RSC Advances



PAPER



Cite this: RSC Adv., 2023, 13, 13303

Received 13th April 2023 Accepted 24th April 2023

DOI: 10.1039/d3ra02472e

rsc.li/rsc-advances

1 Introduction

Magnetite (Fe_3O_4) is a common mineral in different rock types and various metallic deposits. Textural and compositional features of magnetite have been widely employed as petrogenetic indicators and mineral exploration tools, as they are sensitive to the relevant physicochemical conditions.¹⁻⁴ In hightemperature igneous systems and some hydrothermal deposits, sub-solidus exsolution can modify the chemical composition of the magnetite.^{1,5} Titanomagnetite and the intergrowth of ilmenite lamellae within magnetite have been extensively investigated, owing to their widespread distribution in different geological environments.⁵⁻⁷ The importance of this Fe–Ti oxide

Precise determination of major and trace elements in micrometer-scale ilmenite lamellae in titanomagnetite using LA-ICP-MS technique: application of regression analysis to time-resolved signals⁺

Si-Qi Liu,^a Shao-Yong Jiang, ^b*^a Wei Chen, ^b^a Christina Yan Wang,^b Hui-Min Su,^a Yonghua Cao,^b Hao-Xiang Zhang^a and Wen-Tian Li^a

Laser ablation ICP-MS (LA-ICP-MS) is a powerful microbeam technique capable of rapid and precise determination for a large spectrum of trace elements at ppm or sub-ppm levels. Micrometer-scale minerals and inclusions are very common in geological materials, for which direct measurement is restricted by the spot size using LA-ICP-MS (generally 20–50 μ m). In this study, ilmenite lamellae intergrown with magnetite were selected as an example to describe a practical algorithm that applies regression analysis to extract the chemical compositions of binary phases from mixed LA-ICP-MS signals. The method accuracy is confirmed by the agreement between the regressed value for various trace elements in ilmenite exsolutions and their reference values (direct analyses using EPMA and LA-ICP-MS). Results were obtained for most detectable components (Mg, Mn, V, Nb, Ta, Sc, Zr, Hf, Sn, *et c.*) and their relative deviations are within \pm 10%, even for those <10 ppm (such as Hf and W). Relative standard errors on the regressed value were calculated to evaluate the precision of the method, which is mostly within 10%, and the worst up to 25%. Therefore, the algorithm described in this contribution provides a solution for precise determination of trace element compositions for micrometer-scale ilmenite lamellae in titanomagnetite using LA-ICP-MS, and is potentially practical for other geological materials.

system has long been recognized, its thermodynamic mechanism was established and widely applied as a thermaloxybarometer.^{6,8}

However, concentrations for trace elements in the ilmenite lamellae are poorly known. Precise determination of trace elements in these ilmenite exsolutions is hindered by their limited width. Generally, ilmenite exsolutions occur as regular trellis lamellae within the magnetite domains/grains (Fig. 1), <1 to 10 µm in width. Traditional electron probe microanalyses (EPMA) are limited by their relatively high detection limits (commonly hundreds of ppm^{1,9}) despite good spatial resolution. Theoretically, nano-SIMS is capable of in situ trace element analyses on ilmenite exsolution, however, it was rarely applied in previous studies due to its high cost and rigorous matrix effects. Laser ablation ICP-MS (LA-ICP-MS) is a more common technique, and accessible to most laboratories.^{10,11} The LA-ICP-MS allows in situ analysis of a large range of trace elements with lower detection limits at ppm to sub-ppm levels,^{4,10,12} and has been widely applied for studies on magnetite geochemistry.^{2,3,7,13,14} Conventional trace element analyses using LA-ICP-MS employ round laser spots with 20-50 µm diameters. However, ilmenite lamellae in most titanomagnetite grains

[&]quot;State Key Laboratory of Geological Processes and Mineral Resources, Collaborative Innovation Center for Exploration of Strategic Mineral Resources, School of Earth Resources, China University of Geosciences, Wuhan, 430074, P. R. China. E-mail: shyjiang@cug.edu.cn

^bCAS Key Laboratory of Mineralogy and Metallogeny, Guangdong Provincial Key Laboratory of Mineral Physics and Materials, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, Guangzhou, 510640, China

[†] Electronic supplementary information (ESI) available. See DOI: https://doi.org/10.1039/d3ra02472e

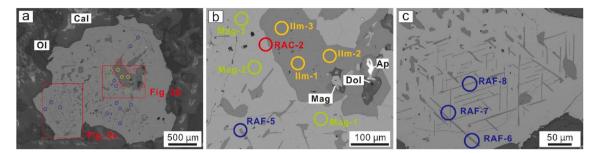


Fig. 1 Back-scattered electron (BSE) images of the selected titanomagnetite grain, in which both granular and lamellae ilmenite are available. (a) A panorama of the titanomagnetite grain, coexisting with calcite and olivine, and the distribution of the LA spots are shown. (b) Positions of direct analyses on magnetite domain (Mag, n = 3), coarse granular ilmenite exsolutions (IIm, n = 3), one of the mixed-sampling analyses for coarse ilmenite (Spot_RAC-2), and one of those for fine ilmenite lamellae (Spot_RAF-5). (c) Textures of micrometer ilmenite lamellae and positions of three spots for regression analysis. Abbreviations: Ap = apatite; Cal = calcite; Dol = dolomite; Ilm = ilmenite; Mag = magnetite; Ol = olivine.

from natural samples are rarely >10 μm in width, and much smaller than the laser pits.

Analyses using smaller laser spots would suffer from diminished instrumental sensitivity and increased downhole elemental fractionation.^{15,16} Alternatively, a 'bulk' analytical strategy has been applied in many studies, in which larger laser spots (50-120 µm) are employed to incorporate both magnetite and ilmenite lamellae, then the obtained compositions are considered representative of the homogenous oxide solid solution prior to the oxidation-driven exsolution.^{2,7,17,18} Analytical procedures using raster mode and lines across the magnetite grain were performed for the same purpose.4,19,20 Obviously, applying 'bulk' analyses, the chemical differentiation of various trace components between the exsolved ilmenite and residual magnetite is obscured. Many trace components in Fe-Ti oxides are temperature- and/or oxygen fugacity-sensitive (including V,²¹ Mn,²² Sn,²³ summarized in ref. 2 and 3). Their re-distribution during the oxy-exsolution process that accompanies cooling is of petrogenetic importance and should not be ignored.

Regression analysis is a statistical approach for evaluating the correlations between variables, and previously has been applied to LA-ICP-MS analyses for the extraction of mineralmatrix mixtures.^{24–26} To estimate the compositions of small zircon grains in a silicate matrix, Rubatto and Hermann set regression equations relative to Zr, and the specific Zr contents were obtained by EPMA (ref. 24). Similarly, to investigate the composition of experimental monazite grains in silicate glass, Stepanov *et al.*²⁵ conducted regression analysis selecting Ce as the explanatory variable, and Ce concentrations obtained by EPMA served as the specific values. Regression analysis was also successfully applied to natural materials, such as fine-grained pyrite in black shales, in which sulfur was used as the explanatory variable and the regressed values were normalized with 100% total requirements (ref. 26).

In this contribution, we describe and test a practical approach to quantifying chemical compositions of micrometerscale ilmenite lamellae in titanomagnetite. Through the application of regression analysis to LA-ICP-MS signals from mixed magnetite–ilmenite composition, concentrations for various components in those ilmenites were calculated with high precision and accuracy (mostly within 10%).

2 Experimental

2.1 Samples

As an example, titanomagnetite grains from calcite carbonatite at the Caotan iron deposit in the lesser Qinling area, China (location and geological descriptions are available in ref. 27 and 28), were selected to illustrate the mathematical algorithm applied to data interpretations. Analyzed samples were prepared as polished thin sections. The morphology and textures of magnetite were investigated using back-scattered electron (BSE) images obtained by a JEOL JCM-7000 scanning electron microscope (SEM), which is fitted with an energy dispersive spectrometry (EDS), at the Collaborative Innovation Center for Exploration of Strategic Mineral Resources, China University of Geosciences (Wuhan). A single magnetite grain was selected for EPMA and LA-ICP-MS analyses, as both coarse granular exsolutions and micrometer-wide ilmenite lamellae are found inside (Fig. 1).

The synthetic basaltic glass GSE-1G, together with USGS basaltic reference glasses BCR-2G, BHVO-2G, and BIR-1G were used as microanalytical reference materials applied in this study.²⁹ The working values for GSE-1G and the USGS reference glasses are from the GeoReM database (http://georem.mpch-mainz.gwdg.de/).

2.2 EPMA

Prior to LA-ICP-MS analyses, major and minor components in the magnetite grain (n = 6), coarse granular exsolutions (n = 2), and micrometer ilmenite lamellae (n = 4) were determined using EPMA. Experiments were carried out using a JEOL JXA-8230 Electron Probe Microanalyzer (EPMA) equipped with five wavelength-dispersive spectrometers (WDS) together with one energy-dispersive spectrometer (EDS) at the Laboratory of Microscopy and Microanalysis, Wuhan Microbeam Analysis Technology Co., Ltd. Iron, Ti, Mg, Mn, V, Si, Ca, Al, Cr, and Nb were analyzed using a focused beam (1 μ m) at 20 kV and 20 nA. The peak overlap of Ti K β on V K α and V K β on Cr K α was corrected, and the correction strategy followed Yang *et al.* (ref. 9). Detailed operating conditions and involved standards for each element are available in ESI Table 1.[†]

2.3 LA-ICP-MS

Trace element analyses were conducted using a single quadrupole ICP mass spectrometer (Agilent 7900, Agilent Technologies, USA) coupled with a 193 nm ArF excimer laser ablation system (ESL 193HE, Elemental Scientific Lasers, USA), located at the Collaborative Innovation Center for Exploration of Strategic Mineral Resources, China University of Geosciences (Wuhan). The carrier gas, helium, was passed through the ablation cell and mixed with argon downstream of the ablation cell. Daily optimization of instrumental performance with NIST SRM 610 involved maximizing the ²³⁸U signal intensity relative to the background, while satisfying low oxide production rates (ThO⁺/ Th⁺ <0.5%), low double-charged ions (Ca^{2+}/Ca^{+} <1.0%), and robust plasma conditions (U^+/Th^+ in a range of 0.95–1.05). The pulse/analog (P/A) factor of the detector was calibrated using a standard tuning solution. A repetition rate of 6 Hz was used together with a spot size of 32 μ m and an energy density of \sim 3.5 J cm⁻². Each spot analysis included a 20 s background acquisition for a gas blank, followed by a 40 s data acquisition from the sample. A total of 57 elements were analyzed: ⁷Li, ⁹Be, ¹¹B, ²³Na, ²⁵Mg, ²⁷Al, ²⁹Si, ³¹P, ³⁹K, ⁴²Ca, ⁴⁵Sc, ⁴⁷Ti, ⁵¹V, ⁵³Cr, ⁵⁵Mn, ⁵⁷Fe, ⁵⁹Co, ⁶⁰Ni, ⁶³Cu, ⁶⁶Zn, ⁷¹Ga, ⁷³Ge, ⁷⁵As, ⁸⁵Rb, ⁸⁸Sr, ⁸⁹Y, ⁹⁰Zr, ⁹³Nb, ⁹⁵Mo, ¹⁰⁷Ag, ¹¹¹Cd, ¹¹⁵In, ¹¹⁸Sn, ¹²¹Sb, ¹³³Cs, ¹³⁷Ba, ¹³⁹La, ¹⁴⁰Ce, ¹⁴¹Pr, ¹⁴⁶Nd, ¹⁴⁷Sm, ¹⁵³Eu, ¹⁵⁷Gd, ¹⁵⁹Tb, ¹⁶³Dy, ¹⁶⁵Ho, ¹⁶⁶Er, ¹⁶⁹Tm, ¹⁷³Yb, ¹⁷⁵Lu, ¹⁷⁸Hf, ¹⁸¹Ta, ¹⁸²W, ²⁰⁸Pb, ²⁰⁹Bi, ²³²Th, and ²³⁸U. The dwell time is 6 ms for each analyzed isotope.

The measurements of quality control (QC) reference material (GSE-1G, in this study) are performed to correct the instrumental time-dependent sensitivity drift. Multiple external standards (GSE-1G, BCR-2G, BIR-1G, BHVO-2G) were used for signal calibration, and the detailed quantification strategy is given in Section 3.2 below. Analyses on sets of external reference materials were performed at the beginning and end of the batch and between every 10 analyses on the samples. Each reference material was analyzed once in a set. To illustrate the regression algorithm and verify its accuracy, measurements for the magnetite grain (Mag, n = 3) and coarse granular exsolutions (Ilm, n = 3) were conducted together with mixed ablation involved in later regression analysis. Mixed ablations were performed at the contact between coarse granular ilmenite and the magnetite domain (regression analysis for coarse exsolutions, RAC, n = 8), and areas where the micrometer ilmenite lamellae are dense (regression analysis for fine exsolutions, RAF, n = 12).

3 Algorithm

The extraction of the compositions of ilmenite lamellae from mixed time-resolved signals involves segmentation of the signal, primary quantification, linear regression analysis of the data, calculation of element concentrations, and estimation of uncertainties. Details for each step are described below, followed by a summary of the data reduction procedure.

Offline data reduction was performed with ICPMSDataCal 10.9,¹⁰ including integration selection of background and analyte signals, time drift correction, and quantitative calibration. The time-resolved signals were carefully checked (Fig. 2). Signal peaks in the background and sample durations were filtered following the 2σ rule. Magnetite free of inclusions yields stable signals (Fig. 2a), whereas synchronous increases in Ti and several other components are indicative of 'contamination' from ilmenite lamellae (Fig. 2b–d).

3.1 Segmentation of the signal

As illustrated by the time-solved drifts (Fig. 2b–d), LA-ICP-MS analyses of titanomagnetite grains with ilmenite lamellae generate highly variable signals. Considering the time for the ablated materials to leave the chamber, reliable sample signals in each spot analysis are actually slightly more than 35 seconds (Fig. 2). For each spot analysis, six short but identical segments were cut from the acquired sample signals and employed for signal quantification. To make better use of the sample signals, six segments overlap as follows (Fig. 2b–d): segment 1 (0–10 s), segment 2 (5–15 s), segment 3 (10–20 s), segment 4 (15–25 s), segment 5 (20–30 s), segment 6 (25–35 s).

Manual segmentations were performed in the 'signal selection' window using the software ICPMSDataCal 10.9. The integration intervals of sample signals were fixed, thus the intervals for all reference materials and unknowns were set the same. The integration interval length of sample signals for each segment was set as 10 seconds, and the starting points for six timeprogressive segments were set at 0 s, 5 s, 10 s, 15 s, 20 s, and 25 s, respectively. To ensure objectivity and avert subjective selections, no extra action was performed during the segmentation.

3.2 Primary quantification and normalization of the analyses

Magnetite (Fe₃O₄) is a volatile-free mineral but contains iron with variable valence states. The three Fe ions in its stoichiometric formula are composed of one Fe²⁺ and two Fe³⁺. Quantitative calibration of segmented signals and direct analyses (magnetite, coarse granular ilmenite, and basaltic glasses) was conducted using the 'total metal oxide' of 100%, applying Fe as the normalizing element but without assuming Fe content as the internal standard. The ratio Fe²⁺/total Fe was set as 0.33 for magnetite, 0.50 for BIR-1G, 0.75 for BCR-2G and BHVO-2G, and 1.0 for ilmenite. For segmented analyses, the Fe²⁺/total Fe ratio was set as 0.50, taking into account the influence of contaminations from ilmenite (Fe²⁺/total Fe = 1) on the exact proportions of ferrous iron in total iron of the ablated materials.

Calibration for each segregated interval followed the MRMC-AYC strategy (multiple reference materials calibration coupled with the ablation yield correction, described in ref. 10). Multiple external reference materials (GSE-1G, BCR-2G, BIR-1G, BHVO-2G) were used for external calibration. Iron for each analysis is chosen as the normalizing element to correct mass-

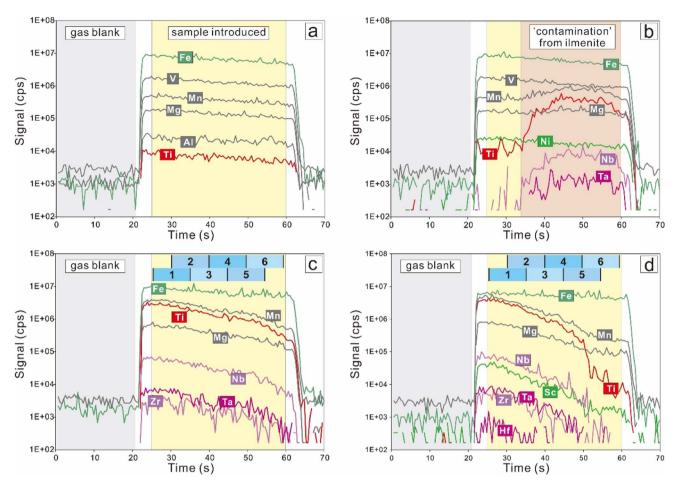


Fig. 2 Examples of time-resolved LA-ICP-MS signals. (a) Stable signals from direct analysis on magnetite (Spot_Mag-1). (b) Variable signals from analysis of relatively clean magnetite on the surface (Spot_RAF-12), and 'contaminations' from ilmenite were figured by synchronous variations of Ti and several other components. (c) Mixed-sampling analysis for micrometer ilmenite exsolutions with decreasing proportions of ilmenite (Spot_RAF-5). (d) Mixed-sampling analysis for coarse ilmenite (RAC-2).

dependent bias. The primary quantification involved two stages. Firstly, primary concentrations for each element were calculated against multiple external reference materials with Fe as the normalizing element:

$$\frac{pC_{\rm sam}{}^{i}}{pC_{\rm sam}{}^{\rm Fe}} = \left(\frac{cps_{\rm sam}{}^{i}}{cps_{\rm sam}{}^{\rm Fe}}\right) \times l^{i} \tag{1}$$

where $cps_{sam}{}^{i}$ is the net count rate (sample signal minus background, after time-dependent instrumental sensitivity drift correction) of element *i*, and $cps_{sam}{}^{Fe}$ is that of Fe. The $pC_{sam}{}^{i}$ is an intermediate variable for element *i*, and $pC_{sam}{}^{Fe}$ is that for Fe, shows the preliminarily-calibrated element concentrations. The l^{i} value represents the calibration factor for element *i* against external multiple reference materials, and was calculated based on the concentration-weighted average:

$$l^{i} = \sum_{s=1}^{n} \left(\left(\frac{cps_{\rm rms}^{i}}{cps_{\rm rms}^{i}s} \right) \times \left(\frac{C_{\rm rms}^{i}}{\sum\limits_{s=1}^{n} C_{\rm rms}^{i}} \right) \right)$$
(2)

where $cps_{rm^s}{}^i$ and $C_{rm^s}{}^{is}$ is the net count rate and concentration of analyte element *i* in reference material rm^s, and *n* is the number of involved reference materials.

The second stage quantification was performed by normalizing the sum of all element oxides to 100 wt% (ref. 10).

$$C_{\rm sam}{}^{i} = p C_{\rm sam}{}^{i} \times k \tag{3}$$

$$\sum_{i=1}^{N} C_{\rm sam}{}^{i} = 100 \text{ wt}\%$$
(4)

where pC_{sam}^{i} is the preliminarily-calibrated element *i* concentration calculated in the first stage. As a unique factor to correct different ablation yields for each analysis, the absolute value of *k* was named the ablation yield correction factor (AYCF, ref. 10), and was calculated as:

$$k \times \sum_{i=1}^{N} p C_{\text{sam}}^{i} = 100 \text{ wt\%}$$
 (5)

3.3 Regression analyses of the data

After the primary quantification of segmented signals, element concentrations for the six segments are available. Titanium contents are variable among different segments, and many components display obvious correlations with Ti contents

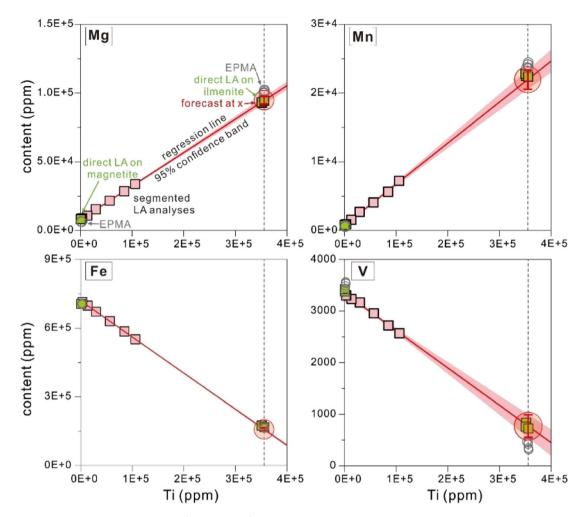


Fig. 3 Correlations between major components (Mg, Mn, Fe, V) and titanium in segmented analyses, and comparisons between regressed values and results from direct measurements. Quantified data from segmented signals were plotted and regressed in binary diagrams (as solid boxes filled with pink), where Ti content (independent variable) is on the horizontal axis and element *i* concentrations are shown on the vertical axis. In each diagram, the regression line is illustrated as red thick solid lines, and the vertical line for the specific independent variable (Ti content from EPMA) is drawn as dashed lines. The pink area on both sides of the regression line represents the confidence band with a confidence level of 95%. The intersection between the regression line and the vertical line for the specific independent variable represents the forecast value of element *i* content in ilmenite exsolutions. The bold red bar was restricted by the upper and lower intersections between the specific vertical line and the confidence band. Results from direct LA-ICP-MS analyses on magnetite and ilmenite are plotted in the diagram as green-filled squares, and compared with the forecast values from regression analysis. Measured values from EPMA are plotted as grey hollow circles in the diagram.

(Fig. 3 and 4). Because Ti is the diagnostic major component in ilmenite, once it is selected as the independent variable, the change of x (the independent variable) will directly reflect the proportions of ilmenite in the ablated materials. Moreover, Fe in the sampled materials has variable valence depending on the proportions of magnetite, which will become a new source of uncertainties. Therefore, Ti is selected as the independent variable in this study. Calculation of necessary statistical factors, quantification of the regression dependencies, and estimation of uncertainties (discussed later) were performed following the notes for the course on rational choice theory given by Robert F. Nau at Duke University (available online: https://people.duke.edu/~rnau/411home.htm, accessed November 2022). Visualization of the regression analysis could illustrate its feasibility in a more intuitive way, and was

performed using the software Origin 2021 in this study (Fig. 3 and 4).

The relation between the element *i* and Ti could be explained by a linear regression model:

$$C_i = A_i \times C_{\mathrm{Ti}} + B_i \tag{6}$$

where C_i is the content of element *i*, C_{Ti} is Ti content in the sampled material. A_i is the slope of the linear regression, B_i is the intercept, and they were calculated as:

$$A_i = R_i \times \frac{\text{St. dev.}(C_i)}{\text{St. dev.}(C_{\text{Ti}})}$$
(7)

$$B_i = \overline{C}_i - A_i \times \overline{C_{\mathrm{Ti}}} \tag{8}$$

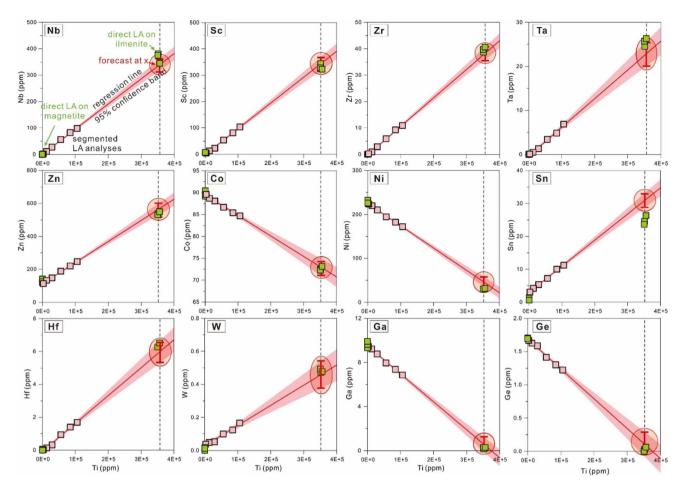


Fig. 4 Correlations between several trace elements and titanium in segmented analyses. Quantified data from segmented signals, regression line, confidence band, the specific value for the explanatory variable, and results from direct LA-ICP-MS analyses are plotted and drawn in the same way as Fig. 3.

where \overline{C}_i is the average element *i* content of the 6 segments, and \overline{C}_{Ti} is the average Ti content, St. dev. (C_i) and St. dev. (C_{Ti}) represent the standard deviation of concentration for the element *i* and Ti. The correlation of *i* contents and Ti contents acquired from segmented signals, R_i in eqn (7), was calculated as:

$$R_{i} = \operatorname{Correl}(C_{\mathrm{Ti}}, C_{i}) = \frac{\sum \left(C_{\mathrm{Ti}} - \overline{C_{\mathrm{Ti}}}\right) \left(C_{i} - \overline{C}_{i}\right)}{\sqrt{\sum \left(C_{\mathrm{Ti}} - \overline{C_{\mathrm{Ti}}}\right)^{2}} \sqrt{\sum \left(C_{i} - \overline{C}_{i}\right)^{2}}}$$
(9)

and the *R*-squared calculated as R_i^2 represents the percent of the variance explained by the model, thus commonly used to evaluate the applicability of the regression analysis.

3.4 Calculation of element concentrations from regression equations

The slope (A_i) and intercept (B_i) in the regression equation were computed from statistical mathematics, and element *i* concentration in ilmenite lamellae was calculated by applying a specified explanatory variable. The average of Ti contents from EPMA on ilmenite (355, 368 ppm, n = 6) was used as the specified independent variable for the regression forecast for concentrations of other components. Then, applying the specified C_{Ti} (Ti content obtained by EPMA, 355, 368 ppm), element *i* content (C_i) calculated from eqn (6) was accepted as that of ilmenite. As illustrated in Fig. 3 and 4, the intersection of the regression line and the specified vertical line (Ti = 355, 368 ppm) represents the chemical composition of ilmenite exsolutions, and the ordinate value of this point is taken as element *i* content (C_i) in ilmenite.

Since Ti content in magnetite (1540 \pm 363 ppm, EPMA results; 263 \pm 5.47 ppm, direct LA-ICP-MS results) is negligible compared to ilmenite, the longitudinal intercept (Ti = 0 ppm) calculated from regression analysis was approximated as element *i* content in magnetite.

3.5 Estimation of uncertainties

Following statistical formulas, statistical uncertainties resulting from regression analysis were calculated as the standard error on the mean:

$$R_{i_{\text{adjusted}}} = 1 - \frac{(1 - R_i^2)(n - 1)}{(n - 2)}$$
(10)

$$\operatorname{SER}_{i} = \sqrt{1 - R_{i_{\text{adjusted}}}^{2}} \times \operatorname{St.} \operatorname{dev.} (C_{i})$$
 (11)

$$VAR.P(C_{Ti}) = \frac{\sum \left(C_{Ti} - \overline{C_{Ti}}\right)^2}{n}$$
(12)

$$SE_{mean}(C_{Ti-EPMA}) = \frac{SER_i}{\sqrt{n}} \times \sqrt{1 + \frac{\left(C_{Ti} - \overline{C_{Ti}}\right)^2}{VAR.P(C_{Ti})}}$$
(13)

$$d_i = \text{SE}_{\text{mean}}(C_{\text{Ti-EPMA}}) \times t_{\alpha/2}(n-2)$$
(14)

adjusted *R*-squared ($R_{i_{adjusted}}$) could be used to determine the standard error of regression (SER_i), *n* is the sample size selected for regression analysis (n = 6, in this study), VAR.P(C_{Ti}) is the variance based on the entire population, and SE_{mean}($C_{Ti-EPMA}$) represents the standard error on the mean of element *i* content for a given value of Ti concentration (acquired by EPMA in this study), $t_{\alpha/2}(n - 2)$ is the two-tailed inverse of the Student's *t*-distribution when the confidence level for two-sided interval was set at α ($\alpha = 95\%$ here).

In Fig. 3 and 4, the pink area on either side of the regression line represents the 95% confidence band, and the distance between its upper or lower limit and the regression line represents the standard error on the mean.

3.6 Data reduction procedure

In this study, time-resolved spectra segmentation and primary quantification were implemented using the software ICPMS-DataCal 10.9 (ref. 10), through which element contents in ablated mixtures were obtained. Then, based on chemical variations among different segments, element *i* concentrations were calculated from regression analysis, while the acquired Ti contents served as the independent variable and Ti contents in ilmenite (EPMA results) were applied as the specified value. The statistical uncertainties were estimated as the standard error on the mean. Algebraic expressions including correlation coefficient, slope, intercept, and standard error of regression are calculated following mathematical formulas.

Analytical results for direct LA-ICP-MS analyses (magnetite, coarse granular ilmenite, and basaltic glasses) are presented in ESI Table 2.† For segmented signals, the spot noted RAC-2 was selected as the example of data reduction procedure (Fig. 2e and f), its segmented LA-ICP-MS results, regression analysis process, and comparisons with the reference values were shown in Fig. 3 and 4, and all arithmetical calculation is available in ESI Table 3.†

4 Results and discussion

4.1 Geochemistry of magnetite and ilmenite

Electron probe micro-analyses (EPMA) are commonly used to determine major and several minor components in magnetite, and LA-ICP-MS is applicable to trace element analyses. The comparison of the results obtained through the two techniques are discussed to confirm the feasibility of the primary quantification strategy for LA-ICP-MS analyses applied in this study. The full list of EPMA analytical results is given in ESI Table 1.† Iron (70.5–71.6 wt%), Ti (1070–2130 ppm), V (3280–3570 ppm), Mg (6360–6760 ppm), and Mn (540–760 ppm) in magnetite yield values above their detection limits. Coarse granular ilmenite and micrometer ilmenite lamellae are compositionally homogeneous in concentrations for Fe (16.2–17.5 wt%), Ti (35.3–35.6 wt%), Mg (9.6–10.2 wt%), Mn (540–760 ppm), V (313–462 ppm), and Nb (126–419 ppm, close to or slightly higher than the detection limits at about 128 ppm).

Since magnetite and ilmenite are both iron-rich oxides, four basaltic glasses (GSE-1G, BCR-2G, BIR-1G, BHVO-2G) containing similarly high Fe contents are used as the external standards for compositional calibration, in order to keep matrix-matching as close as possible. Chen *et al.*⁴ have confirmed the feasibility

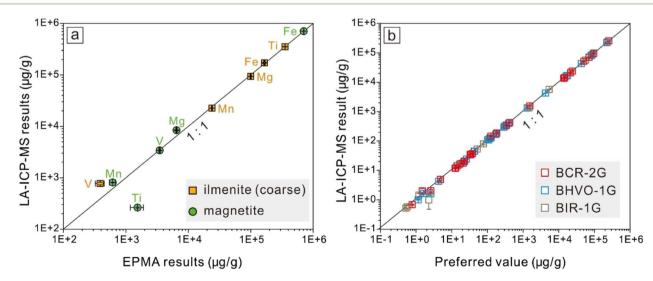


Fig. 5 (a) Cross-comparison of the EPMA and LA-ICP-MS analytical data for Fe, Ti, Mg, Mn, V in magnetite and ilmenite. (b) Comparison of LA-ICP-MS analytical results for the international reference materials (BCR-2G, BHVO-1G, and BIR-2G) with their preferred values (from the GeoReM database, http://georem.mpch-mainz.gwdg.de/). The 1:1 line is provided for comparison.

Table 1 Summary of analytical results (µg g⁻¹) for major and trace elements of coarse granular ilmenite exsolutions and the magnetite domain

	Ilmenite				Magnetite			
	EPMA Value	$\frac{n=6}{2\text{SD}}$	LA-ICP-MS Value	$\frac{n=3}{2\text{SD}}$	EPMA Value	$\frac{n=6}{2\text{SD}}$	LA-ICP-MS Value	$\frac{n=3}{2\text{SD}}$
Ti	355 368	1245	352 785	2384	1540	363	263	5.47
Fe	165 568	4412	172 284	3862	706 363	1275	705 622	243
Mg	99 953	2150	93 383	685	6468	136	8425	149
Mn	23 955	319	22 541	170	615	71.1	811	8.57
V	386	62.2	774	47.2	3478	98.6	3402	17.9
Zn			539	7.74			128	9.26
Nb			367	15.7			0.07	0.05
Sc			332	9.12			6.79	0.12
Со			72.6	0.37			89.7	0.51
Zr			39.7	0.76			0.07	0.02
Ni			30.4	0.77			228	3.09
Та			25.5	0.74			n.a	n.a
Sn			24.9	1.12			1.07	0.33
Al			11.3	2.37			138	7.64
Hf			6.36	0.11			n.a	n.a
W			0.48	0.01			0.01	0.01
Ga			0.25	0.06			9.62	0.24

of LA-ICP-MS and the multiple reference materials for external calibration strategy applied to magnetite geochemical analyses. In this study, the cross-comparison between LA-ICP-MS data and EPMA results indicates a good agreement, other than drifts for Ti contents in magnetite and V contents in ilmenite (Fig. 5, and data are summarized in Table 1). The V content of ilmenite obtained by LA-ICP-MS (774 \pm 47.2 ppm) should be accepted as the best estimate, as it is just slightly higher than its detection limits using EPMA (313–462 ppm, with its detection limits at ~80 ppm). Compared to EPMA results (1070–2130 ppm), Ti contents in magnetite obtained through LA-ICP-MS (263 \pm 5.47 ppm) have a narrower range, and both values are negligible relative to Ti content in ilmenite (35.3–35.6 wt%).

Although magnetite is compatible with a variety of lithophile and siderophile elements,¹⁻³ a single magnetite crystal could not incorporate all the components we selected for LA-ICP-MS analyses. As illustrated by direct LA analyses on coarse granular ilmenite and the magnetite domain in this study, a number of components are found below their respective detection limits, including Be, B, Si, P, Ca, Cr, Cu, As, Rb, Sr, Y, Ag, Cd, In, Sb, Cs, Pb, Bi, and most rare earth elements. Hafnium and tantalum are available in ilmenite exsolutions (Ta = 25.5 \pm 0.74 ppm, $Hf = 6.36 \pm 0.11$ ppm), but not detected in magnetite. Meanwhile, Ga and Ge are compatible in magnetite (Ga = 9.63 \pm 0.24 ppm, Ge = 1.70 \pm 0.003 ppm), but not in ilmenite. Some of these elements are available in Fe-Ti oxides from other locations (e.g. Cr, Ca, Si), while most of them are actually less compatible with magnetite and rarely reported by previous studies, thus of obscure petrogenesis significance.1-3

Among the various trace elements in magnetite and ilmenite, 25 elements (Mg, Al, Si, P, Ca, Sc, Ti, V, Cr, Mn, Co, Ni, Cu, Zn, Ga, Ge, Y, Zr, Nb, Mo, Sn, Hf, Ta, W, and Pb) are commonly listed and interpreted to be diagnostic for genetic studies and the exploration of metallic deposits. A multielement variation diagram proposed by Dare *et al.*² has been widely accepted, on which all 25 elements were plotted and normalized to bulk continental crust. In magnetite and ilmenite in this study, Si, Ca, Y, P, Pb, Cu, Mo, and Cr are too low to be detected. Germanium is available in magnetite, but not compatible with ilmenite exsolutions. The remaining 16 components have >0.1 ppm concentrations in coarse ilmenite exsolutions and their homogeneity could be demonstrated by repeated analyses (listed in ESI Table 2[†]).

4.2 Application of regression analysis

With the application of regression analysis to time-solved signals from mixed-sampling ablation, chemical compositions for coarse granular ilmenite (RAC, n = 8, results listed in ESI Table 4[†]) and micrometer ilmenite lamellae (RAF, n = 12, results listed in ESI Table 5[†]) were obtained. As the results from direct LA analyses could serve as the reference value, regression analysis was employed to extract the composition of coarse granular ilmenite from mixed LA-ICP-MS signals and the method availability could be illustrated by the consistency between the forecast results and their reference values. Then, our algorithm was applied to micrometer ilmenite lamellae, whose trace element contents have no direct reference value for comparison. Variations in concentrations for several major (Mg, Mn) and trace elements (Sc, Nb, Ta, Zr, Hf, Sn) from repeated analyses using regression analysis are shown in Fig. 6, with reference values for RAC plotted for comparison.

Consistent with EPMA results, Mg and Mn contents obtained *via* mixed-sampling applied regression analysis show similar variations between the coarse granular ilmenite (RAC, n = 8, red squares with red bars) and micrometer ilmenite lamellae (RAF, n = 12, blue squares with blue bars, Fig. 6). Most of the

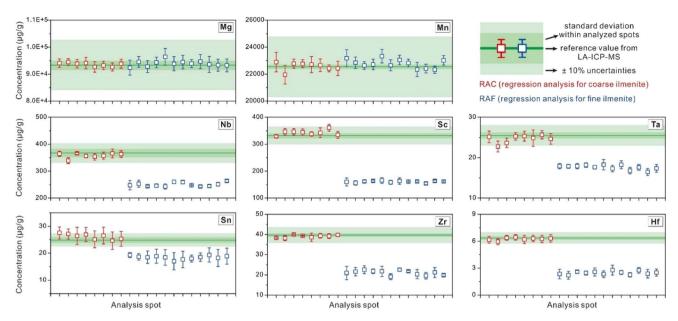


Fig. 6 Variations in major and trace elements of the coarse granular ilmenite (RAC, n = 8) and micrometer fine ilmenite lamellae (RAF, n = 12). Each point represents the forecast value from regression analysis using our algorithm, and the bars represent its statistical uncertainties at 95% confidence level (*i.e.*, 2 times standard deviation). The solid green line represents the average of direct analyses on granular ilmenite (n = 3), the medium green area represents its standard deviations.

estimated Mg contents using regression analysis deviate from the preferred value within 2%, and the worst one doesn't exceed 4% (Spot_RAC-5). Results for Mn contents also meet the reference value within 4%, and most of the deviations are less than 2%. Concentrations for trace elements (Sc, Nb, Ta, Zr, Hf, Sn) in coarse granular ilmenite are generally higher than those in micrometer ilmenite lamellae, but both types show their own internal consistency (Fig. 6). Such discrepancies in trace element compositions should be explained by geological factors, including temperature and oxygen fugacities, and possibly indicates the different conditions under which these two types of exsolutions occur and equilibrate with the magnetite grain. Compared to results from direct LA analyses as the reference value, element contents in coarse granular ilmenite calculated from regression analysis (RAC, n = 8) have relative deviations mostly within $\pm 10\%$ (Fig. 6).

As an example, the time-solved signal drifts for Spot_RAC-2 were presented (Fig. 2d), and its segmented results, regression analysis process, and comparisons with the reference value were shown in Fig. 3 and 4. To demonstrate the accuracy of our algorithm, the element contents in magnetite and coarse ilmenite obtained by direct LA analyses (as the reference value, shown as green squares in Fig. 3 and 4) are plotted together with the forecast values from regression analyses. The forecast values for Mg, Mn, Sc, V, Co, Zr, and W meet their reference value within $\pm 5\%$ relative deviation, and those for Fe, Nb, Sn, Hf, Ta have agreement within $\pm 10\%$. Accurate determination for components at lower contents (<10 ppm) is practicable using regression analysis, like W (0.46 \pm 0.04 ppm, *versus* 0.48 ppm as reference) and Hf (5.95 \pm 0.32 ppm, *versus* 6.36 ppm) in this study.

4.3 Analytical and statistical uncertainties

Errors are part of measurements, and effective error estimations are critical for geochemical analyses.³⁰ In this study, the uncertainty of the final results consists of two types, one is common for all LA-ICP-MS measurements (analytical uncertainty), and the other is brought by the additional algorithm performed in this work (statistical uncertainty). Analytical uncertainties were managed and ameliorated by powerful instrumental conditions and robust data reduction strategies.^{10,31} Instrumental conditions were optimized and maintained at acceptable levels by daily maintenance, involving maximization of the sensitivity and restraining the interferences from by-products like oxides and double charges. To support the data quality of Fe-Ti oxides, the international reference materials BCR-2G, BHVO-1G, and BIR-2G were analyzed as unknown, and the analytical results show good agreement with their preferred value (Fig. 5b and ESI Table 2[†]).

The distance between the regression result and its confidence limits (the standard error on the mean) indicates whether the regression analysis has captured a reliable forecast. Therefore, the standard error on the mean (SE) was calculated to represent the statistical uncertainty for each spot analysis, and the results' precision was expressed as the relative standard error ('SE'/'value'). For mixed-sampling analyses on coarse granular ilmenite (RAC), relative deviations from their preferred values (direct LA results) were calculated to represent the results' accuracy. Results from EPMA for those micrometer ilmenite lamellae could serve as reference values for their forecast values for major components (Mg, Mn), but the accuracy of their results for trace elements is difficult to evaluate. The statistical uncertainties for the regression results for

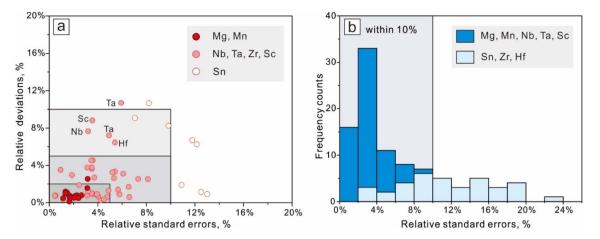


Fig. 7 Distributions for statistical uncertainties. (a) Scatter diagram of relative standard error (%) versus relative deviation from preferred value (absolute value, %) for analyses on coarse granular ilmenite (RAC, n = 8). (b) Histograms of relative standard error (precision, %) for analyses on micrometer ilmenite lamellae (RAF, n = 12).

different ilmenite exsolutions are calculated and listed in ESI Table 4† (for RAC, n = 8) and ESI Table 5† (for RAF, n = 12), and their distributions are shown in Fig. 7.

The relative deviations for all RAC measurements from their reference values (direct LA results) are better than or close to 10% (Fig. 7a, the two outside points have accuracy <11%). Most results meet their preferred values within \pm 5%, except for several measurements on Sc, Nb, Ta, Hf, and Sn. Concentrations for major components (Mg, Mn) constrained by regression analysis meet their reference values within \pm 2.5%, and their relative standard errors are below 3.2%. The regression results for Sn have the worst precision, but their relative errors on the forecast value were still within 13%. The relative standard errors on contents for other trace components (Nb, Ta, Zr, Hf, Sc) are lower than 9%. Results for element contents in micrometer ilmenite lamellae (RAF, Fig. 7b) have statistical uncertainties within 24%, and those for Mg, Mn, Nb, Ta, and Sc are better than 10%.

In this study, analytical uncertainties were optimized by applicable instrumental conditions and data reduction strategies, and the primary data quality was supported by the agreement between analytical results for the basaltic glasses and their preferred value. Statistical uncertainties resulting from regression procedures are expressed by their relative standard errors on the forecast and are mostly below 10%. The method accuracy was illustrated by the agreement between RAC results and their reference value (all within $\pm 11\%$, and the majority within $\pm 5\%$).

5 Conclusions

Ilmenite lamellae are common in titanomagnetite, but their trace element concentrations were rarely reported owing to their limited width. This work presents a practical strategy to determine element contents (including those <10 ppm) in micrometer-scale ilmenite lamellae in titanomagnetite using LA-ICP-MS. Regression analysis was applied to interpret the

mixed magnetite–ilmenite composition obtained by LA-ICP-MS, through which trace element concentrations in those ilmenite lamellae were obtained with high precision and accuracy (mostly within 10%). After precise estimations for their contents in ilmenite lamellae, the re-distribution of trace elements during the oxy-exsolution could be illustrated in a more quantitative way. Furthermore, fine-grained minerals are very common in many geological materials, and the algorithm presented here provides a potential solution for precise and accurate analyses of their chemical compositions.

Data availability

All the data used in this paper are either listed as Tables in the main text or as Appendix ESI Tables attached to this paper, and are also available from the corresponding author upon request.†

Author contributions

All authors listed contributed to the study conception and design. SQL: methodology, formal analysis, investigation, writing – original draft; SYJ: conceptualization, resources, supervision, funding acquisition, writing – review & editing; WC: methodology, investigation; CYW: methodology, investigation; HMS: methodology, investigation; YHC: methodology, investigation; HXZ: methodology, investigation; WTL: methodology, investigation. All authors read and approved the final manuscript.

Conflicts of interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

Acknowledgements

This research was supported by the National Natural Science Foundation of China (grant no. 42030811) and the special fund from the State Key Laboratory of Geological Processes and Mineral Resources, China University of Geosciences (no. MSFGPMR03-2). We thank Prof. Shui-Yuan Yang and Mr Sheng-Bin Guo for their assistance with EPMA analyses, Dr Jue Lu from CUG and William Minarik from McGill University for useful comments during writing of the original manuscript. The authors would also like to thank the Editor Dr Derya Kara Fisher, Prof. Don Davis, and anonymous reviewers for their helpful comments and constructive suggestions which have improved the quality of this paper.

References

- 1 C. Dupuis and G. Beaudoin, *Miner. Deposita*, 2011, **46**(4), 319–335.
- 2 S. A. S. Dare, S. Barnes, G. Beaudoin, J. Méric, E. Boutroy and C. Potvin-Doucet, *Miner. Deposita*, 2014, **49**(7), 785–796.
- 3 P. Nadoll, T. Angerer, J. L. Mauk, D. French and J. Walshe, *Ore Geol. Rev.*, 2014, **61**, 1–32.
- 4 W. Chen, Y. C. Ying, T. Bai, J. J. Zhang, S. Y. Jiang, K. D. Zhao, D. Shin and J. Kynicky, *Ore Geol. Rev.*, 2019, **10**7, 30–40.
- 5 H. Hu, D. Lentz, J. W. Li, T. McCarron, X. F. Zhao and D. Hall, *Econ. Geol.*, 2015, **110**(1), 1–8.
- 6 T. Hou, R. Botcharnikov, E. Moulas, T. Just, J. Berndt, J. Koepke, Z. C. Zhang, M. Wang, Z. P. Yang and F. Holtz, *J. Petrol.*, 2020, **61**(11–12), 1–24.
- 7 L. P. Zeng, X. F. Zhao, C. Spandler, H. Hu, B. Hu, J. W. Li and Y. Hu, *Econ. Geol.*, 2022, **117**(4), 923–942.
- 8 A. F. Buddington and D. H. Lindsley, *J. Petrol.*, 1964, 5(2), 310–357.
- 9 S. Y. Yang, R. X. Zhang, S. Y. Jiang and J. Xie, *Geostand. Geoanal. Res.*, 2018, 42(1), 131–137.
- 10 Y. S. Liu, Z. C. Hu, S. Gao, D. Günther, J. Xu, C. G. Gao and H. H. Chen, *Chem. Geol.*, 2008, 257(1-2), 34-43.
- 11 W. Zhang, Z. C. Hu, L. P. Feng, Z. C. Wang, Y. S. Liu, Y. T. Feng and H. Liu, *J. Earth Sci.*, 2022, 33(1), 67–75.
- 12 P. Nadoll and A. E. Koenig, J. Anal. At. Spectrom., 2011, 26, 1872–1877.

- 13 S. A. S. Dare, S. Barnes and G. Beaudoin, *Geochim. Cosmochim. Acta*, 2012, **88**, 27–50.
- 14 J. L. Knipping, L. D. Bilenker, A. C. Simon, M. Reich, F. Barra,
 A. P. Deditius, M. Wälle, C. A. Heinrich, F. Holtz and
 R. Munizaga, *Geochim. Cosmochim. Acta*, 2015, **171**, 15–38.
- 15 T. J. Ver Hoeve, J. S. Scoates, C. J. Wall, D. Weis and M. Amini, *Chem. Geol.*, 2018, **483**, 201–217.
- 16 N. Lv, K. Y. Chen, Z. A. Bao, K. Wu, D. B. Lei and H. L. Yuan, *At. Spectrosc.*, 2021, 42(1), 51–61.
- 17 M. J. Brzozowski, I. M. Samson, J. E. Gagnon, R. L. Linnen and D. J. Good, *Miner. Deposita*, 2021, **56**(3), 601–618.
- 18 M. J. Brzozowski, I. M. Samson, J. E. Gagnon, D. J. Good and R. L. Linnen, *Miner. Deposita*, 2021, 56(4), 621–642.
- 19 X. W. Huang, É. Boutroy, S. Makvandi, G. Beaudoin, L. Corriveau and A. F. De Toni, *Miner. Deposita*, 2019, 54(4), 525–552.
- 20 X. W. Huang, A. Sappin, É. Boutroy, G. Beaudoin and S. Makvandi, *Econ. Geol.*, 2019, **114**(5), 917–952.
- 21 M. J. Toplis and A. Corgne, *Contrib. Mineral. Petrol.*, 2002, **144**(1), 22–37.
- 22 E. P. Reguir, A. R. Chakhmouradian, N. M. Halden, P. Yang and A. N. Zaitsev, *Can. Mineral.*, 2008, **46**(4), 879–900.
- 23 N. A. Durasova, I. D. Ryabchikov and V. L. Barsukov, *Int. Geol. Rev.*, 1986, 28(3), 305–311.
- 24 D. Rubatto and J. Hermann, *Chem. Geol.*, 2007, **241**(1–2), 38–61.
- 25 A. S. Stepanov, J. Hermann, D. Rubatto and R. P. Rapp, *Chem. Geol.*, 2012, **300–301**, 200–220.
- 26 A. S. Stepanov, L. V. Danyushevsky, R. R. Large, I. Mukherjee and I. A. Zhukova, *Am. Mineral.*, 2020, **105**(6), 820–832.
- 27 C. W. Wei, C. Xu, A. R. Chakhmouradian, M. Brenna, J. Kynicky and W. L. Song, *J. Petrol.*, 2020, **61**(2), 1–26.
- 28 H. J. Jiang, C. S. Yang, D. Q. Wang, H. Zheng, J. Li and H. Y. Chen, *J. Earth Sci.*, 2022, **33**(1), 193–204.
- 29 K. P. Jochum, M. Willbold, I. Raczek, B. Stoll and K. Herwig, Geostand. Geoanal. Res., 2005, 29(3), 285–302.
- 30 R. Dybkaer, Scand. J. Clin. Lab. Invest., 1995, 55(2), 97.
- 31 D. Günther and B. Hattendorf, *TrAC, Trends Anal. Chem.*, 2005, 24(3), 255–265.