Data Article

Early-Middle Pleistocene benthic turnover and oxygen isotope stratigraphy from the Central Mediterranean (Valle di Manche, Crotone Basin, Italy): Data and trends

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A R T I C L E   I N F O

A B S T R A C T

Ostracod faunal turnover and oxygen isotope data (foraminifera) along the Valle di Manche (VdM) section are herein compiled. Specifically, the material reported in this work includes quantitative palaeoecological data and patterns of ostracod fauna framed within a high-resolution oxygen isotope stratigraphy ($\delta^{18}O$) from *Uvigerina peregrina*. In addition, the multivariate ostracod faunal stratigraphic trend (nMDS axis-1 sample score) is calibrated using bathymetric distributions of extant molluscs sampled from the same stratigraphic intervals along the VdM section. Data and analyses support the research article "Dynamics of benthic marine communities across the Early-Middle Pleistocene boundary in the

Mediterranean region (Valle di Manche, Southern Italy): biotic and stratigraphic implications" Rossi et al. [1].

Specifications Table

Subject area	*Earth Science*
More specific subject area	*Palaeoecology and Oxygen Isotope Stratigraphy*
Type of data	*Tables, Figures and Text file*
How data were acquired	*Field and dissecting microscope observations. Isotope ratio mass spectrometry*
Data format	*Raw and analysed*
Experimental factors	
Experimental features	
Data source location	*San Mauro Marchesato (Crotone, Southern Italy)*
Data accessibility	*The data are available with this article*

Value of the data

- Valle di Manche (VdM) is a key-section within the Mediterranean Basin as it straddles the Early-Middle Pleistocene boundary and contains a record of the Matuyama–Brunhes reversal. The abundance data of benthic organisms here presented complement the available documentation for the VdM section.
- The multidisciplinary approach adopted provides a viable strategy for quantifying stratigraphic and palaeontological patterns, which allowed for an improved reconstruction of depositional environments.
- The data here presented could be compared to other Mediterranean siliciclastic successions that record Early-Middle Pleistocene high frequency sea level fluctuations.

1. Data

We report data from ostracod fauna (39 samples, > 3600 valves; Appendix 1) and stable isotope data from the benthic foraminifera *Uvigerina peregrina* sampled at high resolution along the 38m-thick investigated interval of the Valle di Manche section (Crotone Basin, Southern Italy [2,3]).

2. Experimental design, materials and methods

Concerning the ostracod fauna, each valve was counted as one individual (Appendix 1). *Uvigerina peregrina* specimens were picked from the > 150 μm coarse fraction of 229 sediment samples (Table 2 in [3]), which were previously disaggregated using distilled water.

2.1. Unconstrained gradient analysis

Detrended correspondence analysis (DCA) and non-metric multi-dimensional scaling (nMDS) are two widely employed indirect ordination methods in palaeoecology. As both ordination techniques
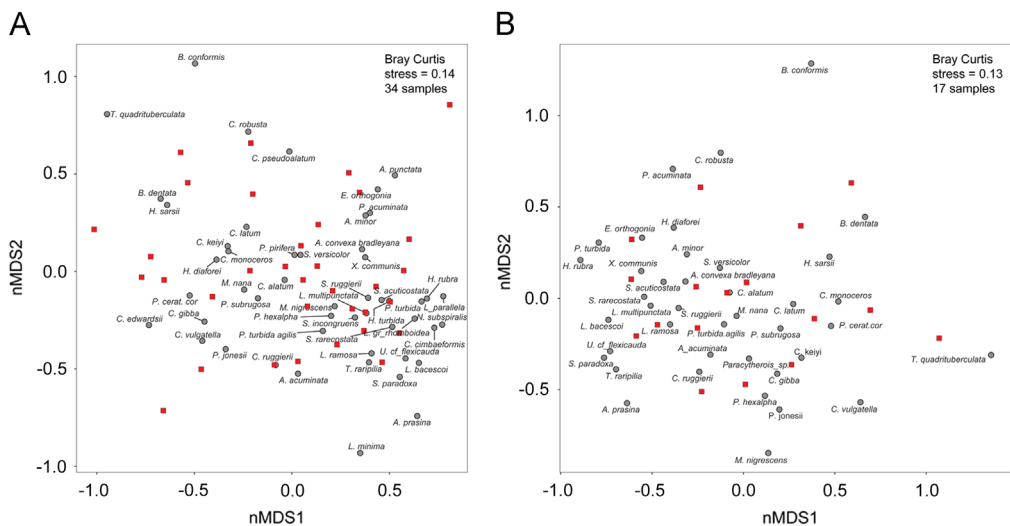
Fig. 1. Non-metric multidimensional scaling outputs performed on data matrices with different taxonomic and numerical resolution. A) Samples ≥ 20 specimens and species recorded in more than one sample (i.e., 34 samples/51 species matrix). B) Samples ≥ 20 specimens and species recorded in more than two samples (i.e., 17 samples/34 species matrix; Fig. 1B). Square and circle symbols represent sample and species, respectively.

have different strengths and weaknesses, the best approach is to use both methods as a crosscheck on the robustness of the outputs [4,5]. Faunal counts were log-transformed to prevent distortion due to very abundant species. Then, DCA and nMDS were performed on a set of abundance matrices derived varying sample and taxon thresholds. In this work, we focus on nMDS outputs (2-dimensions and based on Bray-Curtis distance; Fig. 1, Table 1A), as for DCA outputs we refer to [1]. Stratigraphic plots of nMDS and DCA axis-1 sample scores are also displayed (Table 1A; Fig. 2 A, C, E and B, D, F respectively). Ordination analyses were performed in R 3.3.2 [6] with "vegan" package and PAST software [7].

2.2. Ostracod and mollusc faunal trends along Valle di Manche (VdM) section

Reduced Major Axis (RMA) regression was performed to explore the relationship between ostracod and mollusc faunal composition along the Valle di Manche section (Table 1). The multiple DCA and nMDS axis 1 sample scores obtained from ostracods (Table 1A) were correlated via RMA to the scores previously obtained from DCA on the mollusc matrix (see [8]; Table 1A). All analyses returned high and significant correlation coefficients (Table 1B).

2.3. Oxygen isotope stratigraphy and age model

Between 10 and 15 specimens of *U. peregrina* were analysed in order to reduce statistical variability. After being lightly crushed, to remove organic contaminants, the selected specimens were soaked in hydrogen peroxide (3%). Then, analytical grade acetone was added, and the samples cleaned ultrasonically, after which the excess liquid was removed. All stable isotope analyses were carried out on an automated continuous flow carbonate preparation GasBench II device, attached to a Thermo Scientific Delta V Advantage Isotope Ratio Mass Spectrometer. Measurements of $\delta^{18}O$ were determined relative to the Vienna Peedee belemnite (VPDB) standard, with an analytical precision that is better than 0.1‰.

The chronology for the Valle di Manche section was developed by tuning the *Uvigerina peregrina* $\delta^{18}O$ signal to the stacked planktonic oxygen isotope record derived from the Mediterranean Sea [9,10]. In the initial stages, we produced an alternative age model by making use of the time scale of

Table 1

A - Sample information and major axis sample scores obtained from non-Metric Multidimensional Scaling (nMDS) and Detrended Correspondence Analysis (DCA) on Valle di Manche ostracod and mollusc datasets. A1) Mollusc sample label. A2) DCA axis 1 sample score; A3) Stratigraphic offset with respect to the adjacent ostracod sample. A4) Ostracod sample label. A5–6) nMDS axis 1 sample score obtained from a reduced ostracod matrix (employing absolute—Abs and relative—Rel abundances) comparable to the mollusc one (i.e., 17 samples see Scarponi et al., 2014). Stress values = 0.19 and 0.16, respectively. A7–8) As for A5–6 but employing DCA. A9–10) nMDS axis 1 sample score obtained from the 51×34 ostracod matrix employing absolute—Abs and relative—Rel abundances. Stress values = 0.20 and 0.19, respectively. A11–12) As for A9–10 but employing DCA. B. Linear correlation (RMA) coefficients (r—Pearson) and p-values (α=0.05) between ordination of ostracod matrices (i.e., DCA- 1 or nMDS-1) and mollusc DCA axis 1 sample score (after [8]).

A) Ordination analyses and sample information from the Valle di Manche section

after Scarponi et al. (2014)			Ostracod samples	Matrix 17 samples				Matrix 34 samples			
Label	DCA-1	S-offset	Label	nMDS-1		DCA-1		nMDS-1		DCA-1	
		(cm)		Abs	Rel	Abs	Rel	Abs	Rel	Abs	Rel
1)	2)	3)	4)	5)	6)	7)	8)	9)	10)	11)	12)
Bk22	196	20	SMA50	-0.24433	-0.25854	22	0	0.121	0.128	15	31
Bk21	117	0	SMA42	-0.05252	-0.04989	87	73	0.012	0.014	94	74
Bk20	95	-40	SMA38	0.14789	0.084843	143	127	-0.067	-0.047	135	137
Bk19	122	0	SMA30	-0.10886	-0.10742	67	57	0.051	0.052	79	80
Bk18	67	-10	SMA18	-0.03471	-0.02284	109	61	0.043	0.042	83	88
Bk17	0	40	SMA10	0.41237	0.41683	264	237	-0.307	-0.282	255	238
Bk16	9	-30	SMA8	0.31321	0.33418	206	218	-0.245	-0.237	235	195
Bk15	51	0	SMA4	0.16096	0.18789	173	151	-0.133	-0.160	174	152
Bk14	90	40	SMA-8	0.073954	0.078073	142	106	-0.051	-0.072	118	102
Bk13	98	-20	SMA-14	-0.16361	-0.16797	91	41	0.055	0.048	80	76
Bk12	223	10	SMB14	-0.18192	-0.21069	29	11	0.145	0.145	20	8
Bk11	198	10	SMB20	-0.14185	-0.15454	56	14	0.090	0.095	38	56
Bk9	164	30	SMB40	-0.24981	-0.20625	0	30	0.117	0.096	45	26
Bk8	80	20	SMB52	-0.04612	-0.05474	96	70	0.026	0.019	91	99
Bk7	59	10	SMB56	-0.02318	0.030985	84	157	0.037	0.025	134	137
Bk6	4	0	SMB60	0.35609	0.32507	181	257	-0.224	-0.229	267	236
Bk5	272	60	SMB76	-0.21757	-0.22499	2	5	0.136	0.136	24	20

B) Linear correlation: ordination axis 1 ostracod-sample scores *vs.* DCA axis 1 mollusc-sample score

Ostracod (17 samples matrix) *vs.* Mollusc matrix

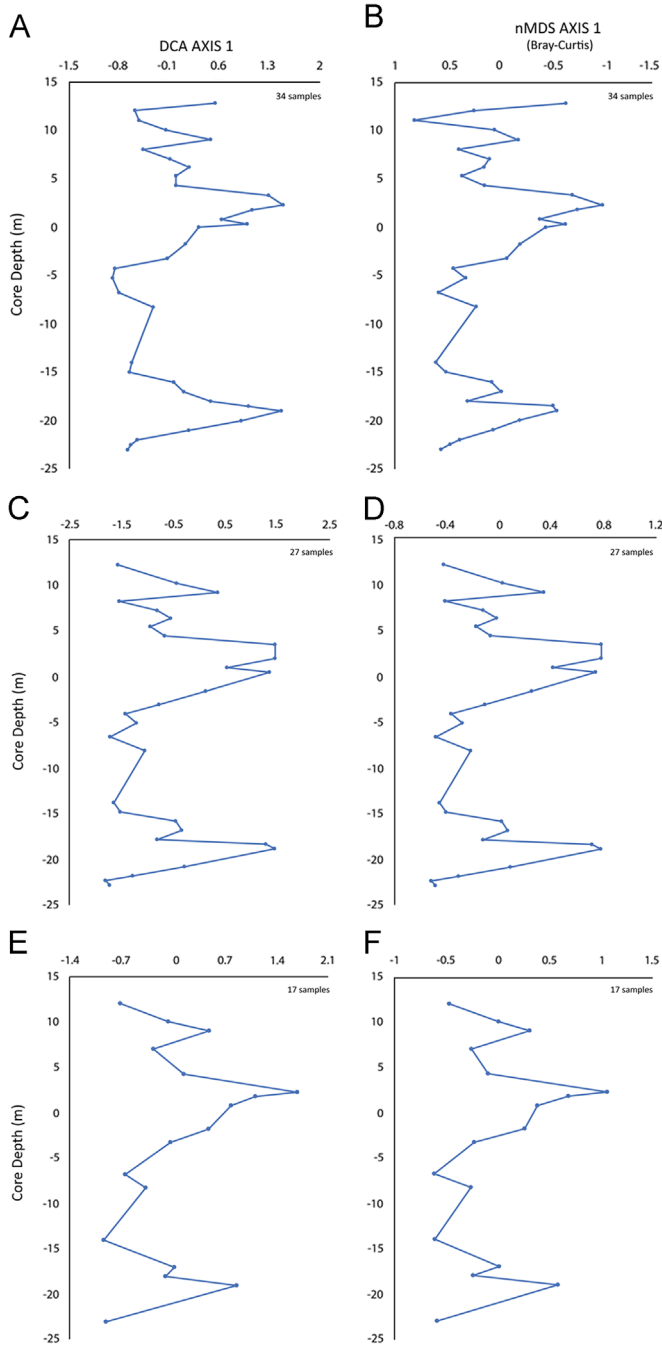|		Ostracod (34 samples matrix) *vs.* Mollusc matrix	
nMDS-1 absolute abundance	r = -0.844, p«0.05	nMDS-1 absolute abundance	r = 0.849, p«0.05
nMDS-1 relative abundance	r = -0.873, p«0.05	nMDS-1 relative abundance	r = 0.864, p«0.05
DCA-1 log-transformed raw values	r = -0.881, p«0.05	DCA-1 log-transformed raw value	r = 0.894, p«0.05
DCA-1 relative abundance	r = -0.880, p«0.05	DCA-1 relative abundance	r = -0.905, p«0.05

Fig. 2. Multiple stratigraphic plots of Detrended Correspondence Analysis (A, C, E) and non-Metric Multidimensional Scaling (B, D, F) axis 1 sample scores. A-B) Sample ≥ 20 specimens and species singletons excluded. C-D) Sample ≥ 25 specimens and species occurrence ≥ 5 samples. E-F) Ostracod dataset comparable (in sample size and sampling resolution) to the mollusc dataset reported in [13]; sample size ≥ 20 specimens and species singletons excluded.

Table 2

A) Sample information and ostracod DCA sample axis 1 score obtained from the 51 species/34 samples matrix of Valle di Manche section (DCA performed with PAST 3.11). B) Bathymetric calibration of ostracod samples. Reduced major axis regression coefficients: slope a=0.46884; intercept b=24.175; r= -0.92; p=7.87 10^{-6}; standard error of the estimates=14.3 m. C) Pearson linear correlation coefficient (r) and p (uncorr.) values (α=0.05) between DCA 1 sample scores and % of sand in each sample are shown. Regression models performed with PAST 3.11.

Label	A) Ostracod Samples: age, grain size and DCA score						B) Water depth
	Position (m)	Age (ka)	Sample weight (gr)	Sand fraction (> 63 µm)		DCA1 sample score	Water depth (m)
				(gr)	(%)		
SMA53	12.81	741.8	46.9	2.67	5.7	164	101
SMA50	12.06	744.4	48.0	6.14	12.8	15	31
SMA46	11.06	747.8	50.8	1.44	2.8	4	26
SMA42	10.06	751.2	46.9	8.57	18.3	94	68
SMA38	9.06	754.6	48.8	3.26	6.7	135	87
SMA34	8.06	758.0	47.4	9.01	19.0	34	40
SMA30	7.06	761.2	45.4	14.39	31.7	79	61
SMA26	6.21	764.0	47.7	5.13	10.8	55	50
SMA22	5.31	767.0	54.9	1.48	2.7	41	43
SMA18	4.31	770.1	55.0	3.33	6.1	83	63
SMA14	3.31	773.4	57.2	3.20	5.6	242	138
SMA10	2.31	777.5	55.3	3.43	6.2	255	144
SMA8	1.81	780.0	56.2	5.11	9.1	235	134
SMA4	0.81	784.5	55.0	1.89	3.4	174	106
SMA2	0.31	786.3	58.5	1.52	2.6	238	136
SMA-1	0.00	787.5	51.8	2.58	5.0	151	95
SMA-8	−1.75	794.0	46.6	3.91	8.4	118	79
SMA-14	−3.25	795.6	46.4	6.79	14.6	80	62
SMB4	−4.25	796.7	54.9	11.41	20.8	0	24
SMB8	−5.25	797.7	57.1	7.82	13.7	2	25
SMB14	−6.75	799.3	56.2	3.52	6.3	20	34
SMB20	−8.25	800.9	55.1	4.81	8.7	38	42
SMB40	−14.00	811.6	54.0	4.96	9.2	45	45
SMB44	−15.00	813.8	50.1	19.76	39.4	20	34
SMB48	−16.00	827.3	54.5	21.30	39.1	66	55
SMB52	−17.00	839.4	55.5	6.05	10.9	91	67
SMB56	−18.00	846.3	55.0	18.22	33.1	134	87
SMB58	−18.50	850.6	53.8	6.11	11.4	240	137
SMB60	−19.00	855.8	56.4	6.79	12.0	267	149
SMB64	−20.00	861.9	53.5	1.30	2.4	204	120
SMB68	−21.00	863.6	54.9	6.80	12.4	162	100
SMB72	−22.00	865.3	54.9	4.57	8.3	52	49
SMB74	−22.50	866.1	50.4	4.05	8.0	25	36
SMB76	−23.00	867.0	54.6	8.54	15.6	24	35

C) DCA score *vs*. % of sand - linear correlation

r = 0.291 r^2 = 0.085 p = 0.094

Konijnendijk and collaborators [11], which is also based on a stacked and averaged suite of oxygen isotope records from the eastern Mediterranean, in this case from benthic foraminifera. This initial tuning approach was based on the assumption that the correlation of the benthic δ^{18}O signal from the VdM succession to a benthic record from the Mediterranean region appeared to be a more advisable choice than the use of a planktonic δ^{18}O stack as a tuning target. However, the benthic δ^{18}O from VdM and the benthic δ^{18}O stack of [11] have little in common at either low or high frequency, as the suite of cores used by [11] reflects the dynamics of different (i.e. deeper) water masses. Serious discrepancies between the dataset from VdM and the benthic δ^{18}O stack in the time interval from ca. 860 to 815 ka (MIS 21), lead to difficulties in developing a tuned timescale (see Figure 10 in [3]). This is an interval
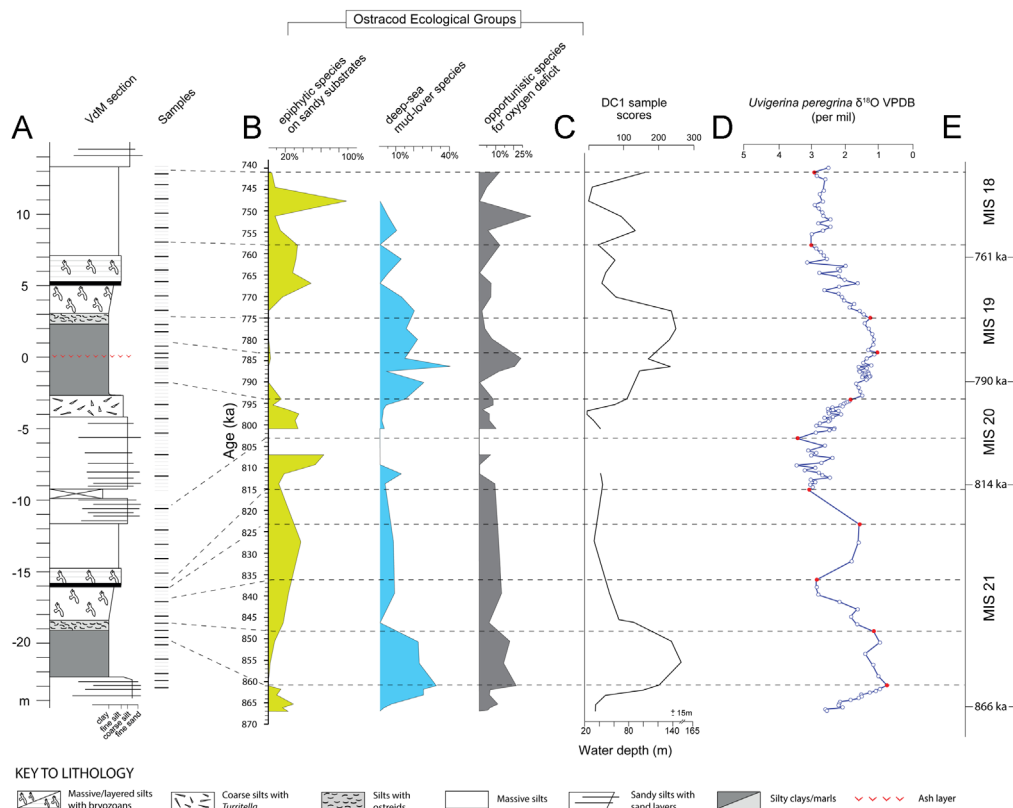
Fig. 3. Data summary of the high-resolution chronostratigraphic and palaeoenvironmental inferences retrieved at Valle di Manche (VdM). A) Physical stratigraphy of VdM section along with location of the 229 collected samples, in bold the 39 samples analysed for the ostracod fauna. B) Ostracod ecological groups distinguished on the basis of different ecological preferences, in terms of substrate and oxygen conditions, of the species recorded along the VdM section. C) Stratigraphic pattern in DCA-calibrated water depth based on the 34×51 ostracod matrix (see also Fig. 2A). D) *U. peregrina* oxygen isotope stratigraphy of the VdM section. E) Marine Isotope Stages (MIS) straddling the Early-Middle Pleistocene transition. Red dots represent the control points employed for reconstructing the VdM section age model. Panel A is plotted versus stratigraphic depth. Panels B-E are plotted versus age.

when some sources of uncertainty arise in the time scale developed by [11], as changes in insolation forcing are generally relatively small between 700–950 ka, no sapropel layers are present, and proxies lack a characteristic pattern to tie to insolation, making the resulting chronology dubious [12]. For these reasons, this initial age model was rejected.

On the other hand, transfer of the time scale by Wang and collaborators [9] proved very straightforward. As each version of the age model was developed, the age of every sample was estimated by linear interpolation between the control points. We closely monitored changes in sedimentation rate when defining age-depth correlations. If substantial changes in sedimentation rates were generated by the use of specific age controls, we evaluated whether the implied changes in the flux of biogenic and/or detrital sediment were reasonable and justified within the geological setting of the VdM section. According to our age model, the studied record spans the time interval from ca. 870 ka to 740 ka (Table 2A and Fig. 3). For more information on *U. peregrina* oxygen isotope data, we refer to [3].

2.4. Environmental proxies calibration

Sand percentages within samples (a proxy for substrate texture) is interpreted as a driver of ostracod turnover along sedimentary successions. In this work, sand percentage was plotted against DCA axis-1 sample scores (Table 2A) via linear correlation (least squares) to evaluate the role of substrate in driving ostracod faunal changes (Table 2C). Sand fraction includes both biotic and abiotic grains $> 63\,\mu m$ (Table 2A).

A linear correlation model (RMA) was also applied for bathymetry estimates of ostracod samples (Table 2B). Given the lack of quantitative water-depth information on ostracods species here recovered, water-depth calibrations rely on bathymetry inferences available for mollusc species retrieved in concomitance or proximity of the horizons sampled for ostracods (Table 1A column 3).

Sample-level bathymetry was calculated via the weighted average of a sub-set of extant mollusc species for which optimum bathymetry values were known (see Appendix 2 in [8]). Among the 14 extant taxa reported in [8], all cemented species (i.e., *Anomia ephippium, Heteronomia squamula*) were excluded from calibration, as they commonly show low association between ordination scores and bathymetry [13,14]. Then, a RMA regression between sample-level bathymetry estimates and DCA axis-1 ostracod sample scores was calculated (Table 2B).

Information collected at Valle di Manche and relative climatic, environmental and chronostratigraphic inferences are summarised in Fig. 3.

Funding sources

Transparency document. Supporting information

Supplementary data associated with this article can be found in the online version at http://dx.doi.org/10.1016/j.dib.2018.02.017.

Appendix A. Supporting information

Supplementary data associated with this article can be found in the online version at http://dx.doi.org/10.1016/j.dib.2018.02.017.

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