



## RESEARCH LETTER

10.1029/2021GL097472

## Key Points:

- Atmospheric response in the Indo-Gangetic Plain varied according to seasonal changes and emissions reductions due to COVID-19 lockdown
- NO<sub>3</sub> production was mainly affected by emission changes, while NO<sub>3</sub> sinks were sensitive to both emissions and seasonal changes
- Nitryl chloride, a photolytic chlorine radical source not previously considered in the inland Indo-Gangetic Plain, may be up to 5.5 ppbv

## Supporting Information:

Supporting Information may be found in the online version of this article.

## Correspondence to:

Y. Rudich,  
[yinon.rudich@weizmann.ac.il](mailto:yinon.rudich@weizmann.ac.il)

## Citation:

Meidan, D., Brown, S. S., Sinha, V., & Rudich, Y. (2022). Nocturnal atmospheric oxidative processes in the Indo-Gangetic Plain and their variation during the COVID-19 lockdowns. *Geophysical Research Letters*, 49, e2021GL097472. <https://doi.org/10.1029/2021GL097472>

Received 11 JAN 2022  
Accepted 28 MAR 2022

## Author Contributions:

**Conceptualization:** D. Meidan, S. S. Brown, V. Sinha, Y. Rudich  
**Data curation:** D. Meidan, V. Sinha  
**Formal analysis:** D. Meidan, S. S. Brown  
**Funding acquisition:** V. Sinha, Y. Rudich  
**Investigation:** D. Meidan, S. S. Brown, V. Sinha, Y. Rudich  
**Methodology:** D. Meidan, S. S. Brown, V. Sinha, Y. Rudich  
**Project Administration:** Y. Rudich

© 2022. The Authors.

This is an open access article under the terms of the [Creative Commons Attribution-NonCommercial-NoDerivs License](#), which permits use and distribution in any medium, provided the original work is properly cited, the use is non-commercial and no modifications or adaptations are made.

# Nocturnal Atmospheric Oxidative Processes in the Indo-Gangetic Plain and Their Variation During the COVID-19 Lockdowns

D. Meidan<sup>1</sup> , S. S. Brown<sup>2,3</sup> , V. Sinha<sup>4</sup> , and Y. Rudich<sup>1</sup>

<sup>1</sup>Department of Earth and Planetary Sciences, Weizmann Institute of Science, Rehovot, Israel, <sup>2</sup>NOAA Chemical Sciences Laboratory, Boulder, CO, USA, <sup>3</sup>Department of Chemistry, University of Colorado, Boulder, CO, USA, <sup>4</sup>Department of Earth and Environmental Sciences, Indian Institute of Science Education and Research Mohali, Mohali, India

**Abstract** This study investigates selected secondary atmospheric responses to the widely reported emission change attributed to COVID-19 lockdowns in the highly polluted Indo-Gangetic Plain (IGP) using ground-based measurements of trace gases and particulate matter. We used a chemical box-model to show that production of nighttime oxidant, NO<sub>3</sub>, was affected mainly by emission decrease (average nighttime production rates 1.2, 0.8 and 1.5 ppbv hr<sup>-1</sup> before, during and relaxation of lockdown restrictions, respectively), while NO<sub>3</sub> sinks were sensitive to both emission reduction and seasonal variations. We have also shown that the maximum potential mixing ratio of nitryl chloride, a photolytic chlorine radical source which has not been previously considered in the IGP, is as high as 5.5 ppbv at this inland site, resulting from strong nitrate radical production and a potentially large particulate chloride mass. This analysis suggests that air quality measurement campaigns and modeling explicitly consider heterogeneous nitrogen oxide and halogen chemistry.

**Plain Language Summary** The Indo-Gangetic Plain (IGP) is one of the most polluted regions on earth, with poor air quality affecting the majority of the Indian population. The atmospheric chemistry that transforms major regional emissions into harmful secondary pollutants is complex. Here, we quantify, for the first time, several important oxidative processes and show the potential for substantial oxidation of biogenic volatile organic compounds and the production of chlorine through unconventional chemistry in the IGP. We further show how these chemical cycles varied due to the emission reductions as a result of COVID-19 lockdown, findings that will serve to define their sensitivity to future emission changes in the region.

## 1. Introduction

Indo-Gangetic Plain (IGP) is home to almost half the Indian population. However, it accounts for only one fifth of India's land area, making it one of the most densely populated places on Earth. Known for its fertile lands, the IGP is a region with intense farming and agricultural activities. Due to the unique topography of the region, air pollution can either persist, a process relevant mostly during wintertime, or be impacted by long-range transport, which occurs mostly during the summer (pre-monsoon; Pawar et al., 2015). These factors result in frequent high-pollution events and rise in premature deaths related to poor air quality (David et al., 2019).

High concentrations of air pollutants in India and their impacts on the atmosphere and on human health have been extensively studied (Ghude et al., 2016; V. Kumar et al., 2016; Mohan & Saranya, 2019; Ojha et al., 2012; Yadav et al., 2014). Tropospheric ozone concentrations are influenced by the concentrations of ozone precursors (e.g., nitrogen oxides [NO<sub>x</sub> = NO + NO<sub>2</sub>] and volatile organic compounds), high exposure being detrimental to human health and agriculture (Y. Chen et al., 2021; Gaudel et al., 2018; Lefohn et al., 2018; Lu et al., 2020; Sharma et al., 2019; B. Sinha et al., 2015). Among the sources of anthropogenic emissions of NO<sub>x</sub> and volatile organic compounds (VOCs) in the IGP region are industry, heavy traffic, power generation and agricultural biomass burning and residential fuel usage. Particulate matter smaller than 2.5 μm in diameter (PM<sub>2.5</sub>) arises from primary emissions or forms in situ by secondary chemical reactions of gas-phase precursors such as NO<sub>x</sub>, ozone, VOCs and sulfur dioxide (SO<sub>2</sub>). 2016 World Health Organization (WHO) report (World Health, 2016) states that PM<sub>2.5</sub> levels in Delhi during the years 2011–2015 were the highest measured in mega-cities around the world and a large fraction of it (>50%) is derived from oxidation of precursors (Gani et al., 2019). More than 80% of Indian population suffers from poor air quality, with conditions deviating from WHO standards for PM<sub>2.5</sub> exposure. In the IGP, for example, PM<sub>2.5</sub> levels often exceed 160 μg/m<sup>3</sup> (Ravishankara et al., 2020). David et al. (2019) investigated

**Resources:** S. S. Brown, V. Sinha, Y. Rudich  
**Software:** D. Meidan  
**Supervision:** S. S. Brown, V. Sinha, Y. Rudich  
**Validation:** S. S. Brown  
**Visualization:** D. Meidan  
**Writing – original draft:** D. Meidan  
**Writing – review & editing:** D. Meidan, S. S. Brown, V. Sinha, Y. Rudich

impacts of emissions from different regions in India on premature death related to  $PM_{2.5}$  exposure. They estimated that anthropogenic air pollution from the IGP region is responsible for 24%–31% of deaths in other regions in India, while about 70% of deaths are caused by local emissions.

Atmospheric oxidation capacity is controlled by concentrations and sources of the major oxidants OH,  $O_3$  and  $NO_3$ . Chlorine radicals may also initiate tropospheric oxidation cycles, and recent observations in polluted regions have shown  $ClNO_2$  to be an important photolytic precursor (Baker et al., 2016; Q. Chen et al., 2022; Haskins et al., 2019; Jeong et al., 2019; Wang et al., 2020; Xia et al., 2021). Nocturnal aqueous-phase reaction of particulate chloride with  $N_2O_5$  produces  $ClNO_2$  that affects next-day chemistry through photolysis (Osthoff et al., 2008). Sources of soluble chloride in the IGP include biomass burning, coal combustion, industrial waste incineration and to a lesser extent long-range transport of sea salt aerosol (Gunthe et al., 2021; Saiz-Lopez & von Glasow, 2012; Wang et al., 2019).

In order to limit the spread of COVID-19, governments imposed lockdowns in early 2020 that restricted mobility and public activities such as schools and businesses. India imposed a strict lockdown from the end of March until the beginning of May 2020. We studied a timeframe that can be regarded as three different periods: between 1 and 22 March 2020 (herein “before lockdown”), between 23 March–14 April 2020 (herein “strict lockdown”), and finally between 15 April–1 May 2020 (herein “lockdown relaxation”). The main differences between the two lockdown periods were that during lockdown relaxation, agricultural activities were approved (mainly harvesting) and some mobility restrictions were lifted (leading to slightly increased transportation emissions). Due to the lockdown, atmospheric concentrations of many anthropogenic pollutants, such as  $NO_x$ , anthropogenic VOC (AVOC) and  $SO_2$ , decreased. Forster et al. (2020) analyzed national mobility data and found an average decrease of >70% in mobility in India during the strict lockdown, resulting in  $NO_x$  emission decrease. Such sudden and unprecedented emission reductions present a unique opportunity to investigate chemical responses under different chemical scenarios.

In this study, we modeled the impact of lockdown restrictions on selected atmospheric oxidation pathways. These pathways are superimposed on changing seasonality of the chemical regime in winter-to-spring transition in the highly polluted IGP region. Elucidating  $NO_3$  production and consumption pathways can help our understanding of the state of secondary atmospheric processes in the IGP under different atmospheric chemical regimes imposed by the lockdown.

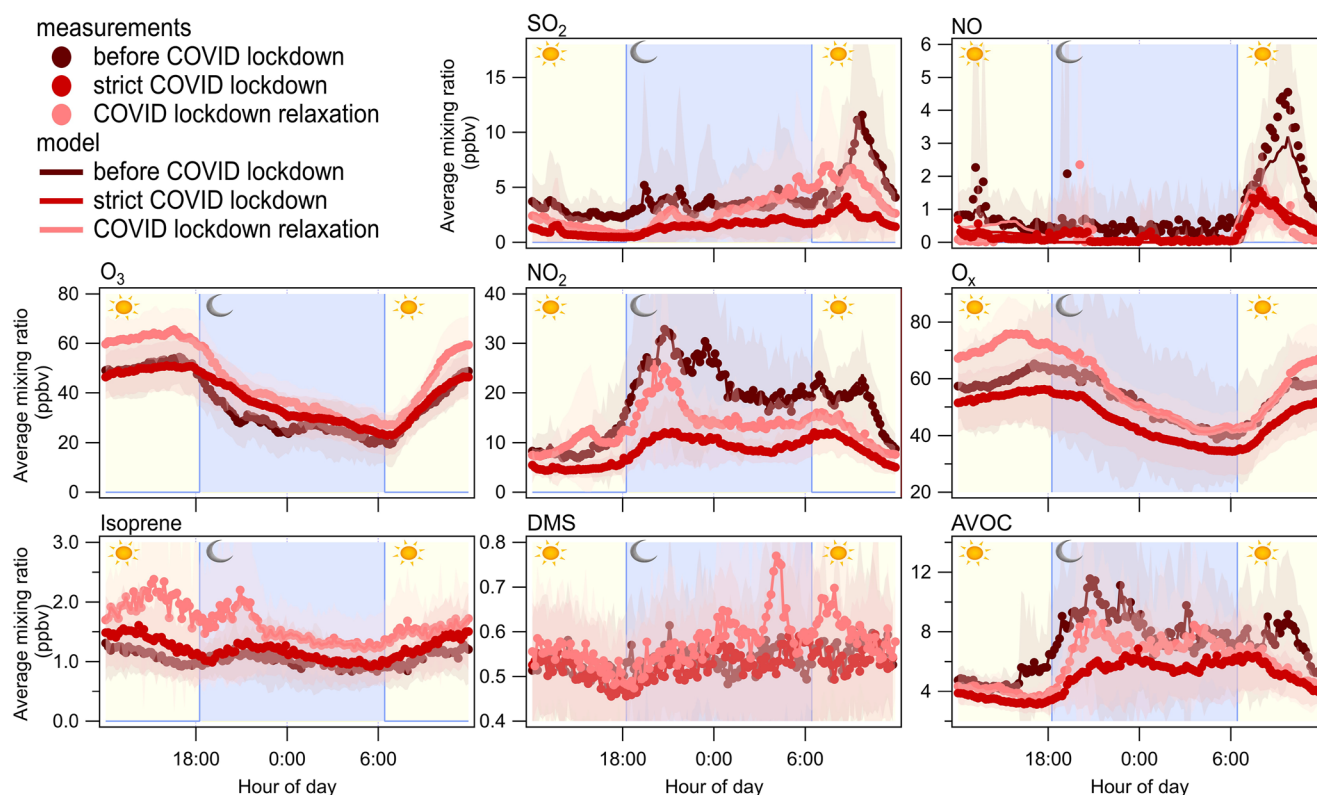
## 2. Materials and Methods

### 2.1. Site Description

The measurement site is located on the outskirts of Mohali, a city in northwest IGP, in Indian state of Punjab. The measurement facility is housed in the Indian Institute of Science Education and Research (IISER; 30.667°N–76.729°E, 310 m a.s.l.) campus. The facility consists of a high sensitivity proton transfer reaction quadrupole mass spectrometer (PTR-QMS), an air quality station equipped with gas analyzers for the detection of trace levels (pptv–ppbv range) of  $O_3$ ,  $NO_x$ , CO,  $SO_2$ ,  $PM_{2.5}$  and  $PM_{10}$  aerosol mass concentrations and a meteorological station for wind direction, wind speed, ambient temperature, relative humidity and solar radiation measurements (Met One Instruments Inc.). Inlets are located at 20 m above ground level. These instruments and their QA/QC have been comprehensively described in previous works reporting multi-year data from the facility (Chandra & Sinha, 2016; V. Kumar et al., 2020; V. Sinha et al., 2014). Calibrations performed before and after the lockdown were consistent with the general instrumental sensitivities and zero drifts. Overall uncertainty for all VOCs reported in this work was less than 20%, except for formaldehyde for which it was estimated to be less than 30% due to lack of calibration gas standard (see V. Kumar et al., 2020). For other trace gases and particulate matter (Pawar et al., 2015) it was less than 10%.

### 2.2. Model Description

The Framework for zero-Dimensional Atmospheric Modeling (FOAM), based on Master Chemical Mechanism (MCM) 3.3.1, was used to explore the effect of COVID-19 lockdown conditions on atmospheric oxidation processes associated with  $NO_3$  (Jenkin et al., 2015; Wolfe et al., 2016). Observations of trace gases and particulates from the Mohali site were used to calculate the average diel cycles of species used to constrain the model in



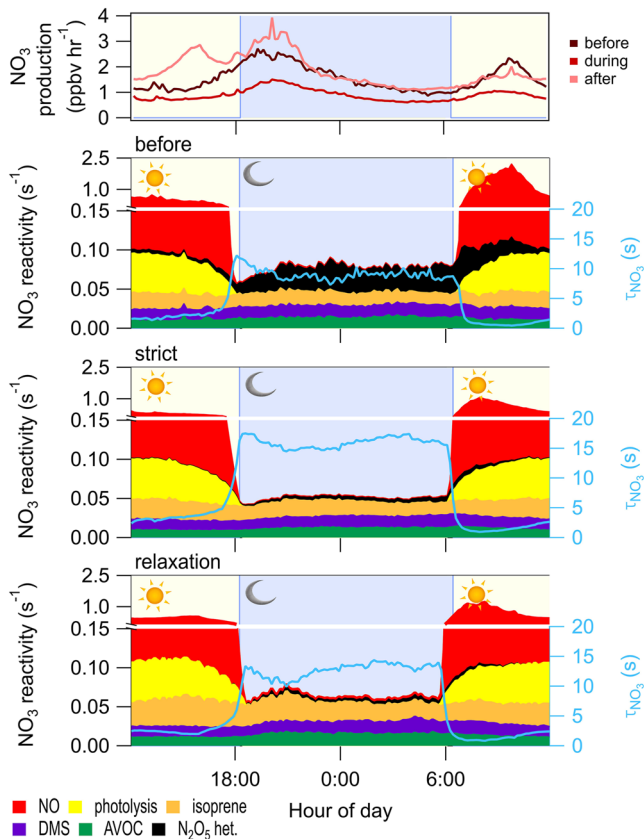
**Figure 1.** Diel cycles, at 10-min time resolution, of the three periods: before lockdown (brown), during strict lockdown (red), and during lockdown relaxation (pink). Dots represent measurements and solid lines represent simulation results. Shaded areas represent standard deviation of the measurements. Changes in isoprene and temperature between lockdown periods are driven by seasonality of the meteorology. During the lockdown, all pollutant mixing ratios decreased, except for ozone, which increased, especially during nighttime.

ten-minute time steps, same resolution as the model. The model was set to simulate a diel cycle with constrained ozone,  $\text{NO}_x$ , CO,  $\text{SO}_2$ , isoprene, DMS, 4 AVOCs, 11 aromatic VOCs, 6 oxygenated VOCs, relative humidity and temperature (Figure 1, Table S1 in Supporting Information S1, additional details on AVOC speciation can be found in Supporting Information S1). Photolysis rates were calculated in F0AM based on the trigonometric solar zenith angle function provided by the MCM. The method provides an upper limit of the photolysis rates since it does not account for surface albedo, overhead ozone column, cloud and aerosol extinction or enhancement, which may alter the photolysis rates (Wolfe et al., 2016). Modeled  $\text{NO}_x$  was constrained to the measurements, but partitioning between NO and  $\text{NO}_2$  was allowed to vary.

In addition to the three lockdown periods (i.e., before lockdown, strict lockdown, and lockdown relaxation), approximately 17% of the before period measurements showed excess NO ( $>5$  ppbv) and low ozone ( $<30$  ppbv; Figure S1 in Supporting Information S1). These datapoints were treated as a separate category, since chemical evolution during events that end in full ozone titration differed considerably from that of non-titrated events (Figure S2 in Supporting Information S1) and are further explored in Supporting Information S1.

Average mixing ratios of measured species across 3 hours prior to sunset (15:30–18:30) provided the initial input for an atmospheric chemistry model intended to represent the residual layer. This approach mimics the transition from a well-mixed boundary layer to a two-layered nighttime boundary layer consisting of a surface layer (or nocturnal boundary layer), impacted by continuous emissions, and an overlying residual layer, consisting of late-afternoon chemical composition that is isolated from the surface emissions.

$\text{NO}_3$  is produced from reaction of  $\text{NO}_2$  with ozone (Brown & Stutz, 2012) and is in equilibrium with its reservoir species  $\text{N}_2\text{O}_5$ , which undergoes heterogeneous uptake to form nitric acid ( $\text{HNO}_3$ ) and  $\text{ClNO}_2$ . The first-order rate coefficient for  $\text{N}_2\text{O}_5$  aerosol uptake is generally expressed as:



**Figure 2.**  $\text{NO}_3$  production rate and reactivity toward anthropogenic volatile organic compound (AVOC), NO, isoprene, photolysis, heterogeneous uptake and DMS during the period before lockdown (top), during the strict lockdown (middle) and after relaxation of the lockdown (bottom). Cyan line is the modeled  $\text{NO}_3$  lifetime.  $\text{NO}_3$  has the shortest lifetime before the lockdown, when the overall emissions were higher. Reduction in aerosol surface area dramatically decreased heterogeneous loss of  $\text{NO}_3$  and increased its atmospheric lifetime.

During the strict lockdown,  $\text{NO}_x$  mixing ratios decreased and  $\text{NO}_3$  production decreased correspondingly relative to before the lockdown. After the lockdown, a late evening  $\text{NO}_2$  peak resulted in a  $\text{NO}_3$  production peak, a feature that appeared only slightly in the prior periods.

Figure 2 shows  $\text{NO}_3$  reactivity toward important atmospheric species and its modeled lifetime.  $\text{NO}_3$  reactivity toward these species is calculated by:

$$k^{\text{NO}_3} = \sum k_i [\text{X}]_i \quad (4)$$

where,  $k_i$  is the rate coefficient ( $\text{cm}^3 \text{s}^{-1}$ ) for reaction between a species X of concentration  $[\text{X}]_i$  and  $\text{NO}_3$ .

Steady-state lifetime,  $\tau_{\text{NO}_3}$ , is given by the inverse of  $k^{\text{NO}_3}$  from Equations 2 and 4. Lifetime is normally calculated with observed  $\text{NO}_3$ . However, in the absence of such measurements at this site (or at any location in the IGP to date),  $(\text{NO}_3)$  in Equations 3 and 5 served as modeled concentration of  $\text{NO}_3$ :

$$[\text{NO}_3] = \tau_{\text{NO}_3} * P(\text{NO}_3) \quad (5)$$

where,  $P(\text{NO}_3)$  is the production rate of  $\text{NO}_3$  ( $\text{ppbv s}^{-1}$ ), as described in Equations 1 and 3.

$$k_{\text{het. N}_2\text{O}_5} = \frac{\gamma \bar{c} \text{SA}}{4} \quad (1)$$

where,  $\gamma$  is the aerosol uptake coefficient,  $\bar{c}$  the mean molecular speed of  $\text{N}_2\text{O}_5$ , and SA the aerosol surface area.  $\text{N}_2\text{O}_5$  measurements that are used to constrain this parameter are not available at this site. Therefore, an uptake coefficient sensitivity test is described in details in Figure S3 of Supporting Information S1. The model uses a single representative uptake coefficient of  $\gamma_{\text{N}_2\text{O}_5} = 10^{-2}$  (Chang et al., 2011; R. Kumar et al., 2014; McDuffie et al., 2018).

The first-order rate coefficient for the  $\text{NO}_3$  loss through  $\text{N}_2\text{O}_5$  aerosol uptake is then (Brown et al., 2003):

$$k_{\text{het. NO}_3} = K_{\text{eq}} [\text{NO}_2] k_{\text{het. N}_2\text{O}_5} \quad (2)$$

where,  $K_{\text{eq}}$  is the equilibrium constant between  $\text{NO}_2$ ,  $\text{NO}_3$  and  $\text{N}_2\text{O}_5$ .

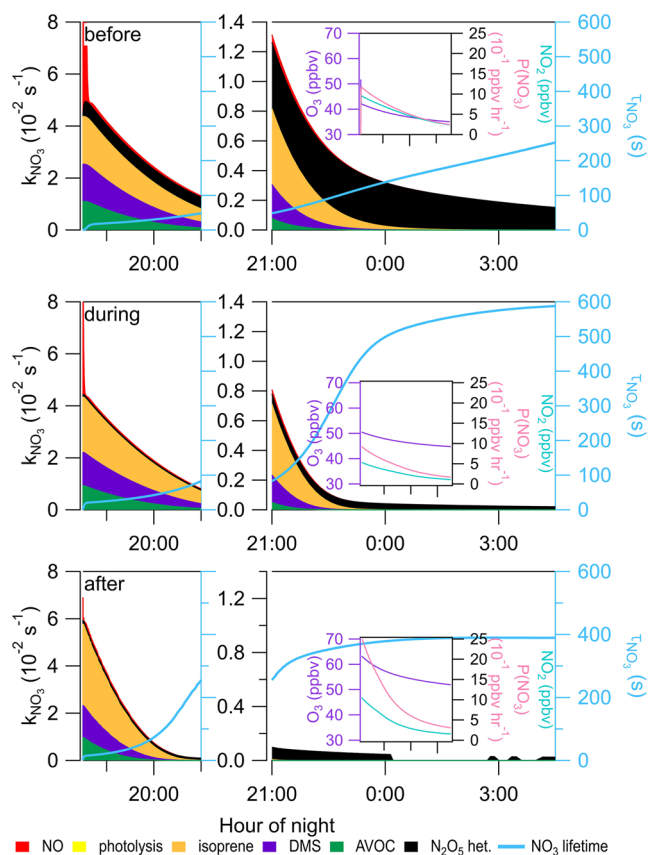
Aerosol surface area was calculated from measured  $\text{PM}_{2.5}$  and  $\text{PM}_{10}$  after assuming a typical size distribution in the IGP according to the relevant season and correction for hygroscopic growth (Pawar & Sinha, 2020). Figure S4 in Supporting Information S1 shows resultant surface area and calculated heterogeneous rate coefficients for  $\text{NO}_3$  and  $\text{N}_2\text{O}_5$ .

### 3. Results

$\text{NO}_3$  is formed from the oxidation of  $\text{NO}_2$  by ozone. It then reacts with  $\text{NO}_2$  to establish an equilibrium with  $\text{N}_2\text{O}_5$ , which can act as a reservoir for  $\text{NO}_3$  or as a loss pathway, by uptake onto aerosols. Since  $\text{NO}_3$  is photolabile, it typically serves as an oxidant only during the night hours. Uptake of  $\text{N}_2\text{O}_5$  to aerosols leads to  $\text{HNO}_3$  and  $\text{ClNO}_2$  production, with a yield that depends on the available chloride (Bertram & Thornton, 2009; Roberts et al., 2009).  $\text{HNO}_3$  may partition to the aerosol phase as inorganic nitrate.

During daytime, reaction with NO and photolysis destroy  $\text{NO}_3$  and recycle it to  $\text{NO}_x$ . The  $\text{NO}_3$  production rate is calculated by:

$$P(\text{NO}_3) = k_{\text{O}_3+\text{NO}_2} [\text{O}_3] [\text{NO}_2] \quad (3)$$



**Figure 3.** Nighttime  $\text{NO}_3$  reactivity toward NO, isoprene, photolysis, heterogeneous uptake, DMS and anthropogenic volatile organic compound (AVOC) during the period before lockdown (top), during strict lockdown (middle) and after lockdown relaxation (bottom). Residual layer is split into the first (until 21:00, left plots) and second (from 21:00, right plots) part of the night to illustrate differences in reactivity scales. Ozone and  $\text{NO}_2$  mixing ratios and production rate of  $\text{NO}_3$  are shown in the inset for the full night period. Cyan line reflects  $\text{NO}_3$  lifetime.

$\text{NO}_3$  during the day ( $0.08 \text{ ppbv hr}^{-1}$ ), and more than twofold after lockdown relaxation ( $0.25 \text{ ppbv hr}^{-1}$ ). Fuchs et al. (2017) and Williams et al. (2016) measured the reactivity of the main daytime oxidant, OH, in polluted areas in north China plain and megacity Beijing, respectively. They have found that during the daytime an average of  $\sim 10\%$  of the OH reactivity is toward isoprene. This is comparable to  $8\text{--}10\%$   $\text{NO}_3$  reactivity toward isoprene in the afternoon once lockdown is imposed.

Surface measurements at night are often not representative of chemistry within the entire boundary layer structure, which consists of both a nocturnal boundary layer and overlying residual layer. Following the approach of Baasandorj et al. (2017), the simulation of chemistry in the residual layer is initialized with atmospheric mixing ratios of the late afternoon, but the model of this layer remained isolated from further surface emissions. Figure 3 shows the calculated  $\text{NO}_3$  reactivity,  $\text{O}_3$  and  $\text{NO}_2$  mixing ratios and modeled production rate of  $\text{NO}_3$  predicted for the residual layer, which could not be directly sampled from this measurement site. Predicted production rate of  $\text{NO}_3$  decreased as surface  $\text{NO}_2$  (and, to a lesser extent,  $\text{O}_3$ ) was consumed.  $\text{O}_3$  mixing ratios in the residual layer increased between the three periods, likely as a result of seasonality in the photochemistry.  $\text{NO}_2$  mixing ratios, by contrast, were impacted by COVID-19 lockdown, marked by a clear decrease. Consequently, during the strict lockdown, modeled  $\text{NO}_3$  production rate was the slowest of the three periods, with fastest  $\text{NO}_3$  production rates occurring upon lockdown relaxation, when  $\text{O}_3$  mixing ratios were the highest.

Production rate of nitrate radicals as shown in Equations 1 and 3 can be used as a useful metric for comparing the extent of nighttime chemistry in different megacities. For example, Brown et al. (2017) determined  $\text{P}(\text{NO}_3)$  from measurements in a tower in the Asian megacity of Seoul, South Korea, in May to early June 2015. Median production rate in Seoul 2015 exceeded 90th percentile  $\text{P}(\text{NO}_3)$  from nighttime research aircraft flights in a similar altitude range in Los Angeles, a comparably sized U.S. megacity. Figure S5 in Supporting Information S1 shows nighttime measurements from Seoul and Los Angeles with  $\text{P}(\text{NO}_3)$  as calculated in our study. 10th percentile  $\text{P}(\text{NO}_3)$  of all three studied periods in Mohali exceeded the 90th percentile production rate measured in Los Angeles. The lockdown relaxation period was measured during a similar time of year as the measurements made in Seoul and Los Angeles. However, median nighttime  $\text{P}(\text{NO}_3)$  in Mohali were the largest during the 2020 lockdown relaxation period ( $1.5 \text{ ppbv hr}^{-1}$ ), compared to Seoul in 2015 ( $1.3 \text{ ppbv hr}^{-1}$ ) and Los Angeles in 2010 ( $0.3 \text{ ppbv hr}^{-1}$ ). These high  $\text{NO}_3$  production rates are comparable to daytime chemical oxidation rates in polluted regions (Volkamer et al., 2010; Young et al., 2012) and lead to a significant nighttime oxidative capacity.

Figure 2 shows that diel profiles of  $\text{NO}_3$  lifetime is similar in all periods. NO and sunlight are the main  $\text{NO}_3$  daytime consumers, longest  $\text{NO}_3$  lifetime occurs at sunset and shortest lifetime occurs at sunrise. Before the lockdown, nighttime  $\text{NO}_3$  lifetime decreased rapidly (from  $\sim 12 \text{ s}^{-1}$  to  $\sim 8 \text{ s}^{-1}$  in under 3 hr) due to heterogeneous uptake by high loading of particulate matter. At the onset of the lockdown, significant decreases in particulate matter concentrations in the IGP region (Singh et al., 2020) eliminated this rapid loss of  $\text{NO}_3$  by heterogeneous uptake, resulting in longer  $\text{NO}_3$  lifetime. NO morning peak consumed most of the  $\text{NO}_3$  (up to 90% Figure S6 in Supporting Information S1) during all periods. As the sun sets, NO mixing ratio decreased and the main  $\text{NO}_3$  reactions are with isoprene, DMS and AVOC.

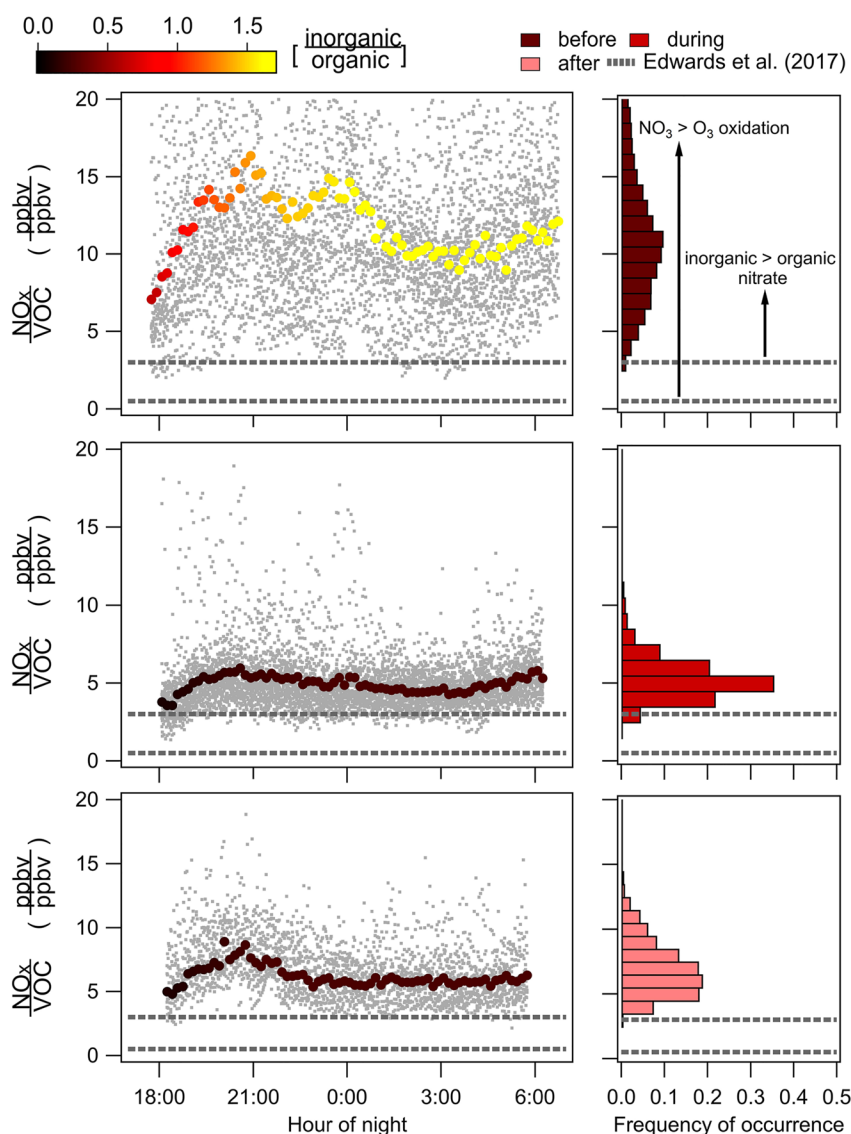
Figure S7 in Supporting Information S1 shows daytime ratio of  $\text{NO}_3$  reactivity toward isoprene, DMS and AVOC to the overall  $\text{NO}_3$  reactivity and predicted amount of produced  $\text{NO}_3$  that reacted with isoprene, DMS and AVOC. Daytime  $\text{NO}_3$  oxidation of these species increases from morning, and reached maximum daytime values in the afternoon. As seen in Figure 2, before the lockdown,  $\text{NO}_3$  oxidation of these three species was approximately the same. However, upon strict lockdown, isoprene consumed more produced

While  $\text{NO}_3$  in the surface and residual layers exhibits similar reactivity at the beginning of the night, predicted  $\text{NO}_3$  lifetime in the residual layer increases as the residual layer VOCs are consumed in the absence of contact with surface-based emissions. Residual layer modeled reactivity of  $\text{NO}_3$  toward  $\text{NO}$  is negligible, since  $\text{NO}$  is rapidly consumed by ozone within  $\sim 2\text{--}5$  min. After the lockdown relaxation, isoprene mixing ratios rose due to higher temperatures and solar flux, as well as higher ozone mixing ratios. Therefore, consumption of the constituents in the residual layer was fastest after the lockdown relaxation. However, predicted  $\text{NO}_3$  lifetime reached the highest values during the strict lockdown, when surface emissions and particulate matter loading were at their lowest.

Integrated production of  $\text{NO}_3$  from  $\text{NO}_x$  over the course of one night in the residual layer exceeds initial mixing ratios of reactive VOCs such as DMS and isoprene, as discussed below. Therefore, loss of  $\text{NO}_3$  and  $\text{N}_2\text{O}_5$  transitions from  $\text{NO}_3 + \text{VOC}$ -dominated to heterogeneous uptake in the residual layer as reactive VOCs are consumed. This transition affects production of inorganic nitrate species  $\text{HNO}_3$  and  $\text{ClNO}_2$ . Figure S8 in Supporting Information S1 shows a comparison of  $\text{HNO}_3$  and  $\text{ClNO}_2$  production potential in the residual and surface layers, here shown as the maximum  $\text{ClNO}_2$  production using an assumed  $\text{ClNO}_2$  yield,  $\phi(\text{ClNO}_2) = 1$ . In this simulation,  $\text{ClNO}_2$  and  $\text{HNO}_3$  are equal. Regardless of the assumed  $\text{ClNO}_2$  yield, total amount of produced inorganic nitrate ( $\text{ClNO}_2 + \text{HNO}_3$ ) would be equal to twice that of  $\text{ClNO}_2$ . Before the lockdown, nighttime maximum  $\text{ClNO}_2$  production in the residual layer (3 ppbv) was about half that in the surface layer (5.5 ppbv). During the strict lockdown, maximum  $\text{ClNO}_2$  production in both layers was similar ( $\sim 1$  ppbv). However, with the lifting of some lockdown restrictions, maximum  $\text{ClNO}_2$  production in the residual layer exceeded production in the surface layer (2.5 and 1.5 ppbv, respectively) due to a faster  $\text{N}_2\text{O}_5$  uptake rate.

$\text{VOC}/\text{NO}_x$  ratios are a useful diagnostic for probing secondary pollutant formation regimes from both photochemistry and nighttime chemistry. The influence of this ratio on daytime ozone photochemistry is well established (Council, 1991; V. Kumar & Sinha, 2021; Seinfeld, 1989). During the day, this ratio indicates the competition between ozone production and consumption through photolysis of  $\text{NO}_2$ , oxidation of VOC,  $\text{NO}$  and  $\text{NO}_2$  reaction with peroxy radicals, scavenging of ozone by  $\text{NO}$  etc. During the night, this ratio indicates the competition between  $\text{NO}_3$ - and  $\text{O}_3$ -dominant regimes of VOC oxidation.  $\text{VOC}/\text{NO}_x$  ratio also determines the competition between inorganic and organic nitrate production during the night and day. At low  $\text{VOC}/\text{NO}_x$  ratios, nighttime production of inorganic nitrate ( $\text{HNO}_3$  and  $\text{ClNO}_2$  from  $\text{N}_2\text{O}_5$  uptake) increases while the level of organic nitrate (the major product of VOC oxidation by  $\text{NO}_3$ ) decreases (Romer Present et al., 2020). The competition between inorganic nitrate and organic nitrate determines the aerosol yield and composition by changing the reaction pathways (Huang et al., 2021; Petit et al., 2021).

Using nighttime measurements from the residual layer in the southeast US, Edwards et al. (2017) calculated the transition between  $\text{NO}_3$ - and  $\text{O}_3$ -dominant regimes, which occurs at  $\text{NO}_x$ /isoprene ratio of approximately 0.5. Transition between inorganic and organic nitrate production occurs at  $\text{NO}_x$ /isoprene ratio of approximately 3. These transition points are shown in Figure 4, which also displays nighttime  $\text{NO}_x$ /VOC ratio in the simulation and in the measurements during the three studied periods. Before the lockdown, inorganic nitrate production was the main sink for  $\text{NO}_3$ . During the lockdown, nitrate production changed from inorganic- to organic-dominant, although the measurements rarely cross the line calculated by Edwards et al. (2017). In all three periods,  $\text{NO}_x$ /VOC ratio never reached the transition point to an  $\text{O}_3$ - (rather than  $\text{NO}_3$ ) dominated nighttime chemistry (Edwards et al., 2017). Consequently, decreasing  $\text{NO}_x$  as occurred during the lockdown did not dramatically affect isoprene oxidation, which was always dominated by reactions with  $\text{NO}_3$  (Figure S9 in Supporting Information S1). This also affected the maximum potential for  $\text{ClNO}_2$  production (Figure S8 in Supporting Information S1), which decreased from nighttime maximum production of 5.5 ppbv before the lockdown to less than 1 ppbv during the strict lockdown, followed by an increase to 1.5 ppbv upon lockdown relaxation. This maximum  $\text{ClNO}_2$  production changes have the capability to affect photochemistry and next-day ozone production, however, this depends on available chloride. Gunthe et al. (2021) measured chemical composition of  $\text{PM}_{10}$  in Delhi and found that during episodes with high chloride, its concentration in  $\text{PM}_{10}$  may reach  $22.1 \pm 13.7 \mu\text{g m}^{-3}$  (equivalent to  $15 \pm 9$  ppbv), whereas during the entire period, measured chloride was  $5.9 \pm 9.1 \mu\text{g m}^{-3}$  (equivalent to  $4 \pm 6$  ppbv). Gani et al. (2019) also measured the chemical composition of  $\text{PM}_{10}$  in Delhi and found that diel cycle of chloride in  $\text{PM}_{10}$  has a morning peak of up to  $50 \mu\text{g m}^{-3}$  (equivalent to 33 ppbv) during winter. Nighttime winter concentrations of chloride in  $\text{PM}_{10}$  varied between 20 and  $30 \mu\text{g m}^{-3}$  (equivalent to 13–20 ppbv). Reaching such concentrations in an inland location such as the IGP indicates there is sufficient chloride to sustain the maximum potential for



**Figure 4.** Left— $\text{NO}_x/\text{volatile organic compound (VOC)}$  ratio during nighttime from the model. Color scale shows the ratio between inorganic and organic nitrate production. Gray dots reflect the spread of measured data during the night. Right— $\text{NO}_x/\text{VOC}$  ratio histograms of measured nighttime data. Data is shown for the three periods studied: before COVID-19 lockdown (top), during strict lockdown (middle) and after lockdown relaxation (bottom). Dashed gray lines are the transition points modeled by Edwards et al. (2017) (0.5 -  $\text{NO}_3$ -to- $\text{O}_3$  dominant regime; 3 - inorganic-to-organic-nitrate-production dominant regime) as is shown in the upper right panel.

$\text{ClNO}_2$  production in the region. Therefore, calculated maximum  $\text{ClNO}_2$  production may be close to actual levels, and this potential chlorine chemistry requires further investigation in the IGP.

#### 4. Conclusions

Effects of emission reductions during COVID-19 lockdowns on atmospheric composition and air quality have been studied in many places around the world and will continue to be of scientific and public interest for years to come. Lockdown emissions reductions combined with seasonal variations provide a unique opportunity to investigate selected oxidation processes at a site in the Indo-Gangetic Plain, a highly polluted region, where nighttime oxidation has rarely, if ever, been investigated. Our work quantifies, for the first time in India, rates of several atmospheric oxidation processes. These processes include the consumption of VOCs by  $\text{NO}_3$  radicals, and potential production of  $\text{ClNO}_2$ . These processes were affected differently by changes related to the transition from

winter to spring and to emission changes due to COVID-19 pandemic. Nighttime chemistry was highly influenced by reduction in emissions related to COVID-19 restrictions. During the lockdown, nighttime  $\text{NO}_3$  average production rates and  $\text{ClNO}_2$  production decreased by 50% and 80%, respectively, accompanied by a clear shift in nitrate production from an inorganic- to an organic-nitrate-production-dominant regime.

Extreme pollution levels associated with the IGP lead to increased rates of unconventional atmospheric chemistry processes, such as those investigated here. This manuscript identifies several processes that have not been studied in this region previously and were likely modified by pollutant emission changes during COVID-19 lockdowns. Future work to quantify these processes, their contribution to particulate matter and ozone chemistry in the IGP and their dependence on emissions will be of substantial interest to the air quality scientific and policy communities.

## Data Availability Statement

The observational data used for this study is available online at <https://data.mendeley.com/datasets/svfv9nwc43/draft?a=5ba5e802-386f-410b-b46f-8107d67f6158>.

## References

- Baasandorj, M., Hoch, S. W., Bares, R., Lin, J. C., Brown, S. S., Millet, D. B., et al. (2017). Coupling between chemical and meteorological processes under persistent cold-air pool conditions: Evolution of wintertime  $\text{PM}_{2.5}$  pollution events and  $\text{N}_2\text{O}_5$  observations in Utah's Salt Lake Valley. *Environmental Science & Technology*, 51(11), 5941–5950. <https://doi.org/10.1021/acs.est.6b06603>
- Baker, A. K., Sauvage, C., Thorenz, U. R., van Velthoven, P., Oram, D. E., Zahn, A., et al. (2016). Evidence for strong, widespread chlorine radical chemistry associated with pollution outflow from continental Asia. *Scientific Reports*, 6(1), 36821. <https://doi.org/10.1038/srep36821>
- Bertram, T. H., & Thornton, J. A. (2009). Toward a general parameterization of  $\text{N}_2\text{O}_5$  reactivity on aqueous particles: The competing effects of particle liquid water, nitrate and chloride. *Atmospheric Chemistry and Physics*, 9(21), 8351–8363. <https://doi.org/10.5194/acp-9-8351-2009>
- Brown, S. S., An, H., Lee, M., Park, J.-H., Lee, S.-D., Fibiger, D. L., et al. (2017). Cavity enhanced spectroscopy for measurement of nitrogen oxides in the Anthropocene: Results from the Seoul tower during MAPS 2015. *Faraday Discussions*, 200(0), 529–557. <https://doi.org/10.1039/c7fd00001d>
- Brown, S. S., Stark, H., & Ravishankara, A. R. (2003). Applicability of the steady state approximation to the interpretation of atmospheric observations of  $\text{NO}_3$  and  $\text{N}_2\text{O}_5$ . *Journal of Geophysical Research*, 108(D17), 4539. <https://doi.org/10.1029/2003jd003407>
- Brown, S. S., & Stutz, J. (2012). Nighttime radical observations and chemistry. *Chemical Society Reviews*, 41(19), 6405–6447. <https://doi.org/10.1039/c2cs35181a>
- Chandra, B. P., & Sinha, V. (2016). Contribution of post-harvest agricultural paddy residue fires in the N.W. Indo-Gangetic Plain to ambient carcinogenic benzenoids, toxic isocyanic acid and carbon monoxide. *Environment International*, 88, 187–197. <https://doi.org/10.1016/j.envint.2015.12.025>
- Chang, W. L., Bhavsar, P. V., Brown, S. S., Riemer, N., Stutz, J., & Dabdub, D. (2011). Heterogeneous atmospheric chemistry, ambient measurements, and model calculations of  $\text{N}_2\text{O}_5$ : A review. *Aerosol Science and Technology*, 45(6), 665–695. <https://doi.org/10.1080/02786826.2010.551672>
- Chen, Q., Xia, M., Peng, X., Yu, C., Sun, P., Li, Y., et al. (2022). Large daytime molecular chlorine missing source at a suburban site in East China. *Journal of Geophysical Research: Atmospheres*, 127(4), e2021JD035796. <https://doi.org/10.1029/2021JD035796>
- Chen, Y., Beig, G., Archer-Nicholls, S., Drysdale, W., Acton, W. J. F., Lowe, D., et al. (2021). *Avoiding high ozone pollution in Delhi, India*. Faraday Discussions.
- Council, N. R. (1991). *Rethinking the ozone problem in urban and regional air pollution* (p. 524). The National Academies Press.
- David, L. M., Ravishankara, A. R., Kodros, J. K., Pierce, J. R., Venkataraman, C., & Sadavarte, P. (2019). Premature mortality due to  $\text{PM}_{2.5}$  over India: Effect of atmospheric transport and anthropogenic emissions. *GeoHealth*, 3(1), 2–10. <https://doi.org/10.1029/2018gh000169>
- Edwards, P. M., Aikin, K. C., Dube, W. P., Fry, J. L., Gilman, J. B., De Gouw, J. A., et al. (2017). Transition from high- to low- $\text{NO}_x$  control of night-time oxidation in the southeastern US. *Nature Geoscience*, 10, 490. <https://doi.org/10.1038/ngeo2976>
- Forster, P. M., Forster, H. I., Evans, M. J., Gidden, M. J., Jones, C. D., Keller, C. A., et al. (2020). Current and future global climate impacts resulting from COVID-19. *Nature Climate Change*, 10(10), 913–919. <https://doi.org/10.1038/s41558-020-0883-0>
- Fuchs, H., Tan, Z., Lu, K., Bohn, B., Broch, S., Brown, S. S., et al. (2017). OH reactivity at a rural site (Wangdu) in the north China plain: Contributions from OH reactants and experimental OH budget. *Atmospheric Chemistry and Physics*, 17(1), 645–661. <https://doi.org/10.5194/acp-17-645-2017>
- Gani, S., Bhandari, S., Seraj, S., Wang, D. S., Patel, K., Soni, P., et al. (2019). Submicron aerosol composition in the world's most polluted megacity: The Delhi Aerosol Supersite study. *Atmospheric Chemistry and Physics*, 19(10), 6843–6859. <https://doi.org/10.5194/acp-19-6843-2019>
- Gaudel, A., Cooper, O. R., Ancellet, G., Barret, B., Boynard, A., Burrows, J. P., et al. (2018). Tropospheric Ozone Assessment Report: Present-day distribution and trends of tropospheric ozone relevant to climate and global atmospheric chemistry model evaluation. *Elementa: Science of the Anthropocene*, 6, 39. <https://doi.org/10.1525/elementa.291>
- Ghude, S. D., Chate, D. M., Jena, C., Beig, G., Kumar, R., Barth, M. C., et al. (2016). Premature mortality in India due to  $\text{PM}_{2.5}$  and ozone exposure. *Geophysical Research Letters*, 43(9), 4650–4658. <https://doi.org/10.1002/2016gl068949>
- Gunthe, S. S., Liu, P., Panda, U., Raj, S. S., Sharma, A., Darbyshire, E., et al. (2021). Enhanced aerosol particle growth sustained by high continental chlorine emission in India. *Nature Geoscience*, 14(2), 77–84. <https://doi.org/10.1038/s41561-020-00677-x>
- Haskins, J. D., Lopez-Hilfiker, F. D., Lee, B. H., Shah, V., Wolfe, G. M., DiGangi, J., et al. (2019). Anthropogenic control over wintertime oxidation of atmospheric pollutants. *Geophysical Research Letters*, 46(24), 14826–14835. <https://doi.org/10.1029/2019GL085498>
- Huang, W., Yang, Y., Wang, Y., Gao, W., Li, H., Zhang, Y., et al. (2021). Exploring the inorganic and organic nitrate aerosol formation regimes at a suburban site on the North China Plain. *The Science of the Total Environment*, 768, 144538. <https://doi.org/10.1016/j.scitotenv.2020.144538>



- Jenkin, M. E., Young, J. C., & Rickard, A. R. (2015). The MCM v3.3.1 degradation scheme for isoprene. *Atmospheric Chemistry and Physics*, 15(20), 11433–11459. <https://doi.org/10.5194/acp-15-11433-2015>
- Jeong, D., Seco, R., Gu, D., Lee, Y., Nault, B. A., Knote, C. J., et al. (2019). Integration of airborne and ground observations of nitryl chloride in the Seoul metropolitan area and the implications on regional oxidation capacity during KORUS-AQ 2016. *Atmospheric Chemistry and Physics*, 19(19), 12779–12795. <https://doi.org/10.5194/acp-19-12779-2019>
- Kumar, R., Barth, M. C., Madronich, S., Naja, M., Carmichael, G. R., Pfister, G. G., et al. (2014). Effects of dust aerosols on tropospheric chemistry during a typical pre-monsoon season dust storm in northern India. *Atmospheric Chemistry and Physics*, 14(13), 6813–6834. <https://doi.org/10.5194/acp-14-6813-2014>
- Kumar, V., Beirle, S., Dörner, S., Mishra, A. K., Donner, S., Wang, Y., et al. (2020). Long-term MAX-DOAS measurements of NO<sub>2</sub>, HCHO, and aerosols and evaluation of corresponding satellite data products over Mohali in the Indo-Gangetic Plain. *Atmospheric Chemistry and Physics*, 20(22), 14183–14235. <https://doi.org/10.5194/acp-20-14183-2020>
- Kumar, V., Sarkar, C., & Sinha, V. (2016). Influence of post-harvest crop residue fires on surface ozone mixing ratios in the N.W. IGP analyzed using 2 years of continuous in situ trace gas measurements. *Journal of Geophysical Research: Atmospheres*, 121(7), 3619–3633. <https://doi.org/10.1002/2015jd024308>
- Kumar, V., & Sinha, V. (2021). Season-wise analyses of VOCs, hydroxyl radicals and ozone formation chemistry over north-west India reveal isoprene and acetaldehyde as the most potent ozone precursors throughout the year. *Chemosphere*, 131184. <https://doi.org/10.1016/j.chemosphere.2021.131184>
- Lefohn, A. S., Malley, C. S., Smith, L., Wells, B., Hazucha, M., Simon, H., et al. (2018). Tropospheric ozone assessment report: Global ozone metrics for climate change, human health, and crop/ecosystem research. *Elementa: Science of the Anthropocene*, 6, 27. <https://doi.org/10.1525/elementa.279>
- Lu, X., Zhang, L., Wang, X., Gao, M., Li, K., Zhang, Y., et al. (2020). Rapid increases in warm-season surface ozone and resulting health impact in China since 2013. *Environmental Science and Technology Letters*, 7(4), 240–247. <https://doi.org/10.1021/acs.estlett.0c00171>
- McDuffie, E. E., Fibiger, D. L., Dubé, W. P., Lopez-Hilfiker, F., Lee, B. H., Thornton, J. A., et al. (2018). Heterogeneous N<sub>2</sub>O<sub>5</sub> uptake during winter: Aircraft measurements during the 2015 WINTER campaign and critical evaluation of current parameterizations. *Journal of Geophysical Research: Atmospheres*, 123(8), 4345–4372. <https://doi.org/10.1002/2018jd028336>
- Mohan, S., & Saranya, P. (2019). Assessment of tropospheric ozone at an industrial site of Chennai megacity. *Journal of the Air & Waste Management Association*, 69(9), 1079–1095. <https://doi.org/10.1080/10962247.2019.1604451>
- Ojha, N., Naja, M., Singh, K. P., Sarangi, T., Kumar, R., Lal, S., et al. (2012). Variabilities in ozone at a semi-urban site in the Indo-Gangetic Plain region: Association with the meteorology and regional processes. *Journal of Geophysical Research*, 117(D20), D20301. <https://doi.org/10.1029/2012jd017716>
- Osthoff, H. D., Roberts, J. M., Ravishankara, A. R., Williams, E. J., Lerner, B. M., Sommariva, R., et al. (2008). High levels of nitryl chloride in the polluted subtropical marine boundary layer. *Nature Geoscience*, 1(5), 324–328. <https://doi.org/10.1038/ngeo177>
- Pawar, H., Garg, S., Kumar, V., Sachan, H., Arya, R., Sarkar, C., et al. (2015). Quantifying the contribution of long-range transport to particulate matter (PM) mass loadings at a suburban site in the North-Western Indo-Gangetic Plain (NW-IGP). *Atmospheric Chemistry and Physics*, 15(16), 9501–9520. <https://doi.org/10.5194/acp-15-9501-2015>
- Pawar, H., & Sinha, B. (2020). Humidity, density, and inlet aspiration efficiency correction improve accuracy of a low-cost sensor during field calibration at a suburban site in the North-Western Indo-Gangetic plain (NW-IGP). *Aerosol Science and Technology*, 54(6), 685–703. <https://doi.org/10.1080/02786826.2020.1719971>
- Petit, J. E., Dupont, J. C., Favez, O., Gros, V., Zhang, Y., Sciare, J., et al. (2021). Response of atmospheric composition to COVID-19 lockdown measures during spring in the Paris region (France). *Atmospheric Chemistry and Physics Discussions*, 2021, 1–24. <https://doi.org/10.5194/acp-21-17167-2021>
- Ravishankara, A. R., David, L. M., Pierce, J. R., & Venkataraman, C. (2020). Outdoor air pollution in India is not only an urban problem. *Proceedings of the National Academy of Sciences*, 117(46), 28640–28644. <https://doi.org/10.1073/pnas.2007236117>
- Roberts, J. M., Osthoff, H. D., Brown, S. S., Ravishankara, A. R., Coffman, D., Quinn, P., & Bates, T. (2009). Laboratory studies of products of N<sub>2</sub>O<sub>5</sub> uptake on Cl<sup>-</sup> containing substrates. *Geophysical Research Letters*, 36(20), L20808. <https://doi.org/10.1029/2009gl040448>
- Romer Present, P. S., Zare, A., & Cohen, R. C. (2020). The changing role of organic nitrates in the removal and transport of NOx. *Atmospheric Chemistry and Physics*, 20(1), 267–279. <https://doi.org/10.5194/acp-20-267-2020>
- Saiz-Lopez, A., & von Glasow, R. (2012). Reactive halogen chemistry in the troposphere. *Chemical Society Reviews*, 41(19), 6448–6472. <https://doi.org/10.1039/c2cs35208g>
- Seinfeld, J. H. (1989). Urban air pollution: State of the science. *Science*, 243(4892), 745–752. <https://doi.org/10.1126/science.243.4892.745>
- Sharma, A., Ojha, N., Pozzer, A., Beig, G., & Gunthe, S. S. (2019). Revisiting the crop yield loss in India attributable to ozone. *Atmospheric Environment X*, 1, 100008. <https://doi.org/10.1016/j.aeoa.2019.100008>
- Singh, V., Singh, S., Biswal, A., Kesarkar, A. P., Mor, S., & Ravindra, K. (2020). Diurnal and temporal changes in air pollution during COVID-19 strict lockdown over different regions of India. *Environmental Pollution*, 266, 115368. <https://doi.org/10.1016/j.envpol.2020.115368>
- Sinha, B., Singh Sangwan, K., Maurya, Y., Kumar, V., Sarkar, C., Chandra, B. P., & Sinha, V. (2015). Assessment of crop yield losses in Punjab and Haryana using 2 years of continuous in situ ozone measurements. *Atmospheric Chemistry and Physics*, 15(16), 9555–9576. <https://doi.org/10.5194/acp-15-9555-2015>
- Sinha, V., Kumar, V., & Sarkar, C. (2014). Chemical composition of pre-monsoon air in the Indo-Gangetic Plain measured using a new air quality facility and PTR-MS: High surface ozone and strong influence of biomass burning. *Atmospheric Chemistry and Physics*, 14(12), 5921–5941. <https://doi.org/10.5194/acp-14-5921-2014>
- Volkamer, R., Sheehy, P., Molina, L. T., & Molina, M. J. (2010). Oxidative capacity of the Mexico City atmosphere—Part I: A radical source perspective. *Atmospheric Chemistry and Physics*, 10(14), 6969–6991. <https://doi.org/10.5194/acp-10-6969-2010>
- Wang, X., Jacob, D. J., Eastham, S. D., Sulprizio, M. P., Zhu, L., Chen, Q., et al. (2019). The role of chlorine in global tropospheric chemistry. *Atmospheric Chemistry and Physics*, 19(6), 3981–4003. <https://doi.org/10.5194/acp-19-3981-2019>
- Wang, X., Jacob, D. J., Fu, X., Wang, T., Breton, M. L., Hallquist, M., et al. (2020). Effects of anthropogenic chlorine on PM<sub>2.5</sub> and ozone air quality in China. *Environmental Science & Technology*, 54(16), 9908–9916. <https://doi.org/10.1021/acs.est.0c02296>
- Williams, J., Keßel, S. U., Nölscher, A. C., Yang, Y., Lee, Y., Yáñez-Serrano, A. M., et al. (2016). Opposite OH reactivity and ozone cycles in the Amazon rainforest and megacity Beijing: Subversion of biospheric oxidant control by anthropogenic emissions. *Atmospheric Environment*, 125(1), 112–118. <https://doi.org/10.1016/j.atmosenv.2015.11.007>
- Wolfe, G. M., Marvin, M. R., Roberts, S. J., Travis, K. R., & Liao, J. (2016). The Framework for 0-D atmospheric modeling (F0AM) v3.1. *Geoscientific Model Development*, 9(9), 3309–3319. <https://doi.org/10.5194/gmd-9-3309-2016>
- World Health, O. (2016). *Ambient air pollution: A global assessment of exposure and burden of disease*. World Health Organization.

- Xia, M., Peng, X., Wang, W., Yu, C., Wang, Z., Tham, Y. J., et al. (2021). Winter ClNO<sub>2</sub> formation in the region of fresh anthropogenic emissions: Seasonal variability and insights into daytime peaks in northern China. *Atmospheric Chemistry and Physics*, *21*(20), 15985–16000. <https://doi.org/10.5194/acp-21-15985-2021>
- Yadav, R., Sahu, L. K., Jaaffrey, S. N. A., & Beig, G. (2014). Distributions of ozone and related trace gases at an urban site in Western India. *Journal of Atmospheric Chemistry*, *71*(2), 125–144. <https://doi.org/10.1007/s10874-014-9286-9>
- Young, C. J., Washenfelder, R. A., Roberts, J. M., Mielke, L. H., Osthoff, H. D., Tsai, C., et al. (2012). Vertically resolved measurements of nighttime radical reservoirs in Los Angeles and their contribution to the urban radical budget. *Environmental Science & Technology*, *46*(20), 10965–10973. <https://doi.org/10.1021/es302206a>

### References From the Supporting Information

- Brown, S. S., An, H., Lee, M., Park, J.-H., Lee, S.-D., Fibiger, D. L., et al. (2017). Cavity enhanced spectroscopy for measurement of nitrogen oxides in the Anthropocene: Results from the Seoul tower during MAPS 2015. *Faraday Discussions*, *200*(0), 529–559. <https://doi.org/10.1039/C7FD00001D>
- McDuffie, E. E., Fibiger, D. L., Dubé, W. P., Lopez-Hilfiker, F., Lee, B. H., Thornton, J. A., et al. (2018). Heterogeneous N<sub>2</sub>O<sub>3</sub> uptake during winter: Aircraft measurements during the 2015 WINTER campaign and critical evaluation of current parameterizations. *Journal of Geophysical Research: Atmospheres*, *123*(8), 4345–4372. <https://doi.org/10.1002/2018JD028336>
- Yu, C., Wang, Z., Xia, M., Fu, X., Wang, W., Tham, Y. J., et al. (2020). Heterogeneous N<sub>2</sub>O<sub>3</sub> reactions on atmospheric aerosols at four Chinese sites: Improving model representation of uptake parameters. *Atmospheric Chemistry and Physics*, *20*(7), 4367–4378. <https://doi.org/10.5194/acp-20-4367-2020>