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# Constructing a model including the cryptic sulfur cycle in Chesapeake Bay requires judicious choices for key processes and parameters <sup>☆</sup>



Rui Jin <sup>a,\*</sup>, Marie-Aude Pradal <sup>a</sup>, Kaley Hantsoo <sup>a</sup>, Anand Gnanadesikan <sup>a</sup>,  
Pierre St-Laurent <sup>b</sup>, Christian J. Bjerrum <sup>c</sup>

<sup>a</sup> Department of Earth and Planetary Sciences, Johns Hopkins University, Baltimore, MD 21218, United States

<sup>b</sup> Virginia Institute of Marine Science, William & Mary, Gloucester Point, VA 23062, United States

<sup>c</sup> Department of Geoscience and Natural Resource Management, University of Copenhagen, Copenhagen 1165, Denmark

## ARTICLE INFO

### Method name:

SNP\_BUR\_DOM

### Keywords:

Modeling predictions

Hypoxia

Particle sinking

Nitrification

Stoichiometry

Optics absorption

## ABSTRACT

A new biogeochemical model for Chesapeake Bay has been developed by merging two published models – the ECB model of Da et al. (2018) that has been calibrated for the Bay but only simulates nitrogen, carbon and oxygen and the BioRedoxCNPS model of al Azhar et al. (2014) and Hantsoo et al. (2018) that includes cryptic sulfur cycling. Comparison between these models shows that judicious choices are required for key processes and parameters. This manuscript documents the sources of differences between the two published models in order to select the most realistic configuration for our new model.

- This study focuses on three sets of differences—processes only included in ECB (burial and dissolved organic matter), processes only included in BioRedoxCNPS (explicit dynamics for hydrogen sulfide, sulfate and nitrite, light attenuation that does not include CDOM or sediments), and differences in parameters common to the two codes.
- Sensitivity studies that highlight particular choices (absorption by dissolved organic matter, nitrification rates, stoichiometric ratios) are also shown.
- The new model includes sulfur cycling and has comparable skill in predicting oxygen as ECB, but also has improved simulation of nitrogen species compared with both original codes.

## Specifications table

|  |   |
|--|---|
| Subject area:                          | Environmental Science   |
| More specific subject area:            | Oceanographic modeling  |
| Name of your method:                   | SNP_BUR_DOM   |
| Name and reference of original method: | ECB-Da et al. (2018), BioRedoxCNPS-al Azhar et al. (2014), Hantsoo et al. (2018)  |
| Resource availability:                 | ( <a href="https://github.com/RuiJinSZ/Merged-Model-MethodsX/">13:italic</a> ) <a href="https://github.com/RuiJinSZ/Merged-Model-MethodsX/">https://github.com/RuiJinSZ/Merged-Model-MethodsX/</a> ( <a href="https://github.com/RuiJinSZ/Merged-Model-MethodsX/">13:italic</a> ) |

<sup>☆</sup> Related research article Jin, R., Pradal, M., Hantsoo, K., Gnanadesikan, A., St-laurent, P., 2023. Comparing two ocean biogeochemical models of Chesapeake Bay with and without the sulfur cycle instead highlights the importance of particle sinking, burial, organic matter, nitrification and light attenuation. *Ocean Model.* 182, 102175. <https://doi.org/10.1016/j.ocemod.2023.102175>

\* Corresponding author.

E-mail address: [ruijin@jhu.edu](mailto:ruijin@jhu.edu) (R. Jin).

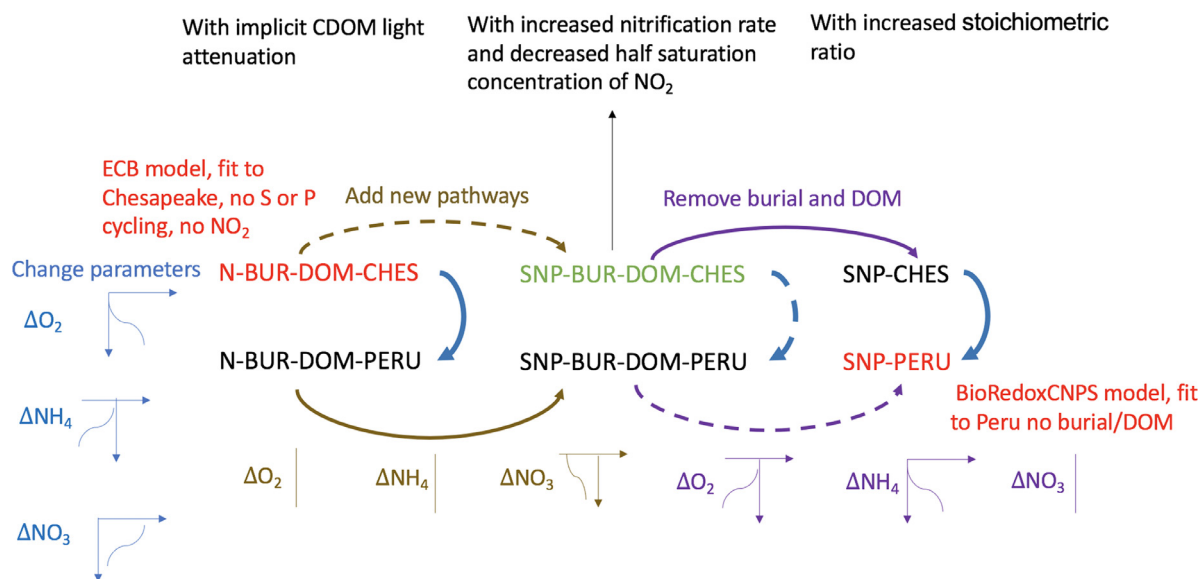
<https://doi.org/10.1016/j.mex.2023.102253>

Received 14 March 2023; Accepted 10 June 2023

Available online 11 June 2023

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**Fig. 1.** Diagram illustrating the results presented in this paper. Arrows show pathways for transforming the ECB code as configured in Da et al. [5] (N\_BUR\_DOM\_CHES) to the BioRedoxCNPS code as configured in Hantsoo et al. [8] (SNP\_PERU). Two previously published models are shown in red. The reference run from the merged model in this paper is shown in green. The other runs are shown in black. The thick blue arrows indicate the differences in parameters common to both models. The purple lines indicate removing burial and DON. The brown lines indicate adding sulfur and phosphorus cycling. The resulting changes to oxygen, ammonium and nitrate profiles are shown schematically next to each set of changes. Three sensitivity studies using SNP\_BUR\_DOM\_CHES as a control are shown at the top of the diagram.

## Method details

### Background

Over recent decades, much effort has been made to study hypoxia in Chesapeake Bay, including the development of a number of numerical models. Chesapeake Bay is a prominent estuary located on the East Coast of the United States, encompassing parts of Maryland and Virginia. It is the largest estuary in the country and plays a crucial role in providing habitat for numerous species, contributing to fisheries, and serving as a valuable recreational resource for surrounding communities. Da et al. [5] implemented a nitrogen-based model (referred to in that paper as ECB, in this paper as N\_BUR\_DOM\_CHES) that can simulate oxygen skillfully in Chesapeake Bay. This model is currently used for hindcast studies of the Bay at the Virginia Institute of Marine Science and is also used to evaluate the impact of changes in nutrient delivery to the Bay. However, as shown both here and in the companion manuscript [10], this model has large biases when simulating certain nitrogen species. Additionally, it does not simulate a number of chemical species and processes which have been hypothesized to play an important role in the Bay. One such chemical species is hydrogen sulfide, which a number of studies have found to be present in the water column of the Bay [4,13]. There is evidence from other ecosystems that simulating sulfide oxidation as well as cryptic sulfur cycling (which involves tightly coupled sulfate reduction and sulfide oxidation without a buildup of hydrogen sulfide) might help to mitigate biases in simulating nitrogen species [3]. Bacteria capable of metabolizing hydrogen sulfide have been shown to be common in Chesapeake Bay with their abundances increasing in the presence of hypoxia [1]. Additionally, N\_BUR\_DOM\_CHES does not simulate nitrite, which means that it cannot simulate anammox or sulfide oxidation by nitrite. Al Azhar et al. [2] constructed a model (referred to as BioRedoxCNPS in that paper) that includes hydrogen sulfide and nitrite and simulates sulfate reduction, sulfide oxidation by nitrite, nitrate and oxygen, anammox, two-stage denitrification and heterotrophic denitrification within the Peru Upwelling. This model matches observations of oxygen, nitrite, nitrate and phosphate as well as being broadly consistent with (low-precision) observational estimates of sulfate reduction, anammox and denitrification. An updated version of this code (with a number of improvements) was published by Hantsoo et al. [8].

This paper documents the construction of a model including the sulfur cycle that can simulate oxygen, nitrate and ammonium skillfully in Chesapeake Bay. As the first step, the model from Al Azhar et al. [2] calibrated for the Peru Upwelling is implemented in the same physical model of Chesapeake Bay as the previously published N\_BUR\_DOM\_CHES model of Da et al. [5]. The new model (referred to here as SNP\_PERU), improves some fields but not others relative to N\_BUR\_DOM\_CHES. This motivated us to develop a merged version of the two biogeochemical codes, run them in the same physical model and configure them so as to highlight specific differences between the two published models (see Fig. 1, also the graphical abstract). Based on this analysis, a specific model configuration is picked as a basis for further development. Next, a set of sensitivity studies was conducted to further improve the model skill. Finally, we propose a reference model that can simulate oxygen, hypoxic volume, nitrate and ammonium more skillfully

**Table 1**

Summary of how the classes of model differences (common parameters, pathways seen one of the original codes) map onto our model simulations.

| Model                  | Include S, NO <sub>2</sub> ,<br>Annamox | Implicit CDOM in<br>light attenuation | Common Parameters | Burial and DOM | Remineralization:<br>NO <sub>2</sub> to NH <sub>4</sub> |
|------------------------|---|---------------------------------------|-------------------|----------------|---|
| N_BUR_DOM_CHES (ECB)   | N                                       | Y                                     | CHES              | Y              | Implicit ∞  |
| SNP_BUR_DOM_CHES       | Y                                       | N                                     | CHES              | Y              | slow  |
| SNP_BUR_DOM_PERU       | Y                                       | N                                     | PERU              | Y              | slow  |
| SNP_PERU(BioRedoxCNPS) | Y                                       | N                                     | PERU              | N              | slow  |
| N_BUR_DOM_PERU         | N                                       | Y                                     | PERU              | Y              | Implicit ∞  |
| SNP_CHES               | Y                                       | N                                     | CHES              | N              | slow  |

than either of the previously published models that can serve as a basis for future studies. A comparison table summarizing the different versions of the model can be found in the companion Ocean Modelling manuscript [10] as Table 2.

### Physical component of the model

The physical component of our model is based on the Regional Ocean Modeling System (ROMS, [17]) which uses an orthogonal curvilinear grid with variable horizontal resolution. The highest resolution (430 m) is found in the northern Bay, the lowest resolution (~10 km) in the southern end of the Mid-Atlantic Bight, and the average grid spacing within the Chesapeake Bay is 1.7 km. The map displaying the model bathymetry as well as the locations of monitoring stations can be found as Fig. 1 in Jin et al. [10]. Governing equations are discretized over a stretched terrain-following s-coordinate with 20 vertical levels. To interpolate between a higher resolution in the surface and the bottom boundary layers in deeper waters and relatively constant resolution in shallow waters, the stretching function of Shchepetkin & McWilliams [16] was used with values  $\theta_s=6.0$  and  $\theta_b=4.0$  (standard values in this version of ROMS) with an  $hc = 10$  m.

Tidal constituents were adopted from the Advanced Circulation (ADCIRC) model [12] and from observed nontidal water levels from Duck, NC and Lewes, DE [15]. They were imposed on the model at the open boundary. Atmospheric forcings, including winds, air temperature, relative humidity, pressure, precipitation, short-wave radiation and longwave radiation, were obtained from the North American Regional Reanalysis (originally described in [14]).

The MPDATA 3-D advection scheme [18,19] was used for tracers. MPDATA 3-D is a third-order upstream advection scheme that ensures that advection does not generate spurious maxima or minima while minimizing numerical diffusion (this is particularly important for biogeochemical tracers). Momentum is advected with a third-order centered difference scheme in the horizontal and fourth-order centered difference in the vertical. The vertical turbulent mixing scheme and background mixing coefficients for both momentum and tracers were all set to the same values as in Feng et al. [6].

### Biogeochemical cycling (BGC) codes and simulation setups

We examined the behavior of three biogeochemical codes (ECB, BioRedoxCNPS, and our merger of the two: SNP\_BUR\_DOM), which we implemented using two parameter sets for phytoplankton growth, coagulation and sinking governed by equations in common to the two codes. One parameter set is taken from the Da et al. [5] model of the Chesapeake and the other is taken from the al Azhar et al. [2] model of the Peru upwelling system as modified by Hantsoo et al. [8]. Our experimental design thus contains six core simulations with two parameter sets implemented for each code. We use a nomenclature that makes it evident what *chemical species* (N vs. S, N and P) are cycled, whether the model includes *burial and DOM*, and whether the *parameter set* is taken from the Peru simulation or the Chesapeake simulation. N\_BUR\_DOM\_CHES thus describes ChesROMS\_ECB which cycles N but not S or P, includes burial and DOM, and uses parameters from Da et al. [5] for (among other things) sinking and phytoplankton growth. By contrast SNP\_PERU describes our implementation of BioRedoxCNPS, with sulfur and phosphate cycling but no burial or DOM and uses parameters from Hantsoo et al., [8]. A more complete description of each simulation is given below and the differences are summarized in Table 1.

#### N\_BUR\_DOM\_CHES

The BGC code in the N\_BUR\_DOM\_CHES simulation is the same as the code used in Da et al. [5], which is derived from a nitrogen-based ecosystem code [7]. This code includes a simplified nitrogen cycle with 8 nitrogen pools (listed with model acronyms): nitrate (NO<sub>3</sub>), ammonium (NH<sub>4</sub>), phytoplankton (P), zooplankton (Z), semilabile and refractory dissolved organic nitrogen (DONsl and DONrf) and small and large nitrogen detritus (SDeN and LDeN). Additionally, the code simulates semilabile and refractory DOC (DOCsl and DOCre), inorganic suspended solids (ISS), chlorophyll (Chl), dissolved inorganic carbon (DIC), alkalinity (Alk), and dissolved oxygen (O<sub>2</sub>). As implemented in the ChesROMS\_ECB model, phytoplankton growth is limited by nitrogen and light and the dominant phytoplankton loss is via coagulation and sinking. The resuspension of fractions of phytoplankton and large detritus into small detritus upon reaching the bottom is influenced by the near-bottom turbulent velocities. Some fraction of the remaining benthic flux is buried permanently following the formula of Henrichs and Reeburgh [9] while the remainder undergoes remineralization.

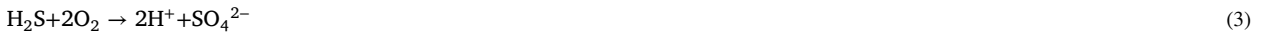
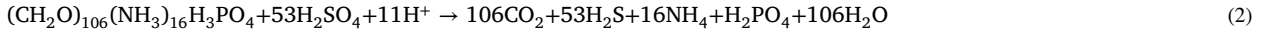
$$f_{bur} = \min(0.75, 0.023 * \text{carbon flux to the bottom}^{0.5797}) \quad (1)$$

The carbon flux to the bottom is in units of mmol C/m<sup>2</sup>/day. This means that when the flux of carbon is much less than ~400 mmol C/m<sup>2</sup>/day most of the flux is remineralized. At low levels of flux to the bottom, doubling the flux results in a tripling of burial.

In this model, there are three pathways involved in transforming the organic material to inorganic nitrogen: 1. Solubilization of excreted materials produces DON. Both DON and detrital materials are remineralized to NH<sub>4</sub>, 2. using oxygen if it is available and 3. nitrate (resulting in denitrification) if it is not.

### SNP\_PERU

The second biogeochemical simulation, SNP\_PERU, uses the code developed by al Azhar et al. [2] and updated by Hantsoo et al. [8] to capture interactions between the cycles of nitrogen, phosphorus and sulfur in the Peru coastal ocean upwelling system. Like the ECB code, this code was also derived from the BGC code of Fennel et al. [7], and it has previously been referred to as BioRedoxCNPS [2] and Fennel\_CNPS [8]. We refer to the unaltered version of this code implemented in the ChesROMS physical model domain with BGC parameters from Hantsoo et al. [8] as the SNP\_PERU simulation. This code adds new explicit kinetic processes to the Fennel BGC code, described by the following set of equations:



1. Sulfate is reduced to H<sub>2</sub>S during organic matter remineralization when other oxidants (oxygen and nitrate) are limiting (Eq. 2). Sulfide is reoxidized to sulfate 2. by oxygen (Eq. 3), 3. by nitrate reduction to nitrite through chemolithoautotrophic nitrate reduction (Eq. 4) or 4. by nitrite reduction to N<sub>2</sub> gas through sulfide-driven denitrification (Eq. 5). When the water is anoxic, ammonium can also be oxidized by nitrite through anammox to produce N<sub>2</sub> gas. The SNP simulations used in this study thus include six state variables not included in N\_BUR\_DOM\_CHES: nitrite (NO<sub>2</sub>), sulfate (SO<sub>4</sub>), hydrogen sulfide (H<sub>2</sub>S), phosphate (PO<sub>4</sub>) and small and large detrital phosphorus (SDeP, LDeP). Autotrophic nitrogen fixation by diazotrophs (which was included in the original study of al Azhar et al. [2]) was turned off in our simulations as it resulted in numerical instability and is not expected to play a major role in Chesapeake nitrogen dynamics given the excess of fixed nitrogen over phosphorus. It is notable that there are no sedimentary burial processes in the SNP code so that all organic materials hitting the bottom are remineralized. Thus, in comparison to N\_BUR\_DOM\_CHES, SNP\_PERU has two new *pathways* (anammox and sulfide-driven denitrification) by which nitrogen is lost to the system, but it simultaneously neglects the loss of nitrogen via burial. Additionally, dissolved organic materials are not included in this model. Finally, as described in Table 1 of the companion manuscript, although the equations for phytoplankton growth, grazing, coagulation, and detrital sinking can be cast in identical forms in SNP\_PERU and N\_BUR\_DOM\_CHES, many of the parameters within these equations are different in these two models. In particular, grazing and remineralization rates in N\_BUR\_DOM\_CHES have an exponential dependence on temperature with a Q<sub>10</sub> of 2.4 taken from Lomas et al. [11] while those in SNP\_PERU do not (corresponding to a Q<sub>10</sub> of 1).

An additional difference between the N\_BUR\_DOM\_CHES (ChesROMS\_ECB) and SNP (BioRedoxCNPS) codes is the parameterization of penetrating photosynthetically active radiation (PAR). In N\_BUR\_DOM\_CHES, PAR is attenuated by water, suspended sediments and implicitly by colored dissolved organic materials (via a dependence on salinity) but not by chlorophyll. In SNP\_PERU it is attenuated by water and chlorophyll alone. The vertical attenuation of PAR is described by the equation

$$\frac{\partial}{\partial z} \text{PAR} = k_{\text{PAR}} * \text{PAR} \quad (6)$$

where

$$k_{\text{PAR}} = 0.04 + 0.024 * \text{Chl} \quad (7)$$

Where *Chl* is the concentration of chlorophyll in mg m<sup>-3</sup>.

### SNP\_CHES

With the exception of temperature dependencies for grazing and remineralization, the code in SNP\_CHES is the same as in SNP\_PERU. However, in any equations which are also in common with N\_BUR\_DOM\_CHES, all common parameters were set to the values in N\_BUR\_DOM\_CHES. We also adopted the temperature dependences from the N\_BUR\_DOM\_CHES simulation.

### N\_BUR\_DOM\_PERU

In parallel, we ran N\_BUR\_DOM\_PERU by replacing common parameters in the N\_BUR\_DOM\_CHES code with PERU parameters, including setting Q<sub>10</sub> to 1 for grazing and remineralization. Thus, comparing SNP\_PERU (original BioRedoxCNPS) to SNP\_CHES (BioRedoxCNPS with parameters from ChesROMS\_ECB) or N\_BUR\_DOM\_CHES (original ChesROMS\_ECB) to N\_BUR\_DOM\_PERU (ChesROMS\_ECB with parameters from BioRedoxCNPS, see Table 1 in the companion manuscript for list of parameters that differ) helps to

**Table 2**  
Biochemical parameters used in SNP\_BUR\_DOM\_CHES.

| Parameter   | SNP_BUR_DOM_CHES | Unit  |
|---|------------------|---|
| half-saturation concentration of O <sub>2</sub> in oxic mineralization                              | 0.3              | mmol O <sub>2</sub> m <sup>-3</sup>                     |
| half-saturation concentration of NO <sub>3</sub> in nitrate reduction                               | 15               | mmol N m <sup>-3</sup>                                  |
| half-saturation concentration of NO <sub>2</sub> in denitrification                                 | 30               | mmol N m <sup>-3</sup>                                  |
| Threshold value below which nitrite does not participate in denitrification                         | 2                | mol N m <sup>-3</sup>                                   |
| half-saturation concentration of O <sub>2</sub> inhibition in nitrate reduction and denitrification | 1                | mmol O <sub>2</sub> m <sup>-3</sup>                     |
| half-saturation concentration of O <sub>2</sub> inhibition in sulfate reduction                     | 0.1              | mmol O <sub>2</sub> m <sup>-3</sup>                     |
| half-saturation concentration of NO <sub>3</sub> inhibition in sulfate reduction                    | 4                | mmol N m <sup>-3</sup>                                  |
| constant rate of sulfide oxidation by NO <sub>3</sub>   | 0.93             | d <sup>-1</sup>   |
| constant rate of sulfide oxidation by NO <sub>2</sub>   | 0.33             | d <sup>-1</sup>   |
| constant rate of sulfide oxidation by O <sub>2</sub>  | 0.93             | d <sup>-1</sup>   |
| half-saturation concentration of O <sub>2</sub> in sulfide oxidation                                | 1                | mmol O <sub>2</sub> m <sup>-3</sup>                     |
| half-saturation concentration of NO <sub>3</sub> in sulfide oxidation                               | 2.9              | mmol N m <sup>-3</sup>                                  |
| half-saturation concentration of NO <sub>2</sub> in sulfide oxidation                               | 6                | mmol N m <sup>-3</sup>                                  |
| half-saturation concentration of O <sub>2</sub> inhibition in sulfide oxidation                     | 0.1              | mmol O <sub>2</sub> m <sup>-3</sup>                     |
| constant rate of anammox rate   | 0.07             | d <sup>-1</sup> (mmol N m <sup>-3</sup> ) <sup>-1</sup> |
| maximum rate of aerobic ammonium oxidation  | 0.05             | d <sup>-1</sup>   |
| maximum rate of aerobic nitrite oxidation   | 0.005            | d <sup>-1</sup>   |
| half-saturation concentration of O <sub>2</sub> in nitrification                                    | 1                | mmol O <sub>2</sub> m <sup>-3</sup>                     |
| radiation inhibition threshold of ammonium to nitrite oxidation                                     | 0.0095           | W m <sup>-2</sup>                                       |
| radiation inhibition threshold of nitrite to nitrate oxidation                                      | 0.0364           | W m <sup>-2</sup>                                       |
| light intensity at which inhibition is half-saturated for ammonium to nitrite oxidation             | 0.036            | W m <sup>-2</sup>                                       |
| light intensity at which inhibition is half-saturated for nitrite to nitrate oxidation              | 0.074            | W m <sup>-2</sup>                                       |
| Small detritus remineralization rate N-fraction   | 0.03             | d <sup>-1</sup>   |
| Small detritus remineralization rate C-fraction   | 0.03             | d <sup>-1</sup>   |
| Large detritus remineralization rate N-fraction   | 0.01             | d <sup>-1</sup>   |
| Large detritus remineralization rate C-fraction   | 0.01             | d <sup>-1</sup>   |
| Q <sub>10</sub>   | 2.4              | Null  |
| phytoplankton growth rate at 0°C  | 0.69             | d <sup>-1</sup>   |
| chlorophyll to phytoplankton N maximum ratio  | 0.053            | mgChl mgC <sup>-1</sup>                                 |
| initial slope of planktonic growth to light curve   | 0.125            | (W m <sup>-2</sup> ) <sup>-1</sup> d <sup>-1</sup>      |
| half-saturation concentration for uptake of NO <sub>3</sub> by phytoplankton                        | 0.5              | mmol N m <sup>-3</sup>                                  |
| half-saturation concentration for uptake of NH <sub>4</sub> by phytoplankton                        | 0.5              | mmol N m <sup>-3</sup>                                  |
| stoichiometry of P to N in phytoplankton and zooplankton  | 1/16             | dimensionless   |
| Stoichiometry of O <sub>2</sub> production/consumption to NO <sub>3</sub> uptake/remineralization   | 8.625            | dimensionless   |
| Stoichiometry of O <sub>2</sub> production/consumption to NH <sub>4</sub> uptake/remineralization   | 6.625            | dimensionless   |
| half-saturation concentration for uptake of PO <sub>4</sub> by phytoplankton (kNO <sub>3</sub> /16) | 0.03125          | mmol P m <sup>-3</sup>                                  |
| excretion rate due to basal metabolism  | 0.1              | d <sup>-1</sup>   |
| excretion rate due to phytoplankton assimilation  | 0.1              | d <sup>-1</sup>   |
| assimilation efficiency   | 0.75             | dimensionless   |
| maximum phytoplankton grazing rate  | 0.6              | (mmol N m <sup>-3</sup> ) <sup>-1</sup> d <sup>-1</sup> |
| phytoplankton mortality   | 0.15             | d <sup>-1</sup>   |
| zooplankton mortality   | 0.025            | d <sup>-1</sup>   |
| half saturation of phytoplankton ingestion  | 2                | (mmol N m <sup>-3</sup> ) <sup>-2</sup>                 |
| aggregation parameter   | 0.005            | d <sup>-1</sup>   |
| sinking velocity of phytoplankton   | 0.1              | m d <sup>-1</sup>                                       |
| sinking velocity of small detritus  | 0.1              | m d <sup>-1</sup>                                       |
| sinking velocity of large detritus  | 5                | m d <sup>-1</sup>                                       |

distinguish the differences that can be attributed to biological parameters (e.g. phytoplankton growth rate) within identical pathways from the differences caused by changing the biogeochemical pathways themselves (e.g. adding anammox).

### SNP\_BUR\_DOM\_PERU

Since the biological model from al Azhar et al. [2] was developed for an open-ocean/coastal upwelling system rather than an estuary with strong forcing from riverine runoff and significant rates of organic matter burial, we modified the SNP code by adding the resuspension and burial code that was used in ChesROMS\_ECB. We also added dissolved organic matter cycling, extending the ECB code which simulated DON and dissolved organic carbon (DOC) to include dissolved organic phosphorus (DOP). Including burial without DOM cycling resulted in an excessive fraction of the nutrients delivered to the model being buried in the river mouths. We denote this merged code as SNP\_BUR\_DOM, and we denote the simulation made with this new code as SNP\_BUR\_DOM\_PERU when biological constants in common with SNP\_PERU are set to those in the latter model.

### SNP\_BUR\_DOM\_CHES

For the simulation SNP\_BUR\_DOM\_CHES, the code is identical to that of SNP\_BUR\_DOM\_PERU. However, in the equations which are identical to those in N\_BUR\_DOM\_CHES, all parameters are set to the values in the latter simulation. The full set of parameters used in this model is listed in Table 2.

### Pairing simulations to isolate sources of the differences between SNP\_PERU and N\_BUR\_DOM\_CHES

With our six simulations, we can isolate which differences between SNP\_PERU and N\_BUR\_DOM\_CHES contribute to the different simulated results. Differences between SNP\_PERU and SNP\_CHES or N\_BUR\_DOM\_PERU and N\_BUR\_DOM\_CHES are due to differences in parameters common to both simulations. Differences between SNP\_BUR\_DOM\_PERU and SNP\_PERU (or SNP\_BUR\_DOM\_CHES and SNP\_CHES) are purely due to the inclusion of DOM and burial/resuspension of organic matter. Differences between SNP\_BUR\_DOM\_PERU and N\_BUR\_DOM\_PERU (or SNP\_BUR\_DOM\_CHES and N\_BUR\_DOM\_CHES) are due to differences in whether we include sulfur and phosphorus cycling, or to differences in the optical scheme used to parameterize the penetration of shortwave radiation.

#### Developing a new configuration based on SNP\_BUR\_DOM\_CHES–Sensitivity studies and final configuration

As shown below, SNP\_BUR\_DOM\_CHES produces a simulation that improves nitrate and ammonium relative to N\_BUR\_DOM\_CHES with the new sulfur pathways, but only at the cost of reducing the fidelity of the oxygen simulation and creating a high bias for nitrite. Starting with this simulation we then show that we can improve individual fields by:

1. Changing the optics to include implicit CDOM and explicit sediment absorption: In this case, we change the optics in SNP\_BUR\_DOM\_CHES to the formulation in N\_BUR\_DOM\_CHES which implicitly relates the diffuse attenuation coefficient to refractory DOC delivered down rivers.

$$k_{PAR} = \max(1.4 + 0.063 * TSS - 0.057 * S, 0.6) \quad (8)$$

where TSS represents total suspended solids. Note that as S becomes very small, the attenuation becomes very large.

2. Making nitrification and denitrification easier: In this case, we decrease half-saturation concentration of NO<sub>2</sub> in denitrification from 30 mmol N m<sup>-3</sup> to 1 mmol N m<sup>-3</sup> and increase nitrification rate constant governing the transformation of NO<sub>2</sub> to NO<sub>3</sub> from 0.005 day<sup>-1</sup> to 0.05 day<sup>-1</sup>. We also remove 2 mmol/m<sup>3</sup> threshold for nitrite to participate in denitrification.
3. Increasing the stoichiometric ratio between N and O in organic matter: In this case, we increase oxygen:nitrate from 8.625 to 9.625 and oxygen:ammonium from 6.625 to 7.625, for reasons justified below (See section “motivating the sensitivity cases: O<sub>2</sub>:N Stoichiometry”).

In the companion paper, we showed some of the impacts of these three changes on nitrogen partitioning (Fig. 9) and model fidelity for oxygen, ammonium and nitrate at CB3.3 (Table 6). In this manuscript we expand the evaluation to other sites, consider nitrite biases and also evaluate the impacts of making all three changes simultaneously.

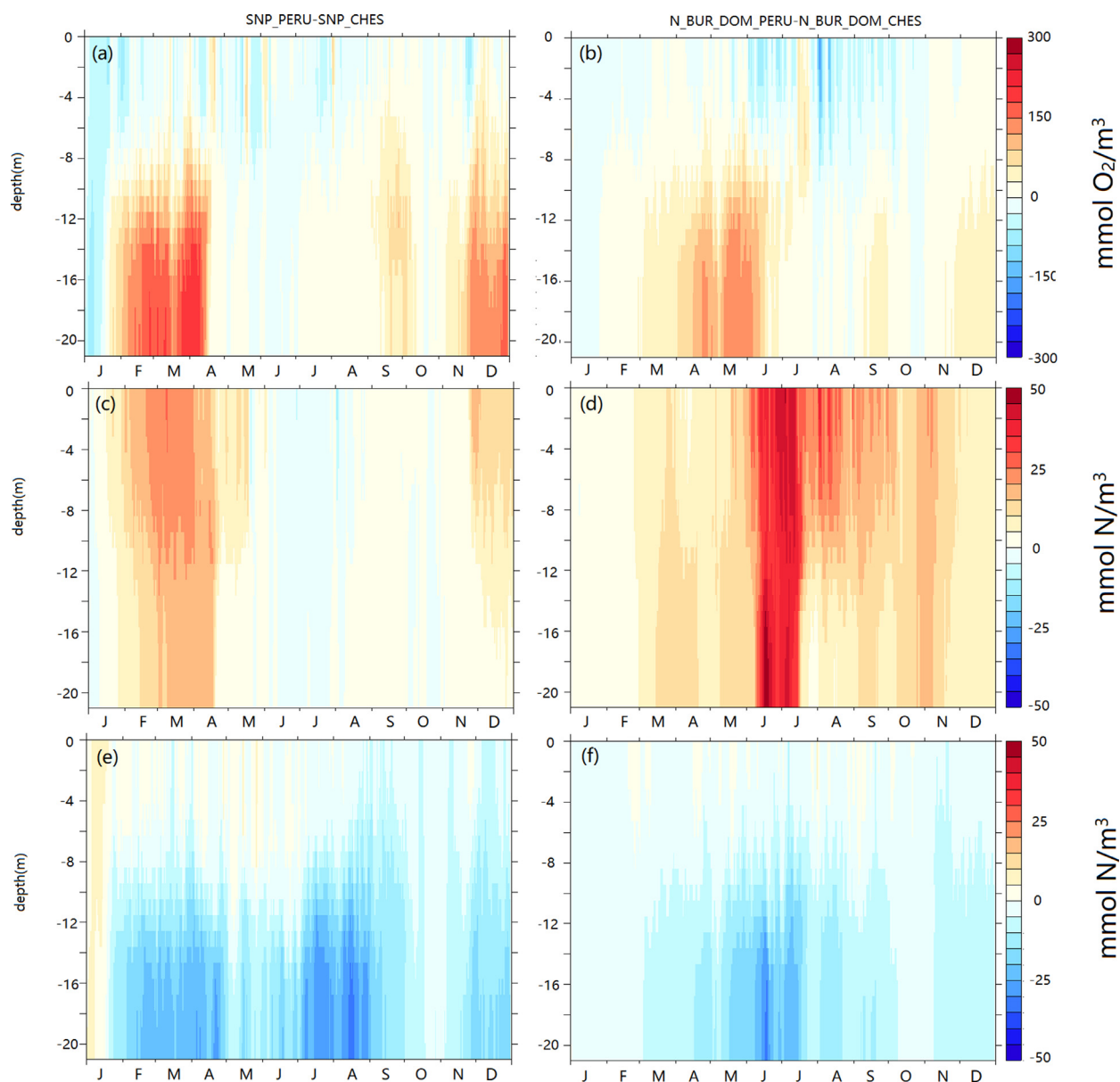
#### Evaluation of model configurations

At station CB3.3, significant differences between simulations of oxygen, ammonium and nitrate using the two published codes are found (right-hand column of Fig. 4 in companion paper). As illustrated in Fig. 1, we now break down the differences between the original models into the three categories described in the companion Ocean Modelling manuscript [10] demonstrating that it does not matter (at least qualitatively) in which order changes are made.

#### Impact of using the PERU parameter set vs CHES parameter set in the two BGC codes

While there are many differences between the biogeochemical cycles in the two published codes, parameters such as growth rates and sinking speeds of detritus that are found in both codes also differ. These common parameters would be expected to have effects on our model results. To quantify this effect, we compare two pairs of models: SNP\_PERU minus SNP\_CHES (left column of Fig. 2) and N\_BUR\_DOM\_PERU minus N\_BUR\_DOM\_CHES (right column of Fig. 2). This comparison isolates the differences contributed by changing common parameters from their values in Da et al. [5] to the values in Hantsoo et al. [8] and vice versa.

Switching parameters from CHES values to PERU values does not explain the differences in the third column in Fig. 4 of companion paper; in fact, the changes seen have the opposite sign. Qualitatively similar changes are seen in the two pairs of simulations. Oxygen becomes higher from near the bottom to around 8 m in depth. Nitrate gets higher while ammonium becomes lower. SNP\_PERU minus SNP\_CHES shows more extreme change for oxygen and ammonium with more moderate change for nitrate compared to N\_BUR\_DOM\_PERU minus N\_BUR\_DOM\_CHES. SNP\_PERU has much more oxygen than SNP\_CHES from late January to mid-April and late November to end of December from near the bottom to 10 m in depth, with relative increases of up to 200 mmol O<sub>2</sub>/m<sup>3</sup>. SNP\_CHES extends the hypoxic zone at CB3.3C through much of the year. In both pairs, using PERU parameters leads to a lower oxygen concentration near the surface, especially during the summer months. From late January to mid-April as well as in December, nitrate in SNP\_PERU is up to 25 mmol N/m<sup>3</sup> higher than SNP\_CHES. This can be explained in terms of the higher levels of oxygen in SNP\_PERU reducing denitrification rates, allowing nitrate to persist longer for the PERU parameters relative to the CHES parameters. Nitrate in N\_BUR\_DOM\_PERU is always higher than N\_BUR\_DOM\_CHES, especially from early June to mid-July, by up to 50 mmol N/m<sup>3</sup>. For ammonium, SNP\_PERU is consistently up to 20 mmol N/m<sup>3</sup> lower than SNP\_CHES from near the bottom to a depth of 10 m. N\_BUR\_DOM\_PERU is also lower than N\_BUR\_DOM\_CHES, but the largest differences occur only in June. The middle column of Fig. 6 in



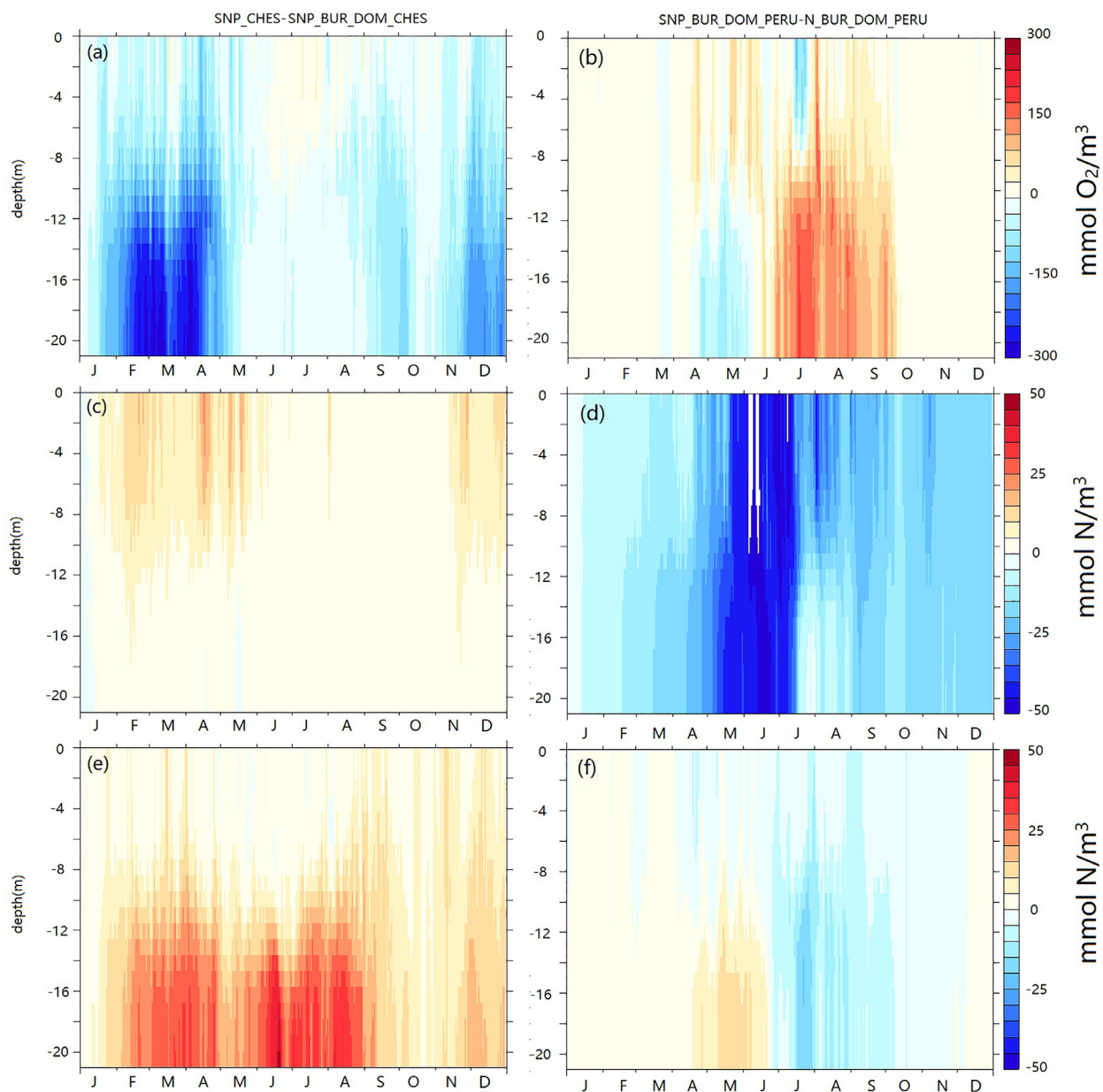
**Fig. 2.** Evaluation of the consistency of the effect of changing common parameters. Modeled oxygen (a, b), nitrate (c, d) and ammonium (e, f) for SNP\_PERU minus SNP\_CHES (left column, a, c and e) and N\_BUR\_DOM\_PERU minus N\_BUR\_DOM\_CHES (right column, b, d and f) at coincident times and locations at the Bay Bridge station (CB3.3C) during 2017.

the companion paper also demonstrates similar changes when transitioning from SNP\_BUR\_DOM\_PERU to SNP\_BUR\_DOM\_CHES. This indicates that the effects of changing common parameters are robust across various model configurations.

#### *Impact of adding burial and DOM to the SNP code*

Next, we turn to the differences between the simulations induced by adding or removing burial of organic matter and cycling of dissolved organic matter, processes which are not included in the original SNP code of al Azhar et al. [2]. Differences between SNP\_CHES versus SNP\_BUR\_DOM\_CHES (Fig. 3 left column) and SNP\_PERU versus SNP\_BUR\_DOM\_PERU (right column in Fig. 6 of the companion paper) in oxygen, nitrate and ammonium are shown. We choose to show the impacts of *removing* burial and DOM cycling so as to make it easier to visually attribute the differences between the original models to different sources.

For both pairs of simulations, removing dissolved organic matter and burial processes generally more than balances the oxygen and ammonium changes caused by changes in common parameters and thus plays a leading role in explaining the differences between

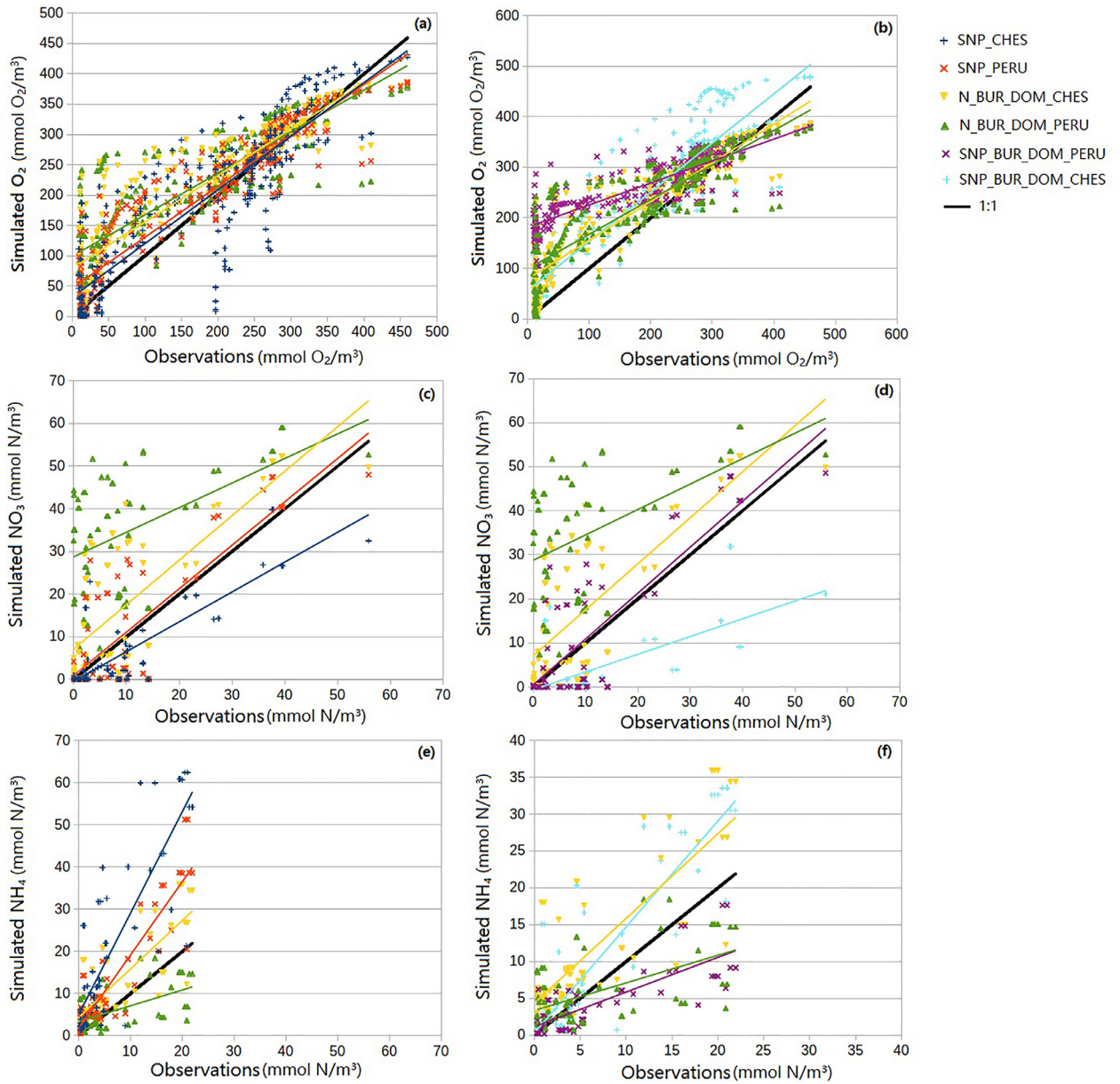


**Fig. 3.** Evaluation of the consistency of the effect of removing DOM and burial and the consistency of the effect of changing pathways and optics. Modeled oxygen (a, b), nitrate (c, d) and ammonium (e, f) for SNP\_CHES minus SNP\_BUR\_DOM\_CHES (left column, a, c and e) and SNP\_BUR\_DOM\_PERU minus N\_BUR\_DOM\_PERU (right column, b, d and f) at coincident times and locations at the Bay Bridge station (CB3.3C) during 2017.

the original models. Both pairs of simulations show decreases in oxygen and increases in ammonium concentrations from the bottom to around 8 m in depth, although the time period during which the decrease is seen is different in the two comparisons.

Oxygen in SNP\_CHES is lower than SNP\_BUR\_DOM\_CHES for most of the year, with significant differences appearing from mid-January to early May and late November to late December. During the summer months, oxygen in SNP\_CHES is slightly higher than SNP\_BUR\_DOM\_CHES near the surface. Larger difference values for SNP\_PERU versus SNP\_BUR\_DOM\_PERU are found from early April to early October. For the most part, surface oxygen concentrations during summertime in SNP\_PERU are slightly higher than SNP\_BUR\_DOM\_PERU. SNP\_CHES shows much higher values of ammonium than SNP\_BUR\_DOM\_CHES from mid-February to late August, while in SNP\_PERU the higher values appear from late May to mid-August. For nitrate, SNP\_CHES is almost always higher than SNP\_BUR\_DOM\_CHES with largest differences appearing near the surface from late January to mid-May and mid-November to late December. However, from late April to mid-May nitrate in SNP\_PERU is slightly lower than SNP\_BUR\_DOM\_PERU. The differences





**Fig. 4.** Simulated versus observed oxygen ( $\text{mmol O}_2/\text{m}^3$ ) (a,b), nitrate ( $\text{mmol N}/\text{m}^3$ ) (c, d) and ammonium ( $\text{mmol N}/\text{m}^3$ ) (e, f) at coincident times and locations from SNP\_CHES (blue), SNP\_PERU (orange), N\_BUR\_DOM\_CHES (yellow), N\_BUR\_DOM\_PERU (green), SNP\_BUR\_DOM\_PERU (purple) and SNP\_BUR\_DOM\_CHES (light blue). Left column shows cases that can be used to compare CHES (dark blue cross, yellow triangle) and corresponding PERU (orange cross, green triangle) parameter sets. Right column shows cases that can be used to evaluate impact of adding new pathways and changing optics (compare light blue with green, purple with yellow). Solid black lines show 1:1 line, colored lines show linear trend. Note that the scales differ between (a) and (b), (e) and (f) in order to make the differences between simulations more visible.

in nitrate are much smaller than the increases resulting from changing common parameters and so do not explain the differences between the original configurations seen in Fig. 4 of our companion paper.

*Impact of the effects of nutrient cycling and optics changes between the two BGC codes: Coupled sulfur, nitrogen and phosphate cycling*

We now turn to the differences induced by adding the pathways for sulfur and phosphorus cycling, explicitly modeling nitrite and anammox and changing the optics in al Azhar et al. [2] but not changing burial or dissolved organic matter cycling. Differences between SNP\_BUR\_DOM\_PERU versus N\_BUR\_DOM\_PERU simulations of oxygen, nitrate and ammonium are shown in right column

**Table 3**

Error metrics for the model suite compared with observations. A perfect model would have  $R^2=1$  and bias=0. Values of  $R^2<0$  are associated with large biases, which result in the error variance being larger than the sample variance. Because the biases of nitrite dominate the  $R^2$  we only report biases here.

|                  | $R^2$ /bias for O <sub>2</sub> | $R^2$ /bias for NH <sub>4</sub> | $R^2$ /bias for NO <sub>3</sub> | bias for NO <sub>2</sub> |
|------------------|--------------------------------|---------------------------------|---------------------------------|--------------------------|
| N_BUR_DOM_CHES   | 0.72/36.44                     | -0.32/5.32                      | -0.29/7.49                      | NA                       |
| N_BUR_DOM_PERU   | 0.59/41.61                     | 0.27/-0.69                      | -4.77/24.94                     | NA                       |
| SNP_CHES         | 0.75/10.66                     | -8.17/14.14                     | 0.62/-3.19                      | 5.16                     |
| SNP_PERU         | 0.85/17.39                     | -1.13/6.58                      | 0.46/1.08                       | 10.20                    |
| SNP_BUR_DOM_CHES | 0.59/51.23                     | -0.03/3.14                      | 0.20/-6.02                      | 4.33                     |
| SNP_BUR_DOM_PERU | 0.19/78.95                     | 0.46/-2.28                      | 0.49/0.86                       | 8.48                     |

of Fig. 3. Differences between SNP\_BUR\_DOM\_CHES versus N\_BUR\_DOM\_CHES can be seen in left column of Fig. 6 in our companion paper.

Adding more complex nutrient cycling and changing the optics produces large decreases in nitrate—explaining why we see decreases in this field in Fig. 4 of our companion paper, but produces smaller changes in oxygen and ammonium. Similar changes for the two pairs of simulations are seen in nitrate and ammonium. Relative to the original ChesROMS\_ECB code, the SNP code decreases nitrate concentration: large decreases (up to 50 mmol N/m<sup>3</sup>) appear from early May to mid-July for SNP\_BUR\_DOM\_PERU minus N\_BUR\_DOM\_PERU, and from late January to early June for SNP\_BUR\_DOM\_CHES minus N\_BUR\_DOM\_CHES. The changes in pathways thus appear to dominate the differences in nitrate seen in the two published models.

For ammonium, SNP\_BUR\_DOM\_PERU predicts values up to 15 mmol N/m<sup>3</sup> higher than N\_BUR\_DOM\_PERU from early May to early June from bottom to 14 m in depth but up to 30 mmol N/m<sup>3</sup> lower in July. Similar changes can be observed in SNP\_BUR\_DOM\_CHES minus N\_BUR\_DOM\_CHES but the range is less extreme. The changes in nutrient cycling and optics are important for determining the timing of the differences in ammonium seen in Fig. 4 of our companion paper but are not the dominant drivers of these differences.

In contrast to nitrate and ammonium, the differences in oxygen induced by adding nutrient cycling and changing the optics depend more on the base simulation. From the bottom to 12 m in depth, oxygen in SNP\_BUR\_DOM\_PERU is lower than N\_BUR\_DOM\_PERU from late April to early June, while from early June to early October, oxygen in SNP\_BUR\_DOM\_PERU becomes higher than N\_BUR\_DOM\_PERU. During the same period and at the same location, SNP\_BUR\_DOM\_CHES and N\_BUR\_DOM\_CHES only exhibit minor differences. During the summer months near the surface, SNP\_BUR\_DOM\_PERU is mostly higher than N\_BUR\_DOM\_PERU while SNP\_BUR\_DOM\_CHES is mostly lower than N\_BUR\_DOM\_CHES. Overall, these differences are smaller than those associated with the previous pairs of experiments.

### Evaluating the realism of the model simulations-Is there a best simulation?

Evaluation of simulations for oxygen, nitrate and ammonium from the two published models can be found in our companion paper in section 3.1. Here we also list the  $R^2$  and bias of nitrite and metrics from SNP\_BUR\_DOM\_CHES and SNP\_BUR\_DOM\_PERU for a more complete comparison which serves as the basis for constructing the best model. All of the SNP simulations show an improvement for the simulation of nitrate at CB3.3.

However, examining the  $R^2$  and biases for oxygen, nitrate and ammonium across the models listed in Table 3 demonstrates that the “best” model is not the same for each variable. Large biases play a significant role in decreasing  $R^2$ : SNP\_CHES has a high ammonium bias of 14.14 with an  $R^2$  of -8.17 while N\_BUR\_DOM\_PERU has a high nitrate bias of 24.94 with an  $R^2$  of -4.77. In terms of  $R^2$  averaged across the three variables and also low biases for nitrogen variables, SNP\_BUR\_DOM\_PERU produces the best simulation at CB3.3C. However, the results come at the cost of a degradation of the simulation of oxygen. A tradeoff can be seen between nitrate/ammonium and oxygen simulations among the six simulations. We will return to the implications of this result in the following section.

By examining scatter plots comparing observations (horizontal axis) to the modeled values (vertical axis) across these sets of simulations (Fig. 4), we can see more details about which mismatches contribute to the  $R^2$  differences, and whether this remains consistent across simulations.

The top row shows the model-data mismatch for oxygen. We can look at the impact of changing parameter sets by comparing SNP\_CHES (yellow, Fig. 4a) with SNP\_PERU (orange, Fig. 4a), N\_BUR\_DOM\_CHES (blue, Fig. 4a) with N\_BUR\_DOM\_PERU (green, Fig. 4a) and SNP\_BUR\_DOM\_CHES (light blue, Fig. 4b) with SNP\_BUR\_DOM\_PERU (purple, Fig. 4b). All the models generally over-predict oxygen with the worst mismatch in the 50-200 mmol O<sub>2</sub>/m<sup>3</sup> range. Switching from PERU to CHES parameters reduces this mismatch across all three pairs, with the trend lines for SNP\_PERU, N\_BUR\_DOM\_PERU and SNP\_BUR\_DOM\_PERU (orange, blue, purple) lying above those for SNP\_CHES, N\_BUR\_DOM\_CHES and SNP\_BUR\_DOM\_CHES (yellow, green, light blue). However, at higher values of oxygen the trends reverse. Which parameter set is used modulates the impact of adding new pathways (illustrated in Fig. 4b). SNP\_BUR\_DOM\_PERU has more oxygen than N\_BUR\_DOM\_PERU at the low end of the range but less at the high end, while the reverse is true for SNP\_BUR\_DOM\_CHES with respect to N\_BUR\_DOM\_CHES. Adding dissolved organic matter and burial processes slightly increases the overestimation of oxygen relative to observations in the 50-200 mmol O<sub>2</sub>/m<sup>3</sup> range.

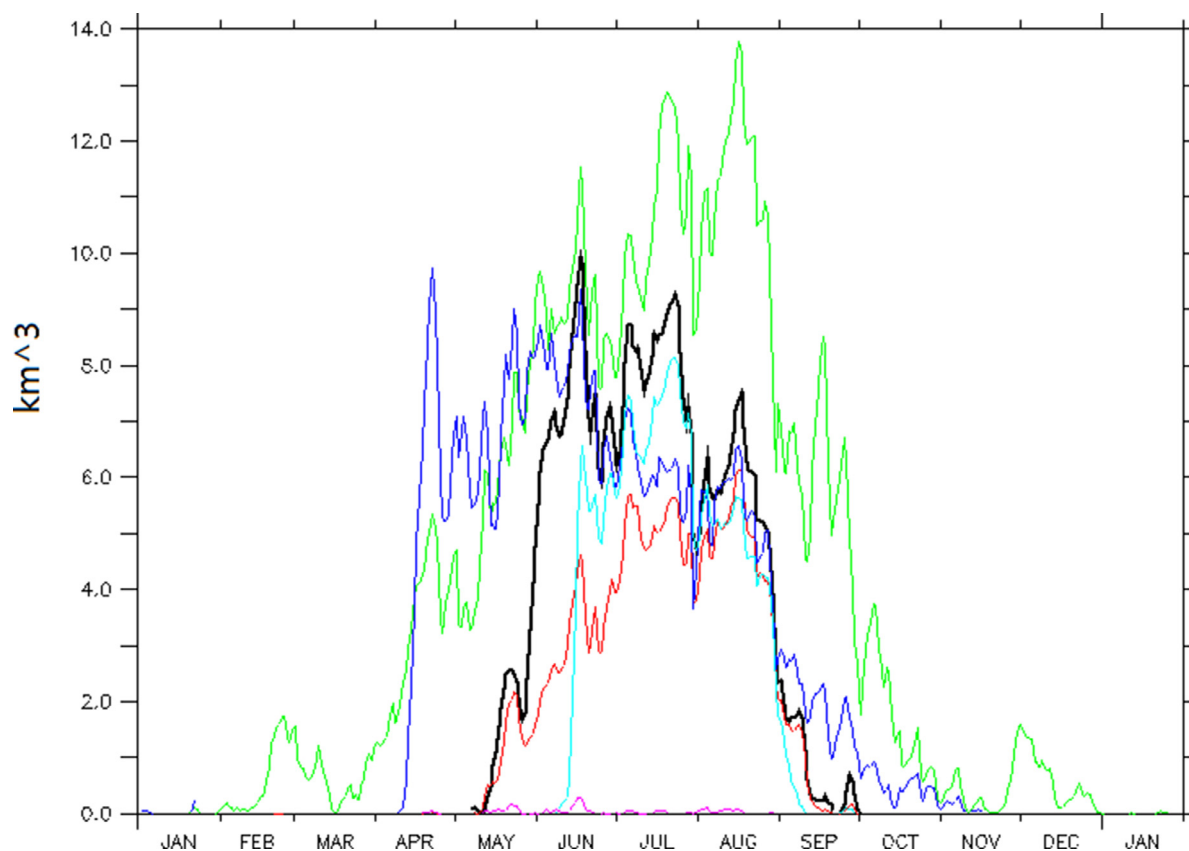


Fig. 5. Hypoxic volume for six runs: N\_BUR\_DOM\_CHES (VIMS hindcast, black), SNP\_CHES (green), SNP\_BUR\_DOM\_CHES (red), SNP\_PERU (dark blue), N\_BUR\_DOM\_PERU (light blue) and SNP\_BUR\_DOM\_PERU (purple).

For nitrate (middle row) and ammonium (bottom row) the changes are clearer and more consistent across the range of observed values. Holding other factors constant, the PERU parameter set lies above the corresponding CHES parameter set for almost all nitrate samples and below it for almost all ammonium samples. However, for nitrate the ranges over which the changes occur are not the same. N\_BUR\_DOM\_PERU largely increases nitrate at the low end of the range relative to N\_BUR\_DOM\_CHES while the SNP\_PERU/SNP\_BUR\_DOM\_PERU simulations see the increase more at the upper end of the range relative to SNP\_CHES/SNP\_BUR\_DOM\_CHES. Adding dissolved organic matter and burial processes lowers both the nitrate and ammonium concentrations. Adding pathways generally lowers nitrate (Fig. 4d) and has a relatively small impact on ammonium (Fig. 4f).

#### Hypoxic volume time series

An alternative diagnostic of model accuracy of interest to policy makers is the total volume of hypoxic waters (oxygen concentrations less than  $62.5 \text{ mmol/m}^3$ ). As previously noted, N\_CHES\_BUR\_DOM has been used as to hindcast hypoxia within the Bay and so large deviations from what it predicts (black line Fig. 5) are suspect. As shown in Fig. 5, SNP\_CHES (green line Fig. 5) produces far too much hypoxia and SNP\_BUR\_DOM\_PERU (purple line, Fig. 5) produces far too little. SNP\_PERU (dark blue line, Fig. 5) produces too much hypoxia early in the spring and the first appearance of hypoxia in N\_BUR\_DOM\_PERU (light blue line, Fig. 5) is too late in the summer. SNP\_BUR\_DOM\_CHES has a simulation similar to VIMS hindcast in terms of when hypoxia appears and disappears, but the magnitude of peak hypoxic volume is significantly smaller.

#### $R^2$ results for other sites

Extending the model simulations to two other sites in Chesapeake Bay (CB2.2 at the northern end of the hypoxic zone, Table 4 and CB5.3 at the southern end of the hypoxic zone, Table 5) further shows the necessity of constructing a better model that can skillfully simulate biogeochemical fields throughout the Bay. N\_BUR\_DOM\_CHES does not do a good job at predicting the annual cycle of many fields at these sites, specifically nitrate and ammonium at CB5.3 and all three fields at CB2.2. While other model configurations can improve the simulation of individual species at individual sites, no one simulation produces an improvement for oxygen, ammonium and nitrate at all sites, and in general improving the simulation of ammonium makes the simulation of oxygen worse and vice versa.

**Table 4**

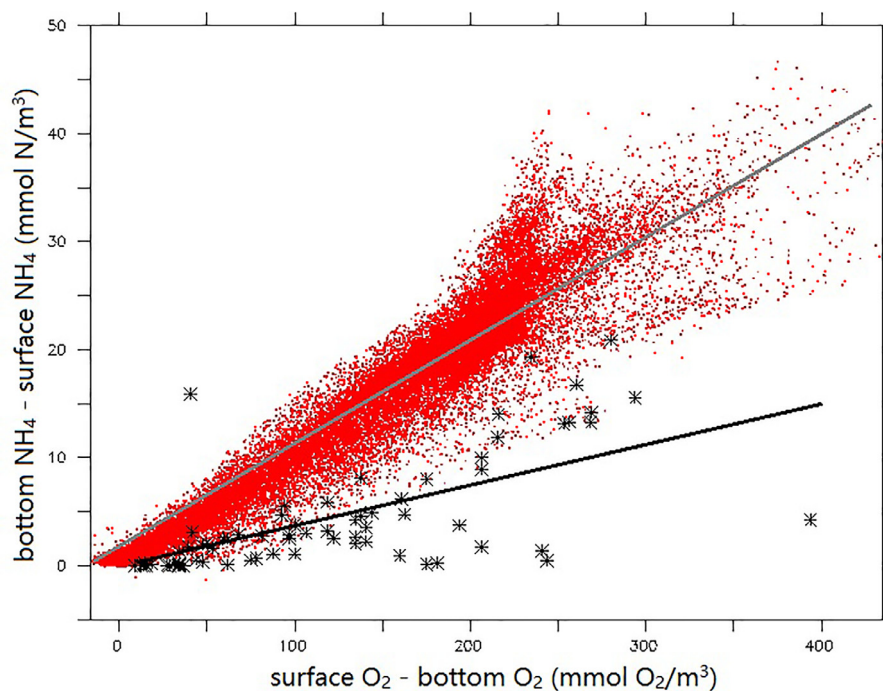
Error metrics for the model suite compared with observations at CB2.2.

|          | N_BUR_DOM_CHES | SNP_PERU | SNP_CHES | N_BUR_DOM_PERU | SNP_BUR_DOM_CHES | SNP_BUR_DOM_PERU |
|----------|----------------|----------|----------|----------------|------------------|------------------|
| Oxygen   | 0.28           | 0.73     | -0.17    | -0.23          | -2.47            | -1.41            |
| Nitrate  | -0.03          | 0.11     | 0.02     | -22.66         | -3.82            | -7.57            |
| Ammonium | -1.49          | -6.99    | -5.72    | 0.01           | -0.27            | -0.36            |

**Table 5**

Error metrics for the model suite compared with observations at CB5.3.

|          | N_BUR_DOM_CHES | SNP_PERU | SNP_CHES | N_BUR_DOM_PERU | SNP_BUR_DOM_CHES | SNP_BUR_DOM_PERU |
|----------|----------------|----------|----------|----------------|------------------|------------------|
| Oxygen   | 0.82           | 0.47     | 0.78     | 0.50           | 0.23             | -0.03            |
| Nitrate  | -7.29          | -4.54    | -0.43    | 0.45           | -0.59            | -0.20            |
| Ammonium | -9.48          | -9.37    | -9.13    | 0.72           | 0.27             | -0.46            |



**Fig. 6.** Modeled surface oxygen minus bottom oxygen ( $\text{mmol O}_2/\text{m}^3$ ) vs bottom  $\text{NH}_4$  minus surface  $\text{NH}_4$  ( $\text{mmol N}/\text{m}^3$ ) from SNP\_BUR\_DOM\_CHES (red dots) vs observations (black symbols) from 2017. The grey line shows the trend for the simulation while the black line shows the trend for the observations.

### Motivating the sensitivity cases

**$\text{O}_2$ :N Stoichiometry:** One potential reason for the tradeoff between our simulations of ammonium and oxygen is stoichiometry. As shown in Fig. 6, during the summertime, the vertical gradient of oxygen tightly tracks the ammonium vertical gradient with a slope of 0.08 in SNP\_BUR\_DOM\_CHES. However, when we look at observations, we see a shallower relationship between oxygen gradient and ammonium gradient, with a slope of 0.04 (Fig. 6 solid black line). This suggests a systematic problem in how our models are constructed. Given that the slope (0.08) between gradient of oxygen and gradient of ammonium of our simulation is around two times higher than the corresponding slope (0.04) of observations, this suggests that nitrogen should be more efficient at removing oxygen, in other words the  $\text{O}_2$ :N stoichiometric ratio should be higher than in the previously published models. We make a relatively small change in our sensitivity study, increasing oxygen:nitrate from 8.625 to 9.625 and oxygen:ammonium from 6.625 to 7.625.

**Optics:** The different representation of light attenuation in SNP\_BUR\_DOM\_CHES versus N\_BUR\_DOM\_CHES also has a large effect in model performance. As noted in the companion paper, in the first half of 2017 SNP\_BUR\_DOM\_CHES has much more primary productivity (30.3  $\text{Gmol N}$  vs. 6.7  $\text{Gmol N}$ ) and burial of nitrogen (2.7  $\text{Gmol N}$  vs. 1.35  $\text{Gmol N}$ ) than N\_CHES\_BUR\_DOM. Analysis of

**Table 6**

Error metrics at CB3.3 for the sensitivity cases compared with observations. A perfect model would have  $R^2=1$  and bias=0. Values of  $R^2<0$  are associated with large biases, which result in the error variance being larger than the sample variance. Because biases dominate the error for nitrite we focus on biases here.

|   | $R^2$ /bias for $O_2$ | $R^2$ /bias for $NH_4$ | $R^2$ /bias for $NO_3$ | bias for $NO_2$ |
|---|-----------------------|------------------------|------------------------|-----------------|
| SNP_BUR_DOM_CHES                          | 0.59/51.23            | -0.03/3.14             | 0.20/-6.02             | 4.33            |
| Higher $O_2$ :N ratio                     | 0.58/51.23            | -0.15/3.35             | 0.19/-6.03             | 3.96            |
| Add CDOM absorption                       | 0.73/35.27            | 0.72/1.39              | 0.49/2.49              | 12.71           |
| Make nitrification/denitrification easier | 0.61/50.22            | -1.46/5.36             | 0.20/-5.91             | 0.86            |
| Reference                                 | 0.78/28.15            | 0.33/5.01              | 0.26/4.36              | 3.78            |

**Table 7**

Error metrics for the model suite compared with observations at CB 2.2 and CB5.3.

| $R^2$ of CB2.2 | SNP_BUR_DOM_CHES | Add CDOM | Easier (de)nitrification | Higher ratio | Reference |
|----------------|------------------|----------|--------------------------|--------------|-----------|
| Oxygen         | -2.47            | 0.40     | -2.30                    | -2.89        | 0.39      |
| Nitrate        | -3.82            | 0.04     | -1.96                    | -1.98        | -1.87     |
| Ammonium       | -0.27            | -1.44    | -1.15                    | -1.10        | -0.19     |
| $R^2$ of CB5.3 | SNP_BUR_DOM_CHES | Add CDOM | Easier (de)nitrification | Higher ratio | Reference |
| Oxygen         | 0.23             | 0.88     | 0.75                     | 0.77         | 0.82      |
| Nitrate        | -0.59            | -3.12    | -0.38                    | -0.37        | -3.93     |
| Ammonium       | 0.27             | -2.43    | -1.42                    | -1.06        | -5.89     |

why this might be the case showed that SNP\_BUR\_DOM\_CHES had a much deeper penetration of light, more primary productivity at depth and more material hitting the bottom. To evaluate the impact of changing the light attenuation we run a simulation in which the chl-only absorption in the SNP run was replaced by the formulation in Da et al. [5] reported in Eq. 8.

Ease of nitrification/denitrification: As also seen in our companion paper, the SNP runs (including SNP\_BUR\_DOM\_CHES) exhibit a high bias for nitrite. However as previously noted, these runs set minimum thresholds for nitrite to denitrify and had a very slow rate of governing the conversion of nitrite to ammonium. We show the impacts of setting the denitrification thresholds to zero, decreasing half-saturation concentration of  $NO_2$  in denitrification from  $30 \text{ mmol N m}^{-3}$  to  $1 \text{ mmol N m}^{-3}$  and increasing the rate constant for nitrite to ammonium conversion so that it matches the conversion rate of nitrate to nitrite ( $0.05 \text{ day}^{-1}$ ). Finally, we developed a new reference simulation that included all three of these changes.

#### Combining multiple changes to improve the simulation: simulation of hydrography at CB3.3

The values for  $R^2$  and bias shown in Table 6 demonstrate that our different changes help to improve performance of SNP\_BUR\_DOM\_CHES in different ways. Increasing the stoichiometric ratios drops the nitrite bias while slightly degrading the simulation of oxygen, ammonium and nitrate. Adding CDOM absorption dramatically increases skills for simulating oxygen, nitrate and ammonium, but it gives unrealistic results for nitrite. Making nitrification easier significantly reduces the bias for nitrite, but at the cost of decreasing  $R^2$  for oxygen, nitrate and ammonium. Combining all those changes (bottom row) gives us a simulation that produces better joint simulations of oxygen, nitrate and ammonium at CB3.3 (as given by the sum of the three  $R^2$  terms) than either N\_BUR\_DOM\_CHES or SNP\_PERU (the two published models). Additionally, it does so while reducing the high nitrite bias seen in the SNP models, though nitrite is still overestimated.

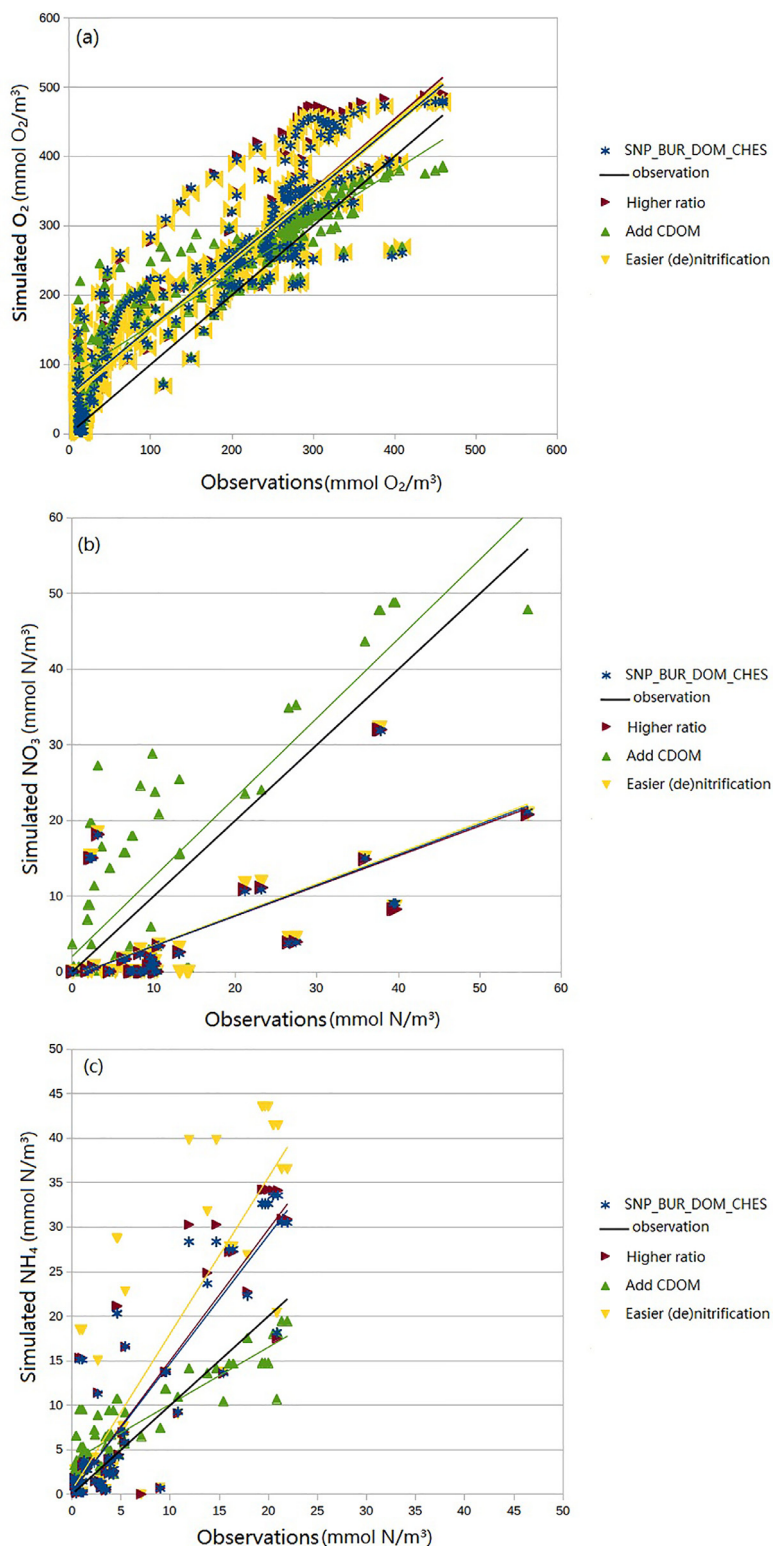
Scatter plots (Fig. 7) show that for oxygen, adding CDOM absorption gives best results comparing with observations. SNP\_BUR\_DOM\_CHES along with the two other cases overpredict oxygen. For nitrate, adding CDOM absorption also gives the most similar results to observations while other cases underestimate nitrate. Finally, adding CDOM absorption leads us to somewhat underestimated ammonium, but the resulting simulation is still better compared to other cases.

#### Combining multiple changes to improve the simulation: Hypoxic volumes

Turning to the predicted hypoxic volume for three cases, Fig. 8 shows that adding CDOM absorption increases hypoxia earlier in summer. Making nitrification/denitrification easier or increasing the stoichiometric ratio both increase hypoxia when compared with SNP\_BUR\_DOM\_CHES.

#### Combining multiple changes to improve the model: $R^2$ from other sites

As seen in Table 7, the changes in hydrographic fields seen at CB3.3 translate to some extent to the northern site. We find that the optics changes improve oxygen and nitrate but make ammonium worse at CB2.2. At CB5.3 changing the optics makes oxygen much better but makes nitrite and ammonium worse. Making nitrification and denitrification easier improves all three fields at CB2.2 (in notable contrast to CB3.3, where they all get slightly worse), improves oxygen and nitrate at CB5.3, but makes ammonium worse.



**Fig. 7.** Simulated versus observed (a) oxygen (mmol  $O_2/m^3$ ) (b) nitrate (mmol  $N/m^3$ ) (c) and ammonium (mmol  $N/m^3$ ) at coincident times and locations from SNP\_CHES\_BUR\_DON (blue) and three sensitivity cases, one with increased  $O_2:N$  stoichiometric ratio (purple triangles), one including CDOM absorption in the optics (green triangle) and one that makes nitrification and denitrification easier (yellow triangle).

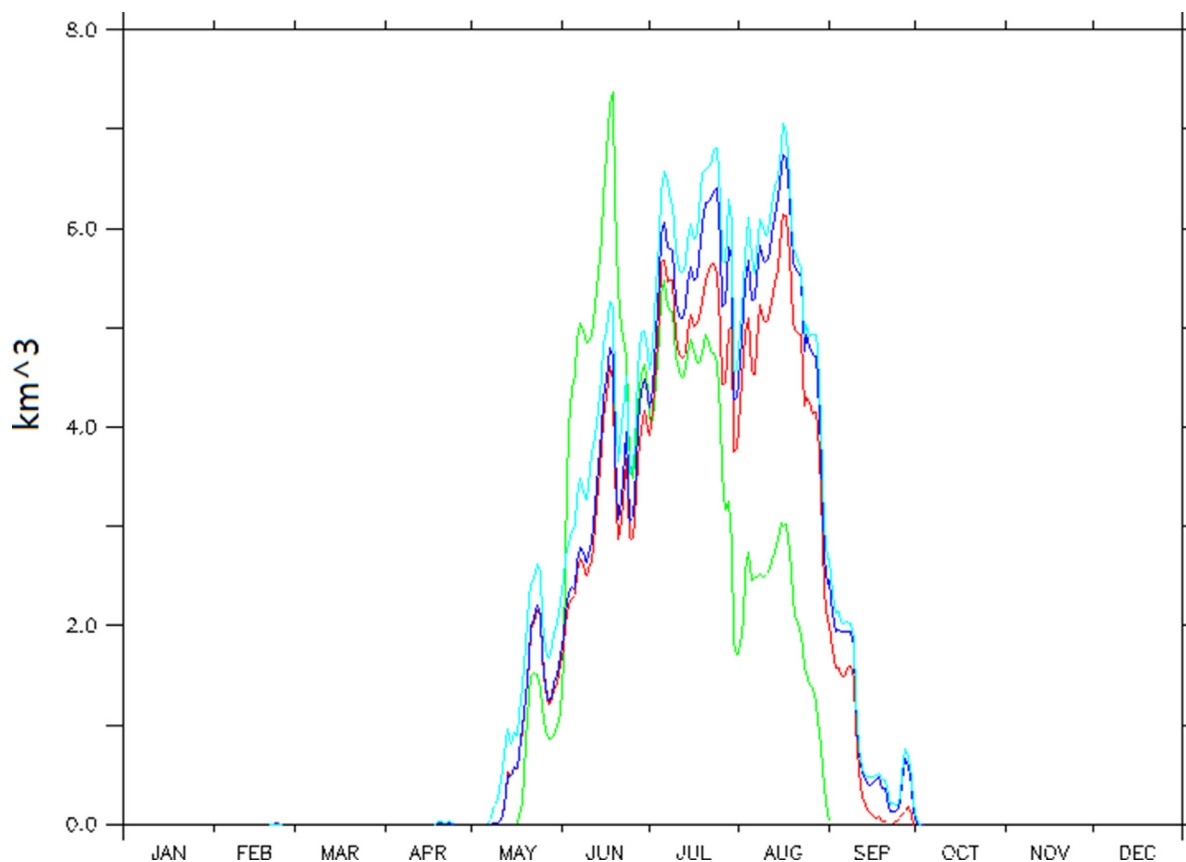


Fig. 8. Hypoxic volume for three sensitivity cases: SNP\_BUR\_DOM\_CHES (red), adding CDOM absorption (green), making nitrification/denitrification easier (dark blue), increasing  $O_2:N$  stoichiometric ratio (light blue).

Increasing the  $O_2:N$  ratio improves nitrate at CB2.2 but makes ammonium and oxygen worse, and improves oxygen and nitrate at CB5.3 while again making ammonium worse.

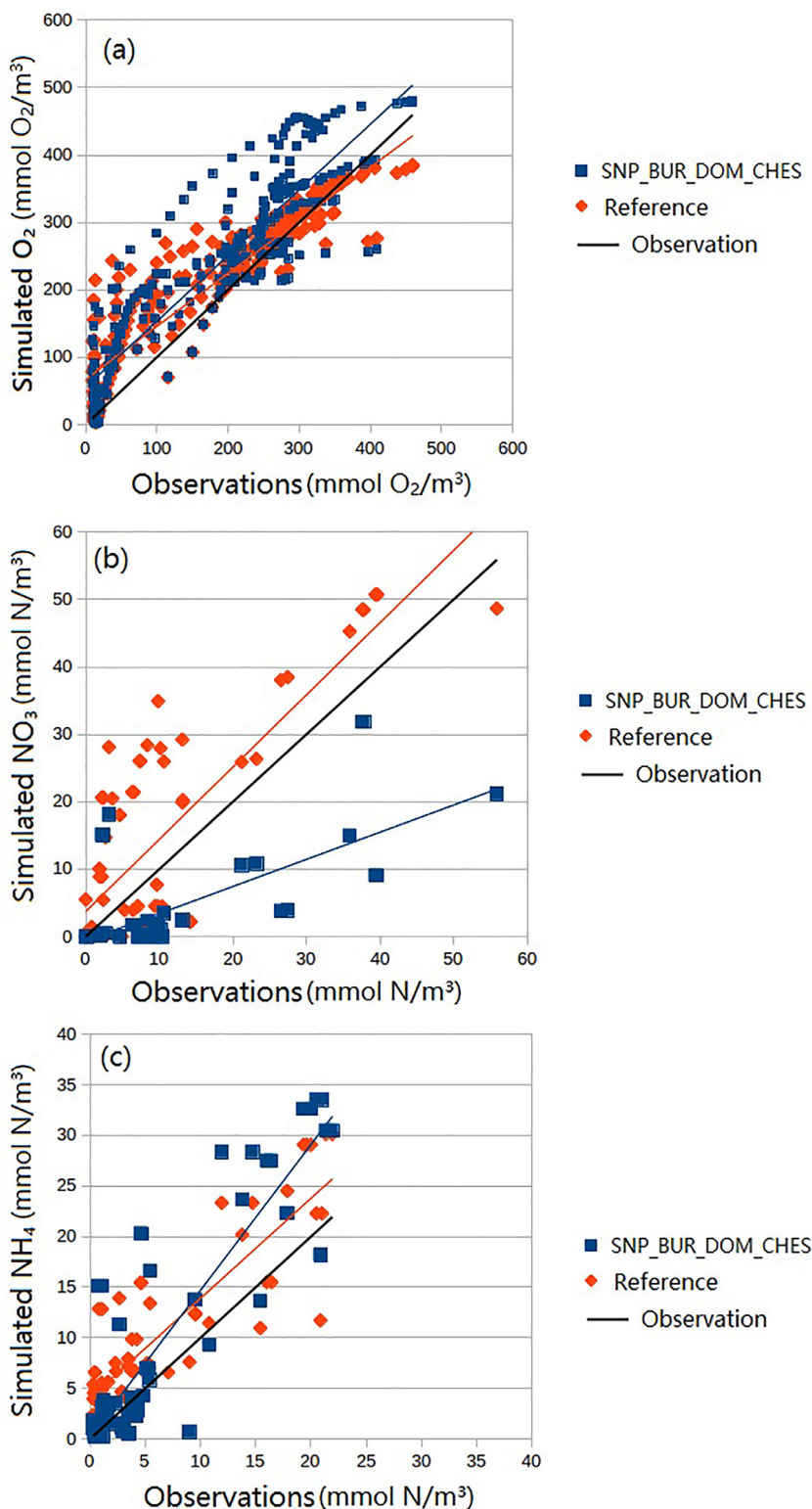
#### Performance of merged (new reference) model

While the bulk statistics make it clear that our new reference run does a better job at simulating oxygen, nitrate and ammonium at CB3.3 than the SNP\_BUR\_DOM\_CHES run on which it is based, it is not clear which parts of the spatiotemporal distribution are improved. Scatter plots (Fig. 9) clearly show that the new reference simulation can simulate oxygen, nitrate and ammonium reasonably well. From 0-200  $mmol O_2/m^3$  range, it still slightly overestimates oxygen, but in 200-500  $mmol O_2/m^3$  range, it simulates oxygen levels that are very close to observations. The new reference run overestimates nitrate but to a much smaller degree than SNP\_BUR\_DOM\_CHES underestimates it. Our reference run also slightly overestimates ammonium in 0-5  $mmol N/m^3$  range but simulates it well in 5-25  $mmol N/m^3$  range.

#### Hypoxic volume in reference run

Finally, we turn to hypoxic volume. As shown in Fig. 10, our new reference run gives a very similar simulation of hypoxic volume as in VIMS hindcast (black), much better than SNP\_BUR\_DOM\_CHES (red), which has a lower volume in the summer.

To conclude, our merged model SNP\_BUR\_DOM\_CHES with CDOM absorption, increased stoichiometric  $O_2:N$  ratio and easier nitrification and denitrification produces a reasonable simulation for oxygen, nitrate, ammonium and hypoxic volume. In most cases it produces lower errors for all three fields relative to the original N\_BUR\_DOM\_CHES model at CB2.2 at the northern edge of the hypoxic zone, CB3.3 in the heart of the hypoxic zone, and CB5.3 at the southern edge of the hypoxic zone though the best simulation is found at CB3.3. The one exception is nitrate at CB2.2, which is overestimated in the reference run at CB2.2. Our new model has a comparable simulation of hypoxia to N\_BUR\_DOM\_CHES, which has been used for hindcasting hypoxia in the Bay. Our new model also has much lower nitrite biases than any of the SNP series of models. While nitrite bias remains a problem, we believe this model can be used as a basis for future work.



**Fig. 9.** Simulated versus observed oxygen (mmol  $O_2/m^3$ ) (a) nitrate (mmol N/ $m^3$ ) (b) and ammonium (mmol N/ $m^3$ ) (c) at coincident times and locations from SNP\_CHES\_BUR\_DON (blue) and new reference simulation which makes all three changes included in previous sensitivity runs (orange).



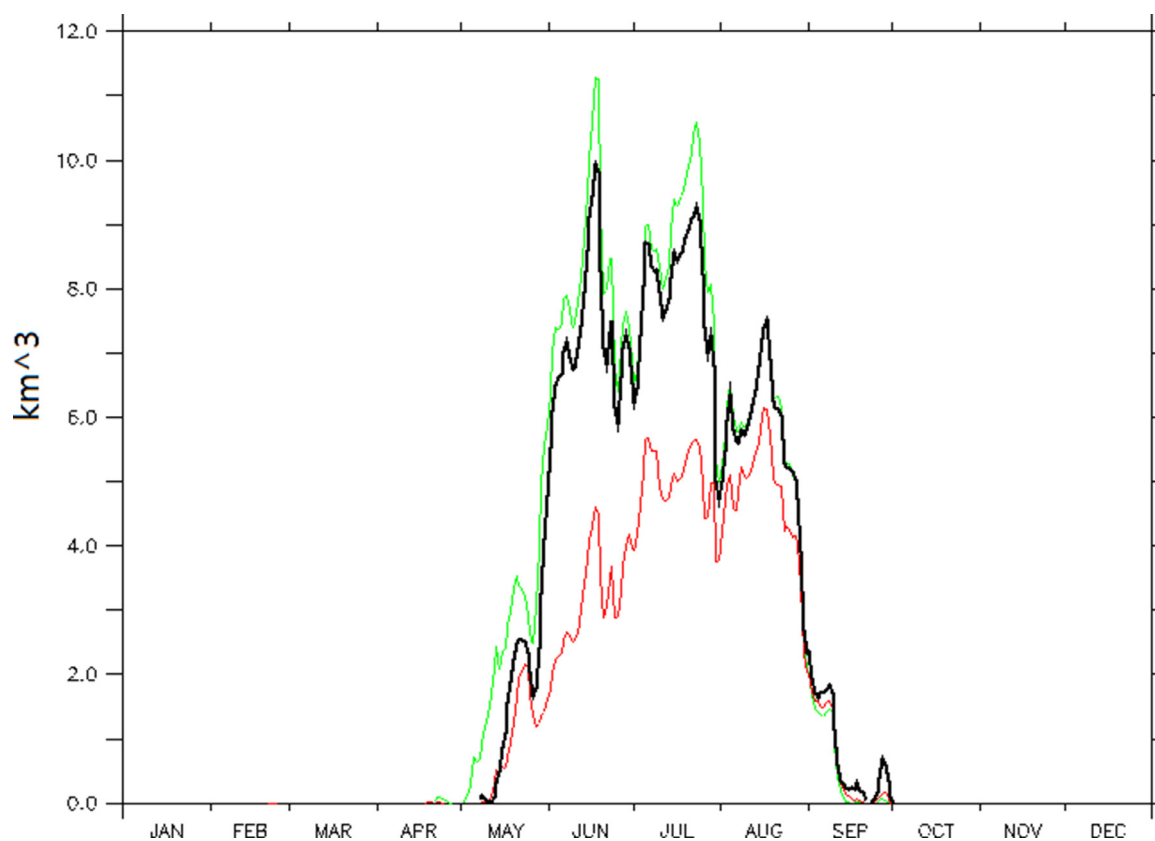


Fig. 10. Hypoxic volume simulated in N\_BUR\_DOM\_CHES (VIMS hindcast, Black), SNP\_BUR\_DOM\_CHES (red) and New reference (green).

### CRedit author statement

**Rui Jin:** Conceptualization, Methodology, Software, Investigation, Validity tests, Visualization, Data curation, Writing- Original draft preparation. **Marie-Aude Pradal:** Conceptualization, software, data curation, writing-original draft, writing-review and editing. **Kalev Hantsoo:** Software, writing-review and editing. **Anand Gnanadesikan:** Conceptualization, Methodology, Visualization, Writing-original draft preparation, writing-review and editing, supervision **Pierre St. Laurent:** Software, data curation, writing-review and editing. **Christian Bjerrum:** Software, writing-review and editing.

### Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

### Data availability

Data will be made available on request.

### Acknowledgments

This work was supported by  
 DOE Office of Science ([SC0019344](#))  
 NOAA Climate Program Office ([NA16OAR4310174](#))  
 The Johns Hopkins Institute for Data Intensive Engineering and Science ([1201600132](#))  
 Danish National Research Foundation ([DNRF 53](#))  
 Villum Foundation (grant [16518](#))

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