

## RESEARCH ARTICLE

# Surface temperature behavior in view of the conversion of tropical dry forest into anthropic uses, northern Minas Gerais–Brazil

Lucas Augusto Pereira da Silva<sup>1</sup>, Andre Medeiros Rocha<sup>2</sup>, Claudionor Ribeiro da Silva<sup>1\*</sup>

**1** Institute of Geography, Program of Post-Graduation in Geography, Federal University of Uberlândia, Uberlândia, Minas Gerais, Brazil, **2** Program of Post-Graduation in Physical Geography, University of São Paulo, São Paulo, São Paulo, Brazil

\* [crs@ufu.br](mailto:crs@ufu.br), [crs.educ@gmail.com](mailto:crs.educ@gmail.com)



## Abstract

Tropical dry forests (TDFs) are essential for environmental dynamics, especially in terms of climate variations. However, several anthropic factors have threatened the integrity of TDFs, and consequently the surface temperature (ST), which is a proxy variable for several environmental processes in TDFs. So, understanding their behavior is crucial. The objective was to analyze the behavior of surface temperature owing to conversion of TDFs into anthropic uses in northern Minas Gerais between 2007 and 2016. In 9 years, dry forests decreased by 22.9%, with pastures as the central driver (counted 93% of change). Between 2007 and 2016, there was an increase in ST by  $1.55 \text{ K} \pm 1.15 \text{ K}$ . When TDFs were converted to pastureland, the increase in ST was  $2.21 \text{ K} \pm 1.39 \text{ K}$  and for crops by  $0.57 \text{ K} \pm 1.24 \text{ K}$ . The remaining TDFs (2016) had an increase in their thermal average of  $1.41 \text{ K} \pm 1.02 \text{ K}$ . This analysis is essential for the adoption of conservation actions for the maintenance of ecological corridors in TDFs, considering their importance in the ecosystem context.

## OPEN ACCESS

**Citation:** Silva LAPd, Rocha AM, Silva CRd (2022) Surface temperature behavior in view of the conversion of tropical dry forest into anthropic uses, northern Minas Gerais–Brazil. PLoS ONE 17(7): e0270991. <https://doi.org/10.1371/journal.pone.0270991>

**Editor:** Julia A. Jones, Oregon State University, UNITED STATES

**Received:** July 10, 2021

**Accepted:** June 21, 2022

**Published:** July 27, 2022

**Copyright:** © 2022 Silva et al. This is an open access article distributed under the terms of the [Creative Commons Attribution License](https://creativecommons.org/licenses/by/4.0/), which permits unrestricted use, distribution, and reproduction in any medium, provided the original author and source are credited.

**Data Availability Statement:** All relevant data are within the paper.

**Funding:** The authors would like to thank the Federal University of Uberlândia and the Postgraduate Program in Geography (PPGEO) for helping to pay the publication fee for this paper.

**Competing interests:** The authors have declared that no competing interests exist.

## Introduction

Subtropical and subtropical and tropical dry broadleaf forests are the most endangered forested biome on the globe [1, 2] and although such fact has ever been known in a form of an alert concerning its conservation status and level of threatening by Janzen [3] in the 1980s, still nowadays, more than 30 years after, tropical dry forests (TDFs) preservation and importance remains a debate topic [4].

As a general incentive to the intense anthropic pressure observed in tropical dry forest areas, its physiological and ecological characteristics are among the main reasons that make it a place favorable to human occupation when compared to the savannas and rain forests domains. TDFs generally occupy flat areas, in fertile soils, with low aluminum content and present physiological and ecological behavior adapted to seasonal droughts [5]. Savannas, on the other hand, occupy more poor, acidic soils, depending on the subtype, and with high aluminum content [6].

Compared to rain forests, dry forests have lower stature, and their ecological processes are time-limited due to its property of deciduousness associated with seasonality. Its net primary productivity, for instance, is lower when compared to rain forests, since TDFs photosynthetic

activity is tied to the rainy season [7]. These characteristics of seasonality, fertile soils, and relatively flat areas are among the factors that make tropical dry forests prone to human activity occupation, in particular agriculture, whose agricultural calendar lifestyle in general better fits with [8].

In terms of conservation, subtropical and tropical dry broadleaf forests ecoregion with more than 50% of their area in protected status (half protected status) accounts only for 2 out of the 56 (3.6%) total existing ecoregions, while 26 of them (46.4%) shows less than 20% in protected status (nature imperiled status) [2]. The aforementioned fact has a relative weight and constitutes a matter of concern, with a view that combined effect of low degree of protection and high anthropic pressure associated with dry forests put at risk the maintenance of ecosystem services those environments provide. TDFs, for example, are largely characterized by providing important ecosystem services such as: carbon storage source [9, 10], biodiversity wealth [8] and hydrological [11] and energy-balance [12–14] regulation.

The disappearance of the listed ecosystem services earns notoriety because such processes interfere with the global air [15] surface [16] temperature, which are parametric variables for measuring the average climatic state of the Earth. In this way, the changes in land use and land cover associated with deforestation and greenhouse gas emissions (GGEs) constitute exemplary anthropogenic radiative forcings that impact global temperature, producing increase (positive forcing) or reduction (negative forcing) [17].

The loss of forest areas, on the one hand, has direct effects on carbon cycle. Forests are one of the largest terrestrial carbon sinks [9], acting as one of the strategies for balancing atmospheric CO<sub>2</sub> content. The replacement of forest by pasturelands or croplands makes such carbon sequestration strategy an unfeasible plan [18]. On the other hand, the loss of forested areas promotes reduction on evapotranspiration rate and surface storage of precipitation. In terms of radiation balance, the deforestation of forest areas has also the effect of increasing the surface albedo and temperature and decreasing surface latent and sensible heat transfer capacity [11–13].

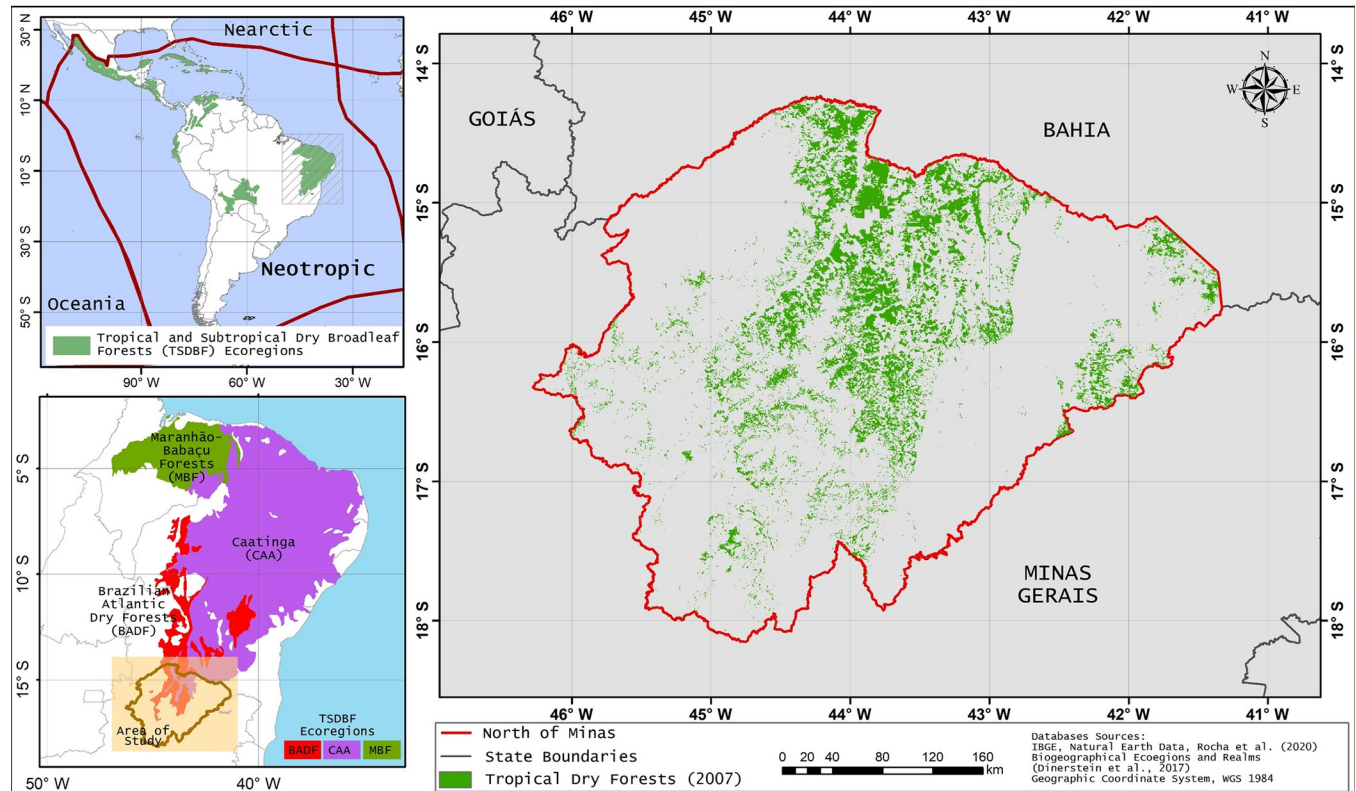
In the context of TDFs, Brazil stands out for sheltering three of the most extensive ecoregions of the neotropics and by such role played, preservation of Brazilian tropical dry forests is of extremely concern. In terms of forest resources, Brazil occupies the second largest forest area on the planet, encompassing 12% of global forests resources (app. 496 million ha) for period 2010–2020 and impressively is the country with the highest annual average loss of forests areas (app. -1.4 million ha /yr) [19]. On the other hand, in terms of dry forests, Brazil stands out in the neotropics for harboring the two ecoregions of TDFs with a larger extent of areas at least degree of protection (nature imperiled category) [2].

In view of the ecosystem services that TDFs play and considering the role that Brazil has in dry forest resources, the current investigation aims to understand the behavior of the surface temperature in the north of Minas Gerais before and after the TDFs conversions to anthropic uses between 2007 and 2016. For this purpose, remote sensing techniques were employed, in order to allow the spatial analysis of extensive areas, to have fast and continuous imaging of the surface, and, mainly, to make products available free of charges. Of equal importance, it is emphasized that remote sensing has often been used as a tool for monitoring and modeling aspects of vegetation [20–22].

## Materials and methods

### Study area

The study area comprises the southern sector of the ecoregions: Caatinga and Brazilian Atlantic dry forests, two of the 56 ecoregions that compose the Tropical and subtropical broadleaf



**Fig 1. Location of TDFs in the north of Minas Gerais, Brazil, South America.**

<https://doi.org/10.1371/journal.pone.0270991.g001>

dry forest biome–neotropic realm (Fig 1). Such area was chosen mainly by 1) possessing historic of TDFs conversion to anthropic uses [23, 24], 2) status of nature imperiled in the categorization scheme of levels of protection for ecoregions presented by Dinerstein et al. [2], 3) compose two of the four most extended ecoregions areas of TDFs in the neotropics, 4) encompassing areas with a high level of annual carbon sequestration [25]. In view of those criteria, the north of the state of Minas Gerais was taken as an area of investigation, being situated between latitudes 14° and 18°S, longitudes 46° and 42°W.

The TDFs had an area of 17.480 km<sup>2</sup> in 2016, distributed over all extensions of the study area. The climatic context of the study area is marked by two climatic types, Aw (wet tropical, with dry winters) and BSh (semi-arid). The rainfall patterns for the region are marked by dry winters and rainy summers, ranging from 700 to 1200 mm, and average temperature ranging from 294 K to 298 K [26]. Morphoclimatically, the area falls within the transition bands between Cerrados and Caatingas [27]. The relief is flat to slightly undulating, composing extensions of the São Francisco valley. The area is predominantly marked by oxisols, nitosols, and neosols [28].

## Database and processing

For the spatial representation of anthropic and natural domains, we used the products provided by Rocha et al. [24], derived from the mapping of land uses and land cover between the north of the state of Minas Gerais and the south of Piauí (the northeast portion of Brazil) for 2007 and 2016. To map the forests in this portion of Brazil, the authors used the MOD13Q1 products (250 meters spatial resolution) from the MODIS sensor [29]. The modeling was done

from supervised classification with the decision tree algorithm [30]. Based on validation processes, the authors obtained an overall accuracy of 89.40% and a kappa coefficient of 86.79% [24].

Originally, the mapping by Rocha et al. [24] has six classes (TDFs, crops, pastures, urban areas, water bodies, and other vegetation), however, for our study, only three were used (TDFs, crops, and pastures). Rocha et al. [24] obtained commission and omission errors for the land use and land cover classes, in which TDFs had 15% and 26.84%, pastureland 28% and 19.86%, and crops 6.4% and 3.14%, respectively.

The MOD11A1 product (tiles H13V10 and H14V10, with a spatial resolution of 1 km) [31] from the MODIS sensor was used to represent the surface temperature dynamics in the study area. This product was provided by the United States Geological Survey. To contemplate our approach, 365 images were obtained (daily time scale) for the two years analyzed (2007 and 2016). The MOD11A1 is a reliable and accurate product for the values of ST because in its generation it considers factors such as i) emissivity, ii) water vapor content in the atmosphere and iii) removal of pixels contaminated by clouds, which is essential for estimation of ST [32].

## Analysis

The territorial limit of the 2007 TDFs was used as a spatial clipping for the subsequent analyzes of this study. Although there may have been a change in the coverage of TDFs before 2007, due to the absence of detailed mappings in the domains of TDFs, we assumed this period to establish the natural domain of TDFs. Therefore, all MOD11A1 images were clipped from the defined limit. These procedures were performed in the QGIS software [33].

To understand the dynamics of ST in land uses and land cover on a monthly scale, the images were compiled in 12 products (January to December) for the two years. In order for record the annual changes in ST in the study area, we obtained the annual averages for 2007 and 2016. To carry out the statistical analyses, we obtained sample points (distance of 1 km between the points) considering the classes of land uses and land cover (Fig 1B). In 2007, for the TDFs domain, 1,500 points were sampled. In 2016, for TDFs and pastures, 1,000 points were obtained, and for crop 123. A descriptive statistical analysis was performed with boxplot plots from the constructed matrix. These analyzes were performed in the R software with the ggplot2 package [34, 35].

In order to obtain the thermal differences by land use and land cover classes, as well as to understand the patterns of increases and decreases and variability of the ST in the TDFs domain, we calculated the difference between the annual STs (Equation 01).

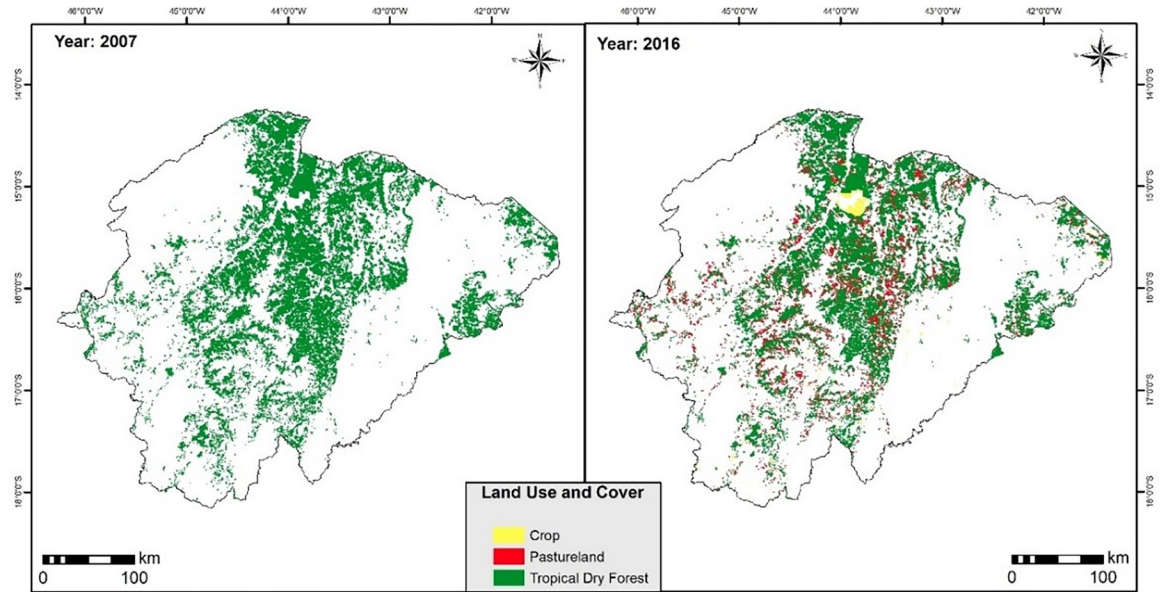
$$\Delta ST = ST_{16} - ST_{07}$$

Where  $\Delta ST$  is the difference between the temperatures for the years 2016 (ST<sub>16</sub>) and 2007 (ST<sub>07</sub>), respectively.

## Results

### Dynamics of land use and coverage in the domain of dry forests

The natural domain of TDFs in 2007 had an area of 17,480 km<sup>2</sup>, with distribution in all parallels of the northern region of Minas Gerais (Fig 2). For the period of 2016, there was a loss of forests, leading to a decrease of 22.9% (4,010.6 km<sup>2</sup>) in the natural domain of TDFs, in a short interval of 9 years (Fig 2). The main uses that replaced forests were, mainly, environments with anthropogenic influence. In this case, pastureland systems (3,730.6 km<sup>2</sup>) and agricultural crop (273.6 km<sup>2</sup>). Pastureland counted for 93% of changes in land use and cover in the TDFs domain.



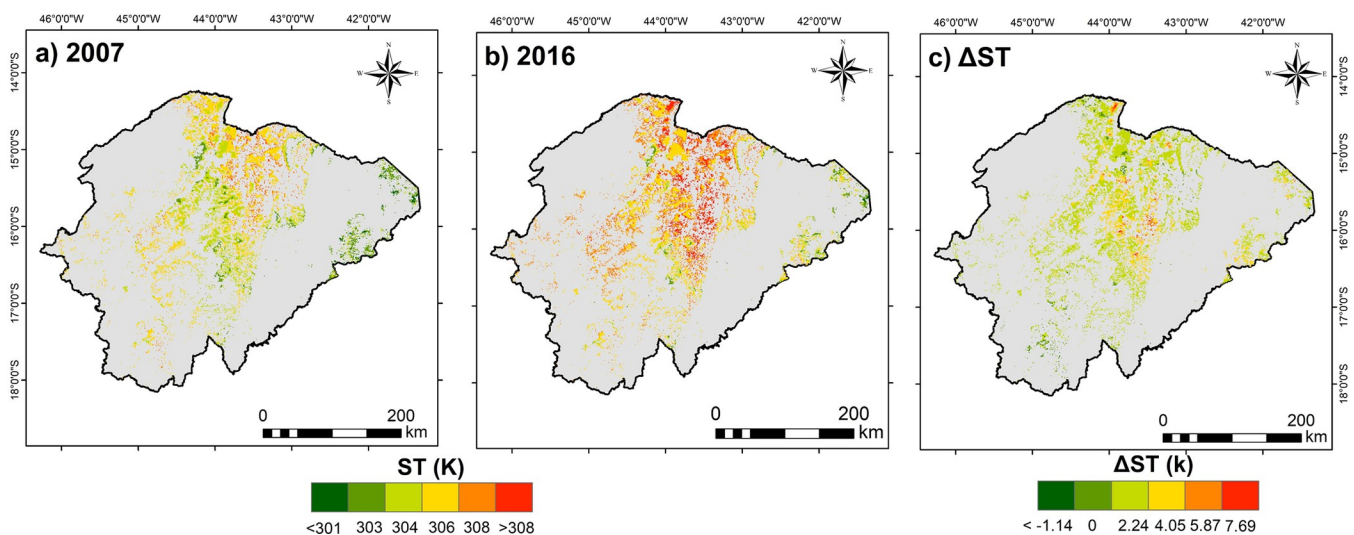
**Fig 2. TDFs and anthropic uses, 2007 and 2016, north of Minas Gerais.**

<https://doi.org/10.1371/journal.pone.0270991.g002>

### Spatial distribution of the surface temperature in the northern Minas Gerais context

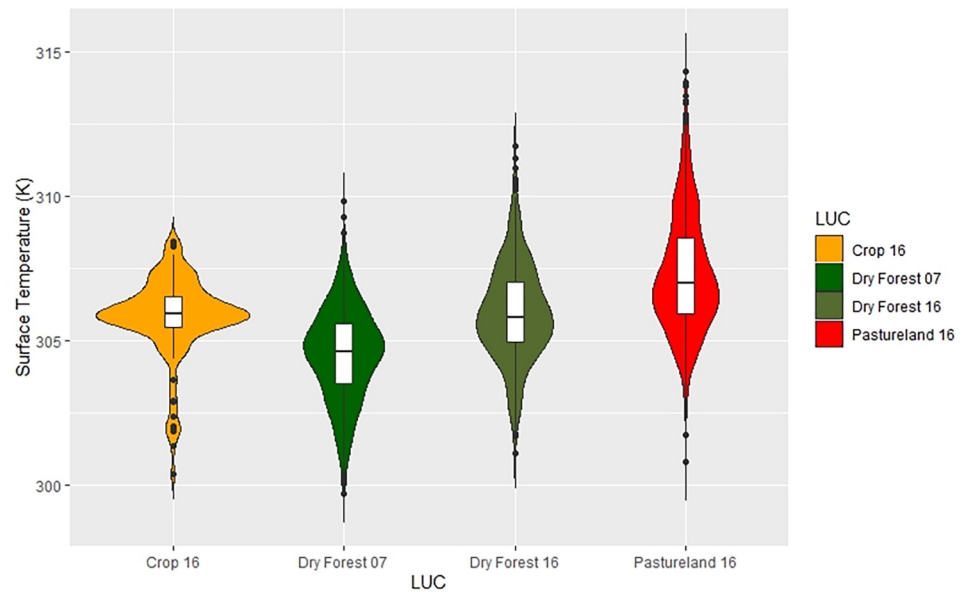
For the spatial domain of TDFs in 2007, the ST mean was  $305\text{ K} \pm 1.61\text{ K}$ , ranging between 301 K to 310 K (Fig 3A). In the 9-year interval (2016, Fig 3B), there was an increase of the order of  $1.55\text{ K} \pm 1.15\text{ K}$  in the mean of ST, with maximum values reaching  $> 7\text{ K}$  (Fig 3C). Considering the changes in the spatial variation of ST between 2007 and 2016 ( $\Delta\text{ST}$ ), most of the study area ( $> 70\%$ ) showed an increase up to 2.24 K.

In 2016 (Fig 3B), advances were observed in the maximum values of ST ( $> 308$ ) in relation to the base period (2007). It is possible to infer a relationship between this behavior with the



**Fig 3. Spatial distribution of surface temperature (ST) in the TDFs domain.** a) Spatial distribution of ST for 2007; b) Spatial distribution of ST for 2016; c) Difference ( $\Delta\text{ST}$ ) between the ST for the years 2016 and 2007 ( $\Delta\text{ST} = 2016 - 2007$ ).

<https://doi.org/10.1371/journal.pone.0270991.g003>



**Fig 4. Boxplot of surface temperature (ST) considering the annual average (from the years 2007 and 2016) for the analyzed land uses and land cover.**

<https://doi.org/10.1371/journal.pone.0270991.g004>

insertion of anthropic uses, mainly pastureland, as this class has advanced to the detriment of TDFs.

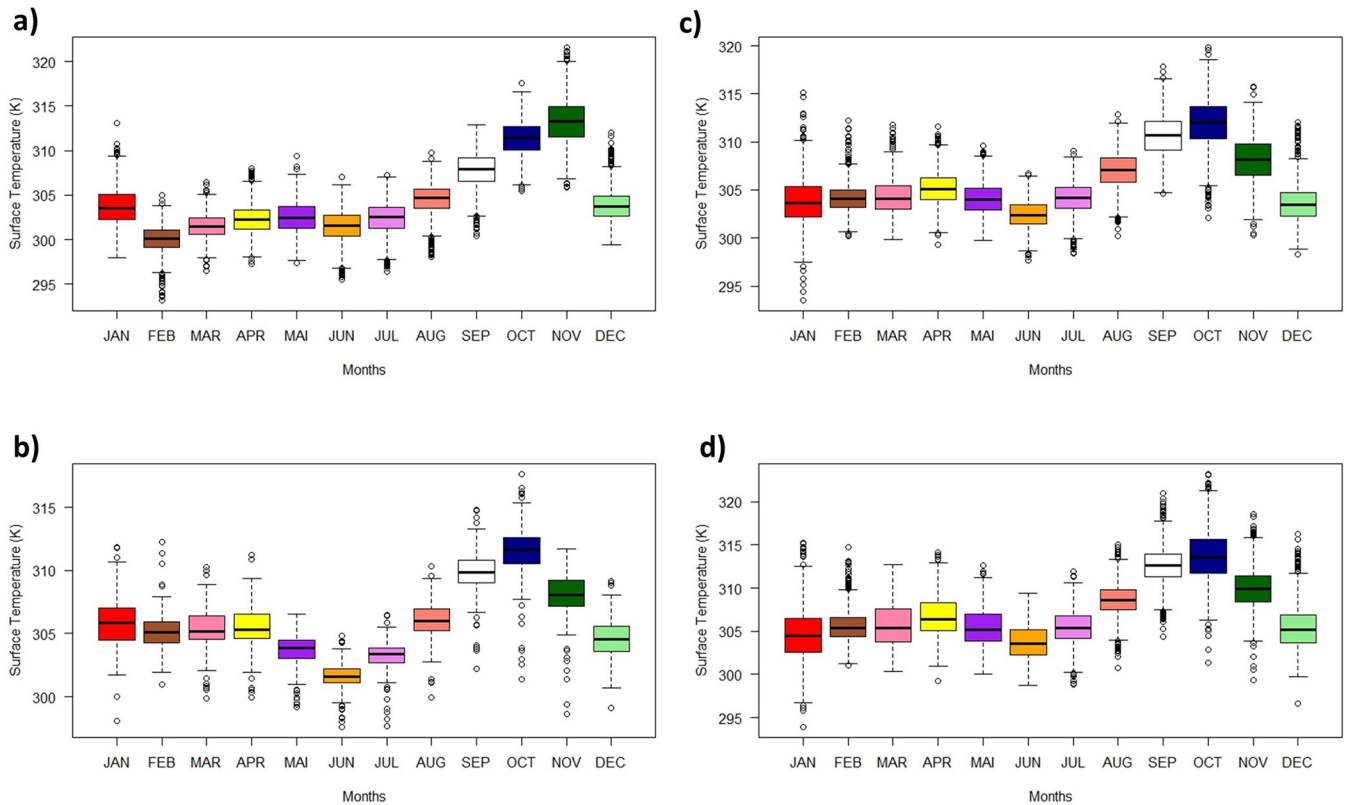
### Surface temperature in tropical dry forest and anthropic uses

In annual terms, in the interval of 9 years (2007 to 2016), with the verified conversions to anthropic uses, there was a substantial increase in ST (Fig 4). When converting forests into pastureland areas, the ST increased by  $2.21 \text{ K} \pm 1.39 \text{ K}$ , while in the conversion to crop there was an increase of  $0.57 \text{ K} \pm 1.24 \text{ K}$ . In this scenario, the remaining forests were warmer, with an increase in ST in the order of  $1.41 \text{ K} \pm 1.02 \text{ K}$ .

Considering the temporal (monthly) variation of the ST for all land uses and land cover, a pattern of behavior was observed, where the highest values occurred in the months of September, October, and November (Fig 5). While the lowest values were for the months of May, June, and July. Among the analyzed domains, in most months, pastureland presented the highest ST values, with a peak in October (average of  $313.8 \text{ K} \pm 10.4 \text{ K}$ , reaching a maximum of  $323.2 \text{ K}$ ). The TDFs areas had the highest monthly ST values, with a peak in November ( $313.3 \text{ K} \pm 8.47 \text{ K}$ , with a maximum of  $321.6 \text{ K}$ ). In 2016, this class had a ST peak in October ( $311.9 \text{ K} \pm 10.21 \text{ K}$ ). The crop class had an ST peak in October, with an average of  $311.14 \text{ K} \pm 28.09 \text{ K}$  and a maximum of  $317.6 \text{ K}$ . Based on the standard deviations of the classes for the months with the highest values of ST, it was evident that in 2007, in the forests domain, ST varied less (more homogeneous) in relation to the others ( $<$  standard deviation), while crop showed greater variability.

### Discussion

In this work, the behavior of ST in the domains of TDFs in the face of land use and land cover conversions in the north of the state of Minas Gerais was evaluated using remote sensing techniques. The main results found were: i) the natural domain of forest decreased by 22% between



**Fig 5. Monthly surface temperature boxplot (January to December) for land uses and land cover in the study area.** a) monthly ST for the TDFs in 2007; b) monthly ST for the TDFs in 2016; c) monthly ST for pastureland in 2016; d) Monthly ST for crop in 2016.

<https://doi.org/10.1371/journal.pone.0270991.g005>

2007 and 2016, mainly due to anthropic uses (pastureland and crop); ii) land use and land cover conversion led to an increase of  $1.55 \text{ K} \pm 1.15 \text{ K}$  in the 9-year interval; iii) the inclusion of pastureland at the expense of forest boosted substantial increases in ST ( $2.21 \text{ K} \pm 1.39 \text{ K}$ ).

Changes in the land use and land cover context for the domain of TDFs have been reported by several authors [23, 24, 36]. Dupin et al. [23] indicated that between 2000 and 2015 there was a net loss in the order of  $3302 \text{ km}^2$ . The authors also concluded that cattle density (located in pastureland areas) had a direct impact on forest loss in the studied period. Regarding crops, Dupin et al. [23] also highlight that the Jaíba irrigation project (irrigated agriculture, mainly for bananas) has a high impact on the loss of TDFs in the study area. Rocha et al. [24], in monitoring land use and land cover in TDFs domains covering the north of the state of Minas Gerais to the south of Piauí (northeast region of Brazil), showed that there was forest loss in the order of  $62975 \text{ km}^2$  (between 2007 and 2016), in which  $> 72\%$  was explained by pastureland areas, and  $> 20\%$  by crop, showing similar behavior to this research, where pastureland predominated in the context of forest conversion, followed by crop.

The TDFs domains are characterized by high levels of biomass and well-structured canopy [37], substantial characteristics for the dynamics of thermal fluctuations in a given ecosystem. These factors condition the interception and absorption of solar radiation incident to the surface, as well as the conversion into water vapor fluxes, which tends to maintain low ST values (as seen in the results, Figs 4 and 5), due to surface wetting [38]. However, spatio-temporal changes in the configuration of these domains tend to quickly infer the thermal properties, especially when there is an insertion of land use with different characteristics to the natural

domain of TDFs, such areas as pastures and crops, which generally have low levels of leaf area, facilitating the penetration and direct contact of solar radiation with the soil surface, consequently heating it and raising the ST, as has been reported in other studies that aimed to understand the impacts of forest systems conversions to pastureland and crop [39, 40].

Santos et al. [41], in the northwest of Brazil (state of Rondônia), showed by remote sensing techniques a difference of ST of 2.5 K in Forest areas when compared to pastureland. Pavão et al. [42], in areas in northern Brazil (Amazonas state), found an increase of 2.4 K in the conversion of forests to pastureland. Similarly, for areas of the Amazon rainforest, Querino et al. [43] observed an increase in surface temperature in the order of 6 K due to the conversion of forests to pastureland. Zhou and Wang [39], in analysis in southwest China, they found an increase of 1.5 K due to the conversion of forests into crop. In a global context, Pongratz et al. [44] observed that from the conversion of forests to pastureland and crop there was an increase in the order of 1.2 and 1.7 K respectively. While for northern Brazil, Silvério et al. [40] mention that the transition from forests to crop leads to an increase in ST in the order of 0.3 K. In addition to the increase in ST observed in view of the conversion of forests into anthropic uses between 2007 and 2016, an increase in the thermal average of the remaining forests was observed in 2016 (Fig 4). Therefore, it is believed that this may be a function of the horizontal transfer of sensible heat fluxes from the land use (pastureland/crop) (which showed higher values of ST) through advection towards the remaining TDFs in 2016. Once forests register lower values of ST, a reestablishment of the thermal balance in the face of the insertion of land use with high  $\Delta$ ST takes place.

In terms of implications, it is understood that the increase in ST due to the conversion of forests into anthropic uses can influence several environmental segments, especially regarding the provision of ecosystem services. ST is a proxy variable for several ecosystem services in forest areas. As forests play an intrinsic role in climate dynamics, mainly in terms of exchanges of energy flows with the atmosphere, such as latent heat and evapotranspiration [45, 46], increases in ST reflect substantial decreases in moisture levels in the surface, as it has an inversely proportional relationship with latent heat and evapotranspiration [47]. Thus, several studies that showed an increase in ST due to the conversion of forests into pastureland and crop, pointed to changes in the standards of latent heat and evapotranspiration. For example, Silvério et al. [40] recorded decreases in latent heat and evapotranspiration in the order of 32% and 35 km<sup>3</sup> (in areas of Brazil), while Lathuilliere et al. [46] observed that there was a decrease in a rate of 16.2 km<sup>3</sup> for evapotranspiration in areas on the western edge of Brazil. For the same region, Pongratz et al. [44] reported a 21% decrease in latent heat, while Sampaio et al. [45], in forest areas in northern Brazil, found that pastureland areas tend to reduce evapotranspiration by ~26%.

Still, in the context of ecosystem service provision, forests have a high potential for carbon (CO<sub>2</sub>) sequestration through net primary productivity [48, 49]. On a global scale, tropical forests store around 40% of terrestrial carbon [50], proving to be a crucial component in mitigating the effects of climate change. Specifically for TDFs, which cover ~42% of tropical forests [51], studies have found that they have high carbon stocks due to assimilation by photosynthetic processes [25]. However, as photosynthetic processes are closely linked to evapotranspiration (especially net primary productivity [52], changes in evapotranspiration can compromise the ability to assimilate atmospheric CO<sub>2</sub>. The exchange of flows between the forests and the atmosphere occurs in an ascending and descending manner so that CO<sub>2</sub> fixation is carried out, it is necessary for the diffusion of water vapor through evapotranspiration to occur [52]. Evapotranspiration and net primary productivity have proportional behavior, therefore, as ST increases in TDFs, water vapor fluxes can decline, subsequently tending to decrease the carbon fixation capacity.



In general, we found that human actions had impacts in the context of changes (increases) in surface temperature in the domains of TDFs in the north of the state of Minas Gerais. However, although TDFs are protected by specific legislation, there is a need for extensive monitoring of forests, especially considering climatic aspects, as they are crucial for the provision of ecosystem services at different scales of approach.

## Conclusions

We found that the conversion of dry forests to pastureland and crop increased the temperature (1.55 K) in the study area (within 9 years). There was a loss of ~4,000 km<sup>2</sup> (22%) of dry forests, most of which (93%) were explained by pastureland. Pastureland were also responsible for the highest values of surface temperature in 2016 (2.21 K).

Surface temperature is a crucial variable in environmental dynamics in the tropical dry forest. Therefore, considering the importance of tropical dry forests for ecosystem segments, we emphasize the need for efficient monitoring of changes in these environments, mainly by including scientific studies on climate variables in the action plans.

## Author Contributions

**Conceptualization:** Lucas Augusto Pereira da Silva, Andre Medeiros Rocha.

**Formal analysis:** Lucas Augusto Pereira da Silva.

**Investigation:** Lucas Augusto Pereira da Silva, Andre Medeiros Rocha.

**Methodology:** Lucas Augusto Pereira da Silva, Andre Medeiros Rocha.

**Supervision:** Claudionor Ribeiro da Silva.

**Writing – original draft:** Lucas Augusto Pereira da Silva, Claudionor Ribeiro da Silva.

**Writing – review & editing:** Lucas Augusto Pereira da Silva, Andre Medeiros Rocha, Claudionor Ribeiro da Silva.

## References

1. Miles L, Newton AC, DeFries RS, Ravilious C, May I, Blyth S, et al. A global overview of the conservation status of tropical dry forests. *Journal of biogeography*, v. 33, n. 3, p. 491–505, 2006. Available from: <https://doi.org/10.1111/j.1365-2699.2005.01424.x>
2. Dinerstein E, Olson D, Joshi A, Vynne C, Burgess ND, Wikramanayake E, et al. An ecoregion-based approach to protecting half the terrestrial realm. *BioScience*, v. 67, n. 6, p. 534–545, 2017. Available from: <https://doi.org/10.1093/biosci/bix014> PMID: 28608869
3. Janzen DH. Tropical dry forest. The most endangered major tropical ecosystem. 1988. Biodiversity. National Academic Press. Washington, DC, p. 130–137.
4. Schröder JM, Rodriguez LPA, Günter S. Research trends: Tropical dry forests: The neglected research agenda?. *Forest Policy and Economics*, v. 122, p. 102333, 2021. Available from: <https://doi.org/10.1016/j.forpol.2020.102333>.
5. Oliveira-Filho AT, Jarenkow JA, Rodal MJN. Floristic Relationships of Seasonally Dry Forests of Eastern South America Based on Tree Species Distribution Patterns. *In: Pennington RT, Lewis GP, Ratter JA. Neotropical Savannas and Seasonally Dry Forests*. Boca Raton: CRC/Taylor & Francis, 2006. p. 159–192.
6. Ratter JA, Ribeiro JF, Bridgewater S. The Brazilian cerrado vegetation and threats to its biodiversity. *Annals of botany*, v. 80, n. 3, p. 223–230, 1997. Available from: <https://doi.org/10.1006/anbo.1997.0469>.
7. Pennington RT, Lewis GP, Ratter JA. An Overview of Plant Diversity, Biogeography and Conservation of Neotropical Savannas and Seasonally Dry Forests. *In: Pennington RT, Lewis GP, Ratter JA. Neotropical Savannas and Seasonally Dry Forests*. Boca Raton: CRC/Taylor & Francis, 2006. p. 1–29.

8. Portillo-Quintero C, Sanchez-Azofeifa A, Calvo-Alvarado J, Quesada M, do Espírito-Santo MM. The role of tropical dry forests for biodiversity, carbon and water conservation in the neotropics: lessons learned and opportunities for its sustainable management. *Regional Environmental Change*, v. 15, n. 6, p. 1039–1049, 2015. Available from: <https://doi.org/10.1007/s10113-014-0689-6>.
9. Gibbs HK, Brown S, Niles JO, Foley JA. Monitoring and estimating tropical forest carbon stocks: making REDD a reality. *Environmental research letters*, v. 2, n. 4, p. 045023, 2007. Available from: <http://dx.doi.org/10.1088/1748-9326/2/4/045023>.
10. Keenan TF, Williams CA. The terrestrial carbon sink. *Annual Review of Environment and Resources*, v. 43, p. 219–243, 2018. Available from: <https://doi.org/10.1146/annurev-environ-102017-030204>.
11. Bonan G. *Ecological climatology: concepts and applications*. Cambridge University Press, 2015.
12. Anderson RG, Canadell JG, Randerson JT, Jackson RB, Hungate BA, Baldocchi DD, et al. Biophysical considerations in forestry for climate protection. *Frontiers in Ecology and the Environment*, v. 9, n. 3, p. 174–182, 2011. Available from: <https://doi.org/10.1890/090179>.
13. Houspanossian J, Noretto M, Jobbágy EG. Radiation budget changes with dry forest clearing in temperate Argentina. *Global change biology*, v. 19, n. 4, p. 1211–1222, 2013. Available from: <https://doi.org/10.1111/gcb.12121> PMID: 23504897
14. Duveiller G, Hooker J, Cescatti A. The mark of vegetation change on Earth's surface energy balance. *Nature communications*, v. 9, n. 1, p. 1–12, 2018. Available from: <https://doi.org/10.1038/s41467-017-02810-8>
15. Barry RG, Chorley RJ. *Atmosphere, weather and climate*. Routledge, 2009.
16. Cao S, Sanchez-Azofeifa A. Modeling seasonal surface temperature variations in secondary tropical dry forests. *International journal of applied earth observation and geoinformation*, v. 62, p. 122–134, 2017. Available from: <https://doi.org/10.1016/j.jag.2017.06.008>.
17. Wolosin M, Harris N. *Tropical Forests and Climate Change: The Latest Science*. World Resources Institute. Washington DC, 2018.
18. Lal R. Carbon sequestration. *Philosophical Transactions of the Royal Society B: Biological Sciences*, v. 363, n. 1492, p. 815–830, 2008. Available from: <https://doi.org/10.1098/rstb.2007.2185> PMID: 17761468
19. FAO. 2020. *Global Forest Resources Assessment 2020: Main report*. Rome. Available from: <https://doi.org/10.4060/ca9825en>
20. Wu J, Chen B, Reynolds G, Xie J, Liang S, O'Brien MJ, et al. Monitoring tropical forest degradation and restoration with satellite remote sensing: A test using Sabah Biodiversity Experiment. *Advances in Ecological Research*, v. 62, p. 117–146, 2020. Available from: <https://doi.org/10.1016/bs.aecr.2020.01.005>.
21. Gao L, Wang X, Johnson BA, Tian Q, Wang Y, Verrelst J, et al. Remote sensing algorithms for estimation of fractional vegetation cover using pure vegetation index values: A review. *ISPRS Journal of Photogrammetry and Remote Sensing*, v. 159, p. 364–377, 2020. Available: <https://doi.org/10.1016/j.isprsjprs.2019.11.018>.
22. Liang Y, Liu L, Huang J. Vegetation Change Detection Using Trend Analysis and Remote Sensing. In: *Integrated Modelling of Ecosystem Services and Land-Use Change*. Springer, Singapore, 2019. p. 39–57. Available from: [https://doi.org/10.1007/978-981-13-9125-5\\_3](https://doi.org/10.1007/978-981-13-9125-5_3).
23. Dupin MG, Espírito-Santo MM, Leite ME, Silva JO, Rocha AM, Barbosa RS, et al. Land use policies and deforestation in Brazilian tropical dry forests between 2000 and 2015. *Environmental Research Letters*, v. 13, n. 3, p. 035008, 2018. Available from: <https://doi.org/10.1088/1748-9326/aaadea>.
24. Rocha AM, Leite ME, do Espírito-Santo MM. Monitoring of Brazilian Seasonally Dry Tropical Forest by Remote Sensing. *Mercator (Fortaleza)*, v. 19, 2020. Available from: <https://doi.org/10.4215/rm2020.e19022>.
25. Castro-Magnani M, Sanchez-Azofeifa A, Metternicht G, Laakso K. Integration of remote-sensing based metrics and econometric models to assess the socio-economic contributions of carbon sequestration in unmanaged tropical dry forests. *Environmental and Sustainability Indicators*, v. 9, p. 100100, 2021. Available from: <https://doi.org/10.1016/j.indic.2021.100100>.
26. Clemente CMS, do Espírito-Santo MM, Leite ME. Estimates of deforestation avoided by protected areas: a case study in Brazilian tropical dry forests and Cerrado. *Landscape Research*, p. 1–14, 2020. Available from: <https://doi.org/10.1080/01426397.2020.1730773>.
27. Ab'Saber AN. *Geomorfologia-Domínios morfoclimáticos do Brasil*. São Paulo: Universidade de São Paulo/Instituto de Geografia, 1970. Portuguese.
28. UFV CETEC, UFLA FEAM. *Mapa de solos do Estado de Minas Gerais*. Universidade Federal de Viçosa; Fundação Centro Tecnológico de Minas Gerais; Universidade Federal de Lavras; Fundação

- Estadual do Meio Ambiente, p. 49, 2010. Available from: [http://www.dps.ufv.br/?page\\_id=742](http://www.dps.ufv.br/?page_id=742). Portuguese.
29. Didan K, Munoz AB, Ramon S, Huete A. MODIS vegetation index user's guide (MOD13 series). University of Arizona: Vegetation Index and Phenology Lab, 2015.
  30. Breiman L, Friedman JH, Olshen RA, Stone CJ. Classification and regression trees. CRC press, 1984.
  31. Wan Z. Collection-6 MODIS land surface temperature products users' guide, ERI. 2013.
  32. Wang K, Li Z, Cribb M. Estimation of evaporative fraction from a combination of day and night land surface temperatures and NDVI: A new method to determine the Priestley–Taylor parameter. *Remote Sensing of Environment*, v. 102, n. 3–4, p. 293–305, 2006. Available from: <https://doi.org/10.1016/j.rse.2006.02.007>.
  33. QGIS Development Team. "QGIS Geographic Information System. Open Source Geospatial Foundation Project". 2021. Available: <http://qgis.osgeo.org>.
  34. Wickham H. GGLOT2: elegant graphics for data analysis. 2016. Available from: <https://ggplot2.tidyverse.org/>.
  35. R Core Team. R: A language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria. 2021. Available from: <https://www.R-project.org/>.
  36. Espírito-Santo MM do Rocha AM, Leite ME Silva JO, Silva LAP Sanchez-Azofeifa GA. Biophysical and socioeconomic factors associated to deforestation and forest recovery in Brazilian tropical dry forests. *Frontiers in Forests and Global Change*, v. 3, p. 141, 2020. Available from: <https://doi.org/10.3389/ffgc.2020.569184>.
  37. Stan K, Sanchez-Azofeifa A. Tropical dry forest diversity, climatic response, and resilience in a changing climate. *Forests*, v. 10, n. 5, p. 443, 2019. Available from: <https://doi.org/10.3390/f10050443>.
  38. Marzahn P, Flade L, Sanchez-Azofeifa A. Spatial estimation of the latent heat flux in a tropical dry forest by using unmanned aerial vehicles. *Forests*, v. 11, n. 6, p. 604, 2020. Available: <https://doi.org/10.3390/f11060604>.
  39. Zhou X, Wang YC. Dynamics of land surface temperature in response to land-use/cover change. *Geographical Research*, v. 49, n. 1, p. 23–36, 2011. Available from: <https://doi.org/10.1111/j.1745-5871.2010.00686.x>.
  40. Silvério DV, Brando PM, Macedo MN, Beck PS, Bustamante M, Coe MT. Agricultural expansion dominates climate changes in southeastern Amazonia: the overlooked non-GHG forcing. *Environmental Research Letters*, v. 10, n. 10, p. 104015, 2015. Available from: <http://dx.doi.org/10.1088/1748-9326/10/10/104015>.
  41. Santos CACD Wanderley RLN, Araújo AL Bezerra BG. Obtenção do saldo de radiação em áreas de pastagem e floresta na Amazônia (estação seca) através do sensor MODIS. *Revista brasileira de meteorologia*, v. 29, n. 3, p. 420–432, 2014. Available: <http://dx.doi.org/10.1590/0102-778620130591>. Portuguese.
  42. Pavão VM, Querino CAS, Beneditti CA, Pavão LL, da Silva Querino JKA, Machado NG, et al. Variação espacial e temporal do saldo de radiação superficial em uma área do sul do Amazonas, Brasil. *Raega-O Espaço Geográfico em Análise*, v. 37, p. 333–352, 2016. Available: <http://dx.doi.org/10.5380/raega.v37i0.42469>. Portuguese.
  43. Querino CAS, Beneditti CA, Machado NG, da Silva MJG, da Silva Querino JKA, dos Santos Neto LA, et al. Spatiotemporal NDVI, LAI, albedo, and surface temperature dynamics in the southwest of the Brazilian Amazon forest. *Journal of Applied Remote Sensing*, v. 10, n. 2, p. 026007, 2016. Available from: <https://doi.org/10.1117/1.JRS.10.026007>.
  44. Pongratz J, Bounoua L, DeFries RS, Morton DC, Anderson LO, Mauser W, et al. The impact of land cover change on surface energy and water balance in Mato Grosso, Brazil. *Earth Interactions*, v. 10, n. 19, p. 1–17, 2006. Available from: <https://doi.org/10.1175/EI1176.1>.
  45. Sampaio G, Nobre C, Costa MH, Satyamurty P, Soares-Filho BS, Cardoso M. Regional climate change over eastern Amazonia caused by pasture and soybean cropland expansion. *Geophysical Research Letters*, v. 34, n. 17, 2007. Available from: <https://doi.org/10.1029/2007GL030612>.
  46. Lathuillière MJ, Johnson MS, Donner SD. Water use by terrestrial ecosystems: temporal variability in rainforest and agricultural contributions to evapotranspiration in Mato Grosso, Brazil. *Environmental Research Letters*, v. 7, n. 2, p. 024024, 2012. Available from: <http://dx.doi.org/10.1088/1748-9326/7/2/024024>.
  47. Liu Z, Ballantyne AP, Cooper LA. Biophysical feedback of global forest fires on surface temperature. *Nature communications*, v. 10, n. 1, p. 1–9, 2019. Available from: <https://doi.org/10.1038/s41467-018-08237-z>

48. Dai Z, Birdsey RA, Johnson KD, Dupuy JM, Hernandez-Stefanoni JL, Richardson K. Modeling carbon stocks in a secondary tropical dry forest in the Yucatan Peninsula, Mexico. *Water, Air, & Soil Pollution*, v. 225, n. 4, p. 1–15, 2014. Available from: <http://dx.doi.org/10.1007/s11270-014-1925-x>.
49. Cao S, Sanchez-Azofeifa GA, Duran SM, Calvo-Rodriguez S. Estimation of aboveground net primary productivity in secondary tropical dry forests using the Carnegie–Ames–Stanford approach (CASA) model. *Environmental Research Letters*, v. 11, n. 7, p. 075004, 2016. Available from: <http://dx.doi.org/10.1088/1748-9326/11/7/075004>.
50. Lewis SL, Lopez-Gonzalez G, Sonké B, Affum-Baffoe K, Baker TR, Ojo LO, et al. Increasing carbon storage in intact African tropical forests. *Nature*, v. 457, n. 7232, p. 1003–1006, 2009. Available from: <https://doi.org/10.1038/nature07771> PMID: 19225523
51. Murphy PG, Lugo AE. Ecology of tropical dry forest. *Annual review of ecology and systematics*, v. 17, n. 1, p. 67–88, 1986.
52. Paul-Limoges E, Wolf S, Schneider FD, Longo M, Moorcroft P, Gharun M, et al. Partitioning evapotranspiration with concurrent eddy covariance measurements in a mixed forest. *Agricultural and Forest Meteorology*, v. 280, p. 107786, 2020. Available from: <https://doi.org/10.1016/j.agrformet.2019.107786>.