

International Journal of *Environmental Research and Public Health*



Article Photochemical Generation of Methyl Chloride from Humic Aicd: Impacts of Precursor Concentration, Solution pH, Solution Salinity and Ferric Ion

Hui Liu *, Yingying Pu, Tong Tong, Xiaomei Zhu, Bing Sun and Xiaoxing Zhang

College of Environmental Science and Engineering, Dalian Maritime University, Dalian 116026, China; py1120181596@dlmu.edu.cn (Y.P.); tongtong@dlmu.edu.cn (T.T.); zhuxm@dlmu.edu.cn (X.Z.); sunb88@dlmu.edu.cn (B.S.); zhangxiaoxing@dlmu.edu.cn (X.Z.)

* Correspondence: liuhui@dlmu.edu.cn

Received: 20 November 2019; Accepted: 9 January 2020; Published: 13 January 2020



Abstract: Methyl chloride (CH₃Cl) is presently understood to arise from biotic and abiotic processes in marine systems. However, the production of CH₃Cl via photochemical processes has not been well studied. Here, we reported the production of CH₃Cl from humic acid (HA) in sunlit saline water and the effects of the concentration of HA, chloride ions, ferric ions and pH were investigated. HA in aqueous chloride solutions or natural seawater were irradiated under an artificial light, and the amounts of CH₃Cl were determined using a purge-and-trap and gas chromatography-mass spectrometry. CH₃Cl was generated upon irradiation and its amount increased with increasing irradiation time and the light intensity. The formation of CH₃Cl increased with an increase of HA concentration ranging from 2 mg L⁻¹ to 20 mg L⁻¹ and chloride ion concentration ranging from 0.02 mol L^{-1} to 0.5 mol L^{-1} . The photochemical production of CH₃Cl was pH-dependent, with the highest amount of CH₃Cl generating near neutral conditions. Additionally, the generation of CH₃Cl was inhibited by ferric ions. Finally, natural coastal seawater was irradiated under artificial light and the concentration of CH₃Cl rose significantly. Our results suggest that the photochemical process of HA may be a source of CH₃Cl in the marine environment.

Keywords: methyl chloride; humic acid; photochemical production; ferric ion

1. Introduction

Methyl chloride, CH₃Cl, is particularly important in the global atmosphere as a major natural source of chlorine and plays a significant role in atmospheric chemical processes such as stratospheric ozone depletion [1]. Besides, as an absorber of infrared radiation, CH₃Cl is of interest for its potential effect on the tropospheric energy balance [2]. The Global Warming Potential (GWP) of CH₃Cl is estimated to be about eight (100-year time integration, $CO_2 = 1$), indicating that CH₃Cl has a direct GWP similar to that of CH₄ [3]. The production of CH₃Cl is dominated by natural sources, but smaller, important anthropogenic sources, such as agricultural fumigation and/or biomass burning, also exist [4]. Natural emissions of CH₃Cl include oceanic sources, terrestrial plants, fungi and biomass burning, etc. [5]. Among these, the ocean has been implicated as one of the main natural sources, where inorganic halogen ions in seawater or its aerosol can transform into organic halogen through biotic or abiotic pathways. There are several reports that organisms such as macroalgae, microalgae and bacteria can release large quantities of halogen hydrocarbons into the atmosphere [6]. However, few studies have been focused on the abiotic sources of halogen hydrocarbons in the ocean.

Photochemical processes occur universally within surface seawater and are considered to be an important abiotic pathway for the formation of organic halogens [7]. There are a number of reports

on the photochemical production of methyl halides (chiefly methyl iodide) in seawater [8,9]. It is supposed that a component of light absorbing molecules, such as marine colored dissolved organic material (CDOM), takes part in the formation of methyl halides. Dissolved organic matter (DOM) in estuaries and coastal areas is a class of natural organic compounds containing molecules derived from biota in terrestrial (allochthonous) and/or marine (autochthonous) ecosystems [10], which contains a large assemblage of complex chemical structures of polyphenols, carboxyls, methoxyls, quinones, carbohydrates and peptide functionalities [11]. As a complex macromolecular mixture, DOM can absorb light energy and induce photochemical reactions. DOM is known to act as the sensitizer or the precursor for various reactive species in aquatic ecosystems, such as hydroxyl radicals (\cdot OH), singlet oxygen ($^{1}O_{2}$) and the excited triplet state of DOM ($^{3}DOM^{*}$), which then participate in the photochemical transformation of organic and inorganic compounds in water [10,12,13].

It is notable that DOM plays an important role in the formation of organic halogens from halide ions and non-halogen compounds. The excited state of DOM, forming upon sunlight irradiation (especially ³DOM*), can oxidize halide ions to generate reactive halogen species, such as $X_2^{\bullet-}/X^{\bullet}$, X_2 and HXO (X = Cl, Br or I), which are the major halogenation reagents [13,14]. In addition, DOM itself can be halogenated, forming non-volatile and volatile org-halogens, i.e., DOM acts as the precursor of organic halogens [15,16]. Research on the photochemical formation of volatile org-halogens in saline water have been concentrated on methyl halogens. Moore [9,17] and Yang et al. [18] have reported that lignin model compounds, such as syringic acid and guaiacol, could act as the carbon source to produce CH₃Cl in saline waters through a photochemical pathway. However, little study was focused on the formation of CH₃Cl from natural DOM. Natural DOM is a complex mixture containing humic acid-like and fulvic acid-like materials, and soluble microbial byproduct-like materials; DOM in coastal and estuarine environments is even more versatile [19]. Therefore, the generation of CH₃Cl from its natural complex precursor in the variable water environment needs further investigation.

In estuarine and coastal waters, due to the mixing of river water (which is rich in terrestrial DOM) and seawater (which has high salinity), sharp increases in DOM and chloride ions occurs simultaneously, which offers a great potential for the production of organic halogens [20]. Considering that humic acid (HA) is a typical terrestrial DOM in water, and HA can stabilize metal ions, e.g., iron, which plays an active role in the photochemical reactions [21], HA was selected as the model DOM in this work. The primary objective of this work was to identify the formation of CH₃Cl from HA in saline solutions under simulated sunlight irradiation, and the influence of some environmental factors on CH₃Cl generation, including chloride ion concentration, sunlight intensity, pH and ferric ion (Fe (III)). Furthermore, the production of CH₃Cl from natural seawater was investigated as well.

2. Materials and Methods

2.1. Reagents

HA was supplied by MP Biomedicals, Inc (Eschwege, Germany). Liquid standards of CH₃Cl (200 μ g mL⁻¹ in methanol) were purchased from AccuStandard (USA). Sodium chloride (NaCl), FeCl₃·6H₂O and other chemicals were reagent grade. Ultrapure water (18 M Ω cm) was obtained from a Millipore water purification unit for preparing all aqueous solutions. Natural coastal seawater was collected from the Northern Yellow Sea near Blackrock Reef (38°55′ N, 121°37′ E), Dalian, China.

2.2. Irradiation Experiments

The reactor used in the experiment was a round bottom sealed quartz tube (3.5 cm o.d.; volume approximately 125 mL) with one outlet (4 mm o.d.) at the middle of the bottom of the tube, where the reaction solution was injected into the reactor by a syringe. The quartz tube reactors were checked for air tightness before irradiation experiments. All quartz tubes were irradiated in a solar simulator (Phchem III, Beijing Newbit Technology Co., Ltd., Beijing, China) containing a 500 W xenon arc lamp and filters to cut off the light with a wavelength below 290 nm. The light intensity of the solar simulator

can be adjusted from 3 mW cm⁻² to 15 mW cm⁻². Details are in Figure S1 in the Supplementary Materials. The aqueous solutions of HA (2–20 mg L⁻¹) and chloride (0.02–0.5 mol L⁻¹) were held in quartz tubes and irradiated for periods of 20 h. The quartz tubes were cooled by a fan during irradiation; the dark controls in tubes wrapped in Al foil were also placed in the solar simulator. All experiments were run in replicate. The light intensity effect was first studied at 3.5 mW cm⁻² and 12 mW cm⁻², separately. Light intensity was set at 12 mW cm⁻² for all other experiments. Then, the concentration of HA, Cl⁻, ferric ions and pH on the production of CH₃Cl were separately investigated. Finally, the generation of CH₃Cl in natural seawater was measured.

2.3. Analysis Methods

The concentration of CH_3Cl was analyzed using a gas chromatography mass spectrometer (GC-MS, Agilent 7890B/5977C, Agilent Technologies, Santa Clara, CA, USA) equipped with a purge-and-trap sample concentrator (Eclipse 4760, OI Analytical, College Station, TX, USA). After irradiation, aqueous samples were withdrawn from the tubes by a glass syringe via a long needle, and then injected into a 25 mL purge tube. Before purging, the tube of each sample was thermostated at 40 °C. After that, the sample was purged with ultrapure nitrogen at 40 mL min⁻¹ for 11 min to facilitate the reproducible extraction of trace gases. Extracted gases were pre-concentrated in a collecting trap of a stainless-steel tube containing Vocarb-3000. CH₃Cl was released from the trap column by heating the trap to 240 $^{\circ}$ C for 1 min, and then the gas was introduced into the GC-MS. After each capture and desorption procedure, the trap column was baked at 250 °C for 10 min to avoid interference between samples. The temperature of the transmission line between the purge-and-trap concentrator and the GC-MS was kept at 200 °C. An Agilent J&W DB-VRX capillary column (length, 60 m; inner diameter, 250 μm; film thickness, 1.4 µm; Agilent Technologies, Palo Alto, CA, USA) was used for the analysis of CH₃Cl. The carrier gas was helium with a flow rate of 1 mL min⁻¹. The inlet worked in a split mode with a split ratio of 20:1 and the temperature was set at 150 °C. The oven temperature was initially set at 32 °C for 6 min and rose to 180 °C at 20 °C min⁻¹ then held at 180 °C for 4 min. An illustration of the purge-and-trap device and the GC-MS is shown in Figure S2. CH₃Cl was quantitatively analyzed by direct external calibration. The selected ion monitor (SIM) mode was used with an m/z of 50 and 52. The detection limit (signal to noise ratio, S/N = 3) of CH₃Cl was 3.2 ng L⁻¹, and the relative standard deviation of replicate analyses (n = 6) of CH₃Cl was within 6%. For qualitative analysis of other volatile gases from HA, scan mode was used with an m/z from 35 to 350, and the structure was identified by comparison with the National Institute of Standards and Technology (NIST) 2014 Library.

2.4. Characteristics of HA and HA-Fe (III)

The elemental composition of HA was analyzed on an element analyzer (Vario EL III, Elementar Analysensysteme GmbH, Langenselbold, Germany). FTIR spectra of HA and the HA-Fe (III) complex were recorded with an FTIR spectrometer (Nicolet iS 5, Thermo Fisher Scientific, Madison, WI, USA).

3. Results and Discussion

3.1. Influence of Light Intensity

It has been found that the formation of methyl iodide under light irradiation was more obvious than its formation in the dark [22], showing that sunlight irradiation is crucial for the production of methyl halide in seawater. The influence of light intensity on the generation of CH₃Cl was first investigated in the solutions containing HA and Cl⁻. DOM as a part of the total organic carbon (TOC) in the open ocean is extremely low, whereas it is much higher in estuaries which can reach dozens of mgC L⁻¹. The elementary composition for HA used in the work is 40.8% C, 3.0% H, 50.6% O and 0.9% N. HA at a concentration of 5 mg L⁻¹ with the corresponding TOC at 2.04 mgC L⁻¹ was first adopted in this experiment. The sunlight intensity of the water surface is about 85 mW cm⁻² at noon on a sunshine day in Dalian, China. Considering the severe attenuation of light in water and the changes of

the solar altitude angle, the average light intensity in the euphotic zone of the water body should be much lower. Thus, the light intensities were set at 3.5 mW cm^{-2} and 12 mW cm^{-2} in this work. It can be seen that the generation of CH₃Cl was significantly increased with increased light intensity and was not detectable in the dark within 20 h (Figure 1). This result is similar to the finding of Moore that CH₃Cl from syringic acid was only produced under illumination [9].



Figure 1. Effect of light intensity on the production of methyl chloride. Reaction conditions: $[Cl^-] = 0.5 \text{ mol } L^{-1}$, $[HA] = 5 \text{ mg } L^{-1}$, pH = 6.5.

3.2. Influence of HA Concentration

Many reports have confirmed the role of DOM in controlling halogen hydrocarbon production in seawater. Manley and Barbero demonstrated that the removal of DOM from natural seawater decreased CHBr₃ production from *Ulva lactuca* sp. [23]. Lin and Manley found a significant correlation between DOM molecular size and CHBr₃ production from seawater treated with bromoperoxidase [24]. The following experiments were performed to study the effect of HA concentration on the production of CH₃Cl through photochemical processes. The concentration of HA, ranging from 2 mg L⁻¹ to 20 mg L⁻¹, almost covers the DOC concentration in coastal and estuarine water. It can be seen from Figure 2A that the amount of CH₃Cl increased with increasing HA concentration. When the HA concentrations were 2 mg L⁻¹, 5 mg L⁻¹, 10 mg L⁻¹ and 20 mg L⁻¹, the maximums of CH₃Cl were 43.5 ng L⁻¹, 49.2 ng L⁻¹, 56.5 ng L⁻¹ and 68.3 ng L⁻¹, respectively. However, the formation of CH₃Cl in the presence of 20 mg L⁻¹ HA was not obviously higher than the lower HA concentrations within 10 h of irradiation. The reason for this may be related to the light shielding effect of HA, which can reduce the absorbance and the effective utilization of photons.



Figure 2. (A) Effect of HA concentration on the production of methyl chloride and (B) after corrected for light shielding effect using Equations (1)–(3). Reaction conditions: light intensity = 12 mW cm^{-2} . [Cl⁻] = 0.5 mol L⁻¹, pH = 6.5.

In order to estimate the light shielding effect of HA on CH_3Cl production, the light shielding correction factor of HA (CF_{HA}) is calculated according to the following equations [25,26]:

$$CF_{HA} = \frac{\sum_{\lambda} I_{0,\lambda} a_{\lambda}}{\sum_{\lambda} I_{0,\lambda} a_{\lambda} S_{\lambda}}$$
(1)

$$S_{\lambda} = \frac{1 - 10^{(-a_{\lambda})z}}{2.303a_{\lambda}z}$$
(2)

where $I_{0,\lambda}$ is the initial irradiation intensity at wavelength λ as displayed in Figure S3; S_{λ} , the light shielding factor, is determined by Equation (2) where z, and a_{λ} (cm⁻¹) refer to light path length and light attenuation factors of HA at wavelength λ , respectively. In this study, CF_{HA} was calculated for the wavelength range of 300–600 nm, wherein HA exhibits considerable absorbance (Figure S3). The CF_{HA} is 1.04, 1.10, 1.18 and 1.36 for 2 mg L⁻¹, 5 mg L⁻¹, 10 mg L⁻¹ and 20 mg L⁻¹ HA, respectively. The concentration of CH₃Cl generated after corrected for light shielding, C_{CH3Cl}, is calculated according to Equation (3).

$$C_{CH3Cl'} = C_{CH3Cl} \times CF_{HA}$$
(3)

The production profile of CH_3Cl after being corrected for light shielding is shown in Figure 2B. The production rates of CH_3Cl in the presence of different concentrations of HA at 10 h were estimated and displayed in Figure 3. It can be seen that the rate constant (k) of CH_3Cl formation after it was corrected for light shielding, k_{CH3Cl} , was almost linear with HA concentration, which indicated that HA acted as the precursor of CH_3Cl and was the limiting factor under this experimental condition.



Figure 3. Effect of concentration of HA on the rate constant (k) of CH₃Cl formation after correction for the light shielding effect. Straight line is the trend line of linear fitting of k_{CH3Cl} and HA concentration. Reaction conditions: [Cl⁻] = 0.5 mol L⁻¹, t = 10 h, pH = 6.5.

However, CH_3Cl production was slow after long term irradiation (>10 h). The reason for this is most likely related to the slow formation of the active precursor of CH_3Cl after long term irradiation. Moore has demonstrated that the photochemical production of CH_3Cl from syringic acid was not directly from the acid itself, but a product of its photolysis (most likely a quinone derivative) [9]. Unfortunately, the nature of this intermediate is unknown. The structure and composition of HA are more complex than that of syringic acid, and thus the situation in this experiment will be more complicated. We proposed that the postulated intermediate that was liable for CH_3Cl production might be generated fast in the first period of irradiation, and then its generation became slow after long term irradiation. Additional work is needed to elucidate the nature of the active intermediate of CH_3Cl production from HA.

3.3. Influence of Chloride

As a variable element in the estuarine area, the concentration of chloride ions in water often affects the results of photochemical reactions [27]. The salinity of sea water around the world is variable with an average value of 34.7, and the salinity range of estuaries is extremely large depending on the mixing of seawater and river water, which is influenced by many factors such as evaporation, precipitation, river runoff and seawater movement. These experiments were carried out with the concentration of Cl⁻ ranging from 0.02 mol L⁻¹ to 0.5 mol L⁻¹ with the corresponding salinity of 1.3‰ to 32.4‰, which can be applied to mimic the salinity changes in estuarine and coastal regions. When Cl⁻ concentrations were 0.02 mol L⁻¹, 0.2 mol L⁻¹ and 0.5 mol L⁻¹, the concentrations of CH₃Cl after 20 h of irradiation were 24.7 ng L⁻¹, 29.5 ng L⁻¹ and 49.2 ng L⁻¹, respectively (Figure 4), indicating that Cl⁻ was the other precursor of CH₃Cl. Estuarine and coastal waters are rich in terrestrial HA and chloride ions; thus, it can be speculated that the photochemical generation of CH₃Cl could occur in estuaries and coastal areas. However, promoting CH₃Cl production by increasing chloride was not as obvious as increasing HA concentration, especially in the first 5 h. We proposed that the limiting factor for CH₃Cl production in our reaction systems is HA, not chloride. Thus, there was no significant increase in CH_3Cl in the first 5 h of irradiation in the presence of different concentrations of Cl^- .



Figure 4. Effect of chloride ion concentration on the production of methyl chloride. Reaction conditions: Light intensity = 12 mW cm⁻², [HA] = 5 mg L⁻¹, pH = 6.5.

3.4. Influence of pH

HA comprises multiple aromatic and aliphatic hydrocarbon structures, and its configuration is often distinct at different pH levels [28,29]. The pH of the solutions containing 5 mg L^{-1} HA and $0.5 \text{ mol } L^{-1} \text{ Cl}^-$ was adjusted using hydrochloric acid or sodium hydroxide from 4.0 ± 0.1 to 10.0 ± 0.1 . Figure 5 shows that the concentration of CH_3Cl at pH 4.0 (26.6 ng L^{-1}) was significantly lower than at pH 6.5 (around 45 ng L^{-1}). As pH increased from 6.5 to 10.0, CH₃Cl formation decreased. The order of CH₃Cl concentration was pH 6.5 > pH 8.0 > pH 10 > pH 4.0, indicating that the most suitable conditions for CH₃Cl generation from HA were near neutral conditions. According to Moore [9,17] and Yang et al. [18], the methoxy group in simple lignin-like molecules can be the source of the methyl group in CH₃Cl produced by a photochemical reaction. HA is a complex mixture partially deriving from the lignin structures of higher terrestrial plants [11] and consequently contains syringyl and guaiacyl lignin groups. The FTIR spectrum of HA (Figure 6) shows the following characteristics: -OH stretching vibration absorption (v_{O-H}) at 3700–2700 cm⁻¹, C-H stretching vibration absorption of aliphatic series (v_{C-H}) at 2923/2851 cm⁻¹, carbonyl C=O absorption ($v_{C=O}$) around 1563 cm⁻¹, bending vibration of alcohols or carboxylic acids (δ_{O-H}) and stretching vibration of phenols (ν_{C-O}) around 1369 cm⁻¹; C-O stretching vibration of alcohols, ethers, phenol, and/or polysaccharides around 1100–1200 cm⁻¹ [30,31]. Thus, it can be seen that HA contains functional groups such as carboxyl, hydroxyl, methoxy and phenolic groups.

According the previous studies on CH₃Cl from lignin-like compounds, the proposed formation pathway was via aromatic ring protonation followed by demethoxylation [9,17,18]. We took the syringyl moiety in HA as an example. The first step was the excitation of the syringyl moiety by the absorption of photons and its subsequent protonation on the benzene ring. The second step involved a nucleophilic attack of Cl⁻ on the methoxy carbon of the syringyl moiety resulting in C_{methyl}–O cleavage and CH₃Cl releasing (Figure 7).



Figure 5. Effect of pH on the production of methyl chloride. Reaction conditions: light intensity = 12 mW cm^{-2} . [Cl⁻] =0.5 mol L⁻¹, [HA]= 5 mg L⁻¹.



Figure 6. FTIR spectra of HA and HA + Fe. The black curve represents spectrum of HA, and the red curve represents the spectrum of HA + Fe with the corresponding concentrations of 5 mg L⁻¹ HA and 100 μ mol L⁻¹ Fe(III).



Figure 7. A possible formation pathway of CH₃Cl from the syringyl moiety of HA under light irradiation.

In the pH range of 6.5 to 10, the fraction of syringyl moieties presenting in their anionic form increased with increasing pH, given that the pKa2 of syringic acid is 9.45 for the ionization of phenolic hydroxy (pKa1 of syringic acid is 4.33 for the ionization of carboxyl). When combined with the lower production of CH₃Cl at pH 10, it is suggested that the ionic form of the syringyl moiety has less photoreactivity than its molecular form when producing CH₃Cl. This result is consistent with the finding of Yang et al. who took guaiacol as the model compound to study CH₃Cl formation [18].

However, Dallin et al.'s result was quite different; they found that CH_3Cl production was inhibited in acidic conditions using syringic acid as the model compound, which demonstrated that the carboxyl form of syringic acid is less reactive than the carboxylate form in producing CH_3Cl , as the carboxyl form was hard to protonate on the benzene ring [17]. HA is an extremely complex mixture, and there may be many groups that can produce CH_3Cl besides the lignin-like moieties; thus, the situation with HA is more complicated and cannot be simply analogized by the single model compound. We proposed that the protonation on the benzene ring by hydrogen ion is the crucial procedure for CH_3Cl production. At pH 4.0, HA was positively charged, and the electrostatic repulsion of H⁺ and the positively charged HA together inhibit the protonation of HA to a certain extent, consequently reducing the generation of CH_3Cl . However, as the pH increased to 10.0, the functional group in HA was in a less photoreactive ionic form, and thus CH_3Cl production was lowered again.

3.5. Influence of Ferric Ions

In surface waters, ferric ions can complex with organic ligands such as DOM, forming Fe(III)–ligand complexes which can undergo ligand-to-metal charge transfer (LMCT) processes and subsequent Fenton reactions upon irradiation [32], leading to the generation of Fe(II) and reactive oxygen species (ROS) such as HO_2/O_2^- , H_2O_2 and $\cdot OH$ [33]. Therefore, Fe(III) often plays an important role in the fate of environmental pollutants. These experiments were carried out with different concentrations of Fe(III) ranging from 20 µmol L⁻¹ to 600 µmol L⁻¹ in the presence of 5 mg L⁻¹ HA. Considering that ferric species are liable to hydrolyze and form precipitates with increasing pH, in order to get a relatively high reactivity of Fe(III) in the solution, the pH of all the samples containing Fe(III) were adjusted to 4.0 \pm 0.1. As shown in Figure 8, the amount of CH₃Cl produced in the presence of Fe(III) was extremely low. The largest amount of CH₃Cl in the sample without Fe(III) reached 26.6 ng L⁻¹ after 20 h of irradiation, whereas it decreased to less than 7.8 ng L⁻¹ with the addition of Fe(III). Therefore, Fe(III) had a significant inhibitory effect on the formation of CH₃Cl.



Figure 8. Effect of Fe(III) on the production of methyl chloride. Reaction conditions: light intensity = 12 mW cm^{-2} . [Cl⁻] = 0.5 mol L⁻¹, [HA] = 5 mg L⁻¹, pH = 4.0.

The FTIR spectrum of HA with Fe(III) is shown in Figure 6. The major changes of the bands of HA + Fe is as follows: the band of C-O stretching vibration around 1100–1200 cm⁻¹ disappeared; the bands of asymmetric and symmetric carboxyl group red-shifted to 1598/1614 and 1404 cm⁻¹, respectively; the band near 1600 cm⁻¹ indicated that some of the carboxylic acid groups were deprotonated through ligand exchange with Fe(III); and the band at 1400 cm⁻¹ could be assigned to –COO–Fe, which was compelling evidence for the HA–Fe(III) complexes [34]. Besides the carboxyl group, phenolic hydroxyl

in HA can also complex with Fe(III). It is proposed that the inhibited effect of Fe(III) may be related to the complexing of Fe(III) with a carboxyl group or hydroxyl group in HA [35]. On the one hand, this complexing process reduces the effective concentration of HA in solution by adsorption and co-precipitation (Figure S4); on the other hand, it consumes the functional groups of HA, and to some extent inhibits the protonation of the aromatic ring and disturbs the reaction between the active sites and chloride ions [36]. It should be noted that the results obtained here under acid conditions may not be widely applied in neutral and alkaline conditions. Complexing of HA–Fe(III) significantly affects the photochemical production of CH_3Cl and the behavior of HA, which needs further investigation.

It is notable that upon analysis of the volatile gas using GC-MS for the sample containing Fe(III) and HA, a significant amount of trichloromethane was confirmed in the scan mode (Figure S5). According to the studies of Liu et al., aromatic halogenated disinfection by-products (DBPs) might be produced from natural organic matter in the presence of chloride during UV irradiation where reactive chlorine species (RCS) acted as the major halogenated reagent, and the resulting DBPs may decompose to aliphatic halogenated DBPs, such as trichloromethane [37]. It can be deduced that there were abundant reactive chlorine species (RCS) in the solution, including radical RCS, e.g., $Cl^{\circ}/Cl_2^{\circ-}$, and non-radical RCS, e.g., HClO and Cl_2 . The generation of the radical and non-radical RCS was closely related to \cdot OH, which could be produced by Fe(III) or Fe(III)–HA (Equation (4)) [38]. The \cdot OH is capable of oxidizing Cl⁻ to form Cl[•]/Cl₂^{\u03ber} (Equations (5–7)), Cl₂ and HClO (Equations (8–9)) [39].

$$Fe(OH)^{2+} \rightarrow Fe^{2+} + OH$$
 (4)

$$^{\circ}OH + Cl^{-} \rightarrow HClO^{--}$$
 (5)

$$HClO^{-} + H^{+} \rightarrow Cl^{-} + H_{2}O$$
(6)

$$Cl' + Cl^- \rightarrow Cl_2^-$$
 (7)

$$Cl_2^- + Cl_2^- \rightarrow Cl_2 + 2Cl^- \tag{8}$$

$$\operatorname{Cl}_{2}^{-} + \operatorname{OH} \rightarrow \operatorname{HClO} + \operatorname{Cl}^{-}$$
 (9)

Although RCS existed abundantly in the solution in the presence of Fe(III), CH_3Cl was significantly decreased by Fe(III). This result could exclude the effect of RCS on CH_3Cl formation. The same result has been observed by Yang et al.: RCS was not the active intermediate to generate CH_3Cl from guaiacol under UV254 irradiation [18].

3.6. Generation of Methyl Chloride in Natural Seawater

To investigate the formation of CH₃Cl in natural seawater, experiments were performed on coastal seawater collected in the Northern Yellow Sea near the Blackrock Reef, Dalian, China. The seawater was first passed through a 0.45 μ m membrane and the DOM from the seawater was retained in water. Then, the seawater sample was exposed to light. The results are shown in Figure 9. The concentration of CH₃Cl after 40 h of irradiation reached 18.8 ng L⁻¹ which was at least three times that of the dark control sample. The concentration of CH₃Cl in the dark control was around 5 ng L⁻¹, which can be attributed to the biotic production by some bacteria and microorganisms in the seawater [40,41]. Of course, this biological process may also produce CH₃Cl under light irradiation through some photosynthetic process in this study. Nevertheless, the light process is a significant pathway for the production of CH₃Cl in natural seawater. Other contributors to the marine source of CH₃Cl include nucleophilic displacement reactions between CH₃X and chloride ions, where CH₃X (X=Br, I) was generated by biological processes of phytoplankton and macroalgae or photochemical process [8,22].



Figure 9. Methyl chloride production in natural seawater from Northern Yellow Sea, China. Reaction conditions: light intensity = 12 mW cm^{-2} .

4. Conclusions

An investigation of CH_3Cl production from HA in aqueous chloride solutions and natural seawater under simulated sunlight irradiation revealed the significant photochemical generation of CH_3Cl in saline waters. The promoting factors for CH_3Cl generation are light intensity, HA and chloride concentration, whereas the inhibiting factor is ferric ions (under acidic conditions). CH_3Cl formation was pH-dependent, with the highest amount generated near neutral conditions. Reactions of the kind described in this paper are expected to contribute to CH_3Cl production in estuarine and coastal waters.

Supplementary Materials: The following are available online at http://www.mdpi.com/1660-4601/17/2/503/s1, Figure S1: Schematic of the device used for irradiation. Figure S2: Illustration of the purge-and trap device and GC-MS. Figure S3: HA absorption and Xenon lamp emission spectrum. Figure S4: UV-vis spectra of HA in the presence of Fe(III). Figure S5: Chromatogram of the gases from the solutions containing 5 mg L⁻¹ HA, 0.5 mol L⁻¹ Cl⁻ and 600 µmol L⁻¹ Fe(III) under irradiation of 15 h, and the MS spectra of trichloromethane.

Author Contributions: Conceptualization, H.L.; methodology, H.L.; formal analysis, H.L. and X.Z. (Xiaomei Zhu); investigation, Y.P. and T.T.; resources, B.S.; writing—original draft preparation, Y.P.; writing—review and editing, H.L. and Y.P.; project administration, X.Z. (Xiaoxing Zhang); funding acquisition, H.L. All authors have read and agreed to the published version of the manuscript.

Funding: This research was funded by the National Natural Science Foundation of China (No. 41576111, 41206095, and 21407018), and Fundamental Research Funds for the Central Universities (No. 3132019329).

Conflicts of Interest: There are no conflicts to declare.

References

- 1. Khalil, M.A.K.; Rasmussen, R. Atmospheric methyl chloride. Atmos. Environ. 1999, 33, 1305–1321. [CrossRef]
- Wallington, T.J.; Pivesso, B.P.; Lira, A.M.; Anderson, J.E.; Nielsen, C.J.; Andersen, N.J.; Hodnebrog, O. CH₃Cl, CH₂Cl₂, CHCl₃ and CCl₄: Infrared spectra, radiative efficiencies, and global warming potentials. *J. Quant.* Spectrosc. Radiat. 2016, 174, 56–64. [CrossRef]
- 3. Grossman, A.S.; Grant, K.E.; Blass, W.E.; Wuebbles, D.J. Radiative forcing calculations for CH₃Cl and CH₃Br. *J. Geophys. Res. Atoms.* **1997**, *102*, 13651–13656. [CrossRef]
- 4. Keppler, F.; Eiden, R.; Niedan, V.; Pracht, J.; Schöler, H.F. Halocarbons produced by natural oxidation processes during degradation of organic matter. *Nature* **2000**, *403*, 298–301. [CrossRef] [PubMed]
- Yokouchi, Y.; Noijiri, Y.; Barrie, L.A.; Toom-Sauntry, D.; Machida, T.; Inuzuka, Y.; Akimoto, H.; Li, H.J.; Fujinuma, Y.; Aoki, S. A strong source of methyl chloride to the atmosphere from tropical coastal land. *Nature* 2000, 403, 295–298. [CrossRef] [PubMed]

- 6. Rhew, R.C.; Miller, B.R.; Weiss, R.F. Natural methyl bro-mide and methyl chloride emissions from coastal saltmarshes. *Nature* **2000**, 403, 292–295. [CrossRef] [PubMed]
- Méndez-Díaz, J.D.; Shimabuku, K.K.; Enumah, J.M.Z.O.; Pignatello, J.J.; Mitch, W.A.; Dodd, M.C. Sunlight-driven photochemical halogenation of dissolved organic matter in seawater: A natural abiotic source of organobromine and organoiodine. *Environ. Sci. Technol.* 2014, 48, 7418–7427. [CrossRef]
- 8. Moore, R.M. Methyl halide production and loss rates from field incubation experiments. *Mar. Chem.* **2006**, 101, 213–219. [CrossRef]
- 9. Moore, R.M. A photochemical source of methyl chloride in saline waters. *Environ. Sci. Technol.* 2008, 42, 1933–1937. [CrossRef]
- McNeill, K.; Canonica, S. Triplet state dissolved organic matter in aquatic photochemistry: Reaction mechanisms, substrate scope, and photophysical properties. *Environ. Sci. Processes Impacts* 2016, 18, 1381–1399. [CrossRef]
- Sandron, S.; Rojas, A.; Wilson, R.; Davies, N.W.; Haddad, P.R.; Shellie, R.A.; Nesterenko, P.N.; Kelleher, B.P.; Paull, B. Chromatographic methods for the isolation, separation and characterisation of dissolved organic matter. *Environ. Sci. Processes Impacts* 2015, *17*, 1531–1567. [CrossRef] [PubMed]
- 12. Glover, C.M.; Rosario-Ortiz, F.L. Impact of halides on the photoproduction of reactive intermediates from organic matter. *Environ. Sci. Technol.* **2013**, *47*, 13949–13956. [CrossRef] [PubMed]
- Brigante, M.; Minella, M.; Mailhot, G.; Maurino, V.; Minero, C.; Vione, D. Formation and reactivity of the dichloride radical (Cl₂⁻) in surface waters: A modelling approach. *Chemosphere* 2014, 95, 464–469. [CrossRef] [PubMed]
- 14. Jammoul, A.; Dumas, S.; D'anna, B.; George, C. Photoinduced oxidation of sea salt halides by aromatic ketones: A source of halogenated radicals. *Atmos. Chem. Phys.* **2009**, *9*, 4229–4237. [CrossRef]
- Hao, Z.; Yin, Y.; Cao, D.; Liu, J. Probing and comparing the photobromination and photoiodination of dissolved organic matter by using ultra-high-resolution mass spectrometry. *Environ. Sci. Technol.* 2017, 51, 5464–5472. [CrossRef] [PubMed]
- Liu, Y.; Thornton, D.C.O.; Bianchi, T.S.; Arnold, W.A.; Shields, M.R.; Chen, J.; Yvon-Lewis, S.A. Dissolved organic matter composition drives the marine production of brominated very short-lived substances. *Environ. Sci. Technol.* 2015, 49, 3366–3374. [CrossRef]
- Dallin, E.; Wan, P.; Krogh, E.; Gill, C.; Moore, R.M. New pH-dependent photosubstitution pathways of syringic acid in aqueous solution: Relevance in environmental photochemistry. *J. Photochem. Photobiol. A Chem.* 2009, 207, 297–305. [CrossRef]
- 18. Yang, Q.; Guo, Y.; Yue, E.; Zhang, S.; Blatchley, E.R., III; Li, J. Methyl chloride produced during UV irradiation of saline water. *J. Hazard. Mater.* **2019**, *384*, 121263. [CrossRef]
- Song, G.; Li, Y.; Hu, S.; Li, G.; Zhao, R.; Sun, X.; Xie, H. Photobleaching of chromophoric dissolved organic matter (CDOM) in the Yangtze River estuary: Kinetics and effects of temperature, pH and salinity. *Environ. Sci. Processes Impacts* 2017, 19, 861–873. [CrossRef]
- Mylon, S.E.; Chen, K.L.; Elimelech, M. Influence of natural organic matter and ionic composition on the kinetics and structure of hematite colloid aggregation: Implications to iron depletion in estuaries. *Langmuir* 2004, 20, 9000–9006. [CrossRef]
- 21. Hopwood, M.J.; Statham, P.J.; Milani, A. Dissolved Fe(II) in a river-estuary system rich in dissolved organic Matter. *Estuar. Coast. Shelf Sci.* **2014**, 151, 1–9. [CrossRef]
- 22. Moore, R.M.; Zafiriou, O.C. Photochemical production of methyl iodide in seawater. *J. Geophys. Res.* **1994**, *99*, 16415–16420. [CrossRef]
- 23. Manley, S.L.; Barbero, P.E. Physiological constraints on bromoform (CHBr₃) production by *Ulva lactuca* (*Chlorophyta*). *Limnol. Oceanogr.* **2001**, *46*, 1392–1399. [CrossRef]
- 24. Lin, C.Y.; Manley, S.L. Bromoform production from seawater treated with bromoperoxidase. *Limnol. Oceanogr.* 2012, *57*, 1857–1866. [CrossRef]
- Kohn, T.; Grandbois, M.; McNeill, K.; Nelson, K.L. Association with natural organic matter enhances the sunlight-mediated inactivation of MS2 coliphage by singlet oxygen. *Environ. Sci. Technol.* 2007, 41, 4626–4632. [CrossRef]
- 26. Leifer, A. *The Kinetics of Environmental Aquatic Photochemistry: Theory and Practice;* American Chemical Society: Washington, DC, USA, 1988.

- 27. Chiron, S.; Minero, C.; Vione, D. Photodegradation processes of the antiepileptic drug carbamazepine, relevant to estuarine waters. *Environ. Sci. Technol.* **2006**, *40*, 5977–5983. [CrossRef]
- 28. Longstaffe, J.G.; Courtier-Murias, D.; Simpson, A.J. The pH-dependence of organofluorine binding domain preference in dissolved humic acid. *Chemosphere* **2013**, *90*, 270–275. [CrossRef]
- 29. Gan, L.; Yan, Z.; Ma, Y.; Zhu, Y.; Li, X.; Xu, J.; Zhang, W. pH dependence of the binding interactions between humic acids and bisphenol AA thermodynamic perspective. *Environ. Pollut.* **2019**, 255, 113292. [CrossRef]
- Karpukhina, E.; Mikheev, I.; Perminova, I.; Volkov, D.; Proskurnin, M. Rapid quantification of humic components in concentrated humate fertilizer solutions by FTIR spectroscopy. *J. Soils Sediments* 2018, 19, 2729–2739. [CrossRef]
- 31. Liu, Q.J.; Li, X.; Tang, J.P.; Zhou, Y.M.; Lin, Q.T.; Xiao, R.; Zhang, M. Characterization of goethite-fulvic acid composites and their impact on the immobility of Pb/Cd in soil. *Chemosphere* **2019**, 222, 556–563. [CrossRef]
- 32. Balmer, M.E.; Sulzberger, B. Atrazine degradation in irradiated iron/oxalate systems: Effects of pH and oxalate. *Environ. Sci. Technol.* **1999**, *33*, 2418–2424. [CrossRef]
- Ou, X.; Quan, X.; Chen, S.; Zhang, F.; Zhao, Y. Photocatalytic reaction by Fe(III)-citrate complex and its effect on the photodegradation of atrazine in aqueous solution. *J. Photochem. Photobiol. A Chem.* 2008, 197, 382–388. [CrossRef]
- Gu, B.H.; Schmitt, J.; Chen, Z.H.; Liang, L.Y.; McCarthy, J.F. Adsorption and desorption of natural organic matter on iron oxide: Mechanisms and models. *Environ. Sci. Technol.* 1994, 28, 38–46. [CrossRef]
- 35. Cieśla, P.; Kocot, P.; Mytych, P.; Stasicka, Z. Homogeneous photocatalysis by transition metal complexes in the environment. *J. Mol. Catal. A Chem.* **2004**, 224, 17–33. [CrossRef]
- 36. Porasso, R.D.; Benegas, J.C.; Van den Hoop, M.G.T.; Paoletti, S. Analysis of trace metal humic acid interactions using counterion condensation theory. *Environ. Sci. Technol.* **2002**, *36*, 3815–3821. [CrossRef]
- 37. Liu, J.; Zhang, X.; Li, Y. Photoconversion of chlorinated saline wastewater DBPs in receiving seawater is overall a detoxification process. *Environ. Sci. Technol.* **2017**, *51*, 58–67. [CrossRef]
- 38. Wu, F.; Deng, N. Photochemistry of hydrolytic iron (III) species and photoinduced degradation of organic compounds. A minireview. *Chemosphere* **2000**, *41*, 1137–1147.
- 39. Zhang, K.; Parker, K.M. Halogen radical oxidants in natural and engineered aquatic systems. *Environ. Sci. Technol.* 2018, 52, 9579–9594. [CrossRef]
- 40. Hirata, M.; Ikeda, M.; Fukuda, F.; Abe, M.; Sawada, H.; Hashimoto, S. Effect of temperature on the production rates of methyl halides in cultures of marine proteobacteria. *Mar. Chem.* **2017**, *196*, 126–134. [CrossRef]
- Sato, N.; Hamamoto, K.; Kurihara, M.; Abe, M.; Hashimoto, S. Methyl halide production by cultures of marine thraustochytrids, *Aurantiochytrium* sp., *Botryochytrium radiatum*, and *Schizochytrium* sp. *Mar. Chem.* 2019, 208, 95–102. [CrossRef]



© 2020 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (http://creativecommons.org/licenses/by/4.0/).