H. F., and J. Smallwood. 1982. Ecology of seed dispersal. Annu. Rev. Ecol. Syst. 13:201–228.

Iacumin, P., L. Bernini, T. Boschetti, and G. P. Viale. 2009. Climatic factors influencing the isotope composition of Italian olive oils and geographic characterisation. Rapid Commun. Mass Spectrom. 23:448–454.

IPCC. 2013. Climate change 2013: the physical science basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change. Cambridge University Press, Cambridge, New York.

Japan Meteorological Agency. 2016. Mean monthly meteorological data at Ogouchi between 1981–2010. Available at http://www.data.jma.go.jp/obd/stats/etrn/view/ nml_amd_ym.php?prec_no=44&block_no=0365&year= &month=&day=&view=p1) (accessed 2 March 2016) (in Japanese).

Koike, S., H. Morimoto, Y. Goto, C. Kozakai, and K. Yamazaki. 2008. Frugivory of carnivores and seed dispersal of fleshy fruits in cool-temperate deciduous forests. J. For. Res. 13:215–222.

Körner, C. 2003. Alpine plant life: functional plant ecology of high mountain ecosystems, 2nd ed. Springer, Heidelberg.

Körner, C., and E. M. Spehn. 2002. Mountain biodiversity: its causes and function: an overview. Pp. 3–20 *in* C. Korne and E. M. Spehn, eds. Mountain biodiversity: a global assessment. Parthenon Publishing, New York.

Masaki, T., K. Takahashi, A. Sawa, T. Kado, S. Naoe, S. Koike, et al. 2012. Fleshy fruit characteristics in a temperate deciduous forest of Japan: how unique are they? J. Plant. Res. 125:103–114.

McEuen, A. B. 2004. Seed dispersal and recruitment limitation across spatial scales in temperate forest fragments. Ecology 31:507–518.

Naoe, S., S. Sakai, A. Sawa, and T. Masaki. 2011. Seasonal difference in the effects of fragmentation on seed dispersal by birds in Japanese temperate forests. Ecol. Res. 26:301–309.

Naoe, S., N. Katayama, T. Amano, M. Akasaka, T. Yamakita, M. Ueta, et al. 2015. Identifying priority areas for national-level conservation to achieve Aichi Target 11: a case study of using terrestrial birds breeding in Japan. J. Nat. Conser. 24:101–108.

Naoe, S., I. Tayasu, Y. Sakai, T. Masaki, K. Kobayashi, A. Nakajima, et al. 2016. Mountain climbing bears protect cherry species from global warming by their vertical seed dispersal. Curr. Biol. 26:R315–R316.

Nathan, R. 2006. Long-distance dispersal of plants. Science 313:786–788.

Nathan, R., and G. G. Katul. 2005. Foliage shedding in deciduous forests lifts up long-distance seed dispersal by wind. Proc. Natl Acad. Sci. USA 102:8251–8256.

Parolo, G., and G. Rossi. 2008. Upward migration of vascular plants following a climate warming trend in the Alps. Basic Appl. Ecol. 9:100–107.

R Core Team. 2014. R: a language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria.

Roden, J. S., and J. R. Ehleringer. 2000. Hydrogen and oxygen isotope ratios of tree ring cellulose for field-grown riparian trees. Oecologia 123:481–489.

Rodrigues, C., M. Brunner, S. Steiman, G. J. Bowen, J. M. F. Nogueira, L. Gautz, et al. 2011. Isotopes as tracers of the Hawaiian coffee-producing regions. J. Agric. Food Chem. 59:10239–10246.

Soepadmo, E., K. M. Wong, and L. G. Saw. 1996. Tree flora of Sabah and Sarawak: volume two. Forest Research Institute Malaysia, Kuala Lumpur.

Sternberg, L. D. L. 2009. Oxygen stable isotope ratios of treering cellulose: The next phase of understanding. New Phytol. 181:553–562.

Terwilliger, V. J., J. L. Betancourt, S. W. Leavitt, and P. K. Van De water. 2002. Leaf cellulose δD and $\delta^{18}O$ trends with elevation differ in direction among co-occurring, semiarid plant species. Geochim. Cosmochim. Acta 66:3887–3900.

Webb-Robertson, B. J., H. Kreuzer, G. Hart, J. Ehleringer, J. West, G. Gill, et al. 2012. Bayesian integration of isotope ratio for geographic sourcing of castor beans. J. Biomed. Biotechnol. 2012:450967.

Yamanaka, T., S. Iizuka, and T. Tanaka. 2006. Water source separation among co-occurring plants: an isotopeecohydrological approach. J. Jpn. Soc. Hydrol. Water Resour. 19:458–464.