



Physiological Response of *Crocosphaera watsonii* to Enhanced and Fluctuating Carbon Dioxide Conditions

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Abstract

We investigated the effects of elevated $p\text{CO}_2$ on cultures of the unicellular N_2 -fixing cyanobacterium *Crocosphaera watsonii* WH8501. Using CO_2 -enriched air, cultures grown in batch mode under high light intensity were exposed to initial conditions approximating current atmospheric CO_2 concentrations (~ 400 ppm) as well as CO_2 levels corresponding to low- and high-end predictions for the year 2100 (~ 750 and 1000 ppm). Following acclimation to CO_2 levels, the concentrations of particulate carbon (PC), particulate nitrogen (PN), and cells were measured over the diurnal cycle for a six-day period spanning exponential and early stationary growth phases. High rates of photosynthesis and respiration resulted in biologically induced $p\text{CO}_2$ fluctuations in all treatments. Despite this observed $p\text{CO}_2$ variability, and consistent with previous experiments conducted under stable $p\text{CO}_2$ conditions, we observed that elevated mean $p\text{CO}_2$ enhanced rates of PC production, PN production, and growth. During exponential growth phase, rates of PC and PN production increased by ~ 1.2 - and ~ 1.5 -fold in the mid- and high- CO_2 treatments, respectively, when compared to the low- CO_2 treatment. Elevated $p\text{CO}_2$ also enhanced PC and PN production rates during early stationary growth phase. In all treatments, PC and PN cellular content displayed a strong diurnal rhythm, with particulate C:N molar ratios reaching a high of 22:1 in the light and a low of 5.5:1 in the dark. The $p\text{CO}_2$ enhancement of metabolic rates persisted despite $p\text{CO}_2$ variability, suggesting a consistent positive response of *Crocosphaera* to elevated and fluctuating $p\text{CO}_2$ conditions.

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Introduction

Anthropogenic emissions and land use change are increasing the concentration of carbon dioxide (CO_2) in the atmosphere and surface ocean waters [1]. The predicted effects of elevated CO_2 partial pressure ($p\text{CO}_2$) and consequent ocean acidification (OA) on marine ecosystems include a potential upregulation of metabolic processes by CO_2 -limited phytoplankton species [2]. Because the carboxylating enzyme ribulose-1,5-bisphosphate carboxylase-oxygenase (RuBisCO) is typically not saturated under ambient surface seawater $p\text{CO}_2$, many phytoplankton groups invest energy in carbon concentrating mechanisms (CCMs) to increase CO_2 concentrations at the catalytic site [3]. Hence, under future OA conditions, phytoplankton with low-affinity RuBisCO could potentially down-regulate CCMs and reallocate energy and elemental resources to allow for increased carbon (C) fixation and growth rates [4]. Indeed, elevated $p\text{CO}_2$ has been shown to stimulate C fixation by select monocultures of phytoplankton and natural assemblages (reviewed in [5]).

Elevated $p\text{CO}_2$ appears to have a particularly strong metabolic enhancement in dinitrogen (N_2)-fixing (diazotrophic) cyanobacteria. Initial laboratory experiments using strain IMS101 of *Trichodesmium*, a group of filamentous, non-heterocystous cyanobacteria, showed that doubling $p\text{CO}_2$ increases N_2 fixation rates by 35–138% and C fixation rates by 23–40% [4,6–8]. Likewise, the unicellular cyanobacterium *Crocosphaera* strain WH8501

displays increased rates of C and N_2 fixation under elevated $p\text{CO}_2$ conditions [9]. *Trichodesmium* and *Crocosphaera* are dominant diazotrophic taxa in oligotrophic open-ocean environments [10], where the bioavailable nitrogen (N) from N_2 fixation can fuel up to half of production exported from the euphotic zone [11]. In theory, a global $p\text{CO}_2$ enhancement of marine N_2 fixation could increase oceanic C uptake and export, producing a negative feedback to climate change [6].

In contrast to laboratory findings, recent field experiments using natural diazotrophic assemblages do not reliably show that raising $p\text{CO}_2$ enhances N- or C-based productivity. Elevating $p\text{CO}_2$ in bottle incubations has been shown to increase N_2 fixation rates by *Trichodesmium* colonies isolated from the Subtropical Atlantic and the Gulf of Mexico [12,13] but not by *Trichodesmium* colonies isolated from the North Pacific Subtropical Gyre [14]. Furthermore, experiments using whole water diazotrophic assemblages from the North and South Pacific gyres have found no relationship between $p\text{CO}_2$ and N_2 fixation rates [15,16]. The inconsistent results from field incubations highlight the importance of assessing how the effect of $p\text{CO}_2$ on diazotrophs is influenced by other factors, including community composition [14,17], physiology, and environmental conditions [18].

One methodological challenge in $p\text{CO}_2$ manipulation studies is producing realistic timescales for $p\text{CO}_2$ perturbations. In nature, seawater $p\text{CO}_2$ varies spatially and temporally: phytoplankton experience $p\text{CO}_2$ fluctuations on diurnal, seasonal, episodic, and

long-term (e.g. OA-driven) timescales [19]. However, most laboratory OA studies grow cultures under stable $p\text{CO}_2$ treatments, allowing cultures to acclimate to a steady $p\text{CO}_2$ for multiple generations. These stable $p\text{CO}_2$ conditions may affect phytoplankton differently than the dynamic $p\text{CO}_2$ experienced in marine ecosystems; for instance, energetic costs associated with resource allocation may be minimized under stable $p\text{CO}_2$ regimes [20].

Predicting the future response of diazotrophic assemblages to OA will require an assessment of how $p\text{CO}_2$ affects marine diazotrophs under variable environmental conditions and physiological states. Here, we present data from experiments tracking the growth, PC and PN production rates of *Crocospaera watsonii* WH8501 cultures bubbled with air at three CO_2 levels (~ 400 , 750, 1000 ppm), while allowing for biologically induced $p\text{CO}_2$ variability in the culture medium resulting from photosynthesis and respiration. This approach contributes to the existing literature on potential responses of marine diazotrophs to future OA, but expands from previous studies by testing the effect of elevated $p\text{CO}_2$ under a dynamic $p\text{CO}_2$ environment. Our results suggest a consistent response of this organism to elevated $p\text{CO}_2$ under variable $p\text{CO}_2$ conditions.

Methods

Culture conditions

Unialgal stock cultures of *Crocospaera watsonii* strain WH8501 were grown in 0.2 μm -filtered, nitrogen-free YBCII medium [21] using 40 $\mu\text{mol L}^{-1}$ K_2HPO_4 . Cultures were not axenic, but heterotrophic bacterial counts were kept at low levels ($1.4\text{--}2.1 \times 10^5$ cells mL^{-1}). Light was provided using cool white fluorescent bulbs set on a 12:12 light/dark cycle. Stock cultures were grown at 24°C and 250 $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$. For the experiment, cultures were grown at 30°C, a temperature which promotes optimal growth of *Crocospaera* in the laboratory [22] and at which high abundances of *Crocospaera* cells have been observed at sea [23]. Incoming irradiance for the experiment was 1000 $\mu\text{mol quanta m}^{-2} \text{s}^{-1}$ as measured by a Biospherical light meter, a level which is saturating but not inhibitory for *Crocospaera* WH8501 [24]. Cultures were stirred at least once a day with magnetic stir bars to minimize cells sticking to the glass. The $p\text{CO}_2$ was manipulated by gently bubbling cultures with commercially prepared air/ CO_2 mixtures of ~ 400 ppm ('low- CO_2 '), ~ 750 ppm ('mid- CO_2 '), and ~ 1000 ppm ('high- CO_2 '). Parent cultures were grown under these CO_2 , light, and temperature conditions for seven days ($\sim 3\text{--}4$ generations) before productivity rates were measured.

Experimental Design

Crocospaera cultures were grown under three CO_2 treatments and monitored over a six-day period (day 0–day 5). Triplicate bottle replicates were used for each CO_2 treatment. Preceding the experiment, 2 L glass bottles were filled with 0.2 μm -filtered media that was pre-equilibrated to target $p\text{CO}_2$ levels. Initial CO_2 equilibration of the media was verified by measuring a stable $p\text{CO}_2$ in outflowing gas (>24 h equilibration) using a LI-840 LICOR gas analyzer (Biosciences). The pH of each replicate was measured and converted through CO2calc (see below) to produce initial $p\text{CO}_2$ values of 404 ± 23 , 724 ± 51 , and 916 ± 34 μatm for low-, mid-, and high- CO_2 treatments, respectively. To initiate the experiment, parent cultures in exponential growth phase were diluted into the pre-equilibrated media, producing initial biomass concentrations of 3.9–4.6 $\mu\text{g chlorophyll } a$ (Chl *a*) L^{-1} (Table 1). Because parent cultures were shifting media $p\text{CO}_2$ through

biological processes, the addition of parent cultures to pre-equilibrated media altered the initial $p\text{CO}_2$ values to 355 ± 14 , 600 ± 21 , and 788 ± 11 μatm (Table S1). Bottles were gently bubbled with air/ CO_2 mixtures at 50 mL min^{-1} throughout the experiment. Each replicate was sampled once daily after the sixth hour of light (L6) from day 0–day 2, then four times daily after the sixth and twelfth hours of light and darkness (L6, L12, D6, and D12) from day 3–day 5. The pH in each replicate was measured at the time of sampling and samples were preserved for particulate carbon (PC), particulate nitrogen (PN), Chl *a*, and flow cytometric cell counts (FCM).

Analytical Measurements

For PC/PN and Chl *a* measurements, three subsamples of 5–50 mL (depending on cell density) were withdrawn from each replicate and filtered onto glass fiber filters (GF/F, Whatman), using pre-combusted GF/F filters for PC/PN. Samples were immediately frozen at -80°C (PC/PN) or -20°C (Chl *a*). PC/PN samples were dried at 60°C overnight, packaged into silver and tin capsules, and analyzed using a Carlo Erba elemental analyzer. Acetanilide (71.09% C and 10.36% N by weight) served as a standard, and filter blanks were $<10\%$ of total C and N content. Chl *a* was extracted in 90% acetone at -20°C for 48 hours and analyzed with a Turner Model 10-AU fluorometer using the acidification method of Strickland and Parsons [25]. On day 0 and day 5 L6 time points, 25 mL samples were withdrawn from GF/F filtrate and immediately frozen for soluble reactive phosphorus (assumed to be equivalent to PO_4) and NH_4 analyses. NH_4 concentrations were measured with a Technicon AutoAnalyzer II, using a modified indophenol blue method [26] and PO_4 via the standard ascorbic acid-molybdate method [25].

Crocospaera and heterotrophic bacterial cell densities were measured using FCM. Two 3-mL subsamples were withdrawn from each replicate, pipetted into 4 mL cryovials, and fixed with paraformaldehyde at a final concentration of 1% (volume volume^{-1}). Samples were inverted and allowed to sit in the dark for ~ 10 minutes before being frozen at -80°C . For analysis of *Crocospaera* cell densities, samples were thawed on ice in the dark then spiked with a known number of 3 μm Polysciences Fluoresbrite yellow-green beads and run on a Becton-Dickinson FASCaliber flow cytometer with a 488 nm laser. *Crocospaera* cells and beads were distinguished from other particulate matter by their side light scatter and fluorescence in orange wavelengths. The bead count determined the volume of sample run, and thus the concentration of *Crocospaera* cells. A similar method was used to enumerate the background heterotrophic bacteria in these cultures. The samples were spiked with Fluoresbrite 1 μm beads, stained with SYBR Green I according to the method of Marie et al. [27], and differentiated by their side light scatter and green fluorescence.

The pH of each replicate was measured directly using a VWR sympHony electrode calibrated with VWR buffers (NBS scale). pH values were converted to $p\text{CO}_2$ by assuming a constant total alkalinity (TA) for the YBCII medium (2500 μM). This TA value was determined by analyzing DIC and $p\text{CO}_2$ of a separate batch of YBCII medium according to the methods of Bandstra et al. [28], then the program CO2calc [29] was used to convert DIC and $p\text{CO}_2$ to TA (with CO_2 constants from Merbach et al. [30] refit by Dickson and Millero [31], and a correction to the NBS scale by the CO2calc program). Finally, CO2calc was used to calculate $p\text{CO}_2$ from our measured pH data and the constant TA.

Our assumption of constant TA of the YBCII medium throughout the experiment is based on the observation that bubbling with air/ CO_2 mixtures perturbs DIC but not TA [32];

Table 1. Time series biomass measurements for cultures of *Crocospaera watsonii* WH8501 grown under three CO₂ treatments.

Time point	Low-CO ₂ treatment				Mid-CO ₂ treatment				High-CO ₂ treatment			
	PC	PN	Cells	Chl <i>a</i>	PC	PN	Cells	Chl <i>a</i>	PC	PN	Cells	Chl <i>a</i>
Day 0 L6	124 (12)	8.8 (1)	-	4.3 (0.6)	120 (4)	8.5 (0.7)	-	4.6 (0.3)	118 (2)	7.6 (0.1)	-	3.9 (0.4)
Day 1 L6	166 (24)	14.2 (2)	5.6E+05 (7.4E+04)	7.3 (1.6)	192 (1)	16.2 (0.2)	8.0E+05 (6.3E+04)	9 (0.4)	172 (21)	13.2 (2)	7.5E+05 (1.9E+04)	7.9 (0.4)
Day 2 L6	374 (145)	22 (4)	7.4E+05 (1.2E+05)	11.6 (1.5)	424 (73)	28.2 (3)	1.2E+06 (1.1E+05)	16.3 (3)	547 (63)	28.3 (2)	1.2E+06 (9.7E+03)	15.1 (1)
Day 3 L6	542 (65)	42.5 (6)	1.4E+06 (2.5 E+05)	24.8 (3)	706 (88)	55.4 (6)	2.4E+06 (1.9E+05)	39.6 (6)	740 (44)	55.9 (4)	2.6E+06 (1.3E+05)	35.3 (6)
Day 4 L6	833 (121)	66 (7)	2.3E+06 (3.1E+05)	31.6 (2.6)	1144 (174)	91.7 (10)	3.6E+06 (1.5E+06)	50.6 (9)	1308 (121)	97.4 (7)	4.8E+06 (2.1E+05)	49.1 (9)
Day 5 L6	1037 (58)	84.2 (3)	3.3E+06 (2.9E+05)	55.4 (2.1)	1589 (186)	130.2 (12)	6.7E+06 (4.4E+05)	82.5 (8)	1852 (65)	155.1 (2)	8.1E+06 (3.8E+05)	101 (2)

Concentrations of particulate carbon ($\mu\text{mol L}^{-1}$; PC), particulate nitrogen ($\mu\text{mol L}^{-1}$; PN), cells ($\# \text{ mL}^{-1}$), and chlorophyll *a* (Chl *a*; $\mu\text{g L}^{-1}$) are provided for L6 time points. Data are mean values from three replicate bottles; standard deviations are presented in parentheses. doi:10.1371/journal.pone.0110660.t001

thus, any change in TA through the experiment was due to biological activity. While the process of N₂ fixation does not affect TA, photosynthesis can have a small effect due to hydrogen ion uptake to balance anionic nutrient (N, P, and S) acquisition [33]. Assuming no inorganic N uptake (as cultures were grown in N-free media) and a 2.4:1 S:P uptake ratio (as in [33]), we estimate that the average PO₄ drawdown of $\sim 9 \mu\text{M}$ observed in our experiment (see *Results and Discussion*) increased TA by an average of $\sim 52 \mu\text{M}$ by day 5, generating a maximum *p*CO₂ error of $\sim 2\%$. In addition, we assume that calcium carbonate (CaCO₃) minerals were not precipitated in our experiment, as the presence of PO₄ has been shown to inhibit CaCO₃ precipitation, even at high CaCO₃ saturation states [34,35].

Rate calculations

Specific growth rates were calculated for each of the biomass parameters measured: cell density, PC, PN, and Chl *a* (Table 2). Growth rates (μ) were determined using Eq. 1,

$$\mu(\text{d}^{-1}) = \frac{\ln(N_T/N_0)}{\Delta T} \quad (1)$$

where N_T is the biomass at day 3, N_0 is the biomass at day 1, and ΔT is the time interval in days. The day 1–day 3 time interval was chosen for growth rate calculations because this was the phase of exponential growth (Fig. 1A).

Carbon-normalized PC and PN production rates (\sim net C and N₂ fixation rates) were calculated using Eq. 2,

$$\text{Production rate} = \frac{(N_T - N_0)/PC_0}{\Delta T} \quad (2)$$

where N_T is the biomass (PC or PN) at the final time point, N_0 is the biomass at the initial time point, PC_0 is the initial PC concentration, and ΔT is the time interval in days. Production rates were calculated for both exponential (day 1–day 3) and early stationary (day 3–day 5) growth phases.

Growth rates, PC and PN production rates were all calculated using data from L6 time points. Day 0 was excluded from these analyses due to missing FCM samples on this day.

Statistics

The effects of *p*CO₂ on growth rates, PC production, PN production, and molar C:N ratios were assessed using the one-way ANOVA. Differences between CO₂ treatments were determined using the Tukey Honest Significance Difference (HSD) test of multiple comparisons. All data reported in this study are averages from triplicate bottles. Statistical tests were run using the program R (<http://www.r-project.org/>).

Results and Discussion

Our study tested how enhanced *p*CO₂ affects the growth, PC and PN production rates of high-density *Crocospaera* cultures. In agreement with a previous study [9], we found that PC production, PN production, and growth rates were all positively correlated with *p*CO₂ (Table 2, Fig. 2). This *p*CO₂ enhancement was observed for *Crocospaera* cultures during both exponential and early stationary growth phases (Fig. 2). The high growth rates and cell densities observed in our study produced a strong diurnal rhythm of C and N metabolism in *Crocospaera* cultures as well as daily *p*CO₂ variability (Fig. 1).

Table 2. Biomass-specific growth rates of *Crocospaera watsonii* WH8501 cultures grown under three CO₂ treatments.

	Specific Growth rate μ (d ⁻¹)			Tukey HSD p -value	
	Low-CO ₂	Mid-CO ₂	High-CO ₂	Low-CO ₂ vs. high-CO ₂	Low-CO ₂ vs. mid-CO ₂
Cell density	0.45 (0.02)	0.54 (0.02)	0.60 (0.02)	<0.001	<0.01
Particulate nitrogen	0.54 (0.02)	0.60 (0.05)	0.71 (0.04)	<0.01	0.2
Particulate carbon	0.58 (0.02)	0.63 (0.06)	0.71 (0.06)	<0.05	0.44
Chlorophyll <i>a</i>	0.60 (0.06)	0.72 (0.05)	0.72 (0.07)	0.12	0.14

Rates were calculated from L6 time points between day 1 and day 3 (exponential growth phase). Tukey HSD p -values are provided for comparisons among CO₂ treatments. Standard deviations of growth rates from three replicate bottles are presented in parentheses.

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Diurnal rhythm in growth and $p\text{CO}_2$

Diazotrophic cyanobacteria employ various mechanisms to separate the oxygen (O₂) evolved through photosynthesis from the enzyme nitrogenase, which catalyzes biological N₂ fixation and is irreversibly inactivated by O₂ [36]. *Crocospaera* circumvents this problem by restricting N₂ fixation to the nighttime, when O₂ is not being produced. The energy needed to fix N₂ is generated photosynthetically in the light and stored primarily as carbohydrate granules [37]; respiration of these organic C reserves fuels N₂ fixation in the dark. The temporal separation and energetic linkage of photosynthesis and N₂ fixation in *Crocospaera* produces a daily pattern in the timing and magnitude of PC and PN production and loss. In our study, cultures were grown under high light conditions, producing high growth rates and an especially pronounced diurnal rhythm.

We observed a strong daily pattern in PC production, PN production, and cell division by *Crocospaera* in all CO₂ treatments (Fig. 1). The PC concentration of *Crocospaera* cultures fluctuated widely between the light and dark periods: PC increased 48–216% in the light (~C fixation) and decreased 17–79% in the dark (~respiration) (Fig. 1B). This substantial dark PC loss is consistent with previous studies of *Crocospaera* and reflects the respiration of carbohydrate reserves to fuel N₂ fixation [37–39]. A fraction of PC may have also been exuded from cells, as it has been shown that *Crocospaera* WH8501 can release ~10% of total C content daily as extracellular polymeric substances [38]; however, dissolved organic C was not measured in our study. The sharp increase in cell concentration following the D12 measurement indicates that cells divided in the first half of the light period (Fig. 1C).

PN production was restricted to the dark period, when PN concentrations increased between 48 and 93% (Fig. 1A). Rates of PN increase approximate net N₂ fixation rates, though we cannot rule out NO₃ or NH₄ utilization as driving some small fraction of PN production. While cultures were grown in N-free media [21], the initial dilution of the parent cultures into fresh media resulted in NH₄ concentrations of 0.3±0.1 μM on day 0. By day 5, NH₄ had increased to 1.1±0.3 μM, supporting the interpretation that the rate of accumulation of inorganic plus particulate N in our cultures was fueled by N₂ fixation. The accumulated NH₄ had presumably been fixed by *Crocospaera* and released from cells; *Crocospaera* have been previously observed to release 23–67% of recently fixed N [38]. PN decreased slightly (1–8%) during the light period of the experiment (Fig 1A), which may indicate daily NH₄ release. The total NH₄ accumulated through the experiment (0.8 μM) represents 0.5–1% of total PN production (day 5–day 0, Table 1).

Together, the high rates of PN production, PC production in the day (~C fixation) and PC loss in the night (~C respiration) led

to large fluctuations in the particulate C:N ratio over the daily cycle: molar C:N ratios in our study ranged from 5.5–22.1 (Fig. 1D). The C:N ratios at D12 time points had a relatively consistent range (~5.5–8) encompassing the 6.6 ratio predicted from Redfield stoichiometry [40]. Elevated C:N ratios at L6, L12, and D6 time points were driven by the accumulation of organic C reserves to fuel dark N₂ fixation and other cellular processes. These daily C:N deviations were independent of $p\text{CO}_2$ treatment. Previous studies have reported less dramatic stoichiometric fluctuations in *Crocospaera*, with daily C:N content ranging from 6.5–8.5 [39] and 5.0–8.8 [38]. The strong daily C:N deviations observed in our study are consistent with metabolic rates at high growth rates (~0.5 d⁻¹) of cultures grown at optimum temperature (30°C, [20]) and saturating incoming irradiance (1000 μmol quanta m⁻² s⁻¹, [24]). Our experiment spanned both exponential and early stationary growth phases. Cultures grew exponentially from day 0 to day 3, at which point growth rates began to decline (Fig. 1A). The shift in growth phase is evident from non-linearity of natural log-normalized PN growth curves at L6 time points (Fig. 1A) and from decreased PC and PN production rates from day 3–day 5 (Fig. 2). Declines in the magnitude of daily PC fluctuations indicate decreased rates of photosynthesis (~positive derivative, in light) and respiration (~negative derivative, in dark) after day 3 (Fig. 1B). *Crocospaera* cultures were grown in an artificial medium initially containing an ample supply of macro and micronutrients [16], and PO₄ remained replete throughout the experiment (PO₄ decreased from 36±1.8 μM on day 0 to 27±1.8 μM on day 5, data not shown). Thus, although we cannot exclude the possibility of limitation by a micronutrient, we hypothesize that the shift to early stationary growth phase resulted from self-shading or a decreased RuBisCO carboxylation efficiency during the second half of the photoperiod, either through direct CO₂ limitation or through competitive inhibition of carboxylation from photorespiration at high O₂:CO₂ ratios [41].

The high *Crocospaera* growth rates and cell densities affected the stability of C chemistry within CO₂ treatments: photosynthesis and respiration produced $p\text{CO}_2$ variability despite continuously bubbling cultures with air/CO₂ mixtures (Fig. 1E). The magnitude of $p\text{CO}_2$ variability increased throughout the experiment, and by day 5, the high-CO₂ treatment had fluctuated between 1216 and 96 μatm (Fig. 1E). In addition, by day 4, all cultures, independent of $p\text{CO}_2$ treatment, reached consistent minimum $p\text{CO}_2$ values close to 100 μatm in measurements taken at the end of the light cycle (L12); hence, ~100 μatm appears to represent a physiological limit for the uptake of inorganic C by this organism. Despite the extreme $p\text{CO}_2$ fluctuations within treatments, mid- and high-CO₂ treatments always had higher $p\text{CO}_2$ values than the low-CO₂ treatment (which fluctuated between 376 and 74 μatm), with the exception of the final time point. To separate the effect of

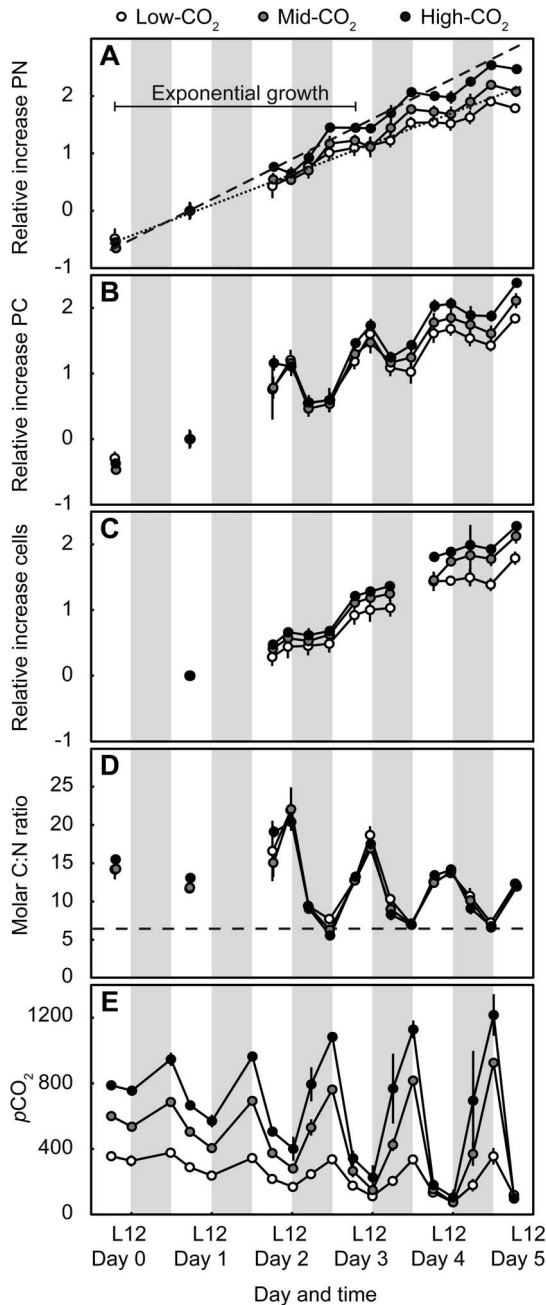


Figure 1. Growth of *C. watsonii* WH8501 batch cultures over a 6-day period under three CO₂ treatments. Shown are concentrations of PN (a), PC (b), and cells (c), molar C:N ratios (d), and $p\text{CO}_2$ in μatm (e) within each treatment. For (a–c), the concentrations for each time point (Table S1) were first normalized to the concentration at the day 1 L6 time point, then ln-transformed. The derived slopes between day 1 L6 and day 3 L6 time points correspond to the exponential growth rates (μ) as shown in Table 2. The lines in (a) represent linear regressions through the day 1, day 2, and day 3 L6 time points for high-CO₂ (dashed line) and low-CO₂ (dotted line) treatments. The regression lines have been extended to the full time period (day 0–5) for visualization of exponential growth (day 0–3 L6 time points) transitioning to early stationary growth (L6 time points after day 3). The dotted line in (d) represents the 6.6 C:N ratio expected from Redfield stoichiometry. Shaded areas represent the dark periods. Error bars represent standard deviations from three replicates.
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inflowing $p\text{CO}_2$ from the possible confounding effect of cell densities, we present growth rates from day 1 to day 3 when $p\text{CO}_2$ fluctuations were less extreme and treatments did not overlap in $p\text{CO}_2$ range (Table 2); the mean $p\text{CO}_2$ levels measured in this time interval were 252, 477, and 665 μatm for the low-, mid-, and high-CO₂ treatments, respectively. The response of *Crocospaera* to elevated $p\text{CO}_2$ combined with daily, biologically induced $p\text{CO}_2$ fluctuations may have ecological implications for bloom scenarios (discussed below).

$p\text{CO}_2$ enhancement of growth, PC and PN production

Consistent with previous studies of *Crocospaera* WH8501 [7] and *Trichodesmium* IMS101 [4,6–8], we found that elevating $p\text{CO}_2$ increased *Crocospaera* growth, PC and PN production rates (Table 2, Fig. 2). Growth in the high-CO₂ treatment was significantly higher than in the low-CO₂ treatment for growth rates specific to PC, PN, and cell density (Table 2). Cell-specific growth rates were also significantly higher in the mid-CO₂ treatment than the low-CO₂ treatment (Table 2). Chl *a* growth rates were not significantly different between CO₂ treatments, possibly because of the large coefficient of variation between replicates (7–10%) (Table 2). Biomass-normalized PC and PN

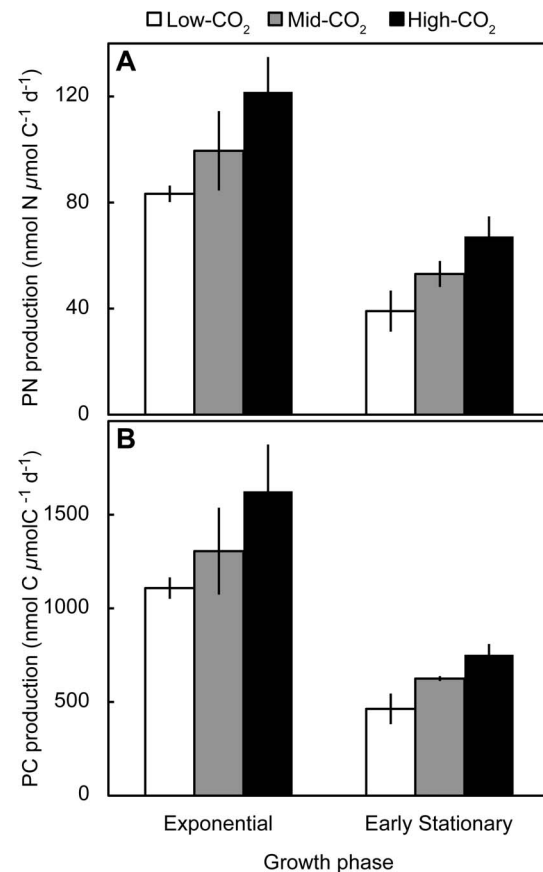


Figure 2. Carbon-normalized PN (a) and PC production rates (b) of *Crocospaera* WH8501 cultures grown under three CO₂ treatments during periods of exponential (day 1–day 3) and early stationary (day 3–day 5) growth phases. Production rates are calculated as increases in PC and PN concentrations (data provided in Table 1) per time normalized to initial PC concentrations within the time interval. Error bars represent standard deviations from three replicates.
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production rates were significantly enhanced in mid-CO₂ and high-CO₂ treatments compared to the low-CO₂ treatment (Fig. 2). The *p*CO₂ enhancement of PC and PN production rates was observed both during exponential (day 1–day 3) and early stationary (day 3–day 5) growth phases (Fig. 2).

The only previous CO₂ manipulation study using *Crocospaera* strain WH8501 tested a *p*CO₂ range of 190–50 μ atm and observed that raising ambient *p*CO₂ (380 μ atm) to 750 μ atm produced 1.2- and 1.4-fold higher rates of C and N₂ fixation, respectively [9]. In our study, PC and PN production rates during exponential growth were both \sim 1.2-fold higher in the mid-CO₂ treatment than the low-CO₂ (Fig. 2). The magnitude of *p*CO₂ enhancement we observed is strikingly similar to those reported by Fu et al. [9], especially considering that the environmental conditions utilized in our study differed from the low light (80 μ mol quanta m⁻² s⁻¹), steady *p*CO₂ conditions of Fu et al. [9]. In our study, including a higher *p*CO₂ treatment displayed even larger enhancements: both PC and PN production rates were \sim 1.5-fold higher in the high-CO₂ treatment than the low-CO₂ treatment. Higher *p*CO₂ treatments would need to be included to determine the threshold *p*CO₂ condition that saturates C and N₂ fixation rates of *Crocospaera* WH8501.

Conclusions and ecological implications

We observed that elevated *p*CO₂ conditions significantly enhanced PC production, PN production, and growth rates of *Crocospaera* strain WH8501. This *p*CO₂ enhancement persisted despite biologically induced *p*CO₂ variability in all treatments. By allowing photosynthesis and respiration to drive *p*CO₂ deviations from target values, our methods contrast with those of many previous OA studies, which often keep cultures optically dilute and/or do not report the measured *p*CO₂ time course for each replicate. Though *p*CO₂ in our study varied within treatments, the mid- and high-CO₂ treatments had higher *p*CO₂ values than the low-CO₂ treatment for nearly all time points (Fig. 1E). Thus, the higher rates of growth, PC and PN production observed in mid- and high-CO₂ treatments can be attributed to the elevated *p*CO₂. Furthermore, the differences between treatments observed in the early stationary growth phase reflect low-end estimates of potential *p*CO₂ enhancements: cell densities were highest in the elevated CO₂ treatments, possibly leading to more severe growth limitations and dampening evidence for CO₂ enhancement.

Our study allows for new insights into the response of *Crocospaera* to enhanced *p*CO₂ under a variable *p*CO₂ environment. Elevated mean *p*CO₂ enhanced the growth of *Crocospaera* cultures despite large *p*CO₂ fluctuations on time-scales of less than a generation time, showing that this organism does not need to be acclimated to a stable *p*CO₂ regime to benefit from elevated *p*CO₂. Testing the response of microbes to CO₂ perturbations on multiple timescales is ecologically relevant, because net community metabolism, temperature and salinity effects on CO₂ solubility, and advective processes cause *p*CO₂ fluctuations on episodic, diurnal, and seasonal timescales. Phytoplankton will experience the long-term OA *p*CO₂ signal superimposed onto this existing *p*CO₂ variability. Furthermore, future OA will increase the DIC:TA ratio of surface waters, leading to a

reduced capacity to buffer processes like photosynthesis and respiration, ultimately increasing the magnitude of *p*CO₂ fluctuations [42].

The *p*CO₂ fluctuations observed in our study are probably more extreme than the natural variability experienced by *Crocospaera* populations in open-ocean habitats. In the North Pacific Subtropical Gyre, surface *p*CO₂ varies by \sim 20–50 μ atm seasonally [43]; mesoscale features in this region may cause biologically induced *p*CO₂ swings of \sim 150 μ atm [44]. However, aggregated cells may experience larger *p*CO₂ swings; for example, *Crocospaera nifH* genes have been observed associated with *Trichodesmium* colonies [14] and could thus experience more extreme *p*CO₂ fluctuations during blooms and subsequent crashes in these concentrated-biomass microhabitats [45]. Regardless of the large magnitude of *p*CO₂ fluctuations employed in our study, our results suggest that elevated mean *p*CO₂ impacts the growth response of *Crocospaera* despite short-term variability.

Overall, our study contributes to the growing literature on the response of marine diazotrophs to elevated *p*CO₂. We observed that growth, PC and PN production rates of *Crocospaera* WH8501 were enhanced under elevated and variable *p*CO₂ in both exponential and early stationary growth phases. It should be noted that a recent study by Garcia et al. [46] found that *Crocospaera* strains WH0401 and WH0402 appear to be fully saturated under present day *p*CO₂ conditions (\sim 400 μ atm); thus, elevated *p*CO₂ seems to have strain-specific effects within *Crocospaera*. Further research investigating how community composition and environmental conditions regulate the response of marine diazotrophs to elevated *p*CO₂ will be key to predicting whether global rates of N₂ fixation will increase under future OA scenarios.

Supporting Information

Table S1 Time series measurements for cultures of *Crocospaera watsonii* WH8501 grown under three *p*CO₂ treatments. Measured pH, calculated *p*CO₂ (μ atm, see Methods) and concentrations of particulate carbon (μ mol L⁻¹; PC), particulate nitrogen (μ mol L⁻¹; PN), cells (# mL⁻¹), and chlorophyll *a* (Chl *a*; μ g L⁻¹) are provided for every time point available. Data are mean values from three replicate bottles; standard deviations are presented in parentheses. Dashes indicate no data available. (DOCX)

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Author Contributions

Conceived and designed the experiments: MRG AEW RML. Performed the experiments: MRG. Analyzed the data: MRG AEW RML. Contributed reagents/materials/analysis tools: AEW RML. Wrote the paper: MRG.

References

- Intergovernmental Panel on Climate Change (IPCC) (2013) Climate change 2013: The physical science basis. Working group I contribution to the fifth assessment report of the Intergovernmental Panel on Climate Change. Available at: <http://www.ipcc.ch/>
- Beardall J, Stojkovic S, Larsen S (2009) Living in a high CO₂ world: impacts of global climate change on marine phytoplankton. *Plant Ecol Div* 2: 191–205.
- Badger MR, Andrews TJ, Whitney SM, Ludwig M, Yellowlees DC, et al. (1998) The diversity and coevolution of Rubisco, plastids, pyrenoids, and chloroplast-based CO₂-concentrating mechanisms in algae. *Can J Bot* 7: 1052–1071.
- Kranz S, Sültemeyer D, Richter KU, Rost B (2009) Carbon acquisition in *Trichodesmium*: The effect of *p*CO₂ and diurnal changes. *Limnol Oceanogr* 54: 548–559.

5. Doney SC, Fabry VJ, Feely RA, Kleypas JA (2009) Ocean acidification: the other CO₂ problem. *Ann Rev Mar Sci* 1: 169–192.
6. Hutchins D, Fu FX, Zhang Y, Warner M, Feng Y, et al. (2007) CO₂ control of *Trichodesmium* N₂ fixation, photosynthesis, growth rates, and elemental ratios: Implications for past, present, and future ocean biogeochemistry. *Limnol Oceanogr* 52: 1293–1304.
7. Barcelos e Ramos JB, Biswas H, Schulz K, LaRoche J, Riebesell U (2007) Effect of rising atmospheric carbon dioxide on the marine nitrogen fixer *Trichodesmium*. *Global Biogeochem Cy* 21: GB2028.
8. Levitan O, Rosenberg G, Setlik I, Setlikova E, Grigel J, et al. (2007) Elevated CO₂ enhances nitrogen fixation and growth in the marine cyanobacterium *Trichodesmium*. *Glob Change Biol* 13: 531–538.
9. Fu FX, Mulholland MR, Garcia NS, Beck A, Bernhardt PW, et al. (2008) Interactions between changing pCO₂, N₂ fixation, and Fe limitation in the marine unicellular cyanobacterium *Crocosphaera*. *Limnol Oceanogr* 53: 2472–2484.
10. Luo YW, Doney S, Anderson L, Benavides M, Berman-Frank I, et al. (2012) Database of diazotrophs in global ocean: Abundances, biomass and nitrogen fixation rates. *Earth Syst Sci Data* 4: 47–73.
11. Karl D, Letelier R, Tupas L, Dore J, Christian J, et al. (1997) The role of nitrogen fixation in biogeochemical cycling in the subtropical North Pacific Ocean. *Nature* 388: 533–538.
12. Hutchins DA, Mulholland MR, Fu F (2009) Nutrient cycles and marine microbes in a CO₂-enriched ocean. *Oceanography* 22: 128–145.
13. Lomas M, Hopkinson B, Losh J, Ryan D, Shi D, et al. (2012) Effect of ocean acidification on cyanobacteria in the subtropical North Atlantic. *Aquat Microb Ecol* 66: 211–222.
14. Gradoville MR, White AE, Böttjer D, Church MJ, Letelier RM. (2014) Diversity trumps acidification: Lack of evidence for carbon dioxide enhancement of *Trichodesmium* community nitrogen or carbon fixation at Station ALOHA. *Limnol Oceanogr* 59: 645–659.
15. Law CS, Breitbarth E, Hoffmann IJ, McGraw CM, Langlois RJ, et al. (2012) No stimulation of nitrogen fixation by non-filamentous diazotrophs under elevated CO₂ in the South Pacific. *Glob Change Biol* 18: 3004–3014.
16. Böttjer D, Karl DM, Letelier RM, Viviani DA, Church MJ (2014) Experimental assessment of diazotroph responses to elevated pCO₂ in the North Pacific Subtropical Gyre. *Glob Biogeochem Cycles* 28: 601–616.
17. Hutchins DA, Fu F-X, Webb EA, Walworth N, Tagliabue A (2013) Taxon-specific response of marine nitrogen fixers to elevated carbon dioxide concentrations. *Nature Geosci* 6: 790–795.
18. Shi D, Kranz SA, Kim JM, Morel FMM (2012) Ocean acidification slows nitrogen fixation and growth in the dominant diazotroph *Trichodesmium* under low-iron conditions *Proc Natl Acad Sci USA* 109: E3094–3100.
19. Joint I, Doney SC, Karl DM (2011) Will ocean acidification affect marine microbes? *ISME* 5: 1–7.
20. Geider RJ, Moore CM, Ross ON (2009) The role of cost-benefit analysis in models of phytoplankton growth and acclimation. *Plant Ecol Div* 2: 165–178.
21. Chen YB, Zehr JP, Mellon M (1996) Growth and nitrogen fixation of the diazotrophic filamentous nonheterocystous cyanobacterium *Trichodesmium* sp. IMS 101 in defined media: Evidence for a circadian rhythm. *J Phycol* 32: 916–923.
22. Fu FX, Yu E, Garcia NS, Gale J, Luo Y, et al. (2014) Differing responses of marine N₂ fixers to warming and consequences for future diazotroph community structure. *Aquat Microb Ecol* 72: 33–46.
23. Moisaner PH, Beinart RA, Hewson I, White AE, Johnson KS, et al. (2010) Unicellular cyanobacterial distributions broaden the oceanic N₂ fixation domain. *Science* 327: 1512–1514.
24. Goebel NL, Edwards CA, Carter BJ, Achilles KM, Zehr JP (2008) Growth and carbon content of three different-sized diazotrophic cyanobacteria observed in the subtropical North Pacific. *J Phycol* 44: 1212–1220.
25. Strickland J, Parsons T (1972) A practical handbook of seawater analysis, 2nd ed. Ottawa, ON: Fisheries Research Board of Canada.
26. US EPA (1983) Methods for chemical analysis of water and waste. Determination of nitrogen as ammonia. Method 350.1.
27. Marie D, Partensky F, Jacquet S, Vaulot D (1997) Enumeration and cell cycle analysis of natural populations of marine picoplankton by flow cytometry using the nucleic acid stain SYBR Green I. *Appl Env Microbiol* 63: 186–193.
28. Bandstra L, Hales B, Takahashi T (2006) High-frequency measurements of total CO₂: Method development and first oceanographic observations. *Mar Chem* 100: 24–38.
29. Robbins L, Hansen M, Kleypas J, Meylan S (2010) CO₂calc—a user-friendly seawater carbon calculator for Windows, Mac OS X, and iOS (iPhone). US Geological Survey Open-File Report 2010–1280, 17 p.
30. Mehrbach C, Culbertson C, Hawley J, Pytkowicz R (1973) Measurement of the apparent dissociation constants of carbonic acid in seawater at atmospheric pressure. *Limnol Oceanogr* 18: 897–907.
31. Dickson A, Millero F (1987) A comparison of the equilibrium constants for the dissociation of carbonic acid in seawater media. *Deep-Sea Res* 34: 1733–1743.
32. Riebesell U, Fabry VJ, Nansson LN, Gattuso J-P (eds) (2010) Guide to best practices for ocean acidification research and data reporting. 260 pp. Luxembourg: Publications Office of the European Union. Available at: <http://www.epoca-project.eu/index.php/guide-to-best-practices-for-ocean-acidification-research-and-data-reporting.htm>
33. Wolf-Gladrow DA, Zeebe RE, Klaas C, Körtzinger A, Dickson AG (2007) Total alkalinity: The explicit conservative expression and its application to biogeochemical processes. *Mar Chem* 106: 287–300.
34. Reddy M (1977) Crystallization of calcium carbonate in the presence of trace concentrations of phosphorous-containing anions: I. Inhibition by phosphate and glycerophosphate ions at pH 8.8 and 25°C. *J Cryst Growth* 41: 287–295.
35. Kranz S, Wolf-Gladrow D, Nehrke G, Langer G, Rost B (2010) Calcium carbonate precipitation induced by the growth of the marine cyanobacteria *Trichodesmium*. *Limnol Oceanogr* 55: 2563–2569.
36. Gallon J (1992) Tansley Review No. 44. Reconciling the incompatible: N₂ fixation and O₂. *New Phytol* 122: 571–609.
37. Dron A, Rabouille S, Claquin P, Chang P, Raimbault V, et al. (2012) Light:dark (12:12 h) quantification of carbohydrate fluxes in *Crocosphaera watsonii*. *Aquat Microb Ecol* 68: 43–55.
38. Dron A, Rabouille S, Claquin P, Le Roy B, Talec A, et al. (2012) Light-dark (12:12) cycle of carbon and nitrogen metabolism in *Crocosphaera watsonii* WH8501: relation to the cell cycle. *Environ Microbiol* 14: 967–981.
39. Mohr W, Intermaggio MP, LaRoche J (2010) Diel rhythm of nitrogen and carbon metabolism in the unicellular, diazotrophic cyanobacterium *Crocosphaera watsonii* WH8501. *Environ Microbiol* 12: 412–421.
40. Redfield AC (1958) The biological control of chemical factors in the environment. *Am Sci* 46: 205–221.
41. Raven JA, Giordano M, Beardall J, Maberly SC (2012) Algal evolution in relation to atmospheric CO₂: carboxylases, carbon-concentrating mechanisms and carbon oxidation cycles. *Phil Trans R Soc B* 367: 493–507.
42. Egleston ES, Sabine CL, Morel FMM (2010) Revelle revisited: Buffer factors that quantify the response of ocean chemistry to changes in DIC and alkalinity. *Global Biogeochem Cy* 24: GB1002.
43. Dore JE, Lukas R, Sadler DW, Church MJ, Karl DM (2009) Physical and biogeochemical modulation of ocean acidification in the central North Pacific. *Proc Natl Acad Sci* 106: 12235–12240.
44. Mahadevan A, Lévy M, Mémery L (2004) Mesoscale variability of sea surface pCO₂: What does it respond to? *Global Biogeochem Cy* 18: GB1017.
45. Flynn KJ, Blackford JC, Baird ME, Raven JA, Clark DR, et al. (2012) Changes in pH at the exterior surface of plankton with ocean acidification. *Nat Clim Change* 2: 510–513.
46. Garcia NS, Fu FX, Hutchins DA (2013) Colimitation of the unicellular photosynthetic diazotroph *Crocosphaera watsonii* by phosphorus, light, and carbon dioxide. *Limnol Oceanogr* 58: 1501–1512.