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OPEN Self-diffusivity, M–S and Fick diffusivity of CO₂ in Na-clay: The influences of concentration and temperature

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Storing CO₂ in underground saline aquifers is an important way to reduce CO₂ emission in atmosphere, where gas/fluid diffusion in clay plays a key role in CO₂ leakage and migration. Various diffusivities, self-diffusivity, Maxwell-Stefan (M-S) diffusivity and Fick diffusivity, in clay interlayer are investigated by molecular dynamics (MD). Self-diffusivity varies with CO₂ concentration, and reaches the maximum value at 2 molecules/unit-cell. High fluid concentration leads to clay swelling, thereby increasing self-diffusivity. However, the fractional free volume of clay explains the trend of CO2 self-diffusivity, which does not decrease with CO₂ concentration monotonously but reaches the maximum when CO₂ concentration reaches 2. Displacement distribution of CO₂ molecules is analysed to explore the microscopic diffusion mechanism, which is characterised by logarithmic normal distribution. The mean value of such distribution further explains the self-diffusivity dependence on CO₂ concentration. M-S and Fick diffusivities of CO₂ are calculated by MD for the first time, both of which increase with increasing CO₂ and H₂O concentration and temperature. Based on self-diffusivity and M-S diffusivity, a quantity representing the coupling strength between CO₂ molecules is presented; it increases firstly with CO₂ concentration but begins to decrease when CO₂ concentration is beyond 2.

Carbon dioxide (CO₂) storage in underground saline aquifers is a promising way to reduce CO₂ emission in the atmosphere. This strategy provides long-term and large-scale storage of CO₂. Clay minerals¹, such as illite, chlorite, kaolinite and montmorillonite (MMT) play an important role in this process. On the one hand, owing to their porous (layered) structure, the clay minerals have remarkable capacity of adsorbing CO₂²⁻⁴. They provide a large amount of space for storing CO₂⁵. On the other hand, clay is almost impermeable, and therefore, the clay-enriched cap rocks show excellent sealing ability to retain injected CO₂⁶⁻⁸. Gas leakage and environmental impact are the main problems found after risk assessment of CO₂ storage. These problems are closely related to fluid (gas) transportation in clay. Gas diffusion in clay is the most important method for gas transportation because of the extremely low permeability of clay.

Many experimental studies have focused on the structure and interaction of CO_2 and clay mineral⁹⁻¹². Loring et al. examined that CO₂ can migrate the interlayer region of MMT based on the in situ X-ray diffraction, magic-angle spinning in nuclear magnetic resonance spectroscopy and attenuated total reflection infrared spectroscopy^{10, 13}. Giesting et al. investigated the impact of CO₂ absorption on Ca-exchanged MMT expansion under different CO₂ pressures¹⁴. In addition, quasi-elastic neutron scattering experiments on hydrated clays have shown that hydrated cation diffusion mobility is probably a complex dynamic process, and the diffusion coefficients of the exchangeable cations were estimated 15. Kozaki et al. determined the apparent diffusion coefficients of Cs+ as functions of the temperature 16. Sánchez et al. discussed that the self-diffusivity of water depends on the temperature and ionic strength of different kinds of clays¹⁷. However, no experiment on diffusivity of CO₂ in clay was reported.

With the current high-performance computational resources, molecular simulations have become powerful tools for understanding the molecular-scale structural^{18, 19}, thermodynamic²⁰, mechanical²¹ and dynamic^{22–25} properties of clay. Grand-canonical Monte Carlo (GCMC) method was used to examine the adsorption of CO₂

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with H_2O^{23} , CH_4^{26-28} and organic molecules 25,29 in clay. Yang *et al.* used molecular dynamics (MD) to study the structure and self-diffusion coefficient (SDC) of CO_2 in uncharged clay-like slit pores 24 . Cygan *et al.* developed CLAYFF force field 30 for clay and three-site flexible potential modes of CO_2^{31} . Myshakin *et al.* showed that the intercalation of CO_2 in clay caused the significant changes in the d-spacing and described that the distribution and diffusion of CO_2 , CO_2^{31} , CO_2^{31} and CO_2^{31} are affected by the number of layers of clay CO_2^{32-34} . Botan *et al.* found that CO_2^{31} affects the diffusion of mobile species in clays CO_2^{31} .

All the above-mentioned MD simulations mainly investigated the SDCs. Two important diffusion coefficients that are closely related to gas transportation, namely, Maxwell–Stefan diffusion coefficient (MDC) and Fick diffusion coefficient (FDC), have not been investigated. In fact, FDC, instead of SDC, is usually experimentally measurable and is used as parameter input in macroscopic gas transportation formulation, which predicts transportation of underground injected CO₂. In this work, a series of MD simulations are performed to explore various diffusion coefficients of CO₂ in clay. The effects of gas concentration, water concentration and temperature on SDC, MDC and FDC are discussed in detail.

Methods

Theory of diffusion. As stated above, diffusion coefficients mainly have three types, SDC, FDC and MDC, which are useful in characterising molecular transport and mobility within porous materials. Self-diffusion means random motions or mixing of particles in the thermodynamic equilibrium. In MD simulation, SDC, D_{iselfs} of component i is computed by Einstein equation³⁵:

$$D_{iself} = \frac{1}{2dn_i} \lim_{t \to \infty} \frac{d}{dt} \sum_{l=1}^{n} (r_{l,i}(t) - r_{l,i}(0))^2,$$
(1)

where n_i is the number of the molecules of component i, d is the dimension of the system and $r_{bi}(t)$ is the position vector of molecule l at time t.

FDC is the particle motions driven by chemical or concentration gradient, resulting in the net mass transport. Fick's law of diffusion defines the transport diffusivity, D_{iFick} , as the proportionality factor between the flux N and the concentration gradient:

$$N = -D_{iFick} \nabla c. \tag{2}$$

Chemical potential gradient is considered as the fundamental driving forces for diffusion^{36–40}. For single-component diffusion in porous material, the transport equation can be expressed as follows:

$$N_i = -L_i \nabla \mu_i. \tag{3}$$

where μ_i is the chemical potential of gas i. Moreover, L is obtained from the MD simulations using⁴¹ the following:

$$L_{i} = \frac{1}{2dVk_{B}T} \lim_{\Delta t \to \infty} \frac{1}{\Delta t} \left\langle \left(\sum_{l=1}^{n_{i}} (r_{l,i}(t + \Delta t) - r_{l,i}(t)) \right)^{2} \right\rangle$$
(4)

where V is the clay volume, k_B is the Boltzmann constant and T is the temperature. The content in <...> represents an average on the cross-displacement correlation function (CDCF). The notation <> represents an average on a number of independent runs of same ensembles starting from different initial conditions.

The Maxwell–Stefan (M–S) theory of diffusion was already available from Maxwell⁴² and Stefan⁴³, in which gas diffusion is driven by chemical gradient. Using the M–S theory, the following expression can be derived for diffusion of the single component in a nanoporous material:

$$N_{i} = -\rho \Theta_{i,sat} D_{iM-S} \frac{\theta_{i}}{k_{B}T} \nabla \mu_{i} , \qquad (5)$$

where ρ is the clay density expressed as the number of unit cells per cubic meter, $\Theta_{i,sat}$ is the saturation loading of the component i, μ_i is the chemical potential expressed in joules per molecule and N_i is the molar flux of the component i expressed in molecules per square meter per second. The fractional occupancy θ_i are defined as follows:

$$\theta_i = \frac{\Theta_i}{\Theta_{i,sat}},\tag{6}$$

where Θ_i is the loading of the component *i*. Combining equation (3) and equation (5), the MDC, D_{iM-S} , is calculated by the following:

$$D_{iM-S} = \frac{k_B T}{\rho \Theta_i} L_i \tag{7}$$

Under isothermal condition, the molecules are considered to move with average velocity subject to a driving force $\nabla \mu_p \mu = \mu_0 + k_B T \ln f$, where f is the fugacity and $\nabla \mu = k_B T \nabla f f$. Thus, equation (2) is equivalent to equation (3).

$$N_{i} = -D_{iFick} \nabla c_{i} = -\rho D_{iFick} \nabla \Theta_{i} = -\rho D_{iFick} \Theta_{i} \frac{\partial \ln \Theta_{i}}{\partial \ln f_{i}} \frac{\nabla \mu_{i}}{k_{B}T}$$

$$(8)$$

Number of CO ₂	Number of H ₂ O	Temperature	d-spacing	Short-name
10 (0.331)	160 (5)	313.15	12.50	C1
16 (0.5)	160 (5)	313.15	12.70	C2
32 (1)	160 (5)	313.15	13.50	C3
48 (1.5)	160 (5)	313.15	14.00	C4
64 (2)	160 (5)	313.15	14.50	C5
80 (2.5)	160 (5)	313.15	14.75	C6
64 (2)	32 (1)	313.15	11.90	H1
64 (2)	96 (3)	313.15	13.00	H3
64 (2)	160 (5)	313.15	14.50	H5
64 (2)	224 (7)	313.15	15.50	H7
64 (2)	288 (9)	313.15	17.25	H9
32 (1)	128 (4)	273.15	12.50	T1
32 (1)	128 (4)	313.15	12.50	T2
32 (1)	128 (4)	353.15	12.50	T3
32 (1)	128 (4)	393.15	12.50	T4
32 (1)	128 (4)	433.15	12.50	T5

Table 1. Simulation parameters of various simulation systems. Numbers in bracket indicate the corresponding CO₂ or H₂O concentration in molecules/cell.

The FDC and MDC are related by

$$D_{iFick} = D_{iM-S}\Gamma, (9)$$

where Γ is the thermodynamic factor

$$\Gamma = \Theta_i \frac{\partial \ln f_i}{\partial \Theta_i}. \tag{10}$$

MD simulation details. The sodium-saturated Wyoming-type MMT is used as the clay crystal model with unit formula $Na_{0.75}[Si_8](Al_{3.25}Mg_{0.75})O_{20}(OH)_4^{44}$, which comprises 2:1 or tetrahedral—octahedral—tetrahedral layers. The substitution of octahedral Al^{3+} by Mg^{2+} and tetrahedral Si^{4+} by Al^{3+} leads to net negative layer charge. The interlayer Na^+ cations are balanced by these negative changes. Three-dimensional periodic boundary condition is applied to a simulation box comprising 32 (8 × 4 × 1) united cells. Clay model with different number of CO_2 and H_2O are allowed to freely swell to their equilibrium d-spacing. All simulated systems are listed in Table 1. Cn (n = 1, 2,...) denotes the system in which the number of H_2O is fixed, and the number of H_2O is represented by n. Hn (n = 1, 3,...) represent systems with different numbers of H_2O . Tn (n = 1, 3,...) represents systems with varying temperature.

In this article, all simulations are performed with the Accelrys Materials Studio software. The CLAFF force field ^{30, 31} (Cygan, Liang *et al.* 2004, Cygan, Romanov *et al.* 2012) is used ^{30–32}. The Ewald summation method is applied for the electrostatic interactions. The atom-based summation method is used in the van der Waals interactions with a cut-off of 12.5 Å. Gas diffusions are simulated as follows: Gases are firstly inserted into the above equilibrium configuration by GCMC method to get a new clay-gas cell. The cell is minimised and equilibrated by NPT and NVT. Then, the cell is run in a NVT ensemble (0.5 ns) followed by a long-time (5 ns) NVE run. The trajectories in the NVE run are saved for diffusion coefficient computation. Integration time step is set to 1 fs. Since diffusion coefficients are based on statistics of numerous molecule trajectories, all the diffusivities in this article are computed by averaging on more than 10 indenpent MD runs to get reliable results. Gas concentration, water component and temperature are used to investigate their impacts on gas diffusions.

Results and Discussion

MD simulation can be used to obtain not only gas diffusion coefficients but also the microscopic details of the diffusion process, which is helpful to understand underlying diffusion mechanism in a confined space. Figure 1 depicts a simulation snapshot of the diffusion of CO_2 and H_2O in MMT clay, from which it is seen that fluid molecules are uniformly distributed in the interlayer of MMT. The direction perpendicular to the walls of interlay is defined as z-axis and the mutually orthogonal x- and y-axes form a plane parallel to the walls. The short length along the z-axis greatly hinders the diffusion of gases along this direction. Thus, only the diffusion behaviour of fluids along the x-y plane is investigated in this study. Unless stated otherwise, the diffusion coefficients in this work refer to the diffusion in the x-y plane.

Self-diffusion coefficients and its Variation with CO₂ concentration. Diffusion coefficients are calculated based on the statistical average of the motion of large amounts of particles, and their accuracy relies on the sample size. For the sake of accuracy, each system is run 10 times independently. The resulting MSDs are

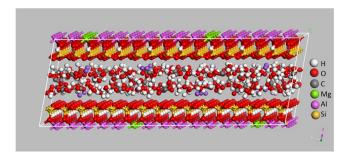


Figure 1. Snapshot of MD simulation of 64 CO₂ and 160 H₂O molecules in MMT cell.

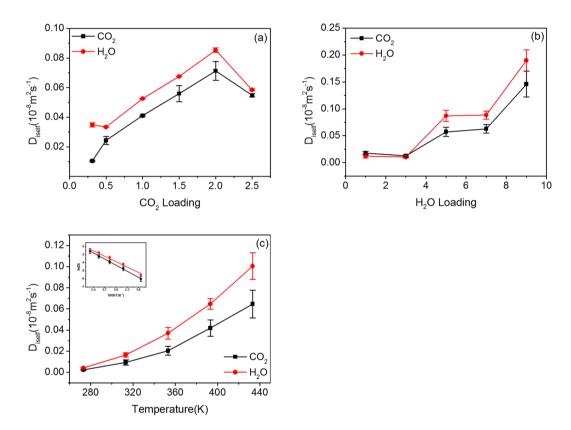
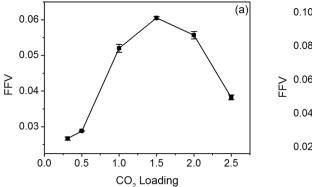


Figure 2. D_{iself} of CO_2 and H_2O in MMT with different systems: (a) CO_2 loading; (b) H_2O loading; (c) temperature. Inset of (c) shows Arrhenius plot of SDC versus temperature.

averaged to calculate SDC. Diffusion coefficients are related to many factors including type and concentration of the gas, size of the pores, temperature and so on. Injection pressure of CO_2 , moisture content and underground temperature are key factors in the geological storage of CO_2 . Therefore, this study focuses on the impacts of CO_2 concentration, water concentration and temperature on CO_2 diffusion coefficient.

Firstly, the influence of CO_2 concentration on its SDC is studied. The number of water molecules is set to 5 molecules per unit cell and the temperature to 313.15 K (approximately the temperature of 1000 meters underground). The Monte Carlo method is used to insert varying amounts of CO_2 molecules into the system. The system is then relaxed to equilibrium through NPT and NVT run, with resulting interlay distances (d-spacing) shown in Table 1. The equilibrated system is then subjected to dynamic run in NVE ensemble to compute SDC, as shown in Fig. 2a. With the increase of CO_2 concentration, the SDCs of CO_2 and CO_2 firstly increase and then decrease. At the CO_2 loading value of 64 (CO_2 system), the SDCs of both reach maximum values. The SDC of CO_2 remains higher than that of CO_2 , which originates from the smaller thermodynamic diameter of CO_2 (2.65 Å) compared with that of CO_2 (3.3 Å).

However, many studies had shown that the SDCs of gases in porous media decrease with the increase of their concentrations $^{36, 38, 45}$, which is different from the aforementioned result in Fig. 2a. To explore the cause of the variation of SDC, we further analyse the effect of CO_2 on the clay structure. As shown in Table 1, the d-spacing of MMT continues to increase with CO_2 concentration, which could provide fluids with larger space for diffusion, resulting in larger diffusion coefficient. SDC is expected to increase with CO_2 concentration



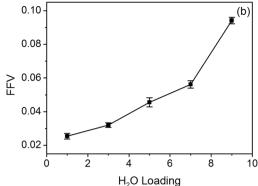


Figure 3. Fractional free volume (FFV) dependence on: (a) CO₂ loading; (b) H₂O loading.

monotonously. However, our simulations reveal a higher SDC in high-concentration system C6 compared to the low-concentration system C5. To further explore the cause of this unusual SDC variation, fractional free volume (FFV) of the system is defined as the ratio of interlayer space unoccupied by fluids to the total volume of MMT:

$$FFV = \frac{V_{free}}{V_{total}},\tag{11}$$

where V_{free} is the free volume and V_{total} is the total volume of the MMT. The FFV can be directly determined by molecular probes^{46,47}. As shown in Fig. 3a, the FFV increases with the CO₂ concentration firstly and peaks at 64 CO₂ molecules, which is in accordance with the position of maximum of CO₂ SDC. This finding indicates two counter-interactions, as follows. Firstly, the increase of CO₂ concentration could cause the expansion of d-spacing in the system and provide fluids with more space for diffusion. Secondly, increasing CO₂ molecules occupy more space, thereby resulting in the decrease of free volumes. The fluid molecules become more crowded and diffuse more slowly. Therefore, fluid SDC depends on the free space in the interlayer. At the initial stage of CO₂ injection, the rapid expansion of d-spacing increases the free volume in the interlayer, which is beneficial for fluid diffusion and increases CO₂ SDCs. The free volume reaches a maximum at 64 CO₂ molecules. The free volume decreases despite the continued expansion of d-spacings, which hinders gas diffusion and decreases the SDCs of CO₂ and H₂O (Fig. 3a).

Variation of SDCs with moisture and temperature. Water content varies with the depth of storage aquifer. A MMT system with 2 CO_2 molecules per unit cell is used to investigate quantitatively the impact of water concentration on the diffusion. Moreover, the temperature is fixed at 313 K. Figure 2b shows the variation of SDCs with H_2O concentration, where the SDCs of CO_2 and CO_2 and CO_3 increase with CO_3 increase with CO_3 increases with moisture content, showing the same trend as the SDCs. The increase of water molecules increases the free volume in the interlayer and supports the diffusion of fluids.

Temperature is an important factor in the diffusion of gases. As shown in Fig. 2c, increasing temperature can accelerate the thermal motion of molecules and increase their kinetic energy, thereby gradually increasing fluid SDC. Calculations show that the relationship between SDC and temperature follows the classical Arrhenius formula:

$$D = D_0 e^{-\frac{E_a}{RT}},\tag{12}$$

where D_0 is a constant, and E_a is the activation energy of diffusion. According to the equation, the activation energy for CO_2 diffusion in MMT is $20.26\,\mathrm{kJ/mol}$ (see inset of Fig. 2c), which is greater than its activation energy of diffusion in pure water $17.8904\,\mathrm{kJ/mol}^{48}$. The MMT layers impose restriction on CO_2 diffusion, causing high activation energy in MMT.

Details such as the paths and displacements of molecules provided by MD simulation can be used to study the microscopic mechanism of gas diffusion. Previous work indicated that the diffusion of gases in porous media (such as polymeric compounds and metal organic frameworks) involves two processes: limited movements within one micropore (small displacements) and jumps among micropores (large displacements)³⁷. Layered MMT provides anisotropic confined space without interconnected micropores. Diffusion of gas in such structure is different from random movements of gas in unlimited environment. We further analyse the displacement distribution of CO₂ molecules, as shown in Fig. 4. All the distributions are seen to have a single peak, and no peak is shown at large displacements. The displacements generally fall within 0.6–1.0 Å. This finding indicates that no inter-pore jumps occur during gas diffusion. Figure 4 also demonstrates that, as CO₂ concentration increases, the displacement peak (i.e. the prevailing displacement) moves towards the left and becomes wider, reaching the maximum width in the C5 system. This finding implies increases of both the prevailing displacement and the number of fast-speed molecules. Therefore, with the increase of CO₂ concentration, SDC of CO₂ increases and reaches a maximum at C5 system. Similarly, the increase of moisture and temperature increases prevailing

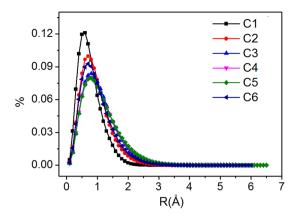


Figure 4. Distributions of displacement, with different systems: (a) CO_2 loading; (b) H_2O loading; (c) temperature.

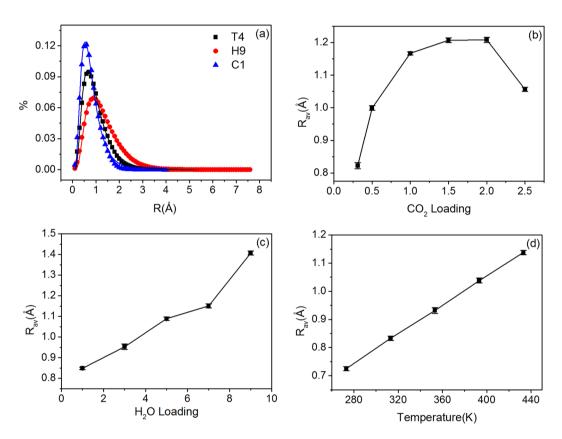


Figure 5. (a) Distributions of displacement and the fitted line in logarithmic normal distribution. The average displacement with different systems: (b) CO_2 loading; (c) H_2O loading; (d) temperature.

displacement and peak width, thereby increasing the SDC (as shown in the Figure above). The displacement distributions can be fitted to logarithmic normal distribution function (LNDF):

$$y = \frac{e^{-\frac{(-a+\log|x|)^2}{2b^2}}}{d\sqrt{2\pi}x}.$$
 (13)

Figure 5a shows some examples, in which curves are observed to be well-fitted to the data ($R^2 > 0.98$). To quantitatively examine the relationship between the displacement distribution and SDC, the mean value of LNDF: $R_{av} = e^{(a+b^2/2)}$, an important characteristic parameter for log-normal distributions, is calculated. The results are shown in Fig. 5b–d. Figure 5b shows the relationship between mean displacement and CO₂ concentration. In this study, R_{av} increases with CO₂ concentration firstly, reaches a maximum at C5 and then decreases,

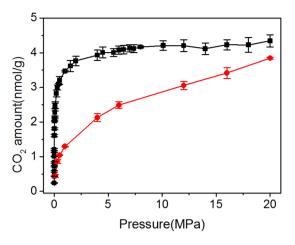


Figure 6. Adsorption isotherms of CO₂ by conventional GCMC method (black) and flexible adsorption method (red) in C1 system.

showing the same trend as the SDC of CO_2 . The average displacement of molecules in unit time (1 ps) reflects the relative speed of diffusion, and hence, larger R_{av} results in faster diffusion speed and a higher SDC. In a word, R_{av} is positively related to the SDCs of CO_2 and H_2O . This finding is further validated by systems with varying moisture contents and temperatures. As shown in Fig. 5c,d, R_{av} increases with moisture and increases linearly with temperature, causing the increase of SDCs of CO_2 and H_2O (Fig. 2).

Variation of MDC and FDC with concentration, moisture content and temperature. To calculate MDC, the Onsager coefficient L should be determined using equation (4) in advance. The CDCF can thus be computed from the slope of the curve, which is used to calculate MDC from equation (7) and FDC using equation (9). The thermodynamic factor in equation (9) requires the adsorption isotherm of CO₂, which can usually be simulated using the GCMC method. The conventional GCMC method assumes a rigid absorbent (µVT ensemble). However, MMT clay swells after adsorbing CO₂. Swollen MMT can continue to adsorb more gas. Therefore, the conventional GCMC is not suitable for calculating adsorption isotherms of CO₂ in MMT. In this study, we adopt a new elastic-adsorption approach to simulate CO₂ isotherm by considering the swelling of MMT. The MMT structure with the initial 224 water molecules is taken as an example (see Fig. 6). Using the conventional GCMC method, we can obtain significantly different adsorption isotherms at different fixed d-spacing values, i.e. the initial and swollen d-spacing. In these isotherms, the curves rapidly approach to saturation at low pressure, and the amount of CO2 adsorption is almost kept constant when the pressure rises above 1 MPa (shown by the black line in Fig. 6), which is inconsistent with experimental results⁴⁹. However, the CO₂ adsorption isotherm obtained by the above elastic-adsorption method increases gradually with the CO₂ pressure, which is more qualitatively coincident with the Langmuir formula and the above-mentioned experiment results. The maximum adsorption θ_{sat} can be obtained by fitting with the Langmuir equation, and then, the thermodynamic factor can be calculated by equation (10). The results show that thermodynamic factor always increases with increasing CO₂ concentration. For low CO₂ concentrations, thermodynamic factor is not influenced by H₂O apparently, whereas at higher CO₂ concentrations, the thermodynamic factor gradually decreases with H₂O concentration.

The effect of CO_2 concentration, H_2O content and temperature on MDC and FDC is also explored. The red dots in Fig. 7 represent the variation of CO_2 MDC, which increase with CO_2 concentration, H_2O content and temperature. The black dots in Fig. 7 represent the FDC. The thermodynamic factor is generally high. Thus, the FDC is usually larger than MDC. Similar to MDC, FDC increases with CO_2 concentration, H_2O content and temperature.

The significance differences of CO_2 diffusivities among CO_2 concentration moisture content and temperature are calculated as 0.033 (SDC), 0.027 (MDC) and 0.012 (FDC) respectively, indicating the reliablity of the above results.

Relationship between SDC and MDC: The coupling effect of CO₂ molecules. Krishna and Paschek *et al.* proved that the difference between MDC and SDC is caused by the self-exchanging coupling between molecules of a diffusing gas⁴⁰. The strength of the coupling between gas molecules can be represented by the self-exchange coefficient D_{ii} , which is calculated by the following formula:

$$\frac{1}{D_{iself}} = \frac{1}{D_{iM-S}} + \frac{\theta_i}{D_{ii}}.$$
(14)

A lower value of D_{ii} indicates stronger correlation and vice versa. The results of Krishna *et al.* indicated that D_{ii} is significantly influenced by the concentration of the gas: Higher concentration leads to stronger correlation⁴⁰. Using the aforementioned SDC and MDC results, the impact of CO₂ concentration on D_{ii} is shown in Fig. 8. With increasing CO₂, D_{ii} firstly increased and then decreased, reaching the maximum at 2 molecules/unit cell. D_{ii} increases with increasing CO₂ molecules, thereby leading to weak correlation among CO₂ molecules. When

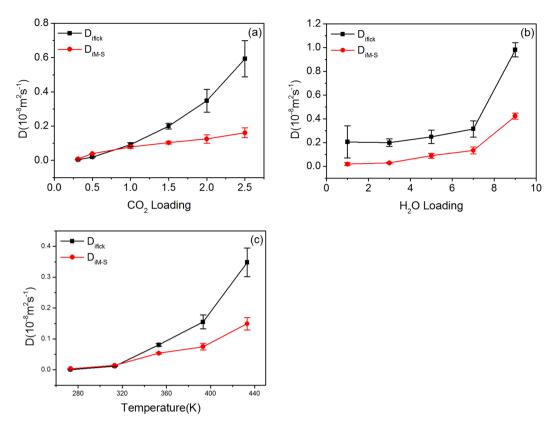


Figure 7. D_{iM-S} and D_{ifick} of CO₂ in different systems: (a) CO₂ loading; (b) H₂O loading; (c) temperature.

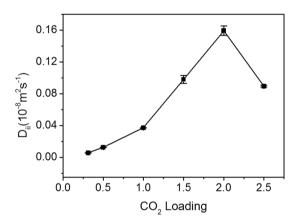


Figure 8. The dependence of D_{ii} on CO_2 loading.

the CO_2 concentration exceeds 2, D_{ii} begins to decrease, resulting in stronger correlation. This result is different from the case described in ref. 40, where D_{ii} monotonically increased with gas concentration in MOF. The coupling between molecules may have been influenced by two factors. On one hand, the correlation is strengthened by increasing the CO_2 concentration. On the other hand, the rising concentration causes the MMT to expand, thereby increasing the volume of the layer and weakening the correlation. The final strength of the coupling may depend on the FFV within the MMT.

Conclusions

In this study, MD is used to study various diffusion coefficients of CO_2 in MMT clay under varying conditions. With increasing CO_2 concentration, the SDCs of CO_2 and H_2O display a peak, respectively. Water content and temperature are positively correlated with the SDC. Many previous studies show that SDCs of gases in porous media usually decrease monotonically with the gas concentration. To further explain the unusual results, the FFV within MMT and the displacement distribution of CO_2 are analysed. FFV has an important effect on the diffusion of gas molecules in MMT. Increasing CO_2 concentration causes the expansion of MMT, thereby increasing the

internal FFV of MMT and providing gas molecules with more space for diffusion. This finding explains the initial increase of the fluid SDC in MMT. FFV begins to decrease when the CO_2 concentration rises above 2, which hinders the gas diffusion and thereby decreasing SDC. Similarly, the increase of water content also causes the MMT to expand and create more free space for CO_2 diffusion, leading to increasing CO_2 SDC. The fluid SDC in MMT depends on the FFV of MMT. Furthermore, the displacement distribution of CO_2 follows log-normal distribution, and the mean of the distribution shows the same trend as the SDC, which provides a good explanation of the effects of CO_2 concentration, moisture and temperature on SDC from another perspective.

Most previous works focused on the self-diffusion coefficients of gases in MMT. FDCs are of practical importance for the geological storage of CO_2 . In this work, the MDC and FDC of CO_2 in MMT are investigated by MD for the first time, both of which increase with CO_2 concentration, moisture and temperature. Using the SDC and MDC, we calculate the self-exchange coefficient D_{ii} , which represents the coupling effect among CO_2 gas molecules. This coefficient firstly increases with CO_2 concentration, reaching a maximum with concentration 2, and then decreases, showing the same trend as FFV. This implies that, with increasing CO_2 concentration, the FFV increases, thereby weakening the correlation between CO_2 molecules until the CO_2 concentration reaches 2. Then, the FFV decreases and the CO_2 correlation coupling strength begins to increase.

References

- 1. Josh, M. et al. Laboratory Characterisation of Shale Properties. J Petrol Sci Eng 88-89, 107-124 (2012).
- 2. Gasparik, M. et al. Geological Controls on the Methane Storage Capacity in Organic-Rich Shales. Int J Coal Geol 123, 34–51 (2014).
- 3. Khosrokhavar, R., Wolf, K. H. & Bruining, H. Sorption of CH₄ and CO₂ on a Carboniferous Shale from Belgium Using a Manometric Set-up. *Int J Coal Geol* **128**, 153–161 (2014).
- 4. Fu, M. H., Zhang, Z. Z. & Low, P. F. Changes in the Properties of a Montmorillonite-Water System during the Adsorption and Desorption of Water: Hysteresis1. Clays & Clay Minerals 38, 485–492 (1990).
- Wang, J. G., Ju, Y., Gao, F., Peng, Y. & Gao, Y. Effect of CO₂ Sorption-Induced Anisotropic Swelling on Caprock Sealing Efficiency. J Clean Prod 103, 685–695 (2014).
- Gaus, I. Role and Impact of CO₂-Rock Interactions during CO₂ Storage in Sedimentary Rocks. Int J Greenhouse Gas Control 4, 73–89 (2010).
- 7. Gernot, R. et al. CO₂ Sorption to Subsingle Hydration Layer Montmorillonite Clay Studied by Excess Sorption and Neutron Diffraction Measurements. Environ Sci Technol 47, 205–211 (2013).
- Abdou, M. I. & Ahmaed, H. E. Evaluation of Low-solids Mud Reological Behavior during Drilling Shale Formation and Their Effect on the Pay Zone Productivity. Petrol Sci Technol 28, 934–945 (2010).
- 9. Saiyouri, N., Tessier, D. & Hicher, P. Y. Experimental Study of Swelling in Unsaturated Compacted Clays. Clay Miner 39, 469–479 (2004).
- Loring, J. S. et al. In Situ Study of CO₂ and H₂O Partitioning between Na–Montmorillonite and Variably Wet Supercritical Carbon Dioxide. Langmuir 30, 6120–6128 (2014).
- Schaef, H. T. et al. In Situ XRD Study of Ca²⁺ Saturated Montmorillonite (STX-1) Exposed to Anhydrous and Wet Supercritical Carbon Dioxide. Int J Greenhouse Gas Control 6, 220–229 (2012).
- 12. Krukowski, E. G. *et al.* FT-IR Study of CO₂ Interaction with Na⁺ Exchanged Montmorillonite. *Appl Clay Sci* **114**, 61–68 (2015).
- Loring, J. S. et al. In Situ Molecular Spectroscopic Evidence for CO₂ Intercalation into Montmorillonite in Supercritical Carbon Dioxide. Langmuir 28, 7125–7128 (2012).
- 14. Giesting, P., Guggenheim, S., Groos, A. F. K. V. & Busch, A. X-ray Diffraction Study of K- and Ca-Exchanged Montmorillonites in CO₂ Atmospheres. *Environ Sci Technol* **46**, 5623–5630 (2012).
- Sobolev, O. et al. Hydration of Na⁺, Ni²⁺, and SM³⁺ in the Interlayer of Hectorite: A Quasielastic Neutron Scattering Study. J Phys Chem C 113, 13801–13812 (2009).
- Kozaki, T., Sato, H., Sato, S. & Ohashi, H. Diffusion Mechanism of Cesium Ions in Compacted Montmorillonite. Eng Geol 54, 223–230 (1999).
- 17. Sánchez, F. G. et al. Self-Diffusion of Water and its Dependence on Temperature and Ionic Strength in Highly Compacted Montmorillonite, Illite and Kaolinite. Appl Geochem 23, 3840–3851 (2008).
- 18. Teich-Mcgoldrick, S. L., Greathouse, J. A., Jové-Colón, C. F. & Cygan, R. T. Swelling Properties of Montmorillonite and Beidellite Clay Minerals from Molecular Simulation: Comparison of Temperature, Interlayer Cation, and Charge Location Effects. *J Phys Chem C* 119, 20881–20891 (2015).
- 19. Lee, M. S., Mcgrail, B. P. & Glezakou, V. A. Microstructural Response of Variably Hydrated Ca-Rich Montmorillonite to Supercritical CO₂. Environ Sci Technol **48**, 8612–8619 (2014).
- 20. Boek, E. S., Coveney, P. V. & Skipper, N. T. Monte Carlo Molecular Modeling Studies Of Hydrated Li-, Na-, And K-Smectites: Understanding The Role Of Potassium As A Clay Swelling Inhibitor. *J.am.chem.soc* 117, 12608–12617 (1995).
- 21. Zhang, W., Hu, H., Li, X. & Fang, Z. Interplay of Montmorillonite-H₂O-scCO₂ System between Mechanical Behavior and Adsorption: Molecular Dynamics. *J Phys Chem C* 119, 21959–21968 (2015).
- Malikova, N., Marry, V., Dufrêche, J. F. & Turq, P. Na/Cs montmorillonite: Temperature Activation of Diffusion by Simulation. Curr Opin Colloid In 9, 124–127 (2004).
- 23. Botan, A., Rotenberg, B., Marry, V., Turq, P. & Noetinger, B. Carbon Dioxide in Montmorillonite Clay Hydrates: Thermodynamics, Structure, and Transport from Molecular Simulation. *J Phys Chem C* 114, 14962–14969 (2010).
- 24. Yang, X. & Zhang, C. Structure and Diffusion Behavior of Dense Carbon Dioxide Fluid in Clay-Like Slit Pores by Molecular Dynamics Simulation. *Chem Phys Lett* **407**, 427–432 (2005).
- Krishnan, M., Saharay, M. & Kirkpatrick, R. J. Molecular Dynamics Modeling of CO₂ and Poly(ethylene glycol) in Montmorillonite: The Structure of Clay–Polymer Composites and the Incorporation of CO₂. J Phys Chem C 117, 20592–20609 (2013).
- Yang, N., Liu, S. & Yang, X. Molecular Simulation of Preferential Adsorption of CO₂ Over CH₄ in Na-montmorillonite Clay Material. Appl Surf Sci 356, 1262–1271 (2015).
- Kadoura, A., Narayanan Nair, A. K. & Sun, S. Adsorption of Carbon Dioxide, Methane, and Their Mixture by Montmorillonite in the Presence of Water. *Micropor Mesopor Mat* 225, 331–341 (2016).
- 28. Jin, Z. & Firoozabadi, A. Methane and Carbon Dioxide Adsorption in Clay-Like Slit Pores by Monte Carlo Simulations. *Fluid Phase Equilibr* **360**, 456–465 (2013).
- 29. Yu, C. H., Newton, S. Q., Norman, M. A., Schäfer, L. & Miller, D. M. Molecular Dynamics Simulations of Adsorption of Organic Compounds at the Clay Mineral/Aqueous Solution Interface. *Struct Chem* 14, 175–185 (2003).
- 30. Cygan, R. T., Liang, J. J. & Kalinichev, A. G. Molecular Models of Hydroxide, Oxyhydroxide, and Clay Phases and the Development of a General Force Field. *J Phys Chem B* **108**, 1255–1266 (2004).
- 31. Cygan, R. T., Romanov, V. N. & Myshakin, E. M. Molecular Simulation of Carbon Dioxide Capture by Montmorillonite Using an Accurate and Flexible Force Field. *J Phys Chem C* **116**, 13079–13091 (2012).

- 32. Makaremi, M., Jordan, K. D., Guthrie, G. D. & Myshakin, E. M. Multiphase Monte Carlo and Molecular Dynamics Simulations of Water and CO, Intercalation in Montmorillonite and Beidellite. J Phys Chem C 119 (2015).
- 33. Myshakin, E. M., Makaremi, M., Romanov, V. N., Jordan, K. D. & Guthrie, G. D. Molecular Dynamics Simulations of Turbostratic Dry and Hydrated Montmorillonite with Intercalated Carbon Dioxide. *J Phys Chem A* 118, 7454–7468 (2014).
- 34. Myshakin, E. M., Saidi, W. A., Romanov, V. N., Cygan, R. T. & Jordan, K. D. Molecular Dynamics Simulations of Carbon Dioxide Intercalation in Hydrated Na-Montmorillonite. *J Phys Chem C* 117, 11028–11039 (2013).
- 35. Einstein, A. On the Movement of Small Particles Suspended in a Stationary Liquid Demanded by the Molecular-Kinetic Theory of Heat. *Annalen der Physik (Leipzig)*. 17, 549–560 (1905).
- Skoulidas, A. I., Sholl, D. S. & Krishna, R. Correlation Effects in Diffusion of CH4/CF4 Mixtures in MFI Zeolite. A Study Linking MD Simulations with the Maxwell–Stefan Formulation. *Langmuir* 19, 7977–7988 (2003).
- 37. Krishna, R. & Wesselingh, J. A. The Maxwell-Stefan Approach to Mass Transfer. Chem Eng Sci 52, 861-911 (1997).
- 38. Krishna, R. & Paschek, D. Verification of the Maxwell-Stefan Theory for Diffusion of Three-Component Mixtures in Zeolites. *Chem Eng J* 87, 1–9 (2002).
- 39. Krishna, R. & Jasper, V. B. Diffusion of Alkane Mixtures in Zeolites: Validating the Naxwell-Stefan Formulation Using MD Simulations. *J Phys Chem B* **109**, 6386–6396 (2005).
- 40. Krishna, R. Diffusion in Porous Crystalline Materials. Chem Soc Rev 41, 3099-3118 (2012).
- 41. Theodorou, D. N., Snurr, R. Q. & Bell, A. T. Molecular Dynamics and Diffusion in Microporous Materials. Comprehensive Supramolecular Chemistry 7, 507–548 (1996).
- 42. Maxwell, J. C. On the dynamical theory of gases. Proceedings of the Royal Society of London 15, 167-171 (1867).
- 43. Stefan, J. uber das Gleichgewich und die Bewegung, insbesondere die Diffusion von Gasmengen. Sitzungsber Akad Wiss Wien. 63, 63–124 (1871).
- 44. Skipper, N. T., Chang, F. R. C. & Sposito, G. Monte Carlo Simulation of Interlayer Molecular Structure in Swelling Clay Minerals. I: Methodology. Clay Clay Miner 43, 285–293 (1995).
- 45. Skoulidas, A. I. & Sholl, D. S. Self-Diffusion and Transport Diffusion of Light Gases in Metal-Organic Framework Naterials Assessed Using Molecular Dynamics Simulations. *J Phys Chem B* **109**, 15760–15768 (2005).
- 46. Connolly, M. L. Solvent-Accessible Surfaces of Proteins and Nucleic Acids. Science 221, 709-713 (1983).
- 47. Ronova, I. A., Rozhkov, E. M., Alentiev, A. Y. & Yampolskii, Y. P. Occupied and Accessible Volumes in Glassy Polymers and Their Relationship with Gas Permeation Parameters. *Macromol Thero Simul* 12, 425–439 (2003).
- 48. Tamimi, A., Rinker, E. B. & Sandall, O. C. Diffusion Coefficients for Hydrogen Sulfide, Carbon Dioxide, and Nitrous Oxide in Water over the Temperature Range 293–368 K. *J Chem Eng Data* 39, 330–332 (1994).
- 49. Busch, A. et al. Carbon Dioxide Storage Potential of Shales. Int J Greenhouse Gas Control 2, 297-308 (2008).

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Author Contributions

H.H. organized the work and prepared the manuscript. Y.X. performed molecular dynamics simulations. X.L. conducted theoretical aspects and analyzed data. All the authors discussed and approved the results and conclusions of this article.

Additional Information

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