

# CORRESPONDENCE

DOI: 10.1038/s41467-018-04943-w

OPEN

# Reply to "On the origin of molecular oxygen in cometary comae"

Y. Yao 1 & K.P. Giapis 1

Laboratory experiments suggest that the molecular oxygen, detected in the coma of comet 67P, is produced in part by abstraction reactions of cometary water ions at exposed surfaces on the nucleus and on the spacecraft. While production rates are likely small relative to water,  $O_2$  formation on the spacecraft, near the spectrometer used to detect  $O_2$ , questions what its measured abundance means, and renders conclusions of a primordial origin premature.

Heritier, Altwegg, Berthelier (HAB) et al. discount the contribution of our proposed Eley-Rideal (ER) reaction mechanism<sup>1</sup> to the observed  $O_2$  abundance<sup>2</sup> in the 67P/G–C coma by positing that: (1) the flux of energetic water-group ions ( $H_2O^+$ ,  $H_3O^+$ , and  $OH^+$ ) hitting the nucleus is not sufficient to produce the observed  $O_2$  signal, and (2) there are no instrumental effects in response to energetic  $O_2^-$  ions and energetic  $O_2$  neutrals entering the DFMS.

The ion flux deficiency was conceded in our paper<sup>1</sup>. However, HAB et al. offer a new argument that shifts the debate. They show that, of the 2 water-group ion populations<sup>3</sup> reaching Rosetta, the (50–300 eV) "accelerated" water ions originating in the extended coma exhibit a peak in flux perfectly out-of-phase with the H<sub>2</sub>O and O<sub>2</sub> densities measured at the spacecraft between March 6 and 23, 2016. This anticorrelation is projected to also hold for any O<sub>2</sub> produced when the accelerated H<sub>2</sub>O<sup>+</sup> ions subsequently reach the nucleus. Assuming that ER reaction products are not trapped on the nucleus surface, the anticorrelation makes a compelling case against the accelerated water ions being the main driver for O<sub>2</sub> production. The culprit flux must indeed be well-correlated with the neutral gas signal at ROSINA-COPS.

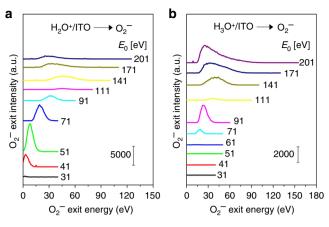
In contrast, the "cold" water ions are correlated with O<sub>2</sub>. Though not mentioned by HAB et al., the anticorrelation does not hold for the more abundant "cold" water-group ions<sup>3</sup>, produced in the space between the 67P nucleus and Rosetta, and arriving at the spacecraft with energies between 10 and 50 eV. Depending on heliocentric distance, these newly formed water ions experience the solar wind convective electric field<sup>3,4</sup>, or the ambipolar electric field<sup>5</sup> of the inner cometary plasma and gain energy as they move away from the nucleus. Upon reaching Rosetta, the negative spacecraft potential accelerates them further to impinge on exposed spacecraft surfaces at energies that can be measured by the RPC-ICA<sup>3</sup> and ion and electron sensor (IES)<sup>4,6</sup>

instruments. Unlike the sporadic arrival of accelerated ions, "the cold population is almost always present" at Rosetta<sup>3</sup> during the entire mission, tracking well the averaged neutral gas density preperihelion and postperihelion<sup>7</sup>. With respect to flux oscillations, Goldstein et al.<sup>6</sup> present timed cold-ion arrival data for September 10, 2014, demonstrating in Fig. 1 that the flux of these ions peaks contemporaneous with the neutral gas density measured by COPS. More intense IES signal is seen for 10-50 eV ions between October 17 and 21, 2014, when the ion and neutral gas peaks are perfectly synchronized, see Fig. 5 in Galand et al.8. Langmuir probe derived ion densities also exhibit peaks perfectly coincidental with the neutral gas density between 14 and 22 October 2014, see Fig. 1 in Edberg et al.9. Remarkably, the October 17-23, 2014 period coincides with the strongest linearity (R = 0.97) seen between O<sub>2</sub> and H<sub>2</sub>O DFMS signals<sup>2</sup>. The correlation holds even closer to perihelion, see Fig. 2 in Volwerk et al. 10 for June 7, 2015. Thus, it appears the cold water-group ions are well-correlated with the H<sub>2</sub>O and O<sub>2</sub> neutral gas densities throughout the mission.

Rosetta emits its own O2. "Cold" water ions possess enough kinetic energy to also drive ER reactions on exposed spacecraft surfaces—the threshold for neutral O2 formation in H2O+ collisions with oxygen atoms on metal surfaces is estimated to be in the 5-8 eV range. These surfaces, include aluminium frame components, photovoltaic (PV) panels, and multi-layer insulation (MLI) protection. The PV panel windows are coated with transparent conductive indium-tin oxide (ITO), while the MLI has a top layer consisting also of conductive ITO (for uniform spacecraft potential) (M.G.G.T. Taylor & A.I. Eriksson, personal communication). Thus, a substantial surface area of ITO is exposed to and bombarded by water-group ions with energies between 10 and 50 eV. Figure 1 presents new results from scattering of energetic H<sub>2</sub>O<sup>+</sup> and H<sub>3</sub>O<sup>+</sup> on ITO surfaces under identical conditions to our original studies on cometary material analogues<sup>1</sup>. As in that case, we find that  $O_2$  is produced readily on ITO, in fact with a lower H<sub>2</sub>O<sup>+</sup> incidence energy threshold than that observed for scattering on  $SiO_x$  or  $FeO_y$  (Al-oxide behaves similarly).  $O_2^+$  and neutral  $O_2$  are also co-produced (not shown) with varying kinetic energies and states of excitation. This experiment suggests that the "cold" water ions bombarding Rosetta produce O<sub>2</sub> in situ, thus populating the gas cloud around

1

<sup>&</sup>lt;sup>1</sup> Division of Chemistry and Chemical Engineering, California Institute of Technology, 1200 E. California Blvd, Pasadena, CA 91125, USA. Correspondence and requests for materials should be addressed to K.P.G. (email: giapis@cheme.caltech.edu)



**Fig. 1** Production of  $O_2^-$  from energetic  $H_2O^+$  and  $H_3O^+$  bombardment of ITO surfaces. Energy distributions of  $O_2^-$  scattered from a thick layer of conductive indium–tin oxide following bombardment by **a**  $H_2O^+$  and **b**  $H_3O^+$  ion beams at various incidence energies ( $\mathcal{E}_O$ ). Scattering geometry: 45° angle of incidence and 45° angle of exit. The ITO layer was deposited on a Cu sample by magnetron sputtering of a commercial high-purity ITO target

the spacecraft with O<sub>2</sub>. Can any of this O<sub>2</sub> be detected by the DFMS? This phenomenon is arguably equivalent to outgassing of the spacecraft, which has been shown<sup>11–13</sup> to lead to detectable signal after many years of space travel, even when the DFMS is not in direct line-of-sight of the outgassing sources (e.g., during spacecraft maneuvers or other payload operations). Indeed, Beth et al.<sup>13</sup> rationalized a false-positive detection of NH<sub>4</sub><sup>+</sup> by the DFMS on the grounds that "the gas cloud around the spacecraft may be contaminated by Rosetta itself." Based on other background gas detection experiments, Schläppi et al.<sup>11</sup> were the first to wonder whether "the spacecraft is surrounded by a significantly denser atmosphere that enhances the collision frequency and thus increases the return flux." Given that the DFMS can detect spacecraft outgassing emissions far away from the comet, we see no reason why some of the O<sub>2</sub> produced locally, while in orbit around the comet, will not make it into the DFMS.

Do ion-Rosetta collisions produce enough O<sub>2</sub>? The argument circles back to ion flux, albeit in H<sub>2</sub>O<sup>+</sup> collisions with Rosetta surfaces. "Cold" water-group ion flux has been measured to be 2 orders of magnitude larger than that of "accelerated" water ions<sup>3</sup> with the caveat that it may be underestimated "due to the limited field of view of the instrument". Though more significant, this flux is still too low (roughly by 100×) to justify the measured O<sub>2</sub> abundance. However, O2 is now produced proximal to the DFMS, expanding the possibility of an instrumental effect. Can locally produced O2 entering the DFMS be ionized more efficiently than cometary O2? An important difference with O2 formed on Rosetta vs. the nucleus is the state of excitation of the molecule. ER reactions produce rovibrationally hot molecules, often also electronically excited (e.g., Rydberg states)<sup>1</sup>. Such excited O<sub>2</sub> molecules (e.g., long-lived low lying singlet states) are more likely to survive the transit time into the DFMS ionizer when produced in its vicinity. Vibrationally and electronically excited O<sub>2</sub> states have lower energy threshold and larger cross-section for electron impact ionization than the ground state<sup>14</sup>. Bottom line, excited O<sub>2</sub> molecules entering an ionizer will produce more detectable O<sub>2</sub><sup>+</sup> ions than ground-state O<sub>2</sub> neutrals.

Why does  $O_2$  appear to follow the  $r^2$  Haser law?  $O_2$  yield in ER reactions depends on both flux and energy of the incident  $H_2O^+$ , where the ion energy is determined effectively by the spacecraft potential. While the "cold" water ion flux follows a 1/r scaling law

 $(r = \text{cometocentric distance})^3$ , the  $O_2$  flux will exhibit a different r scaling because of the convoluted energy dependence. The spacecraft potential is determined by the balance between ions and electrons arriving at its surfaces, whose fluxes depend on cometocentric distance and latitude<sup>15</sup>. As a result, the spacecraft potential exhibits generally a decaying dependence on r, which transfers to the ion energy gained when traversing the sheath. The convoluted ion flux and energy dependencies on cometocentric distance produce a  $1/r^n$  scaling, where n > 1. Thus,  $O_2$  signal may exhibit a scaling closer to the Haser law for entirely different reasons than those assumed by HAB et al.

Has the DFMS been calibrated for  $O_2$ ? None of the published papers  $^{16-18}$  and Ph.D. theses  $^{19,20}$  on DFMS operation and characterization contains any calibration data for  $O_2$ , neither to energetic ions  $(O_2^-, O_2^+)$ , nor to energetic  $O_2$  neutrals, nor to excited states of  $O_2$ . Only background trace amounts of thermal  $O_2$  have been detected  $O_2$ . In his Ph.D. thesis, Schläppi  $O_2$  presents calibration data to energetic  $O_2$  ions, but includes no such experiments with  $O_2^+$  ions. An instrumental effect cannot be ruled out without knowledge of the DFMS response to energetic or excited  $O_2$ .

In conclusion, laboratory scattering experiments of  ${\rm H_2O^+}$  on ITO surfaces suggest that ER reactions may produce  ${\rm O_2}$  on Rosetta surfaces from "cold" water-group ions. Given the prevalence of the cold ion population, this phenomenon resembles intensified spacecraft outgassing. Therefore, some of the in situ produced  ${\rm O_2}$  must contribute to the overall  ${\rm O_2}$  signal detected. The magnitude of the contribution depends not only on the number density but also on the state of excitation of the  ${\rm O_2}$  molecules entering the DFMS. Without instrument calibration, the actual level of cometary  ${\rm O_2}$  cannot be established.

Received: 10 October 2017 Accepted: 5 June 2018 Published online: 03 July 2018

## References

- Yao, Y. & Giapis, K. P. Dynamic molecular oxygen production in cometary comae. Nat. Comm. 8, 15298 (2017).
- Bieler, A. et al. Abundant molecular oxygen in the coma of comet 67P/ Churyumov-Gerasimenko. *Nature* 526, 678–681 (2015).
- Nilsson, H. et al. Evolution of the ion environment of comet 67P/ Churyumov–Gerasimenko. Astron. Astrophys. 583, A20 (2015).
- Fuselier, S. A. et al. Rosina/DFMS and IES observations of 67P: ion-neutral chemistry in the coma of a weakly outgassing comet. Astron. Astrophys. 583, A2 (2015).
- Madanian, H. et al. Suprathermal electrons near the nucleus of comet 67P/ Churyumov–Gerasimenko at 3 AU: model comparisons with Rosetta data. J. Geophys. Res. Space Phys. 121, 5815 (2016).
- Goldstein, R. et al. The Rosetta ion and electron sensor (IES) measurement of the development of pickup ions from comet 67P/Churyumov–Gerasimenko. Geophys. Res. Lett. 42, 3093 (2015).
- Nilsson, H. et al. Evolution of the ion environment of comet 67P during the Rosetta mission as seen by RPC-ICA. MNRAS 469, S252 (2017).
- Galand, M. et al. Ionospheric plasma of comet 67P probed by Rosetta at 3 au from the sun. MNRAS 462, S331 (2016).
- Edberg, N. J. T. et al. Spacial distribution of low-energy plasma around comet 67P/CG from Rosetta measurements. Geophys. Res. Lett. 42, 4263 (2015).
- Volwerk, M. et al. Mass-loading, pile-up, and mirror-mode waves at comet 67P/ Churyumov-Gerasimenko. Ann. Geophys. 34, 1 (2016).
- Schläppi, B. et al. Characterization of the gaseous spacecraft environment of Rosetta by ROSINA. In Proc. of the 3rd AIAA Atmospheric Space Environments Conference (AIAA, 2011).
- Schläppi, B. et al. Influence of spacecraft outgassing on the exploration of tenuous atmospheres with in situ mass spectrometry. J. Geophys. Res. 115, A12313 (2010).
- Beth, A. et al. First in situ detection of the ammonium cometary ion NH<sub>4</sub><sup>+</sup> (protonated ammonia NH<sub>3</sub>) in the coma of 67P/C-G near perihelion. MNRAS 462, S562 (2016).

- Kosarim, A. V. et al. Electron impact ionization cross sections of vibrationally and electronically excited oxygen molecules. *Chem. Phys. Lett.* 422, 513–517 (2006).
- Odelstad, E. et al. Measurements of the electrostatic potential of Rosetta at comet 67P. MNRAS 469, S568–S581 (2017).
- Balsiger, H. et al. ROSINA-Rosetta orbiter spectrometer for ion and neutral analysis. Space Sci. Rev. 128, 745–801 (2007).
- Graf, S. et al. A cometary neutral gas simulator for gas dynamic sensor and mass spectrometer calibration. J. Geophys. Res. 109, E07S08 (2004).
- Hässig, M. M. et al. The capabilities of ROSINA/DFMS to measure argon isotopes at comet 67P/Churyumov–Gerasimenko. *Planet. Space Sci.* 105, 175–178 (2015).
- Schläppi, B. Characterization of the ROSINA Double Focusing Mass Spectrometer. PhD thesis, Universität Bern (2011).
- Hässig, M. M. Sensitivity and Fragmentation Calibration of the ROSINA Double Focusing Mass Spectrometer. PhD thesis, Universität Bern (2013).

### **Acknowledgements**

This work was supported by NSF (Award no. 1202567).

### **Author contributions**

Y.Y. performed the experiments and K.P.G. wrote the reply.

### Additional information

Competing interests: The authors declare no competing interests.

Reprints and permission information is available online at http://npg.nature.com/reprintsandpermissions/

**Publisher's note:** Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.

Open Access This article is licensed under a Creative Commons Attribution 4.0 International License, which permits use, sharing, adaptation, distribution and reproduction in any medium or format, as long as you give appropriate credit to the original author(s) and the source, provide a link to the Creative Commons license, and indicate if changes were made. The images or other third party material in this article are included in the article's Creative Commons license, unless indicated otherwise in a credit line to the material. If material is not included in the article's Creative Commons license and your intended use is not permitted by statutory regulation or exceeds the permitted use, you will need to obtain permission directly from the copyright holder. To view a copy of this license, visit <a href="http://creativecommons.org/licenses/by/4.0/">http://creativecommons.org/licenses/by/4.0/</a>.

© The Author(s) 2018