

Linking Isotope Exchange with Fe(II)-Catalyzed Dissolution of Iron(hydr)oxides in the Presence of the Bacterial Siderophore Desferrioxamine-B

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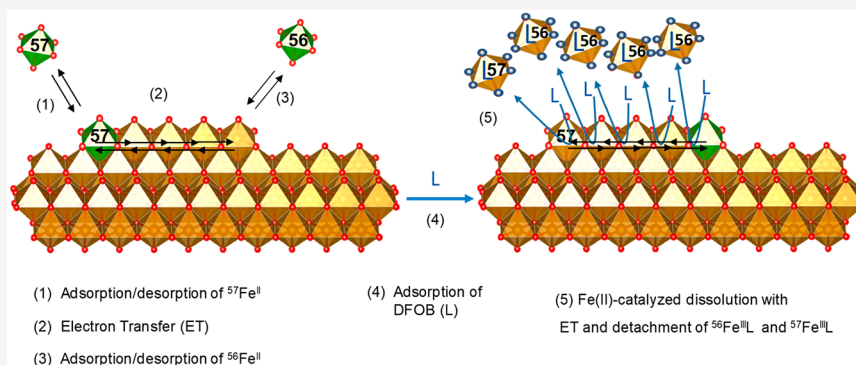
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Supporting Information



ABSTRACT: Dissolution of Fe(III) phases is a key process in making iron available to biota and in the mobilization of associated trace elements. Recently, we have demonstrated that submicromolar concentrations of Fe(II) significantly accelerate rates of ligand-controlled dissolution of Fe(III) (hydr)oxides at circumneutral pH. Here, we extend this work by studying isotope exchange and dissolution with lepidocrocite (Lp) and goethite (Gt) in the presence of 20 or 50 μM desferrioxamine-B (DFOB). Experiments with Lp at pH 7.0 were conducted in carbonate-buffered suspensions to mimic environmental conditions. We applied a simple empirical model to determine dissolution rates and a more complex kinetic model that accounts for the observed isotope exchange and catalytic effect of Fe(II). The fate of added tracer $^{57}\text{Fe}(\text{II})$ was strongly dependent on the order of addition of $^{57}\text{Fe}(\text{II})$ and ligand. When DFOB was added first, tracer ^{57}Fe remained in solution. When $^{57}\text{Fe}(\text{II})$ was added first, isotope exchange between surface and solution could be observed at pH 6.0 but not at pH 7.0 and 8.5 where $^{57}\text{Fe}(\text{II})$ was almost completely adsorbed. During dissolution of Lp with DFOB, ratios of released ^{56}Fe and ^{57}Fe were largely independent of DFOB concentrations. In the absence of DFOB, addition of phenanthroline 30 min after tracer ^{57}Fe desorbed predominantly $^{56}\text{Fe}(\text{II})$, indicating that electron transfer from adsorbed ^{57}Fe to ^{56}Fe of the Lp surface occurs on a time scale of minutes to hours. In contrast, comparable experiments with Gt desorbed predominantly $^{57}\text{Fe}(\text{II})$, suggesting a longer time scale for electron transfer on the Gt surface. Our results show that addition of 1–5 μM Fe(II) leads to dynamic charge transfer between dissolved and adsorbed species and to isotope exchange at the surface, with the dissolution of Lp by ligands accelerated by up to 60-fold.

INTRODUCTION

Iron (Fe) mobilization from poorly soluble Fe(III) phases by natural ligands (e.g., siderophores) is a key Fe acquisition strategy for organisms in Fe-deficient conditions.^{1–14} Siderophore-promoted dissolution rates of Fe(III)(oxyhydr)oxides have been reported to be slow at circumneutral pH, but can be accelerated by synergistic effects of two or more ligands.^{15–21} For example, the dissolution rates of hematite,^{20,22} ferrihydrite,²³ and goethite with bacterial siderophore desferrioxamine-B (DFOB) were accelerated 2–10 times by addition of

oxalate,^{15–17,21} Suwannee river fulvic acid,²⁴ and ascorbate.^{20,25} The synergistic Fe mobilization was also observed in carbonate soils with ascorbate and siderophores.²⁶

Very recently, we have demonstrated that, at circumneutral pH, submicromolar concentrations of added Fe(II) signifi-

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cantly accelerate dissolution rates of lepidocrocite (Lp), goethite (Gt), hematite, and ferrihydrite, with the synthetic ligands ethylenediaminetetraacetate (EDTA) and hydroxybenzyl ethylenediaminediacetic acid (HBED) and also with the bacterial siderophore desferrioxamine-B (DFOB).^{27,28} The catalytic effect (CE) of Fe(II) on the dissolution of Lp at pH 6 with EDTA could be described by a kinetic model in which accelerated dissolution was attributed to electron transfer (ET) to surface Fe(III) followed by detachment of Fe(III)EDTA.

In other studies, recrystallization of Fe(III)(oxyhydr)oxide minerals observed in ⁵⁷Fe isotope tracer studies was attributed to adsorption of added ⁵⁷Fe(II) followed by interfacial ET and bulk conduction.^{29–33} Thus, during recrystallization, mineral growth occurs at the adsorption site and bulk ET leads to Fe(II) release at a distant site with zero net dissolution of the mineral.^{34–40} Although Fe(II) adsorption is pH-dependent, the net interfacial ET per adsorbed Fe(II) was found to be independent of solution pH.^{37,40}

We included these processes in interpreting the release of ⁵⁷Fe at pH 6 in our recent study of Lp dissolution with ⁵⁷Fe(II) and EDTA.²⁷ In that study, however, we did not explicitly include isotope exchange in our kinetic model, nor did we quantify the ratio of ⁵⁷Fe release over ⁵⁶Fe release during dissolution. Here, we expand our previous work by examining isotopic exchange in greater detail at pH 6.0, 7.0, and 8.5 using the dissolution-promoting ligand DFOB. We use an empirical kinetic model to compare the rates of dissolution for the complete dissolution reaction with initial rates that we reported previously.²⁸ We then apply an expanded mechanistic kinetic model to describe isotopic exchange before and during Lp dissolution in the presence and absence of DFOB at pH 6.0 and 7.0. We also conducted some experiments with Gt to assess the effect of the solids (Gt and Lp) on the fate of added ⁵⁷Fe(II). Finally, we conducted experiments with both Lp and Gt in which we added phenanthroline (phen) to desorb Fe(II) and to examine isotopic exchange in the absence of dissolution.

MATERIALS AND METHODS

Chemicals and Solutions. Aqueous solutions were prepared using high-purity doubly deionized (DDI) water (Barnstead Nanopure, >18 MΩ cm). All chemicals used were of analytical grade and are listed in the Supporting Information (SI) (Table S1). The detailed synthesis procedure and characterization of Lp and Gt were described in our recent studies.^{27,28} Surface areas as measured by Brunauer–Emmett–Teller (BET) were 63 m² g^{−1} for Lp and 105 m² g^{−1} for Gt. A 20 mM ⁵⁷Fe(II) stock solution was prepared by dissolving 5.7 mg of ⁵⁷Fe (Sigma-Aldrich, 99% pure in Fe, with an isotopic composition of 95.06% ⁵⁷Fe, 3.04% ⁵⁶Fe, and 1.86% ⁵⁸Fe) in 100 μL of 2.5 M HCl at 50–60 °C for 4 h and dilution to 5.00 mL with DDI H₂O for conducting studies on isotope exchange and dissolution. The final concentration of Fe(II) was verified colorimetrically with phen, and the concentration of ⁵⁷Fe with inductively coupled plasma-mass spectrometry (ICP-MS). Both concentrations were 20 ± 1 mM (standard deviation of triplicate measurements).

Iron(hydr)oxide Dissolution in Batch Experiments with DFOB and Fe(II) or ⁵⁷Fe(II). Batch dissolution experiments were conducted under anoxic conditions at room temperature (22–24 °C) with suspensions of 1125 μM (=0.1 g L^{−1}) Lp or Gt, 20 or 50 μM DFOB, and varied concentrations of Fe(II) (1, 2, and 5 μM). Isotope experiments

were conducted with 2 μM ⁵⁷Fe(II). Most experiments were conducted at pH 7.0 in 3 mM NaHCO₃ solution, purged with a high-purity gas mixture of 2% CO₂ in N₂. Additional experiments were conducted at pH 6.0 (5 mM 2-(*N*-morpholino)ethanesulfonic acid (MES)), pH 7.0 (5 mM 3-(*N*-morpholino)propanesulfonic acid (MOPS)), and pH 8.5 (5 mM piperazine-*N,N'*-bis(2-ethanesulfonic acid) (PIPES)), with 9.5 mM NaCl as inert electrolyte. In the series of connected experiments with different Fe(II) concentrations, each experiment was conducted once. All ⁵⁷Fe(II) isotope experiments were conducted in duplicate. Specifics of the conditions for all experiments are summarized in Table S2.

Briefly, suspensions were prepared by dispersing 10 mg of Lp or Gt in 100 mL of aqueous solution. The 130 mL reaction flasks were sonicated for 10–15 min to obtain homogeneous suspensions. All suspensions were purged with high-purity N₂ (either with or without 2% CO₂) for at least 3 h before initiation of the experiments (*t* = 0). After dissolved O₂ concentrations dropped to below 20 nM (monitored with a PreSens Fibox 4 trace oxygen sensor), experiments were started by addition of small volumes of N₂-purged stock solutions to reach 20 or 50 μM DFOB and 0–5 μM Fe(II). In one set of experiments, Fe(II) was added first, followed by addition of DFOB after 1800 s; in a second set, DFOB was added first, followed by addition of Fe(II) after 1800 s. Samples (2 mL) were drawn periodically, filtered through 0.1 μm nylon filters (Whatman Puradisc 13 syringe filters), and diluted in 1% HNO₃ (Merck, suprapure) for inductively coupled plasma-mass spectrometry (ICP-MS, Agilent 7500cx) analysis. The dissolved Fe concentrations ([Fe]_{diss}) were measured with repeated ICP-MS measurements. At the applied concentrations of up to 5 μM Fe(II), our solutions and suspension were not oversaturated with respect to FeCO₃ (siderite), as shown with speciation calculations in Figure S1 in the SI.

Isotope Exchange and Dissolution. To investigate the fate of added Fe(II), we added ⁵⁷Fe(II) as a tracer 1800 s before or after DFOB addition to Lp or Gt suspension. Mineral dissolution was followed by measuring the concentration of dissolved iron ([⁵⁶Fe]_{diss}) with ICP-MS at mass 56, calibrated to represent the concentration of the total natural abundance iron (5.85% ⁵⁴Fe, 91.75% ⁵⁶Fe, 2.12% ⁵⁷Fe, 0.28% ⁵⁸Fe). This is the concentration of Fe released from the solid during isotope exchange and dissolution.

The dissolved concentration of ⁵⁷Fe ([⁵⁷Fe]_{diss}) was measured with ICP-MS at mass 57. The concentration of ⁵⁷Fe released from the solid during isotope exchange and dissolution (0.0212[⁵⁶Fe]_{diss}) was subtracted from [⁵⁷Fe]_{diss} to obtain the dissolved concentration of tracer ⁵⁷Fe ([⁵⁷Fe]_{tracer,diss}).

Isotope Exchange without Net Dissolution. To assess isotopic exchange at pH 7.0 without net dissolution, we added phen (100 μM) 1800 s after ⁵⁷Fe(II) (2 μM) addition to Lp or Gt suspension. Phen complexes over 99.9% of dissolved Fe(II) as Fe(II)(phen)₃²⁺ (Figure S2). In control experiments, phen did not lead to dissolution of Lp, and Fe(II)(phen)₃²⁺ formed by addition of Fe(II) to an Lp suspension containing 100 μM phen did not, or only very, weakly adsorb to the surface of Lp (Figures S3 and S4). After desorption of Fe(II) with phen, dissolved concentrations of ⁵⁷Fe and ⁵⁶Fe were measured with ICP-MS.

Kinetic Modeling. In the present work and our previous studies, we have used a variety of approaches for kinetic

modeling: (i) empirical modeling based on initial dissolution rates,²⁸ (ii) empirical modeling based on the full time course of dissolution (this work), and (iii) detailed mechanistic modeling (both this and previous work²⁷).

Empirical modeling of Lp dissolution at pH 7.0 with Fe(II) added 1800 s after 20 or 50 μM DFOB was conducted to describe the full time course of the experimentally measured dissolved Fe concentrations. We applied a pseudo-first-order model that assumes that the rate of Fe(II)-catalyzed dissolution (R_C) is proportional to the free ligand concentration $[L]$ in the solution and the concentration of added Fe(II). In short, the increase in $[\text{Fe}]_{\text{diss.}}$ over time is expressed in terms of a rate law with a single adjustable rate coefficient $k_{C,\text{app}}$ [$\text{s}^{-1} \text{M}^{-1}$], the concentration of added Fe(II), and the ligand concentration. The reaction rate decreases as $[\text{Fe}]_{\text{diss.}}$ approaches the concentration of initially added ligand $[L]_0$

$$R_C = \frac{d[\text{Fe}]_{\text{diss.}}}{dt} = [\text{Fe(II)}] \cdot k_{C,\text{app}} \cdot ([L]_0 - [\text{Fe}]_{\text{diss.}, t}) \quad (1)$$

To calculate $[\text{Fe}]_{\text{diss.}}$ as a function of time, we used the integrated form

$$[\text{Fe}]_{\text{diss.}, t} = [\text{Fe}]_{\text{diss.}, t_1} + [L]_{t_1} \cdot (1 - e^{-(t-t_1) \cdot k_{C,\text{app}} \cdot [\text{Fe(II)}]}) \quad (2)$$

where t_1 is the time of Fe(II) addition (between 1800 and 1920 s). $[\text{Fe}]_{\text{diss.}, t_1}$ is the concentration of dissolved Fe at t_1 including the concentration of added Fe(II). $[L]_{t_1}$ is the free ligand concentration at t_1 ($[L]_0 - [\text{Fe}]_{\text{diss.}, t_1}$). Note that, in this model, the contribution of the noncatalyzed dissolution rate (R_L) after the addition of Fe(II) is assumed to be negligible in comparison to R_C .

For detailed mechanistic kinetic modeling, we used the kinetic program ACUCHEM⁴¹ and MATLAB (MATLAB, MathWorks, Inc., Natick, Massachusetts, www.mathworks.com). The model was used to determine whether the proposed mechanisms (i.e., reactions included in the model) could explain the measured data including isotope exchange. The Nelder Mead Simplex Optimization routine in MATLAB was applied to optimize the rate coefficients of rate-determining reactions by minimization of the sum of squared differences between measured data points and model output for several experiments simultaneously.

RESULTS AND DISCUSSION

Dissolution of Lp in a Carbonate Buffer. Dissolution experiments were conducted with Lp (1125 μM) at pH 7.0 in a carbonate buffer for comparison with previous studies in MOPS buffer and to compare previously reported initial dissolution rates²⁸ with rates describing the full course of the dissolution reaction. As shown in Figure 1, addition of either 20 or 50 μM DFOB alone (gray diamonds) resulted in only a very slow increase of dissolved Fe concentrations ($[\text{Fe}]_{\text{diss.}}$) over time. In experiments where 1–5 μM Fe(II) was added after DFOB, accelerated dissolution was observed. Note that the time interval between addition of the two reactants, here DFOB first followed by Fe(II), was always 1800 s.

Because of the low solubility of Fe(III) at pH 7, dissolved Fe measured in the Lp suspensions can be assumed to be Fe(III) complexed by DFOB. The extent of Lp dissolution is ultimately limited by the total ligand concentration ($[L]_T$). With addition of 5 μM Fe(II), $[\text{Fe}]_{\text{diss.}}$ (green squares) increased rapidly, indicating accelerated dissolution, and

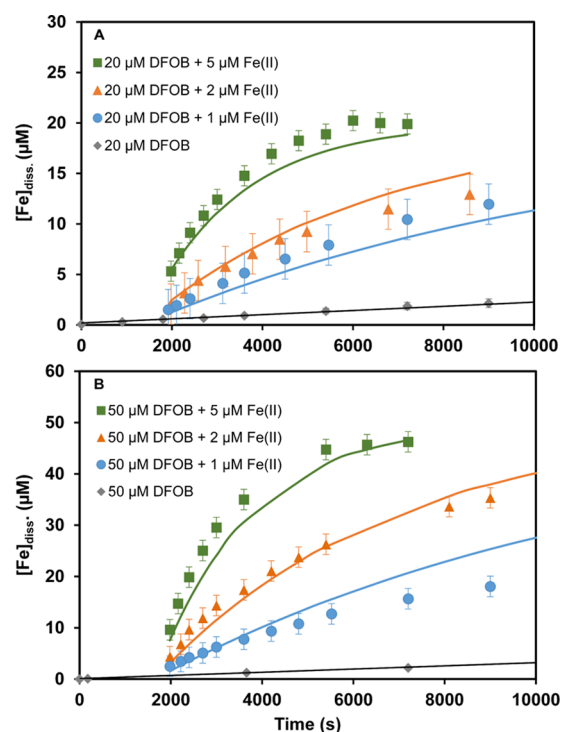


Figure 1. Lepidocrocite (Lp) dissolution after addition of Fe(II) to an Lp suspension (1125 μM) in the presence of (A) 20 μM DFOB and (B) 50 μM DFOB, under anoxic conditions at pH 7.0 (carbonate-buffered). The error bars correspond to the standard deviations of ICP-MS measurements obtained from repeated calibrations. Fe(II) was added 1800 s after DFOB addition. The solid lines represent empirical model fits.

reached a plateau within 8000 s at values approaching the total concentration of DFOB (i.e., 20 μM in Figure 1A and 50 μM in Figure 1B). At lower added Fe(II) concentrations of 1 or 2 μM , Lp dissolution proceeded more slowly (blue circles and orange triangles) and did not reach a plateau within the experimental timeframe.

An empirical kinetic model with a single adjustable parameter $k_{C,\text{app}}$ [s^{-1}] (see Kinetic Modeling in the Materials and Methods section) is sufficient to describe the full time course of the dissolution experiments at both DFOB concentrations and all three concentrations of added Fe(II). The fits to this empirical model are shown by the solid lines in Figure 1. The simple rate expression (2) leads to adequate fits to the experimentally determined concentrations with the optimized value for $k_{C,\text{app}}$ of $93 \text{ M}^{-1} \text{ s}^{-1}$.

This empirical model provides a basis for comparison with the observations reported in our previous study conducted in MOPS buffer. In that study, we defined catalytic effect (CE) as the ratio of R_C to R_L , where, in both cases, the reaction was far from completion. (For the definitions of R_C and R_L , see Kinetic Modeling in the Materials and Methods section.) The comparison of CE values derived from the current study and those reported previously shows that CE is larger in carbonate- than in MOPS-buffered systems (Table 1 and Figure S5). This might be explained by more adsorption of Fe(II) in the presence of carbonate, which forms negatively charged inner-sphere surface complexes.⁴² In a study of the oxidation of adsorbed Fe(II) by CCl_4 ,⁴³ Fe(II) adsorption on Gt was reported to be slightly lower in Good's buffer than in a carbonate buffer. Faster and more complete adsorption of

Table 1. Dissolution Rates of Lepidocrocite and Catalytic Effect of Fe(II), in the Presence of DFOB (20 or 50 μM) with Varied Fe(II) Concentrations^b

experiments	[Fe(II)] (μM)	rate of dissolution ^a R_C and R_L (for [FeII] = 0 μM)		catalytic effect R_C/R_L	buffers (pH)
		(nM s^{-1})	($\text{nmol s}^{-1} \text{m}^{-2}$)		
20 μM DFOB	0	0.25	0.04	1	carbonate (pH 7.0)
20 μM DFOB + 1 μM Fe(II)	1	1.71	0.27	7	
20 μM DFOB + 2 μM Fe(II)	2	3.24	0.51	13	
20 μM DFOB + 5 μM Fe(II)	5	6.70	1.06	27	
20 μM DFOB	0	0.07	0.07	1	MOPS (pH 7.0) (Kang et al., 2019)
20 μM DFOB + 2 μM Fe(II)	2	0.29	0.29	4	
50 μM DFOB	0	0.34	0.05	1	carbonate (pH 7.0)
50 μM DFOB + 1 μM Fe(II)	1	4.50	0.71	13	
50 μM DFOB + 2 μM Fe(II)	2	8.79	1.40	26	
50 μM DFOB + 5 μM Fe(II)	5	20.6	3.18	60	MOPS (pH 7.0)
50 μM DFOB	0	0.20	0.03	1	
50 μM DFOB + 2 μM ⁵⁷ Fe(II)	2	2.03	0.32	10	PIPES (pH 8.5)
50 μM DFOB	0	0.47	0.08	1	
50 μM DFOB + 2 μM ⁵⁷ Fe(II)	2	0.94	0.15	2	

^aThe reported rates of accelerated dissolution were calculated at t_1 , by applying eq 1. There was no significant difference in rates when Fe(II) or ⁵⁷Fe(II) was added. The order of addition of Fe(II) (i.e., before or after DFOB addition) had no significant effect on dissolution rates. The rates of the noncatalyzed dissolution (i.e., ligand alone; R_L) were determined from fits to the data without addition of Fe(II). ^bReported dissolution rates from Kang et al. (2019) were used to compare with the empirical model fits to the measured data of this study. The catalytic effect is defined as the ratio of the rate of dissolution in the presence of ligand to Fe(II) over the rate of dissolution in the presence of the ligand alone.

Fe(II) with carbonate compared to MOPS was also observed experimentally, as shown in Figures S5B, S6A, and S6B.

⁵⁷Fe Isotope Exchange and Lp Dissolution. The use of ⁵⁷Fe as an isotopic tracer provides additional insight into the process of accelerated Lp dissolution. Experiments were conducted at pH 6.0 (MES), 7.0 (carbonate and MOPS), and 8.5 (PIPES) with 2 μM ⁵⁷Fe(II) and 20 or 50 μM DFOB. The order of addition of ⁵⁷Fe(II) and DFOB was varied with one reactant added 1800 s before the other. In all cases, rapid increase of [⁵⁶Fe]^{*}_{diss.}, corresponding to the natural abundance Fe released from Lp (shown as triangles in Figure 2), was observed only after both Fe(II) and DFOB had been added, without regard to the order of addition. In contrast, the order of addition of Fe(II) and DFOB did have a significant effect on the time course of the dissolved concentrations of tracer ⁵⁷Fe ([⁵⁷Fe]_{tracer,diss.}) shown as the open and closed squares in Figure 2.

⁵⁷Fe(II) Added First. When ⁵⁷Fe(II) was added first (i.e., before DFOB), [⁵⁷Fe]_{tracer,diss.} (solid orange squares) decreased over time at pH 6.0 and was nearly or completely undetectable at pH 7.0 and 8.5. At pH 6.0 (Figure 2A), a concurrent increase in [⁵⁶Fe]^{*}_{diss.} (red filled triangles) was observed. The sum of [⁵⁷Fe]_{tracer,diss.} and [⁵⁶Fe]^{*}_{diss.} (green circles) corresponds to the added ⁵⁷Fe(II) concentration (2.0 μM) at $t = 0$ and then decreases (over 1800 s) by 20%. This suggests that some of the added ⁵⁷Fe(II) is immediately adsorbed to the surface and undergoes isotopic exchange with surface-bound Fe(III), resulting in detectable concentrations of dissolved ⁵⁶Fe, which is presumably still in the +II oxidation state. Further adsorption appears to occur more slowly and also to be accompanied by isotopic exchange. These observations agree with our previously reported findings with ⁵⁷Fe(II) and Lp²⁷ and with studies on isotope exchange with Gt^{33,37} and hematite.^{32,38} At pH 7.0 (Figure 2B) and 8.5 (Figure 2C), [⁵⁶Fe]^{*}_{diss.} was nearly or completely undetectable (i.e., before DFOB addition). This is consistent with the immediate and

(nearly) complete loss of tracer ⁵⁷Fe(II) from solution at these pH values.

Upon addition of 50 μM DFOB, accelerated dissolution, shown by increasing [⁵⁶Fe]^{*}_{diss.} (filled purple triangles) over time, was observed at all pH values. Lp dissolution was more rapid and proceeded further toward completion at pH 7.0 than at either pH 6.0 or 8.5, in agreement with our previous findings.²⁸ In all cases, between 50% (pH 8.5) and 85% (pH 7.0) of the ⁵⁷Fe tracer was released back into solution by the end of the experiment.

DFOB Added First. As already discussed for the results shown in Figure 1, addition of DFOB alone causes only very slow Lp dissolution. Upon addition of ⁵⁷Fe(II), increasing [⁵⁶Fe]^{*}_{diss.} (empty purple triangles) was observed over time. The overlap of filled and empty purple triangles in Figure 2B indicate that the order of ⁵⁷Fe(II) and DFOB addition did not significantly affect the rate of accelerated Lp dissolution. In contrast, the behavior of the ⁵⁷Fe(II) tracer was distinctly different when DFOB was added before rather than after ⁵⁷Fe(II). The open orange squares in Figure 2B,C show that, at pH 7.0 and 8.5, 75–100% of the added ⁵⁷Fe tracer is measured as [⁵⁷Fe]_{tracer,diss.} over the entire experimental period. Preequilibration of Lp with DFOB inhibits the adsorption of ⁵⁷Fe(II) and presumably also the isotopic exchange with surface Fe(III) that was observed in the absence of DFOB. A reasonable fit to the data at pH 6.0 and 7.0 was achieved with the detailed, mechanistic kinetic model (lines in Figure 2A,B) as described below.

Kinetic Model. In our recent study, we presented a kinetic model for the dissolution of Lp with EDTA at pH 6.0 that was able to describe the entire course of the dissolution.²⁷ Here, we extended this model to the dissolution of Lp with DFOB and added reactions to describe the observed adsorption and release of ⁵⁷Fe. We also tested possible explanations for the delayed release of ⁵⁷Fe when ⁵⁷Fe(II) was added before the ligand. In particular, the model allows us to estimate the charge delocalization between added ⁵⁷Fe(II) and ⁵⁶Fe(III) in Lp

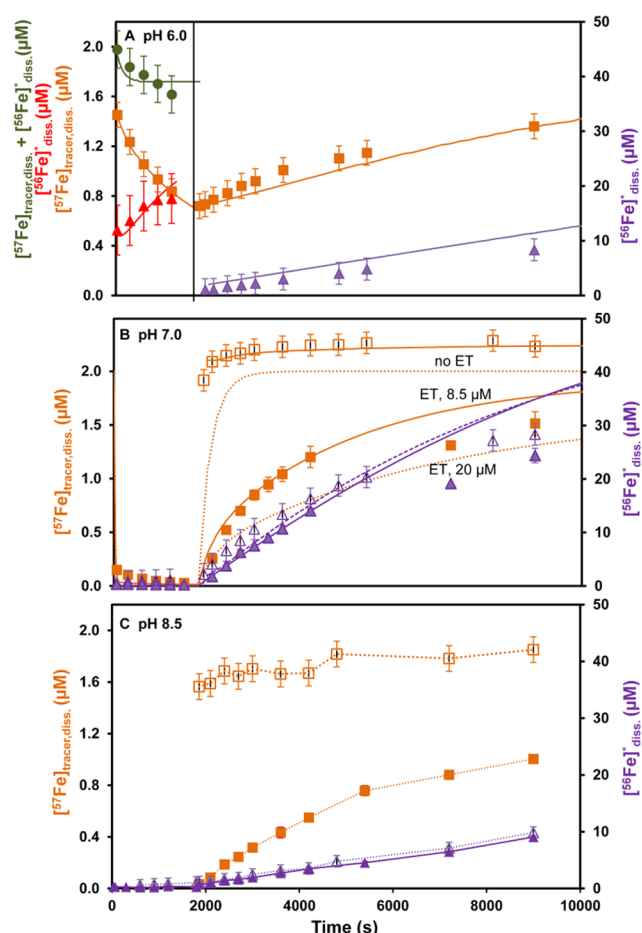


Figure 2. ^{57}Fe isotope exchange and Lp dissolution under anoxic conditions as a function of time at (A) pH 6.0 (MES), (B) pH 7.0 (carbonate buffer), and (C) pH 8.5 (PIPES). $^{57}\text{Fe}(\text{II})$ (2 μM) was added to an Lp suspension (1125 μM) 1800 s before (filled symbols, error bars correspond to the ranges of duplicate experiments) or after (empty symbols, error bars as in Figure 1) 50 μM DFOB addition. Symbols: purple triangles (right axis): concentration of Fe released into solution by Lp dissolution ($[^{56}\text{Fe}]_{\text{diss.}}$); orange squares (left axis): dissolved concentration of tracer ^{57}Fe corrected for the natural abundance of ^{57}Fe in Lp ($[^{57}\text{Fe}]_{\text{tracer,diss.}}$); red triangles (left axis): concentration of $^{56}\text{Fe}^*$ in solution resulting from isotopic exchange of ^{57}Fe with ^{56}Fe in Lp ($t < 1800$ s); green circles (left axis): sum of dissolved concentration of Fe measured as $[^{57}\text{Fe}]_{\text{tracer,diss.}}$ and $[^{56}\text{Fe}]_{\text{diss.}}$ ($t < 1800$ s). Lines: (A, B) Kinetic model fits to the measured data. The solid orange lines show the fits with the optimized concentration of surface sites (8.5 μM) for ET between $\equiv\text{Fe}^{\text{III}}-\text{O}-^{57}\text{Fe}^{\text{II}}$ and $\equiv\text{Fe}^{\text{III}}-\text{O}-\text{Fe}^{\text{II}}$ sites (R6 in the kinetic model). The dotted orange lines show release of ^{57}Fe that is too fast with no ET or too slow with ET between more surface sites (20 μM). In (C), the lines are only shown as visual guides.

before and during the dissolution process. Table 2 lists important reaction steps in the kinetic model describing the dissolution of Lp with DFOB and the acceleration by added $\text{Fe}(\text{II})$.

When DFOB is added before $^{57}\text{Fe}(\text{II})$, it adsorbs reversibly to the surface at $\text{Fe}(\text{III})$ sites (R1) and causes slow ligand-controlled dissolution (R2). Upon addition of $^{57}\text{Fe}(\text{II})$, dissolved $^{57}\text{Fe}(\text{II})$ DFOB is formed quickly (R3). Adsorption of $^{57}\text{Fe}(\text{II})$ DFOB, ET, and detachment of $^{57}\text{Fe}(\text{III})$ DFOB are aggregated in reaction R4-1. We assume that upon adsorption, negative charge is transferred to Lp, due to the strongly

reducing properties of $\text{Fe}(\text{II})$ complexed to DFOB, but we cannot resolve the individual reaction steps experimentally. R4-2 shows an alternative pathway leading to the same overall reaction. After detachment from the surface, $^{57}\text{Fe}(\text{III})$ DFOB does not undergo further reactions.^{44–47} Reactions R4-1 and R4-2 are the dominant reaction pathways for the initial transfer of negative charge when $^{57}\text{Fe}(\text{II})$ is added in the presence of DFOB. Adsorption of $^{57}\text{Fe}(\text{II})$ is minimal in this case, and added ^{57}Fe thus remains in solution, as shown in Figure 2B.

When $^{57}\text{Fe}(\text{II})$ is added first, it is reversibly adsorbed (R5) and charge transfer to $\text{Fe}(\text{III})$ can occur at the surface (R6). We cannot determine rate coefficients for charge transfer, only the distribution of charge on the surface. In the model, we assume charge transfer and equilibration of charge on a time scale of seconds to minutes for Lp and adjust the concentration of the surface sites over which the charge can distribute. In the absence of DFOB, the sequence of adsorption, charge transfer, and desorption (R5–R7) can explain the isotope exchange between $^{57}\text{Fe}(\text{II})$ in solution and $^{56}\text{Fe}(\text{II})$ on the surface (formed by ET from an adsorbed $^{57}\text{Fe}(\text{II})$ to a surface $^{56}\text{Fe}(\text{III})$). The exchange of dissolved $^{57}\text{Fe}(\text{II})$ with $^{56}\text{Fe}(\text{II})$ from the surface was observable in the filtered solutions at pH 6.0, but it was not measurable at pH 7.0 because dissolved equilibrium concentrations were too low. Note that we assumed that charge transfer at the surface is fast ($>1 \times 10^{-1} \text{ s}^{-1}$) and that equilibration of charge between ^{56}Fe and ^{57}Fe surface is reached within minutes or faster for Lp, while adsorption and desorption reactions are slower. Soltis et al.⁴⁸ reported that lifetimes of electrons in deep traps on the surface of ferrihydrite cover a large range and can reach microseconds or longer. In our model, rate coefficients for ET from 1×10^{-1} to $1 \times 10^6 \text{ s}^{-1}$ (which yield good fits with Lp) are thus in the expected range. Charge transfer and distribution of charge is slower for Gt, as discussed below.

Ligand that is added after $\text{Fe}(\text{II})$ or is present in excess over added $\text{Fe}(\text{II})$ (which was always the case in our experiments) can react with surface-bound $\text{Fe}(\text{II})$. Ligand adsorption to surface-bound $\text{Fe}(\text{II})$, charge transfer, and detachment are summarized in reactions R8 and R9. We assumed that charge transfer from a $\text{Fe}(\text{II})$ surface site with adsorbed DFOB to a neighboring $\text{Fe}(\text{III})$ site is rapid (due to strongly reducing properties of $\text{Fe}(\text{II})$ complexed to DFOB). Once formed, the adsorbed $\text{Fe}(\text{III})$ L detaches more quickly than in reaction R2, due to weaker $\text{Fe}(\text{II})-\text{O}-\text{Fe}(\text{III})$ bonds to the neighboring $\text{Fe}(\text{II})$, compared to $\text{Fe}(\text{III})-\text{O}-\text{Fe}(\text{III})$ bonds. The larger rate coefficients associated with reactions R8 and R9 compared to dissolution in the absence of $\text{Fe}(\text{II})$ (R2) account for the $\text{Fe}(\text{II})$ -catalyzed accelerated dissolution. Reactions R6 and R8–R9 explain the release of $^{57}\text{Fe}(\text{III})$ L and $\text{Fe}(\text{III})$ L and their ratios over the course of the dissolution. Finally, reactions R10 and R11 describe the rapid formation of new surface sites.

As explained in Table 2, some equilibrium constants and rate coefficients were fixed (italic font) and others optimized (normal font). Modeling was performed to determine whether the suggested reaction steps could fit the measured data. Values for the rate coefficients and equilibrium constants could not be uniquely determined because of correlations among them. For example, the values for the adsorption equilibrium constants in R1, R5, and R7 cannot be determined independently from the values of the rate coefficients for dissolution in R4, R8, and R9. However, equilibrium constants obtained in previous work²⁸ can be used to constrain fitted values of rate coefficients. The model is very sensitive to the

Table 2. Kinetic Model with List of Reactions and Equilibrium Constants (K) and Rate Coefficients (k)

nr.	reaction ^a	description	K/k^b pH 7	K/k^b pH 6
R1	$\equiv\text{Fe}^{\text{III}} + \text{L} \rightleftharpoons \equiv\text{Fe}^{\text{III}}\text{L}$	adsorption of ligand (L) on surface Fe^{III}	3.0×10^{5c}	3×10^4 to 3×10^{5c}
R2	$\equiv\text{Fe}^{\text{III}}\text{L} \rightarrow \equiv\text{Fe}^{\text{III}} + \text{Fe}^{\text{III}}\text{L}$	nuncatalyzed dissolution	3.5×10^{-5}	n.d. ^d
R3	$^{57}\text{Fe}^{\text{II}} + \text{L} \rightleftharpoons ^{57}\text{Fe}^{\text{II}}\text{L}$	dissolved $^{57}\text{Fe}^{\text{II}}\text{L}$ complex formation	5.3×10^{4e}	3.8×10^{2e}
R4-1	$\equiv\text{Fe}^{\text{III}} + ^{57}\text{Fe}^{\text{II}}\text{L} \rightarrow \equiv\text{Fe}^{\text{II}} + ^{57}\text{Fe}^{\text{III}}\text{L}$	ET from $^{57}\text{Fe}^{\text{II}}\text{L}$ to surface Fe^{III} and detachment of $^{57}\text{Fe}^{\text{III}}\text{L}$	1.4×10^2	200–600
R4-2	$\equiv\text{Fe}^{\text{III}}\text{L} + ^{57}\text{Fe}^{\text{II}} \rightarrow \equiv\text{Fe}^{\text{II}} + ^{57}\text{Fe}^{\text{III}}\text{L}$	ET from $^{57}\text{Fe}^{\text{II}}$ to surface $\text{Fe}^{\text{III}}\text{L}$ and detachment of $^{57}\text{Fe}^{\text{III}}\text{L}$	2.2×10^4	
R5	$\equiv\text{Fe}^{\text{III}} + ^{57}\text{Fe}^{\text{II}} \rightleftharpoons \equiv\text{Fe}^{\text{III}}-\text{O}-^{57}\text{Fe}^{\text{II}}$	adsorption and desorption of $^{57}\text{Fe}^{\text{II}}$ on surface Fe^{III}	7.2×10^6	6.3×10^4
R6	$\equiv\text{Fe}^{\text{III}}-\text{O}-^{57}\text{Fe}^{\text{II}} \rightleftharpoons \equiv\text{Fe}^{\text{III}}-\text{O}-^{56}\text{Fe}^{\text{II}}$	ET between ^{57}Fe and ^{56}Fe surface sites (K_{ex})	$k_{\text{ET}} > 0.1$ $K = 1^f$	$k_{\text{ET}} > 0.1$ $K = 1^f$
R7	$\equiv\text{Fe}^{\text{III}} + \text{Fe}^{\text{II}} \rightleftharpoons \equiv\text{Fe}^{\text{III}}-\text{O}-\text{Fe}^{\text{II}}$	adsorption and desorption of Fe^{II} on surface $^{57}\text{Fe}^{\text{III}}$	7.2×10^6	6.3×10^4
R8	$\equiv\text{Fe}^{\text{III}}-\text{O}-^{57}\text{Fe}^{\text{II}} + \text{L} \rightarrow \equiv\text{Fe}^{\text{II}} + ^{57}\text{Fe}^{\text{III}}\text{L}$	adsorption of L on adsorbed $^{57}\text{Fe}^{\text{II}}$, ET, and detachment	61	<5
R9	$\equiv\text{Fe}^{\text{III}}-\text{O}-^{56}\text{Fe}^{\text{II}} + \text{L} \rightarrow \equiv\text{Fe}^{\text{II}} + ^{56}\text{Fe}^{\text{III}}\text{L}$	adsorption of L on adsorbed Fe^{II} , ET, and detachment	61	<5
R10 ^g	$\equiv\text{Fe}^{\text{II}} + \text{bulk} \rightarrow \equiv\text{Fe}^{\text{III}}-\text{O}-\text{Fe}^{\text{II}}$	re-formation of surface site with adsorbed Fe^{II}	1×10^{10}	1×10^{10}
R11 ^g	$\equiv\text{Fe}^{\text{II}} + \text{bulk} \rightarrow \equiv\text{Fe}^{\text{III}}-\text{O}-^{57}\text{Fe}^{\text{II}}$	re-formation of surface site with adsorbed $^{57}\text{Fe}^{\text{II}}$	1×10^{10}	1×10^{10}
R12 ^h	$\equiv\text{Fe}^{\text{III}}-\text{O}-^{57}\text{Fe}^{\text{II}} + \text{phen} \rightarrow \equiv\text{Fe}^{\text{III}} + ^{57}\text{Fe}^{\text{II}}\text{phen}$	desorption of ^{57}Fe with phen (adjusted to desorption of 51% $\text{Fe}(\text{II})$)	20 (Gt) 120 (Lp)	
		initial concentration of active surface sites ($[\equiv\text{Fe}^{\text{III}}]_0$)	8.5 μM	8.5 μM

^aIn the surface complexation reactions, the type of surface complex (e.g., monodentate and/or bidentate) is not specified. Surface hydroxyl groups are thus omitted and reactions are not balanced for OH^- , H^+ , and H_2O . For example, in R1, $\equiv\text{Fe}^{\text{III}} + \text{L} \rightarrow \equiv\text{Fe}^{\text{III}}\text{L}$ represents the sum of surface complexation reactions, such as $\equiv\text{Fe}^{\text{III}}\text{OH} + \text{HL} \rightarrow \equiv\text{Fe}^{\text{III}}\text{L} + \text{H}_2\text{O}$ and $\equiv\text{Fe}^{\text{III}}(\text{OH})_2 + \text{H}_2\text{L} \rightarrow \equiv\text{Fe}^{\text{III}}\text{L} + 2 \text{H}_2\text{O}$. Bonds between Fe ions consisting of several (μ -oxo) and (μ -hydroxo) bonds are represented in simplified form as $\text{Fe}-\text{O}-\text{Fe}$. The optimized initial concentration of surface sites ($[\equiv\text{Fe}^{\text{III}}]$) in the model was 8.5 μM . ^b ^{56}Fe is abbreviated as Fe. All reactions with Fe and ^{57}Fe isotopes in solution and on the surface include the additional of the four possible permutations. For example, in R4-1: $\equiv\text{Fe}^{\text{III}} + ^{57}\text{Fe}^{\text{II}}\text{L} \rightarrow \equiv\text{Fe}^{\text{II}} + ^{57}\text{Fe}^{\text{III}}\text{L}$ also R4-1b: $\equiv\text{Fe}^{\text{III}} + \text{Fe}^{\text{II}}\text{L} \rightarrow \equiv\text{Fe}^{\text{II}} + \text{Fe}^{\text{III}}\text{L}$, R4-1c: $\equiv\text{Fe}^{\text{III}} + \text{Fe}^{\text{II}}\text{L} \rightarrow \equiv\text{Fe}^{\text{II}} + \text{Fe}^{\text{III}}\text{L}$, and R4-1d: $\equiv\text{Fe}^{\text{III}} + ^{57}\text{Fe}^{\text{II}}\text{L} \rightarrow \equiv\text{Fe}^{\text{II}} + ^{57}\text{Fe}^{\text{III}}\text{L}$. The full list of reactions is provided in Table S3. ^c K and k are fitted equilibrium constants and rate coefficients (M^{-1} , s^{-1} , and $\text{M}^{-1} \text{s}^{-1}$). Numbers in normal font are fitted values; numbers in italic font are values from previous work, from the literature, or are non-rate-determining rate coefficients. ^dValues from Borer et al. 2009. ^eFitted value for pH 7.0, not determined for pH 6.0 (not critical for model fits at pH 6.0). ^fCalculated with DFOB complex formation constants reported by Kim et al., 2010. ^gSee Figure S8 in the SI for more information. ^hRate coefficients (k_{ET}) for exchange of charge between ^{57}Fe and ^{56}Fe surface sites. Note that $\equiv\text{Fe}^{\text{III}}-\text{O}-\text{Fe}^{\text{II}}$ and $\equiv\text{Fe}^{\text{III}}-\text{O}-^{57}\text{Fe}^{\text{II}}$ do not have to be neighboring sites, only rapid charge transfer between the sites has to be possible. ⁱIn the model, bulk-Lp contains only ^{56}Fe (again, abbreviated as Fe) (reactions R10 and R11). This is consistent with the experimental results, in which the measured ^{56}Fe concentrations are scaled to represent the concentrations of the sum of Fe isotopes in Lp. ^jDesorption of $\text{Fe}(\text{II})$ with phen and formation of $\text{Fe}(\text{II})(\text{phen})_3^{2+}$ is abbreviated in this reaction. It is assumed that the first step (formation of $\text{Fe}(\text{II})\text{phen}^{2+}$) is rate-determining and complexation with two more phen is rapid. Phen is in large excess (100 μM) and the change in the concentration of uncomplexed phen does not affect the desorption kinetics.

value of the surface site concentration over which exchange of charge can take place in the time before and during dissolution in reaction R6. Without exchange of charge ($k_{\text{ET}} < 1 \times 10^{-6} \text{ s}^{-1}$), the model predicts a very quick release of ^{57}Fe after addition of the ligand (dotted orange line in Figure 2B), in disagreement with the data. With values of k_{ET} from 1×10^{-1} to 10^6 s^{-1} and a distribution of the negative charge of 2 μM adsorbed $\text{Fe}(\text{II})$ over 8–9 μM surface sites (each charge in average distributed over 4.0–4.5 sites), we obtain good fits (solid orange line, fit with 8.5 μM surface sites). A distribution of charge over more sites (e.g., 20 μM) leads to a slower release of ^{57}Fe than experimentally observed (dotted orange line). Additional plots with model outputs for charge transfer occurring only before addition of DFOB are shown in Figure S9. The best fits were obtained with ET occurring before and during dissolution.

The validity for distribution of each negative charge over an average of 4.0–4.5 sites is also illustrated by the plot of $[\text{Fe}]_{\text{tracer,diss}}$ as a function of $[\text{Fe}]_{\text{diss}}^*$ (Figure 3). The data for pH 7.0 and 8.5 (for both 20 and 50 μM DFOB and for both carbonate and MOPS buffers) collapse onto a single curve. The offset of the data for pH 6.0 corresponds to the presence of dissolved $\text{Fe}(\text{II})$ at the time of DFOB addition. At the start of the dissolution, the ratio of $^{57}\text{Fe}/^{56}\text{Fe}$ (i.e., the slope of the line for pH 7.0) is 0.23, which means that the release of ^{56}Fe is about 4-fold larger than that of ^{57}Fe , independent of DFOB concentrations. During dissolution, the ratio of $^{57}\text{Fe}/^{56}\text{Fe}$ in solution continually decreases because less and less ^{57}Fe is present at the surface, as new surface sites are

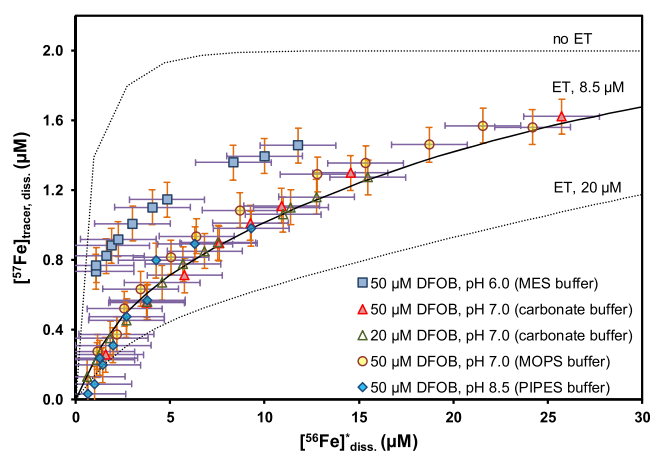


Figure 3. Release of tracer ^{57}Fe in solution as a function of dissolved $^{56}\text{Fe}^*$ (Fe released from Lp) during accelerated Lp dissolution ($t > 1800 \text{ s}$) when 2 μM $^{57}\text{Fe}(\text{II})$ was added 1800 s before DFOB (same data used for 50 μM DFOB treatment at pH 6.0, pH 7.0 (carbonate-buffered), and pH 8.5, as shown in Figure 2). The solid and dotted lines represent the kinetic model fits to the data. The solid line shows the fit to the ^{57}Fe data at pH 7.0 with the optimized concentration of surface sites (8.5 μM) for ET between $\equiv\text{Fe}^{\text{III}}-\text{O}-^{57}\text{Fe}^{\text{II}}$ and $\equiv\text{Fe}^{\text{III}}-\text{O}-\text{Fe}^{\text{II}}$ sites. (R6 in the kinetic model). The dotted lines show release of ^{57}Fe that is too fast with no ET, or too slow with ET over more surface sites (20 μM). Data for 50 μM DFOB buffered with MOPS and 20 μM DFOB buffered with carbonate are shown in SI Figure S6.

formed from bulk ^{56}Fe (R11). The plots in Figure 3 suggest that adsorption of Fe(II) to Lp leads to exchange of charge on the surface before and during dissolution, largely independent of ligand concentrations. As the model fits (solid lines in Figure 2A,B) show, the model is able to explain our observed dissolution rates in the absence and presence of Fe(II) and the isotope exchange and release of ^{57}Fe and ^{56}Fe during dissolution.

The values for the rate coefficient for accelerated dissolution in reactions R8 and R9 (after adsorption of Fe(II) and transfer of negative charge from dissolved Fe(II)DFOB to the surface) of $61\text{ M}^{-1}\text{ s}^{-1}$ are in the range, but somewhat lower than the value for $k_{\text{C,app}}$ of $93\text{ M}^{-1}\text{ s}^{-1}$ obtained in our empirical model, because reactions R4-1 and R4-2 in the more complex model also contribute to the dissolution. The empirical model does overpredict the rate of dissolution at later time points (e.g., results for $1\text{ }\mu\text{M}$ Fe(II) and $50\text{ }\mu\text{M}$ DFOB in Figure 1B). Kim et al.⁴⁹ reported oxidation of Fe(II) in autodecomposition reactions of Fe(II)DFOB, but this reaction is too slow to affect Fe(II)-catalyzed dissolution on the time scale of our experiments and inclusion of oxidation reactions did not lead to significantly improved overall fits. Thus, we attribute the apparent decrease in the dissolution rate at longer times to slow oxidation of Fe(II) by low residual O_2 concentrations.

^{57}Fe Isotope Exchange and Gt Dissolution. Several previous studies^{33,37,50} reported ^{57}Fe (II) isotope exchange with Gt in the absence of ligand over time scales of days to months, leading to ET and isotope exchange between aqueous ^{57}Fe (II) and solid ^{56}Fe . Here, we examine isotope exchange only on shorter time scales of minutes to hours, before and during Gt dissolution.

When ^{57}Fe (II) was added before DFOB, $[^{57}\text{Fe}]_{\text{tracer,diss.}}$ (solid orange squares) decreased over time at pH 7.0 as shown in Figure 4. Since adsorption was almost complete within 1800 s, formed $[^{56}\text{Fe}]_{\text{diss.}}^*$ was too low for accurate determination before addition of DFOB.

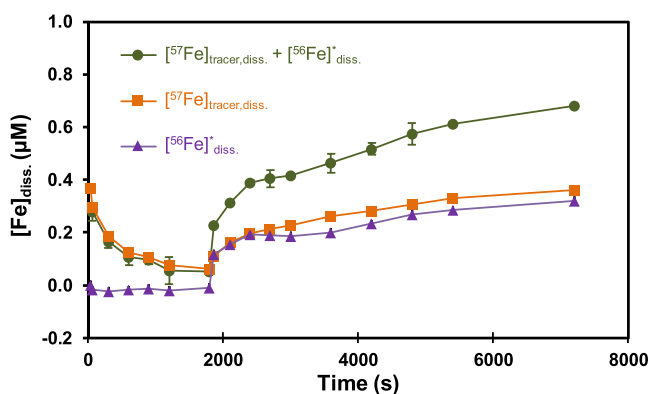


Figure 4. ^{57}Fe isotope exchange and goethite dissolution at pH 7.0 (carbonate-buffered) under anoxic conditions. ^{57}Fe (II) ($2\text{ }\mu\text{M}$) was added to a goethite suspension ($1125\text{ }\mu\text{M}$) 1800 s before $50\text{ }\mu\text{M}$ DFOB addition. The error bars correspond to the range of duplicate measurements. Data points under negative values represent measurements which were below the detection limit. The lines serve as visual guides. Symbols: purple triangles: concentration of Fe released into solution by goethite dissolution ($[^{56}\text{Fe}]_{\text{diss.}}^*$); orange squares: dissolved concentration of tracer ^{57}Fe corrected for the natural abundance of ^{57}Fe in goethite ($[^{57}\text{Fe}]_{\text{tracer,diss.}}$); green circles: sum of dissolved concentration of Fe measured as $[^{57}\text{Fe}]_{\text{tracer,diss.}}$ and $[^{56}\text{Fe}]_{\text{diss.}}^*$ representing the total dissolved Fe concentration.

Addition of $50\text{ }\mu\text{M}$ DFOB led to accelerated dissolution, shown by increasing $[^{56}\text{Fe}]_{\text{diss.}}^*$ (filled purple triangles) and total dissolved Fe (green filled circles). Notably, $[^{57}\text{Fe}]_{\text{tracer,diss.}}$ increased at a similar rate to $[^{56}\text{Fe}]_{\text{diss.}}^*$, in contrast to Lp where $[^{56}\text{Fe}]_{\text{diss.}}^*$ increased by a factor of around 4 faster than $[^{57}\text{Fe}]_{\text{tracer,diss.}}$. This indicates that within our experimental time scale, the extent of isotope exchange was much smaller for Gt than for Lp. As also reported in our previous work,²⁸ dissolution of Gt is slower than that of Lp at pH 7.0, at a rate which is in good agreement with our previously reported rate (Figure S7). We did not attempt to apply our kinetic model to dissolution of Gt because the dissolution was slow and the concentration of dissolved Fe remained below the added Fe(II) concentration. The measured concentrations could thus be the result of desorption and Fe from dissolution of Gt, but the concentrations are too low to separate the two processes. Experiments over longer time scales were not attempted due to slow oxidation of Fe(II) by residual oxygen, which could not be completely prevented in experiments requiring an anoxic atmosphere with 2% CO_2 .

^{57}Fe Isotope Exchange without Net Dissolution. Phen was added after adsorption of ^{57}Fe (II) to desorb Fe(II) and to allow analysis of its isotopic composition. Control experiments were performed with phen added before ^{57}Fe (II).

Lp. As shown in Figure 5A, $2\text{ }\mu\text{M}$ ^{57}Fe (II) added before phen was almost completely adsorbed within 1800 s at pH 7, consistent with the data shown in Figure 2B. Upon addition of $100\text{ }\mu\text{M}$ phen, $1.0\text{ }\mu\text{M}$ total Fe(II) was desorbed, of which $\sim 20\%$ was ^{57}Fe (II) and $\sim 80\%$ was ^{56}Fe (II). Although only 51% of the added Fe(II) was desorbed, the ratio of $[^{56}\text{Fe}(\text{II})]_{\text{diss.}}^*/[^{57}\text{Fe}]_{\text{tracer,diss.}}$ of ~ 4 indicates that extensive charge transfer occurred and the ratio is comparable to the ratio of 4.0–4.5 at the start of accelerated Lp dissolution with DFOB. The solid lines show model fits with ET between $\equiv\text{Fe}^{\text{III}}-\text{O}-^{57}\text{Fe}^{\text{II}}$ and $\equiv^{57}\text{Fe}^{\text{III}}-\text{O}-\text{Fe}^{\text{II}}$ over the optimized concentration of surface sites ($8.5\text{ }\mu\text{M}$). The model correctly reproduces the relative concentrations of desorbed ^{57}Fe and ^{56}Fe . (To account for the incomplete desorption of adsorbed Fe(II) by phen, model output concentrations were multiplied by 0.51 for Lp.) Incomplete desorption can be due to ET to surface sites where Fe(II) is strongly bound and is not desorbed by addition of phen. In a control experiment with ^{57}Fe (II) added after phen, all added ^{57}Fe (open circles) remained in solution.

Gt. Very different results were obtained with Gt, as shown in Figure 5B. ^{57}Fe (II) added before phen was again almost completely adsorbed, with no detectable release of ^{56}Fe . Addition of phen leads to release of adsorbed Fe(II), although more slowly than with Lp. In contrast to Lp, more ^{57}Fe (II) than ^{56}Fe was released. From 4000 to 9000 s, $0.85\text{ }\mu\text{M}$ total Fe(II) was desorbed, of which 76% was ^{57}Fe (II) and 24% was ^{56}Fe (II). The ratio of $[^{56}\text{Fe}(\text{II})]_{\text{diss.}}^*/[^{57}\text{Fe}]_{\text{tracer,diss.}}$ was 0.31, in contrast to the ratio in Lp of 4.0–4.5. These contrasting results show that much less charge transfer from added ^{57}Fe to ^{56}Fe in the solid must have occurred with Gt than with Lp. As shown by the solid lines (model output multiplied by 0.40 to account for incomplete desorption), the model can reproduce the data for Gt if we assume lower rate coefficients for ET in Gt (e.g., $k_{\text{ET}} = 1.2 \times 10^{-4}\text{ s}^{-1}$ in the fit shown) than in Lp. Slow ET would allow only limited charge equilibration in the 1800 s before addition of phen. This contrasts with ET in Lp, where

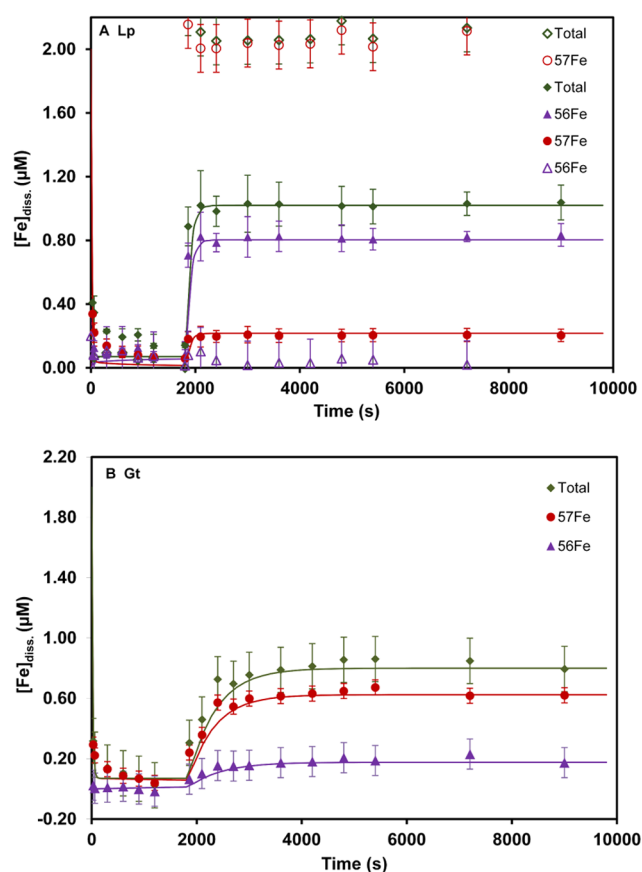


Figure 5. Adsorption and desorption of $^{57}\text{Fe}(\text{II})$ (tracer) and $^{56}\text{Fe}(\text{II})$ on (A) Lp and (B) Gt at pH 7.0 (carbonate-buffered) under anoxic conditions. $^{57}\text{Fe}(\text{II})$ ($2\ \mu\text{M}$) was added to the suspensions ($1125\ \mu\text{M}$) 1800 s before (filled symbols) or after (empty symbols) $100\ \mu\text{M}$ phenanthroline (phen) addition. Data points with negative values represent measurements below the detection limit. The error bars correspond to the range of duplicate measurements. The lines show model fits with ET between $8.5\ \mu\text{M}$ surface sites. Modeled concentrations were multiplied by 0.51 for Lp and by 0.40 for Gt to account for the incomplete desorption of adsorbed $\text{Fe}(\text{II})$ by phen. With Lp, mostly ^{56}Fe was released back into solution after addition of phen. In contrast, mostly ^{57}Fe was released back into solution with Gt. This can be rationalized with rate coefficients ($>1 \times 10^{-1}\ \text{s}^{-1}$) for ET leading to rapid equilibration of charge among surface sites in Lp and slower ET and incomplete equilibration after 1800 s in Gt (the fits shown in (B) were obtained with $k_{\text{ET}} = 1.2 \times 10^{-4}\ \text{s}^{-1}$).

$k_{\text{ET}} > 1 \times 10^{-1}\ \text{s}^{-1}$ leads to full equilibration of charge between ^{57}Fe and ^{56}Fe over $8.5\ \mu\text{M}$ surface sites.

The rate for charge equilibration and isotope exchange on the surface of Lp and Gt is dependent not only on the rate of ET in reaction of R6 but also on the rates of adsorption and desorption of $\text{Fe}(\text{II})$ in reactions R5 and R7. To fit the over 90% adsorption of $\text{Fe}(\text{II})$ within 1–5 min, we used rate coefficients in the range of 1×10^3 to $1 \times 10^4\ \text{M}^{-1}\ \text{s}^{-1}$ (times for adsorption of 63% of the added $\text{Fe}(\text{II})$ of 12–120 s with $8.5\ \mu\text{M}$ surface sites) and of 1×10^{-2} to $1 \times 10^{-3}\ \text{s}^{-1}$ (times for desorption of 63% of adsorbed $\text{Fe}(\text{II})$ of 100–1000 s). We would like to point out that our model is a single-site surface model and thus rests on the assumption that one type of surface site can capture the average properties of many different surface sites with a range of properties regarding adsorption/desorption rates and equilibria and rates of ET between sites. The complexities of atom exchange on the

surface and in the bulk of goethite as probed by atom tomography are presented and discussed in two recent articles.^{51,52} These studies show that atom exchange is spatially heterogeneous and occurs in exchange fronts that can penetrate several nanometers deep into the lattice, and is partly consistent with defect-accelerated exchange. Although our simple model can thus not be expected to capture all features of the experimental results, it explains our overall observations well and supports the proposed mechanisms.

Possible Mechanism for Lp Dissolution and Pathways for ^{57}Fe Isotope Exchange. *DFOB Added First.* In the presence of the ligand, interaction of dissolved $^{57}\text{Fe}(\text{II})\text{DFOB}$ with the surface leads fast transfer of charge to surface $\text{Fe}(\text{III})$ and formation of dissolved $^{57}\text{Fe}(\text{III})\text{DFOB}$. This proposed mechanism is supported by the observation that virtually all of the ^{57}Fe added after DFOB remains in solution. Excess dissolved ligand can then complex surface-bound $\text{Fe}(\text{II})$, causing fast ET to a neighboring $\text{Fe}(\text{III})$ and subsequent detachment of $\text{Fe}(\text{III})\text{DFOB}$. The ET and detachment steps can repeat until all of the free DFOB is consumed.

$^{57}\text{Fe}(\text{II})$ Added First. Charge can be transferred along the surface layer and possibly also into the bulk solid. Addition of DFOB again leads to complexation of $\text{Fe}(\text{II})$ at the surface, fast transfer of charge to neighboring $\text{Fe}(\text{III})$ sites, and detachment of $\text{Fe}(\text{III})\text{DFOB}$. The ratio of $[\text{Fe}]_{\text{tracer,diss.}}$ to $[\text{Fe}]_{\text{diss.}}^*$ corresponds to the fraction of the charge that resides on ^{56}Fe and ^{57}Fe at the Lp surface. The model fits indicate that $^{56}\text{Fe}(\text{II})$ is 4.0–4.5 times more abundant than $^{57}\text{Fe}(\text{II})$ at the surface.

Assuming equal probabilities for charge distribution over available surface sites, this would mean that each unit of negative charge is on average distributed over 4.0–4.5 Fe sites. The release of ^{57}Fe at pH 7.0 with DFOB occurred more slowly than at pH 6.0 with EDTA as reported in our recent study with EDTA, indicating that charge distribution at pH 7.0 before and during dissolution was more extensive than at pH 6.0. This can be rationalized by the nearly complete adsorption of $\text{Fe}(\text{II})$ at pH 7.0, compared to only around 20% adsorption at pH 6.0. Even after dissolution of $38\ \mu\text{M}$ $\text{Fe}(\text{III})$ with DFOB, which is over four times the concentration of the fitted concentration of surface sites, we observed the reappearance of only 85–90% of added ^{57}Fe in solution. The model accounts for this by “dilution” of ^{57}Fe sites through the regeneration of predominantly ^{56}Fe surface sites from the bulk solid.

ENVIRONMENTAL SIGNIFICANCE

We found that in anoxic carbonate-buffered suspensions at pH 7.0, micromolar concentrations of added $\text{Fe}(\text{II})$ can accelerate the rates of Lp dissolution with DFOB up to a factor of 60 under anoxic conditions. Since aquatic systems are generally carbonate-buffered, it is likely that the acceleration of ligand-controlled dissolution by traces of $\text{Fe}(\text{II})$ in natural waters would be at least as large as observed in laboratory experiments with Good’s buffers. Our results are most relevant for soils and surface waters with pH values around 7. Recent analytical developments have demonstrated that strong organic ligands found in the ocean contain compounds with structures similar to DFOB.⁵³ Thus, the mechanisms identified for DFOB may also be important for ligands secreted by marine biota.

The interfacial $\text{Fe}(\text{II})/\text{Fe}(\text{III})$ ET plays a crucial role in numerous natural processes, for example in Fe cycling and bioavailability, and in trace element incorporation and release. We show that rates for ET and isotope exchange during

sorption and accelerated dissolution are very different for Lp and Gt, presumably reflecting the difference in structure and mineralogy. We suggest that future studies are extended to different Fe(III)(oxyhydr)oxide phases. Fe(III) phases formed under natural conditions are often less crystalline and contain co-precipitated cations and anions.⁵⁴ Future work should test if these phases are also susceptible to Fe(II)-accelerated dissolution. The application of ⁵⁷Fe as a tracer in dissolution experiments can provide relevant information about the fate of Fe(II), and the mobility and effect of negative charge on the dissolution of Fe(III) phases in the presence of ligands and low concentrations of Fe(II).

■ ASSOCIATED CONTENT

Supporting Information

The Supporting Information is available free of charge at <https://pubs.acs.org/doi/10.1021/acs.est.9b04235>.

List of chemicals; list of experiments and experimental conditions; kinetic model with full list of reactions; speciation of Fe(II) in carbonate-buffered suspensions; speciation of Fe(II) and control experiments with phenanthroline; ⁵⁷Fe isotope exchange experiments in carbonate- and MOPS-buffered suspensions; Gt dissolution; speciation of Fe and DFOB; and output of kinetic model with different charge distributions between ⁵⁷Fe and ⁵⁶Fe surface sites (PDF)

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