

SCIENTIFIC REPORTS



OPEN

Effect of cation substitution on bridgmanite elasticity: A key to interpret seismic anomalies in the lower mantle

Hiroshi Fukui^{1,2}, Akira Yoneda³, Akihiko Nakatsuka⁴, Noriyoshi Tsujino³, Seiji Kamada^{5,6}, Eiji Ohtani^{5,7}, Anton Shatskiy⁷, Naohisa Hirao⁸, Satoshi Tsutsui⁸, Hiroshi Uchiyama^{2,8} & Alfred Q. R. Baron^{2,8}

Seismological observations show that, in some regions of the lower mantle, an increase in bulk sound velocity, interestingly, occurs in the same volume where there is a decrease in shear velocity. We show that this anti-correlated behavior occurs on cation substitution in bridgmanite by making single crystal elasticity measurements of MgSiO_3 and $(\text{Mg,Fe,Al})(\text{Si,Al})\text{O}_3$ using inelastic x-ray scattering in the ambient conditions. Cation substitution of ferrous iron and aluminum may explain large low shear velocity provinces in the lower mantle.

Bridgmanite, or, *Pbnm*-type magnesium-silicate perovskite, is the dominant mineral in the Earth's lower mantle. Materials with perovskite or related structures also attract broad attention since they can display novel physical properties such as colossal magnetoresistance¹, multiferroicity², and high-temperature superconductivity³. At pressures over 125 GPa (corresponding to depths more than ~2700 km) and at temperature greater than 2500 K, bridgmanite transforms to a post-perovskite (pPv) phase⁴ with the *Cmcm*-type CaIrO_3 structure. It is widely believed that pPv is the main component of the *D''* layer at the bottom of the lower mantle, which is 200 km thick just above the core mantle boundary (~2900 km depth).

In the deep mantle, between 2000 and 2891 km in depth, some regions show an increase in bulk sound velocity ($V_B = \sqrt{K_S/\rho}$), and a decrease in shear wave velocity ($V_S = \sqrt{G/\rho}$): $\Delta V_B > 0 > \Delta V_S$, and others show a decrease in V_B and an increase in V_S : $\Delta V_B < 0 < \Delta V_S$ ^{5,6} (K_S , G , and ρ are adiabatic bulk modulus, shear modulus, and density, respectively). This feature is called an anti-correlated seismic velocity anomaly. It is reported that the phase transformation of $(\text{Mg,Fe,Al})(\text{Si,Al})\text{O}_3$ from *Pbnm*-type to *Cmcm*-type can explain the increase in V_S and decrease in V_B from the average ($\Delta V_B < 0 < \Delta V_S$) in some deeper regions⁷. However, this cannot explain the anomaly in the shallower part of the mantle where the pPv phase is not stable. More importantly, it is difficult to interpret the anti-correlated nature of the anomaly where ΔV_B and ΔV_S have opposite signs. The regions showing this anomaly, which are beneath Africa and the central Pacific, attract attention as large low shear velocity provinces (LLSVPs).

The origin of the LLSVPs is under debate. Thermal heterogeneity has been considered⁸, but exclusively thermal effects are insufficient to explain the LLSVPs because usually both V_B and V_S decrease with temperature. It is thus suggested that the LLSVPs have very different chemical composition from that of the average mantle⁹ due to accumulations of subducted oceanic slabs¹⁰, remnants of Earth's early magma ocean¹¹, or even chemical reactions

¹Center for Novel Material Science under Multi-Extreme Conditions, Graduate School of Material Science, University of Hyogo, 3-2-1 Kouto, Kamigori, Hyogo 678-1297, Japan. ²Materials Dynamics Laboratory, RIKEN SPring-8 Center, RIKEN, 1-1-1 Kouto, Sayo, Hyogo 689-5148, Japan. ³Institute for Planetary Materials, Okayama University, 827 Yamada, Misasa, Tottori 682-0193, Japan. ⁴Graduate School of Sciences and Technology for Innovation, Yamaguchi University, 2-16-1 Tokiwadai, Ube, Yamaguchi 755-8611, Japan. ⁵Graduate School of Science, Tohoku University, 6-3 Aramaki, Aoba, Sendai, Miyagi 980-8578, Japan. ⁶Frontier Research Institute for Interdisciplinary Sciences, Tohoku University, 6-3 Aramaki, Aoba, Sendai, Miyagi 980-8578, Japan. ⁷V. S. Sobolev Institute of Geology and Mineralogy, Siberian Branch, Russian Academy of Sciences, Novosibirsk 630090, Russia. ⁸Research and Utilization Division, Japan Synchrotron Radiation Research Institute, SPring-8, 1-1-1 Kouto, Sayo, Hyogo 689-5198, Japan. Correspondence and requests for materials should be addressed to H.F. (email: fukuih@sci.u-hyogo.ac.jp)

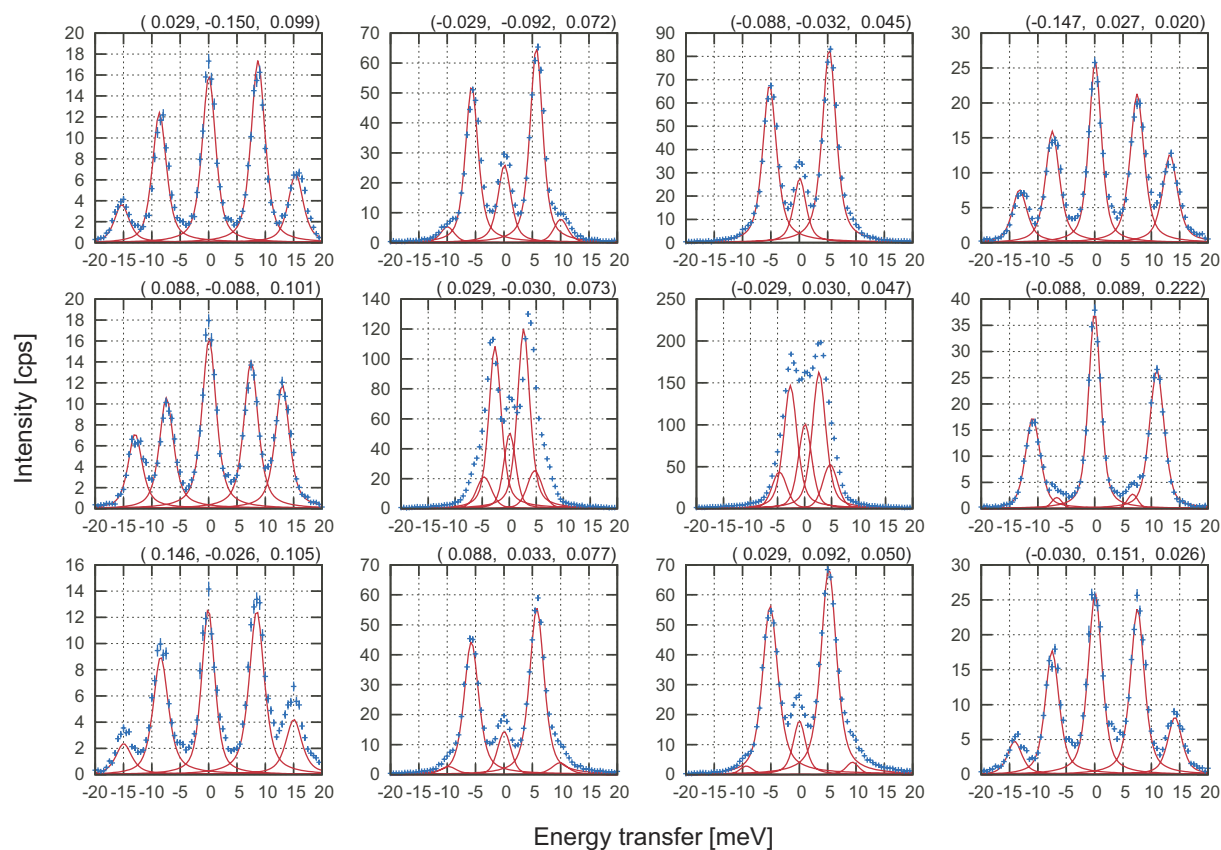


Figure 1. A representative set of IXS spectra of (Fe,Al)-Bdg collected at BL35XU. Twelve spectra can be measured at once. Reduced \mathbf{q} positions in reciprocal lattice unit are shown in parentheses. A total momentum transfer was $(2, -2, 0) + \mathbf{q}$ in reciprocal lattice unit. Each spectrum show one or two pairs of clear phonon signals. Blue crosses show data points. Red lines show pseudo-Voigt functions fitted to peaks.

with the core¹². Recently primordial metallic melt trapped in the mantle was suggested as the nature of LLSVPs¹³. A complicated model², including multiple chemical and thermal effects, can reproduce the distribution of the LLSVPs. But this model requires rather a specific distribution of effects that are not internally well correlated. Houser¹⁴ suggested that slow V_S might be correlated with temperature and chemical anomaly using the parameter set for bridgmanite¹⁵, as was used in ref. 6, but did not discuss the anti-correlated anomaly between V_B and V_S . The theoretical result¹⁵ used in both seismological studies^{6,14} shows anti-correlation only in elastic moduli but not in velocities, and, more importantly, has not yet been experimentally verified.

In order to address these issues, we investigated the elastic properties of single-crystal bridgmanite at ambient conditions. Although Brillouin light scattering (BLS) is frequently used to determine elastic properties of high-pressure minerals, the elasticity of iron-bearing bridgmanite has not been determined by BLS due to its opacity, and its instability against strong optical laser irradiation. We used inelastic x-ray scattering (IXS) technique in this study. We prepared two types of bridgmanite: MgSiO_3 (Mg-Bdg, hereafter) and $\text{Mg}_{0.943}\text{Fe}_{0.045}\text{Al}_{0.023}\text{Si}_{0.988}\text{O}_3$ ((Fe,Al)-Bdg hereafter). Iron in (Fe,Al)-Bdg was confirmed by synchrotron Mössbauer spectroscopy to be in high-spin ferrous state and to occupy a large A site of perovskite structure. The sample characterization and more details of the IXS measurements are given in the Methods section. Elastic stiffness tensors (C_{ij}) for Mg-Bdg, and for (Fe,Al)-Bdg was obtained from analysis of IXS spectra based on Christoffel's equation¹⁶. A typical set of IXS spectra is shown in Fig. 1. The elastic moduli obtained are listed in Table 1 together with literature values^{15,17–23}. The velocity surface plots of Mg-Bdg from the present C_{ij} determined from two sets of IXS measurements are shown in Fig. 2 together with those calculated from BLS results^{17–19}. The patterns of the velocity surfaces are similar to each other: the longitudinal velocity is the fastest along about b axis and minimum along c axis, etc. The absolute values determined using IXS are generally smaller than those from BLS.

The pattern of the velocity surface of (Fe,Al)-Bdg is basically similar to that of Mg-Bdg (red and blue lines in Fig. 2). The present cation substitution affected the velocity surface as follows: 1. V_p along the b and c axes is increased; 2. the average V_S along the b axis is decreased; 3. the difference of V_S along the a and c axes is increased and decreased, respectively. Crystallographic studies²⁴ report that iron substitution enlarges the a axis more than other axes, which is consistent with the present result (see Method section). The large elongation of the a axis probably results in the least change in V_p along a axis. Thus qualitatively the velocity surfaces indicate that elastic anisotropy in bridgmanite increases with the present cation substitution. More quantitatively, the acoustic anisotropy defined by $2 \times (V_{\max} - V_{\min}) / (V_{\max} + V_{\min})$ increases from 8.81% to 8.92% for longitudinal waves, and from

	C_{11}	C_{22}	C_{33}	C_{44}	C_{55}	C_{66}	C_{12}	C_{13}	C_{23}	K_S	G	V_B	V_S	ρ
IXS ^a	450 (7)	498 (9)	420 (8)	190 (6)	168 (5)	148 (4)	117 (8)	125 (8)	137 (9)	236 (4)	166 (2)	7.58 (6)	6.37 (4)	4.103
IXS ^b	448 (6)	519 (5)	438 (5)	178 (3)	180 (4)	138 (3)	124 (5)	113 (8)	160 (5)	244 (3)	165 (1)	7.66 (5)	6.32 (2)	4.140
BLS ¹⁷	515 (5)	525 (5)	435 (5)	179 (4)	202 (3)	175 (4)	117 (5)	117 (5)	139 (6)	246.4	184.2	7.74	6.69	4.108 ^c
BLS ¹⁸	482 (4)	537 (3)	485 (5)	204 (2)	186 (2)	147 (3)	144 (6)	147 (6)	146 (7)	264.0	177.3	8.02	6.57	4.108 ^c
BLS ¹⁹	481 (4)	528 (3)	456 (4)	200 (2)	182 (2)	147 (2)	125 (3)	139 (3)	146 (3)	253 (3)	175 (2)	7.84	6.52	4.112 ^c
LDA ¹⁵	491	554	474	203	176	153	134	139	152	263	178	7.94	6.53	4.174 ^c
	488	543	469	193	173	135	147	148	164	268	168	7.74	6.13	4.471 ^d
GGA ²²	438	488	422	182	163	134	118	122	136	233	160	7.65	6.34	3.985 ^c
GGA-DFPT ²³	494 (2)	511 (2)	426 (2)	193 (1)	176 (1)	151 (1)	109 (1)	137 (1)	149 (1)	247	172	7.88	6.58	3.973 ^c
RUS ²⁰	—	—	—	—	—	—	—	—	—	238	174	7.61	6.51	4.106 ^c
UI ²¹	—	—	—	—	—	—	—	—	—	247	176	7.75	6.54	4.110 ^c
	—	—	—	—	—	—	—	—	—	236	174	7.53	6.47	4.161 ^c

Table 1. Elastic moduli and density of bridgmanite at ambient conditions. The units are GPa, km/s, and g/cm³ for elastic moduli, velocities and density, respectively. ^aPresent study (Mg-Bdg). ^bPresent study ((Fe,Al)-Bdg). ^cMgSiO₃. ^dMg_{0.75}Fe_{0.25}SiO₃. ^eMg_{0.95}Fe_{0.05}SiO₃.

12.4% to 13.4% for transverse waves. The present results experimentally demonstrate that the degree of anisotropy is increased by the present cation substitution.

The Voigt-Ruess-Hill average of bulk and shear moduli calculated from C_{ij} are listed in Table 1. We determine K_S and G to be 236 and 166 GPa, respectively for Mg-Bdg. The value of K_S in the present study is consistent with that determined from RUS²⁰ and that by a calculational study²², but is lower than the other values by ~15 GPa (6%). G is also smaller by ~10 GPa (also 6%) than those in the previous results. These differences correspond to 3% in velocity. The origin of the differences in K_S and G between two techniques should be further investigated. Nevertheless, this study experimentally demonstrated that the present cation substitution in bridgmanite increases K_S and V_B and decreases G and V_S : an anti-correlated behavior.

This anti-correlated behavior in elastic moduli and velocities by cation substitution has not been reported. Bulk and shear moduli and velocities are summarized in Fig. 3 together with previous results. Previously, the effect of Fe substitution was investigated using ultrasonic interferometry²¹ (UI) and calculations^{15,22}. The sample used in the UI study contained not only Fe²⁺ but also Fe³⁺. The results of the UI study disagree with one calculation¹⁵, where Fe²⁺ substituted for Mg²⁺, but rather agree with another²², where Fe³⁺ and Al³⁺ substituted for Mg²⁺ and Si⁴⁺. These results^{21,22} imply that Fe³⁺ substituting for Mg²⁺ decreases both K_S and G . The effect of aluminum substitution was reported using BLS²⁵ and theoretical calculation²¹. An experimental study²⁵ reported that the substitution of only Al decreases both elastic moduli and slightly increases V_B . A theoretical study²² showed that the substitution of only Al decrease V_B and V_S as well as K_S and G . That study²² also investigated the effect of coupled substitution of Fe³⁺ and Al, demonstrating that the effect of this pair substitution is qualitatively the same as that of the substitution of aluminum only.

Water content sometimes reduces elastic moduli. The present samples contain a certain amount of water (140 and 460 ppm). However, it is not known how much water content affects the elasticity of bridgmanite. If water content decreased shear modulus for bridgmanite, e.g. by 0.3 GPa/100 ppm or shear velocity by 0.02 km/s/100 ppm, the present cation substitution for dry bridgmanite would show a positive correlated behavior, or increase both K_S and G . The anti-correlated behavior observed in this study may be due to a combination of the cation substitution and the water content.

We simply consider the effect of iron and aluminum separately though these substitutions may be coupled. Many investigations have been done about the effect of cation substitution on isothermal bulk modulus of bridgmanite by measuring compression curves. It is well known that K_{T0} and K' derived from a compression-curve fitting are strongly correlated. K_{T0} also depends on the pressure range of the measurement, sample conditions, etc. Nevertheless, the relative change in K_{T0} determined by the same technique is reliable. The effect of a small amount of Fe²⁺ substitution on K_{T0} is reported to be positive^{26–29}. This is qualitatively consistent with the theoretical study¹⁵. In contrast, the effect of aluminum on K_{T0} is still controversial; a positive effect (increasing K_{T0}) is reported in some studies^{26,30} and negative effect (decreasing K_{T0}) in others^{30–32}. Based on the BLS studies^{17–19,25} and the theoretical one²¹, the effect of aluminum substitution on K_S can be considered negative. Note that the theoretical study²² also investigated the effect of coupling substitution of Fe³⁺ and Al, demonstrating that the effect of this pair substitution is qualitatively the same as that of the substitution of aluminum only. The effect of only Fe²⁺ on the velocities can be calculated from the present study by subtracting the effect of Al from the BLS results^{17–19,25}, assuming that the effects of Fe²⁺ and Al are independent. This analysis suggests that Fe²⁺ substitution increases both V_B and V_S (Table 2).

We apply the present results to a geochemical and geothermal model to estimate if this effect is sufficient to explain the LLSVPs. We assume a perovskitic lower mantle³³ for simplicity. The seismic anomaly observed in the LLSVPs (+1 and –1% of $\Delta V_B/V_B$ and $\Delta V_S/V_S$, respectively)^{5,6} may, then be explained by variation of Fe²⁺ and Al substitution into bridgmanite at temperature conditions for 2000–2891 km depth (2250–2450 K³⁴). The temperature effects on V_B and V_S were assumed to be independent of pressure and composition (Table 2). The observed anomaly of +/–1% for V_B and –/+1% for V_S corresponds to the compositional variation between MgSiO₃ and Mg_{0.959}Fe_{0.027}Al_{0.028}Si_{0.986}O₃, +/–2.0 atom% of $\Delta Fe/(Mg + Fe + Si + Al)$ and +/–1.3 atom% of Δ

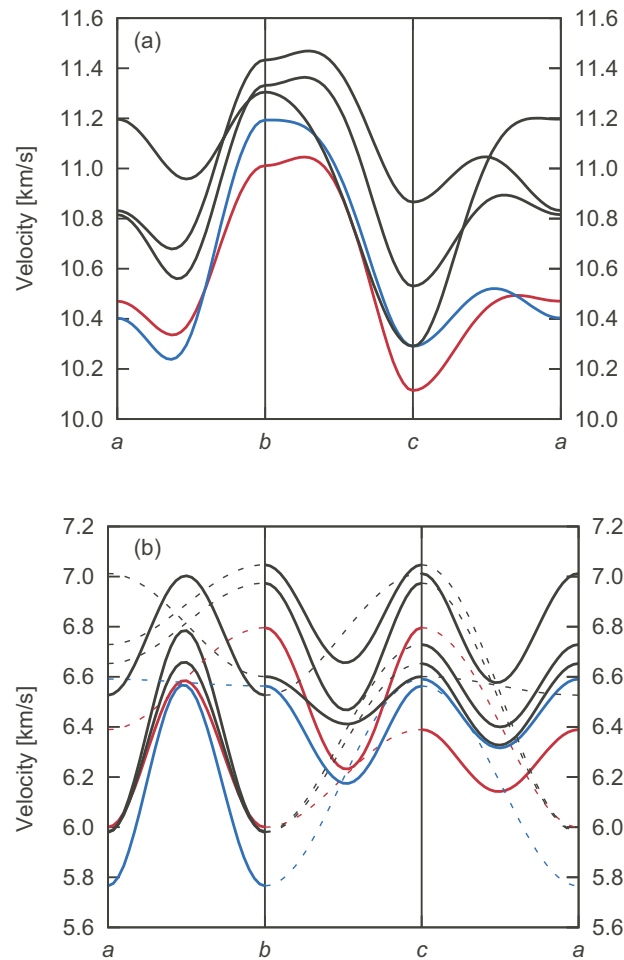


Figure 2. Velocity surface of bridgmanite. Longitudinal (a) and transverse (b) waves from the present study (red lines for Mg-Bdg and blue for (Fe,Al)-Bdg) and by BLS^{17–19} (black lines). Solid and broken lines indicate the polarization vector is in-plane or out-of-plane, respectively, in (b).

Al/(Mg + Fe + Si + Al) in temperature range of 2250–2450 K³⁴. This compositional heterogeneity of bridgmanite then explains the anti-correlated seismic anomaly (Figs 4 and 5AB). This model indicates that cation substitution of a few atomic percent causes an anti-correlated anomaly comparable to that observed in the LSSVPs.

We now consider to include the effect of temperature since the LSSVPs may correlate with local temperature changes. We assume the temperature difference ΔT between the regions with the highest V_B and the average value, i.e. $\Delta T = T(\Delta V_B/V_B = 1\%) - T(\Delta V_B/V_B = 0\%)$. The chemical inhomogeneity, $\Delta X/(Mg + Fe + Si + Al)$ ($X = Fe$ or Al), needed to explain the velocity anomaly is then shown in Fig. 4. Especially when ΔT is about 113 K, the LSSVPs can be explained by only 2.7 atom% of Fe^{2+} substitution without Al variation (Fig. 5AC). More detailed modeling requires ferropericlase and taking the effect of spin transition for these two materials into account.

We have experimentally demonstrated that cation substitutions in bridgmanite enhances elastic anisotropy and causes anti-correlated behavior in elastic wave velocities. This result indicates that seismic anomalies observed in the lower mantle could be explained by chemical heterogeneity in bridgmanite.

Methods

Sample synthesis and characterizations. The single crystals examined in this study were synthesized at 24 GPa and 1500 °C using a Kawai-type multi anvil press (USSA-5000) installed at ISEI, Okayama University³⁵. The isotope ratios of chemical reagents were at natural abundance.

The chemical compositions are confirmed as $MgSiO_3$ and $Mg_{0.943}Fe_{0.045}Al_{0.023}Si_{0.988}O_3$ by an electron microprobe analyzer. The number ratios of $Fe/(Mg + Fe + Si + Al)$ and $Al/(Mg + Fe + Al + Si)$ of this sample are 0.023 and 0.012, respectively. Assuming all Fe is bivalent, the sum of the charge estimated from the EPMA results is -0.003 . This is negligible, taking the uncertainty of the chemical analysis into account. A typical amount of water content of single crystals in the run product was 140 ± 52 and 460 ± 45 ppm according to synchrotron IR absorption analyses³⁵.

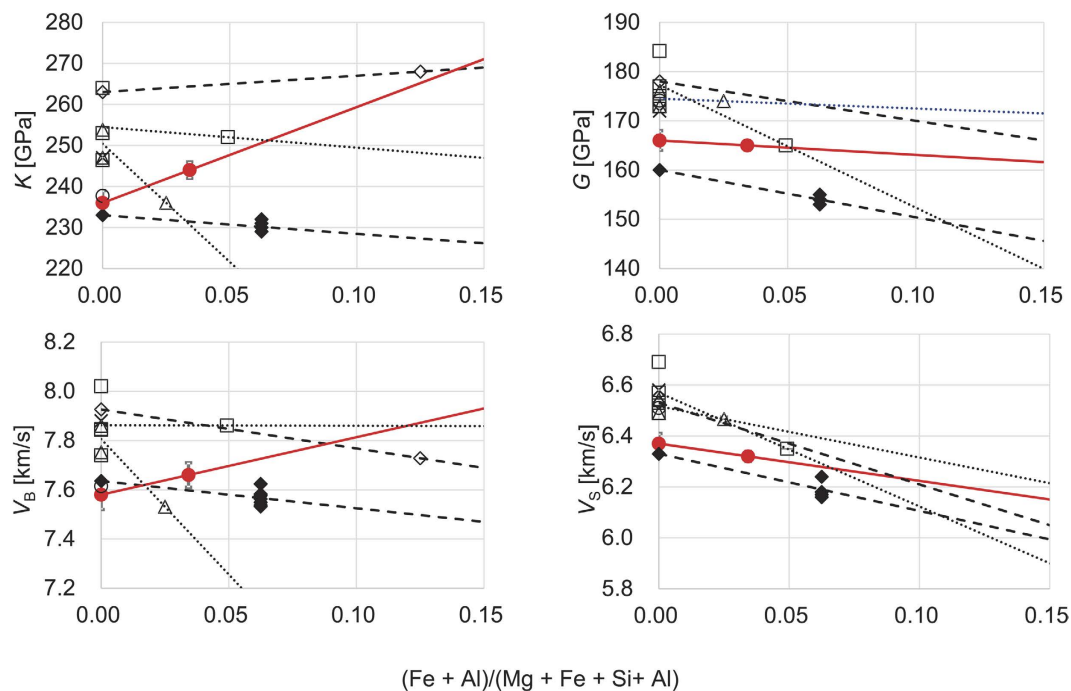


Figure 3. Elastic moduli and elastic velocities of bridgmanite with chemical substitution of Fe and Al for Mg and Si. The present results are shown by solid (red) circles. Open circles, squares, and triangles indicate values determined by RUS²⁰, BLS^{17–19,25} and US^{21,59}. Diamonds indicate theoretical values (open for ferrous iron (LDA)¹⁵ and solid for aluminum and ferric iron (GGA)²²). Crosses indicate values by GGA-DFPT-IXS²³. Lines are guide for eyes (solid (red): present study; dotted: previous experiments; dashed: theory).

dV_B/dFe [km/s] ^{a,*}	5.534
dV_S/dFe [km/s] ^{a,*}	1.040
dV_B/dAl [km/s] ^{b,*}	-0.024
dV_S/dAl [km/s] ^{b,*}	-4.465
dK_S/dT [GPa/K] ^c	-0.021
dG/dT [GPa/K] ^c	-0.028
a [10^{-5} 1/K] ^c	1.98
b [10^{-8} 1/K ²] ^c	0.80

Table 2. Parameters for elastic wave velocity modeling. ^aFrom the present results and BLS^{17–19,25}. ^bFrom BLS^{17–19,25}. ^cFrom Li⁶⁰ (Thermal expansion coefficient is $\alpha = \frac{1}{V} \frac{\partial V}{\partial T} = a + bT$). *Per cation fraction in mole.

The $Fe^{3+}/\Sigma Fe$ ratio of (Fe,Al)-Bdg was evaluated with synchrotron Mössbauer spectroscopy at BL10XU of SPring-8³⁶. An obtained spectrum were analyzed using program MossA³⁷. Without any prejudice, the spectrum seems to consist of two absorption lines with different intensities (Fig. 6). They can be interpreted either as two singlets, as an asymmetric doublet, or as combination of a doublet and a singlet. If the spectrum consists of two singlets, an isomer shift of 1.96(9) mm/s corresponds to that of monovalent high-spin iron. Considering the charge neutrality of the system, it is difficult for Fe^+ to substitute for Mg or Si in perovskite structure. Therefore, the absorption line at 1.96 mm is the higher velocity one of a doublet. Analysis based on an asymmetric doublet gives the isomer shift of 1.05(6) mm/s and quadrupole splitting of 1.8(1) mm/s. These values indicate that iron in this sample was in a divalent high-spin state³⁸ and substitute for magnesium³⁹. The higher intensity at the lower velocity side is attributed either to that the sample was a single crystal or to that iron existed in another state. The former case is more plausible than the latter due to the following reasons. 1. The linewidths determined using two singlets (0.97(16) and 0.76(30) mm/s for lower and higher velocity lines) are consistent within the fitting uncertainty; 2. The line shape of the lower velocity signal looks symmetric and additional singlet/doublet to the lower velocity signal has not improved the fitting quality at all. We off course have paid much attention to possible existence of iron in a trivalent high-spin state, which should give a doublet with an isomer shift of ~ 0.5 mm/s³⁸ and a quadrupole splitting of 0.5–1.0 mm/s³⁹. Since parameter fitting assuming two doublets (one for HS Fe^{2+} and another for HS Fe^{3+}) was not converged, we were not able to detect the amount of ferric iron if existed. The asymmetric doublet is probably attributed to a certain angle between the principal electric field gradient in the Fe site and the incident X-ray beam direction because the sample was a single crystal. The intensity ratio (2.7: 1) indicates the sign of the quadrupole splitting was negative. The linewidth assuming one doublet is 0.93(13) mm/s,

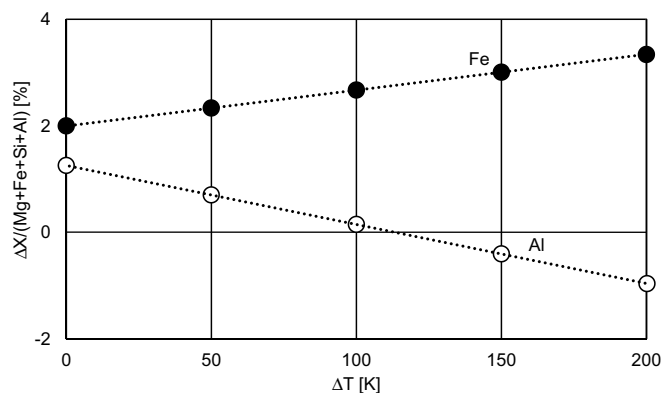


Figure 4. Required $\Delta X/(Mg + Fe + Si + Al)$ ($X = Fe$ or Al) to explain the velocity anomaly of +1% for V_B and -1% for V_S with temperature difference ΔT . Solid and open circles are for Fe and Al, respectively. With increasing ΔT , required ΔFe and ΔAl increases and decreases respectively. When ΔT is around 110 K, the velocity anomaly can be explained only with ΔFe and the distribution of Al should be uniform.

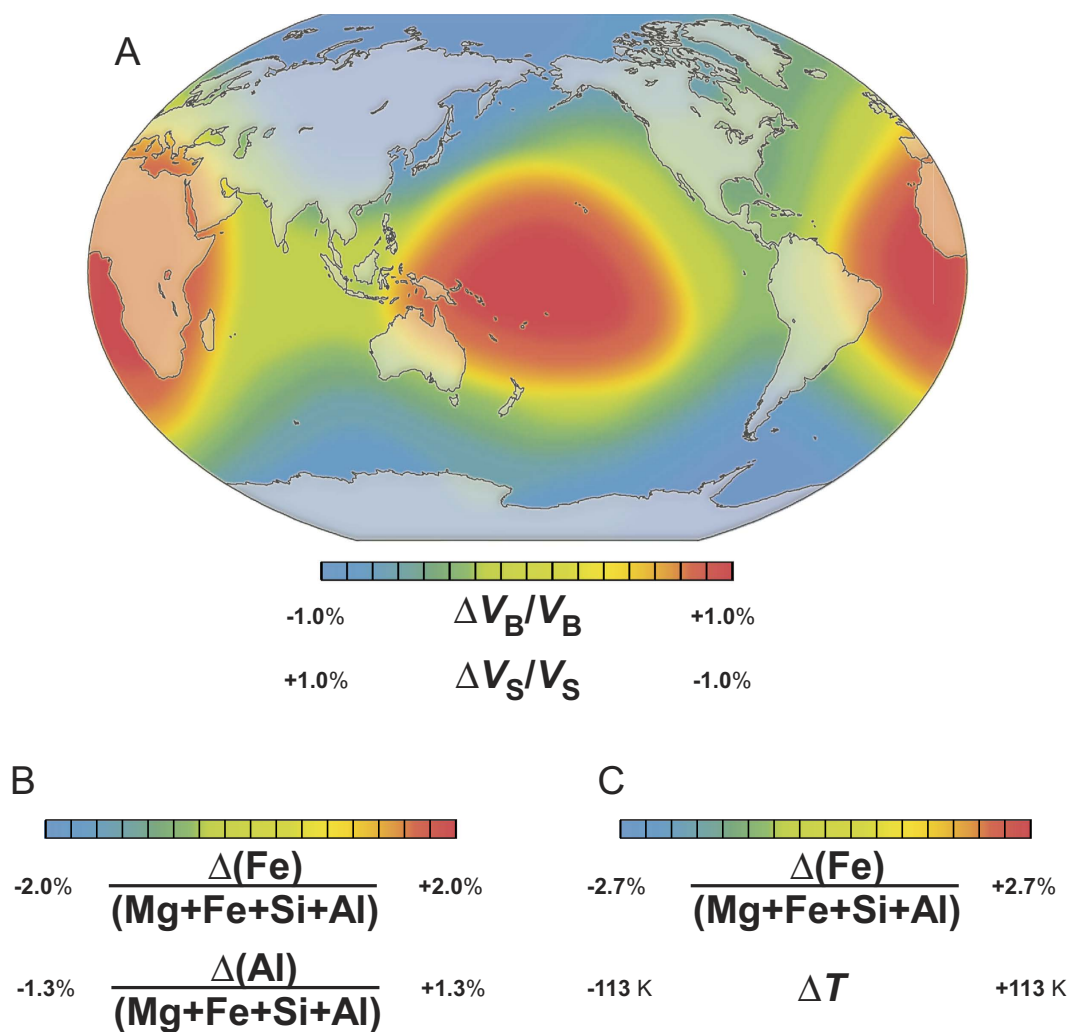


Figure 5. Schematic image of regional variation of seismic velocities, chemical composition of bridgmanite, and temperature variation at depth between 2000 and 2891 km. (A) Red color indicate the anomaly of $\Delta V_B > 0 > \Delta V_S$. The map outline was made using CraftMAP (<http://www.craftmap.box-i.net/>). (B) The case of ΔT (difference from the average temperature) = 0. Fe and Al contents are the highest in the red regions. (C) The case of $\Delta Al/(Mg + Fe + Si + Al) = 0$. Fe content is the highest in the red regions, where temperature is higher than the average temperature by 113 K. The map was colored based on ref. 6.

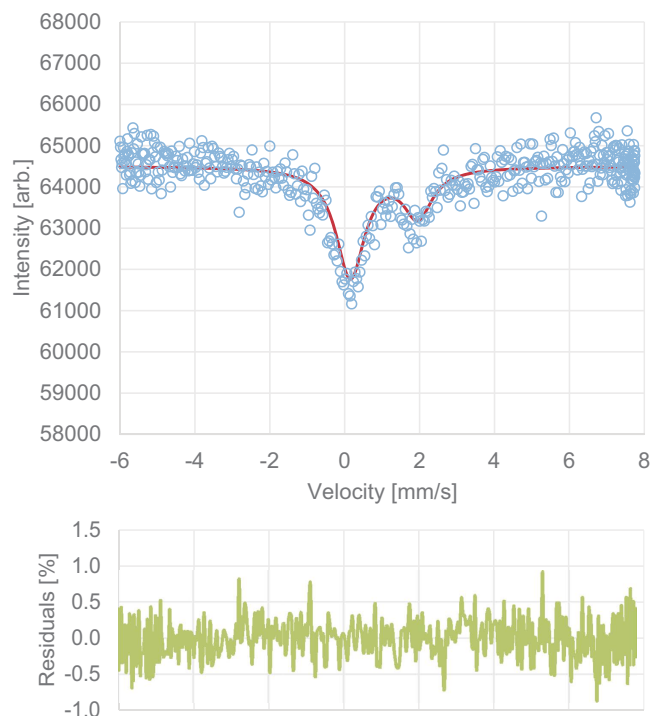


Figure 6. Mössbauer spectrum of ^{57}Fe in (Fe,Al)-Bdg. Blue circles are raw data from which the backgrounds have been subtracted. The red and green lines indicate one doublet fitted to the data and fitting residuals. The isomer shift of 1.05 mm/s and the quadrupole splitting of 1.8 mm/s indicate that the sample contains high-spin Fe^{2+} in the magnesium site^{38,39}.

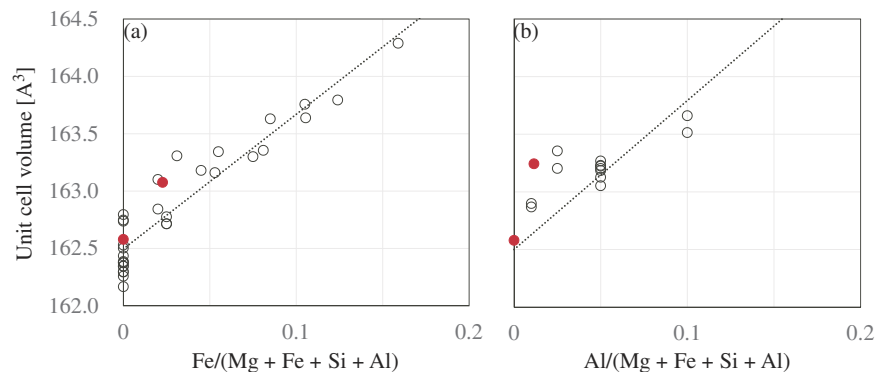


Figure 7. Analytical curves between unit cell volume of bridgmanite and chemical substitutions (a) by iron and (b) by aluminum. The solid (red) circles are the present results. The dotted lines were fitted to literature values shown by open circles^{19–21,24–32,41–53}. (a) $11.741x + 162.5$ and (b) $12.926x + 162.5$, where x is $(\text{Fe} + \text{Al})/(\text{Mg} + \text{Fe} + \text{Si} + \text{Al})$ and V at $x = 0$ is fixed to 162.5 \AA^3 .

which is much broader than a typical energy resolution of the Mössbauer spectrometer at BL10XU (0.43 mm/s). This is perhaps due to variation of the local environment around Fe in this sample given by Mg/Al/Si distribution in neighboring sites, hydrogen, and/or oxygen vacancy. The results of the synchrotron Mössbauer measurement conclude that most iron atoms were in a divalent high-spin state and occupied a magnesium site. Consequently, the simplest substitution model, where the iron substitutes for magnesium and aluminum substitutes for both magnesium and silicon in perovskite structure is consistent with the results of these analyses.

The investigated grains were confirmed to be single domains using a four-circle diffractometer with a laboratory x-ray source at room temperature. The lattice constants a , b , and c were 4.7784(3), 4.9306(4), 6.9005(8) Å and 4.787(1), 4.934(1), 6.904(1) Å for Mg-Bdg and (Fe,Al)-Bdg, respectively. It was reported that the unit cell volume of MgSiO_3 bridgmanite does not change even with 100 ppm water content⁴⁰. An analytical curve drawn by fitting a linear function to literature values^{19–21,24,26–28,30–32,41–53} is shown in Fig. 7. The obtained analytical line for iron substitution is consistent with literature²⁹. The unit cell volumes of Mg-Bdg is larger than the present analytical line by only 0.05%. Since these differences are comparable to the experimental error, the water content of

	C_{11}	C_{22}	C_{33}	C_{44}	C_{55}	C_{66}	C_{12}	C_{13}	C_{23}	K_S	G	V_B	V_S	ρ
BL35XU	445 (6)	489 (9)	417 (6)	187 (4)	159 (3)	148 (5)	117 (8)	122 (5)	137 (7)	233 (3)	163 (2)	7.54 (5)	6.30 (4)	4.103
BL43LXU	452 (4)	502 (4)	428 (6)	185 (4)	181 (4)	142 (2)	122 (4)	127 (6)	140 (6)	240 (2)	167 (1)	7.64 (4)	6.38 (3)	4.103
Whole data	450 (7)	498 (9)	420 (8)	190 (6)	168 (5)	148 (4)	117 (8)	125 (8)	137 (9)	236 (4)	166 (2)	7.58 (6)	6.37 (4)	4.103

Table 3. Elastic moduli and density of Mg-Bdg at ambient conditions determined by IXS. The units are the same as those in Table 1.

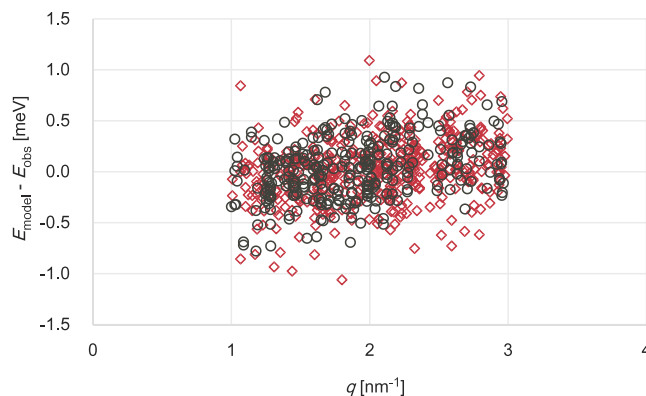


Figure 8. Residuals of the fitting ($\Delta E = E_{\text{model}} - E_{\text{obs}}$). Red diamonds and black circles are for Mg-Bdg and (Fe,Al)-Bdg, respectively.

140 ppm seems to give a negligible effect. In contrast, the unit cell volume of the (Fe-Al)-Bdg is larger than those of the analytical lines by 0.29%. This excess volume may be explained by effect of aluminum and water content. It is known that aluminum incorporation increases the unit cell volume^{26,30–32}. Estimating from the previous results, the value of 0.012 for Al/(Mg + Fe + Al + Si) makes the unit cell volume larger by 0.09%. Although the degree of water effect on the unit cell volume of magnesium silicate perovskite is uncertain, the water content of 460 ppm probably made the unit cell volume larger by 0.20%. The densities of the Mg-Bdg and (Fe,Al)-Bdg are 4103.3 and 4139.5 g/cm³, respectively.

IXS measurement and data analysis. Inelastic X-ray scattering with a single crystal sample in conjunction with an analysis based on Christoffel's equation has been recently used for accurate determination of elastic moduli^{16,54–56}. This technique has been adopted to data along high-symmetry directions about samples at high-pressure and high-temperature conditions^{55,56}. In this study, we did not limit data along high-symmetry directions, but measured rather redundant data at off-symmetry positions to determine C_{ij} precisely and to utilize all measured data with an analyzer array^{16,54} (see Fig. 1). We performed IXS measurements at BL35XU of SPring-8⁵⁷ at 21.747 and 17.794 keV, with which typical energy resolutions were 1.5 and 3.0 meV full-width-half-maximum (FWHM), respectively. 21.747 keV x-ray was used for Mg-Bdg and 17.794 keV for (Fe,Al)-Bdg. The size of the incident X-ray beam was $\sim 70 \mu\text{m}$ in diameter. We performed another measurement for Mg-Bdg to insure the quality of our results. We measured another grain from the same sample growth run at BL43LXU of SPring-8⁵⁸. At BL43LXU, x-ray beam with size of $\sim 20 \mu\text{m}$ and energy of 17.794 keV was used. The energy resolution was 3.0 meV (FWHM).

For each observed phonon mode, the elastic wave velocity was calculated assuming a linear relationship between phonon energy and momentum. Single crystal elasticity at ambient conditions was determined by least-square fitting to the observed velocities using the measured densities. Details of the fitting are given in ref. 16. Phonons with momentum transfers, $|\mathbf{q}|$ from 1 to 3 nm⁻¹ away from Bragg peaks were used for analysis.

The elastic moduli determined from data at BL35XU and BL43LXU are consistent in contrast to different BLS studies which are not so consistent. The individual results are listed in Tables 1 and 3. The different IXS measurement agree to better than 14% (the maximum deviation) whereas those from three BLS studies are spread more (26%, the maximum deviation). Therefore these two sets of IXS data for Mg-Bdg were analyzed as one set (giving 461 modes) to obtain more reliable elastic properties. For (Fe,Al)-Bdg, 319 modes were used. The residuals of the fitting are shown in Fig. 8. There is a slightly linear relationship between ΔE and $|\mathbf{q}|$ observed, probably meaning the assumed linear relationship between ΔE and $|\mathbf{q}|$ is not completely valid at this q range.

References

- Ramirez, A. P. Colossal magnetoresistance. *J. Phys. Cond. Mat.* **9**, 8171 (1997).
- Kimura, T. *et al.* Magnetic control of ferroelectric polarization. *Nature* **426**, 55 (2003).
- Bednorz, J. G. & Müller, K. A. Perovskite-type oxides-The new approach to high- T_c superconductivity. *Rev. Mod. Phys.* **60**, 585 (1998).
- Murakami, M., Hirose, K., Kawamura, K., Sata, N. & Ohishi, Y. Post-perovskite phase transition in MgSiO₃. *Science* **304**, 855 (2004).

5. Masters, G., Laske, G., Bolton, H. & Dziewonski, A. The Relative Behavior of Shear Velocity, Bulk Sound Speed, and Compressional Velocity in the mantle: Implications for Chemical and Thermal Structure. In *Earth's deep interior. Mineral physics and tomography from the atomic to the global scale* vol. 117, edited by Karato, S.-I., Forte, A., Liebermann, R. C., Masters, G. & Stixrude, L. pp. 63–87, American Geophysical Union, Washington, DC (2000).
6. Trampert, J., Deschamps, F., Resovsky, J. & Yuen, D. Probabilistic Tomography Maps Chemical Heterogeneities Throughout the Lower Mantle. *Science* **306**, 853–856 (2004).
7. Hutko, A. R., Lay, T., Revenaugh, J. & Garnero, E. J. Anticorrelated seismic velocity anomalies from post-perovskite in the lowermost mantle. *Science* **320**, 1070, (2008).
8. Davies, D. R., Goes, S., Davies, J. H., Schubert, B. S. A., Bunge, H. P. & Ritsema, J. Reconciling dynamic and seismic models of Earth's lower mantle: The dominant role of the thermal heterogeneity. *Earth Planet. Sci. Lett.* **353**, 253–269 (2012).
9. Kennett, B. L. N., Widiyanto, S. & van der Hilst, R. D. Joint seismic tomography for bulk sound and shear wave speed in the Earth's mantle. *J. Geophys. Res.* **103** B6, 12469–12493 (1998).
10. Tackley, P. J. Living dead slabs in 3-D: The dynamics of compositionally-stratified slabs entering a “slab graveyard” above the core-mantle boundary. *Phys. Earth Planet. Interiors* **188**, 150–162 (2011).
11. Labrosse, S., Hernlund, J. W. & Coltice, N. A crystallizing dense magma ocean at the base of the Earth's mantle. *Nature* **450**, 866–869 (2007).
12. Kellogg, L. H. & King, S. D. Effect of mantle plumes on the growth of D'' by reaction between the core and mantle. *Geophys. Res. Lett.* **20**, 379 (1993).
13. Zhang, Z. *et al.* Primordial metallic melt in the deep mantle. *Geophys. Res. Lett.* **43**, 3693 (2016).
14. Houser, C. Constraints on the Presence or Absence of Post-Perovskite in the Lowermost Mantle From Long-Period Seismology. In *Post-Perovskite: The Last Mantle Phase Transition Geophysical Monograph Series* **174**, pp. 191–216 (2007).
15. Kiefer, B., Stixrude, L. & Wentzcovitch, R. M. Elasticity of (Mg,Fe)SiO₃-Perovskite at high pressure. *Geophys. Res. Lett.* **29**, 1539 (2002).
16. Fukui, H. *et al.* Precise determination of elastic constants by high-resolution inelastic X-ray scattering. *J. Synchro. Radiat.* **15**, 618 (2008).
17. Yeganeh-Haeri, A., Weidner, D. J. & Ito, E. Elasticity of MgSiO₃ in the Perovskite Structure. *Science* **243**, 787 (1989).
18. Yeganeh-Haeri, A. Synthesis and re-investigation of the elastic properties of single-crystal magnesium silicate perovskite. *Phys. Earth Planet. Inter.* **87**, 111–121 (1994).
19. Sinogeikin, S. V., Zhang, J. & Bass, J. D. Elasticity of single crystal and polycrystalline MgSiO₃ perovskite by Brillouin spectroscopy. *Geophys. Res. Lett.* **31**, L06620 (2004).
20. Aizawa, Y. *et al.* Temperature derivatives of elastic moduli of MgSiO₃ perovskite. *Geophys. Res. Lett.* **31**, L01602 (2004).
21. Chantel, J., Frost, D. J., McCammon, C. A., Jing, Z. & Wang, Y. Acoustic velocities of pure and iron-bearing magnesium silicate perovskite measured to 25 GPa and 1200 K. *Geophys. Res. Lett.* **39**, L19307 (2012).
22. Li, L. *et al.* Elasticity of (Mg, Fe)(Si, Al)O₃-perovskite at high pressure. *Earth Planet. Sci. Lett.* **240**, 529 (2005).
23. Wehinger, B. *et al.* Dynamical and elastic properties of MgSiO₃ perovskite (bridgmanite). *Geophys. Res. Lett.* **43**, 2568 (2016).
24. Kudoh, Y., Prewitt, C. T., Finger, L. W., Darovskikh, A. & Ito, E. Effect of Fe on the crystal structure of (Mg,Fe)SiO₃ perovskite. *Geophys. Res. Lett.* **17**, 1481–1484 (1990).
25. Jackson, J. M., Zhang, J. & Bass, J. D. Sound velocities and elasticity of aluminous MgSiO₃ perovskite: Implications for aluminum heterogeneity in Earth's lower mantle. *Geophys. Res. Lett.* **31**, L10614 (2004).
26. Andraut, D., Bolfan-Casanova, N. & Guignot, N. Equation of state of lower mantle (Al,Fe)-MgSiO₃ perovskite. *Earth. Planet. Sci. Lett.* **193**, 501 (2001).
27. Boffa Ballaran, T. *et al.* Effect of chemistry on the compressibility of silicate perovskite in the lower mantle. *Earth Planet. Sci. Lett.* **333–334**, 181 (2012).
28. Lundin, S. *et al.* Effect of Fe on the equation of state of mantle silicate perovskite over 1 Mbar. *Phys. Earth Planet. Interiors* **168**, 97–102 (2008).
29. Dorfman, S. M., Meng, Y., Prakapenka, V. B. & Duffy, T. S. Effects of Fe-enrichment on the equation of state and stability of (Mg,Fe) SiO₃ perovskite. *Earth Planet. Sci. Lett.* **361**, 249 (2013).
30. Yagi, T., Okabe, K., Nishiyama, N., Kubo, A. & Kikegawa, T. Complicated effects of aluminum on the compressibility of silicate perovskite. *Phys. Earth Planet. Interiors* **143–144**, 81–91 (2004).
31. Daniel, I. *et al.* Effect of aluminum on the compressibility of silicate perovskite. *Geophys. Res. Lett.* **31**, L15608 (2004).
32. Zhang, J. & Weidner D. J. Thermal Equation of State of Aluminum-Enriched Silicate Perovskite. *Science* **284**, 782 (1999).
33. Murakami M. *et al.* A perovskitic lower mantle inferred from high-pressure, high-temperature sound velocity data. *Nature* **485**, 90–94 (2012).
34. Brown, J. M. & Shankland, T. J. Thermodynamic parameters in the Earth as determined from seismic profiles. *Geophys. J. Roy. Astron. Soc.* **66**, 579–596 (1981).
35. Shatskiy, A. *et al.* Growth of large (1 mm) MgSiO₃ perovskite single crystals: A thermal gradient method at ultrahigh pressure. *Am. Min.* **92**, 1744 (2007).
36. Hirao, N. *et al.* Development of combination method with SR x-ray diffraction and Mössbauer spectroscopy for high-pressure research. *Special issue of Rev. High Pressure Sci. Tech.* **23**, 2C08 (2013).
37. Prescher, C., McCammon, C. & Dubrovinsky, L. MossA: a program for analyzing energy-domain Mössbauer spectra from conventional and synchrotron sources. *J. Appl. Cryst.* **45**, 329–331 (2012).
38. Gütllich, P. Mössbauer Spectroscopy in Chemistry In *Mössbauer Spectroscopy, Topics in Applied Physics* Vol. 5 (eds Gonser, U.) (Springer-Verlag Berlin Heidelberg New York, 1975).
39. Hsu, H., Blaha, P., Cococcioni, M. & Wentzcovitch, R. M. Spin-State Crossover and Hyperfine Interactions of Ferric Iron in MgSiO₃ Perovskite. *Phys. Rev. Lett.* **106**, 118501 (2011).
40. Litasov, K. *et al.* Water solubility in Mg-perovskites and water storage capacity in the lower mantle. *Earth Planet. Sci. Lett.* **211**, 189 (2003).
41. Ito, E. & Matsui, Y. Synthesis & crystal-chemical characterization of MgSiO₃ perovskite. *Earth Planet. Sci. Lett.* **38**, 443–450 (1978).
42. Yagi, T., Mao, H. K. & Bell, P. M. Structure and crystal chemistry of perovskite-type MgSiO₃. *Phys. Chem. Minerals* **3**, 97–110 (1978).
43. Yagi, T., Mao, H. K. & Bell, P. M. Lattice parameters and specific volume for perovskite phase of orthopyroxene composition (Mg,Fe) SiO₃ Carnegie Inst. Washington Yearbook **78**, 612–613 (1979).
44. Ito, E. & Yamada, H. Stability relations of silicate spinels, ilmenites and perovskites. In Akimoto, S. & Manghani, M. H. Eds *High-Pressure Research in Geophysics* p. 405–419. Center for Academic Publications, Tokyo, Japan (1982).
45. Kudoh Y., Ito, E. & Takeda, H. Effect of Pressure on the Crystal Structure of Perovskite-type MgSiO₃. *Phys. Chem. Minerals* **14**, 350 (1989).
46. Horiuchi, H., Ito, E. & Weidner, D. J. Perovskite-type MgSiO₃: single crystal X-ray diffraction study. *American Mineralogist* **72**, 357–360 (1987).
47. Parise, J. B., Wang, Y., Yeganeh-Haeri, A., Cox, D. E. & Fei, Y. Crystal structure and thermal expansion of (Mg,Fe)SiO₃ perovskite. *Geophys. Res. Lett.* **17**, 2089–2092 (1990).
48. Ross, N. L. & Hazen, R. M. High-pressure crystal chemistry of MgSiO₃ perovskite. *Phys. Chem. Minerals* **17**, 228–237 (1990).

49. Jephcoat, A. P. *et al.* High-resolution synchrotron X-ray powder diffraction and Rietveld structure refinement of two (Mg_{0.95}Fe_{0.05}) SiO₃ perovskite samples synthesized under different oxygen fugacity conditions. *American Mineralogist*, **84**, 214–220 (1999).
50. Dobson, D. P. & Jacobsen, S. D. The flux growth of magnesium silicate perovskite single crystals. *American Mineralogist* **89**, 807–811 (2004).
51. Sugahara, M. *et al.* Reinvestigation of the MgSiO₃ perovskite structure at high pressure. *American Mineralogist* **91**, 533–536 (2006).
52. Katsura, T. *et al.* P-V-T relations of MgSiO₃ perovskite determined by *in situ* X-ray diffraction using a large-volume high-pressure apparatus. *Geophys. Res. Lett.* **36**, L01305 (2009).
53. Tange, Y., Takahashi, E., Nishihara, Y., Funakoshi, K. & Sata, N. Phase relations in the system MgO-FeO-SiO₂ to 50 GPa and 2000 °C: An application of experimental techniques using multianvil apparatus with sintered diamond anvils. *J. Geophys. Res.* **114**, B02214 (2009).
54. Yoneda, A. *et al.* Elastic anisotropy of experimental analogues of perovskite and post-perovskite help to interpret D'' diversity. *Nature Comm.* **5**, 3453 (2014).
55. Antonangeli, D. *et al.* Elasticity of Cobalt at High Pressure Studied by Inelastic X-Ray Scattering. *Phys. Rev. Lett.* **93**, 215505 (2004).
56. Antonangeli, D. *et al.* Elasticity of Hexagonal-Closed-Packed Cobalt at High Pressure and Temperature: A Quasiharmonic Case. *Phys. Rev. Lett.* **100**, 085501 (2008).
57. Baron, A. Q. R. *et al.* X-ray scattering beamline for studying dynamics. *J. Phys. Chem. Solids* **61**, 461 (2000).
58. Baron, A. Q. R. Status of the RIKEN Quantum NanoDynamics Beamline (BL43LXU): The Next Generation for Inelastic X-Ray Scattering. *SPring-8 Inf. Newsl.* **15**, 14 (2010).
59. Li, B. & Zhang, J. Pressure and temperature dependence of elastic wave velocity of MgSiO₃ perovskite and the composition of the lower mantle. *Phys. Earth Planet. Interiors* **151**, 143–154 (2005).
60. Li, B. Characteristics of lateral heterogeneities with thermal and chemical origins in the pyrolytic lower mantle. *Progress in Natural Science* **19**, 1603–1611 (2009).

Acknowledgements

We appreciate help by Longjian Xie, Ryo Watanabe, and Tatsuya Hiratoko during IXS experiments. This work was performed using joint-use facilities of the Institute for Study of the Earth's Interior, Okayama University and supported in part by Grants-in-Aid for Scientific Research (Grant Nos 22224008 and 15H021128 awarded to AY and Nos 22000002 and 15H05748 to EO) from the Japan Society for the Promotion of Science. Inelastic X-ray scattering measurements at BL35XU⁵⁷ were done with the approval of JASRI (Proposal Nos 2012B1196, 2013A1047, 2013B1054, and 2014B1290). Measurements at RIKEN BL43LXU⁵⁸ were made during commissioning time. The synchrotron Mössbauer spectroscopy was performed on the BL10XU at SPring-8 (Proposal No. 2014A0104).

Author Contributions

H.F. planned the present study. A.S. synthesized crystals. A.N. crystallographically characterized synthesized crystals. N.T. analyzed the chemical compositions of the samples. S.K., E.O. and N.H. performed the synchrotron Mössbauer measurements. S.K. analyzed the Mössbauer spectrum with input from S.T., H.F., A.Y., S.T., H.U. and A.Q.R.B. performed the IXS measurements at BL35XU and H.F., A.Y. and A.Q.R.B. at BL43LXU. H.F. analyzed the IXS data with input from A.Y. and A.Q.R.B., H.F., A.Y. and A.Q.R.B. wrote the manuscript.

Additional Information

Competing financial interests: The authors declare no competing financial interests.

How to cite this article: Fukui, H. *et al.* Effect of cation substitution on bridgmanite elasticity: A key to interpret seismic anomalies in the lower mantle. *Sci. Rep.* **6**, 33337; doi: 10.1038/srep33337 (2016).



This work is licensed under a Creative Commons Attribution 4.0 International License. The images or other third party material in this article are included in the article's Creative Commons license, unless indicated otherwise in the credit line; if the material is not included under the Creative Commons license, users will need to obtain permission from the license holder to reproduce the material. To view a copy of this license, visit <http://creativecommons.org/licenses/by/4.0/>

© The Author(s) 2016