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Ontogeny of highly variable ceratitid ammonoids from the Anisian (Middle Triassic)

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ABSTRACT

Ammonoids reached their greatest diversity during the Triassic period. In the early Middle Triassic (Anisian) stage, ammonoid diversity was dominated by representatives of the family Ceratitidae. High taxonomic diversity can, however, be decoupled from their morphologic disparity. Due to its high phenotypic variability, the high diversity of ceratitids of the Anisian of Nevada was initially assumed to be caused by artificial over-splitting. This study aims to contribute data to settle this issue by applying geometric morphometrics methods, using landmarks and semi-landmarks, in the study of ontogenetic cross-sections of ammonoids for the first time. The results reveal that alterations in ontogenetic trajectories, linked to heterochronic processes, lead to the morphologic diversification of the species studied herein. Our knowledge, based on these ontogenetic changes, challenge the traditional treatment of species using solely adult characters for their distinction. This study furthermore demonstrates that the high diversity of the Anisian ammonoid assemblages of Nevada based on the traditional nomenclatoric approach is regarded to be reasonably accurate.

Subjects Biodiversity, Evolutionary Studies, Paleontology, Taxonomy, Zoology **Keywords** Ammonoidea, Ceratitidae, Anisian, Nevada, Geometric morphometrics, Ontogeny, Phenotypic variation, Beyrichitinae, Paraceratitinae, Fossil Hill Member

INTRODUCTION

After the Permian-Triassic mass extinction event, ammonoids flourished and spread globally to become an important part of the marine biota (*House, 1993; Brosse et al., 2013; Brayard & Bucher, 2015; Neige, 2015*). They reached their greatest generic diversity of all time in the Triassic Period (*Brayard et al., 2009; Whiteside & Ward, 2011*). The diversity peak in the late Anisian is dominated by genera of the family Ceratitidae (*Brayard et al., 2009;* supporting material Fig. S2). Not least due to their wide paleogeographic distribution and high diversity as well as abundance in the fossil record, ammonoids are an excellent biostratigraphic tool. This is especially true for members of the family Ceratitidae *Mojsisovics 1879* after which many North American Anisian biostratigraphic zones and subzones are named (*Jenks et al., 2015*; figs. 13.13, 13.14).

The fossil material used in this study was collected in the late Anisian Fossil Hill Member of the Star Peak Basin in north-western Nevada, USA. The studied successions are considered to be the world's most complete low-paleolatitude successions, yielding

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late Anisian ammonoid assemblages (Monnet & Bucher, 2005). The first comprehensive taxonomic work on the Anisian ammonoid communities of the famous fossil locality at Fossil Hill in the Humboldt Range was published by *Smith* (1914) in his monograph on the North American Middle Triassic marine invertebrates. According to the taxonomic practice of his time, he described or listed a total of 110 ammonoid species from Fossil Hill. More recently, Silberling & Nichols (1982), Bucher (1992) and Monnet & Bucher (2005) refined the original alpha taxonomy and the biostratigraphy with contemporaneous methods and reduced the number to 81 valid species (*Brosse et al.*, 2013). However, it is important to note that succeeding assemblages show a progressive shift in morphology; therefore, the cutoff between contiguous species is essentially arbitrary (Silberling, 1962; Silberling & Nichols, 1982). This challenges the taxonomic concept and sheds new light on diversity patterns in general. An increasing number of studies suggest that the seemingly high diversity could in some cases be artificially inflated by taxonomic over-splitting (i.e., Kennedy & Cobban, 1976; Forey et al., 2004; De Baets, Klug & Monnet, 2013; Knauss & Yacobucci, 2014; De Baets et al., 2015). Furthermore, taxonomic diversity and morphological disparity of Triassic ammonoids were probably decoupled (McGowan, 2004; McGowan, 2005; Brosse et al., 2013). At present, only a few studies have investigated trends in morphological disparity of Triassic ammonoids (Monnet, Brayard & Brosse, 2015).

Previous studies have proven that-particularly due to their accretionary planispiral conch growth with conservation of previous growth stages-ammonoids offer a highresolution data set for ontogenetic, developmental and also taxonomic studies. While the study of conch ontogeny has a long history in the study of Paleozoic ammonoids (e.g., Korn & Klug, 2007; Korn, 2010; Monnet, De Baets & Klug, 2011; Naglik et al., 2015), it was only rarely examined on Mesozoic ammonoids (e.g., Rieber, 1962; Tajika et al., 2015; Bischof & Lehmann, 2020). So far, the morphology and ontogeny of ammonoids was mainly assessed using descriptive, comparative or traditional morphometric methods (linear measurements). In his classic work of 1966, Raup introduced traditional geometric parameters for the description of coiled conch morphospace. These "Raupian parameters" were subsequently refined by Korn & Klug (2003), Korn & Klug (2007), Korn (2010), Klug et al. (2015a) and Klug et al. (2015b). However, the shapes of discoidal ammonoids often differ through their characteristic ways of ventral arching and presence or absence of a keel. Both characteristics can hardly be described with linear measurements (*Neige*, 1999). Therefore, the use of traditional morphometric methods might be limited when it comes to distinguishing ceratitid species.

For the first time, the morphology and ontogeny of whorl profiles of the late Anisian ceratitids were analyzed using landmark- and semilandmarks-based geometric morphometric methods (GMM) instead of linear measurements (traditional morphometrics). This is reasoned in the tremendous advantages of GMM over the latter; landmarks and semi-landmarks cover shape variations of complete morphologies, which are sometimes not to be recognized or overseen with linear measurements of traditional morphometrics methods (*Neige*, 1999). In addition, GMM allow the analysis of shape and size separately (*Hammer & Harper*, 2005; *Zelditch*, *Swiderski & Sheets*, 2012;

Polly & Motz, 2016) and do not introduce artifactual patterns of covariation (*Gerber, 2017*), which is often the case when proportions are studied.

The literature on geometric morphometric analyses (landmark-based approaches and Fourier analysis) of molluscs is rather scarce. Although important pioneering works exist, the previous studies are of limited use in an ontogenetic context because they all focus either on the shape of the whole conch or on single (isolated) ontogenetic stages (landmarks: e.g., *Johnston, Tabachnick & Bookstein, 1991; Neige & Dommergues, 1995; Reyment & Kennedy, 1998; Stone, 1998; Neige, 1999; Reyment, 2003; Van Bocxlaer* & Schultheiß, 2010; Knauss & Yacobucci, 2014; Fourier analysis: e.g., *Courville & Crônier, 2005; Simon, Korn & Koenemann, 2010; Simon, Korn & Koenemann, 2011; Korn & Klug, 2012; Klein & Korn, 2014*).

In order to evaluate the hitherto used taxonomic scheme, ontogenetic patterns within the family Ceratitidae and their changes over time were investigated. Working with ontogenetic cross-sections allows the estimation of the relative age of the whorls, which adds an extra dimension to the analysis. The tools presented here are intended to complement traditional descriptions and to evaluate and quantify their results. This study should serve as a general motivation to conduct GMM studies on invertebrates with accretionary planispiral growth.

MATERIALS & METHODS

Geological setting

The ammonoid material derives from the Fossil Hill Member of Fossil Hill in the Humboldt Range and Muller Canyon in the Augusta Mountains (Pershing County), north-western Nevada, USA (Fig. 1) and is stored in the Geosciences Collection of the University of Bremen (GSUB), Germany. The material from the Wilderness Study Area of the Augusta Mountains, Pershing County was collected with permission of the US Department of the Interior, Bureau of Land Management (BLM, Nevada State office, Winnemucca District). The Fossil Hill Member is a succession of alternating layers of mudstone with lenticular limestone and calcareous siltstone beds (see Fig. 2). The rich and diverse fossil content consists primarily of halobiid bivalves and ammonoids. Detailed geological and stratigraphic descriptions were published by *Nichols & Silberling (1977)*, *Silberling & Nichols (1982)* and *Monnet & Bucher (2005)*.

Studied specimens

The fossil material comprises 72 ammonoid specimens of the family of Ceratitidae (*Mojsisovics, 1879*). These represent twelve species in seven genera (Fig. 2, Table 1) that either belong to the subfamily Beyrichitinae (*Spath, 1934*) or Paraceratitinae (*Silberling, 1962*). Most of the studied species show high intraspecific variation with overlapping morphologies (see Table 1 and Figs. 3–5). Members of these genera (*Gymnotoceras, Frechites* and *Parafrechites* in particular) are sometimes hard to differentiate. They mainly differ in the ventral conch outline, ornamentation, adult ribbing and maximum growth size. The younger the individuals are, the greater the similarities. Despite their complicated taxonomy, all selected species are index fossils of the late Anisian Fossil Hill Member (see Fig. 2). It was assumed that the individual species have similar coiling rates (i.e., the



Figure 1Location of the study area in NW Nevada, USA. The Fossil Hill and the Muller Canyon locali-
ties are marked. Figure adapted from Bischof & Lehmann (2020), Fig. 1.Full-size IDOI: 10.7717/peerj.10931/fig-1

individual species develop the same number of whorls in the course of their life). The total number of volutions developed by the species varies between five and a half and seven (see Table 1—Total number of volutions).

Preparation and data acquisition

We prepared high-precision cross-sections intersecting the protoconch of each specimen, following the methods by *Korn (2010)*, *Klug et al. (2015a)* and *Klug et al. (2015b)*. Subsequently, we scanned the polished surfaces in high resolution with a flat screen scanner to ensure that all pictures have the same scale. Thereafter, the scan images were digitized. CT scan images of Anisian ammonoids from Nevada do not provide sufficient contrast of the internal structures for a reliable analysis (*Bischof & Lehmann, 2020*).

Based on the digitized cross-sections, we performed a 2D landmark-based geometric morphometrics analysis. The landmarks were retrieved in tpsDig2 v.2.31 (*Rohlf, 2010*). Sixteen landmarks were digitized per half whorl (i.e., whorl stage), which resulted in 176 landmarks per specimen (16 landmarks on 11 half whorls; Fig. 6). This set of landmarks consists of two single (1, 2) and 7 pairs of landmarks (3–16), of which eight are sliding semi-landmarks. Whereas landmarks are discrete anatomical loci (i.e., point of highest curvature of venter), sliding semi-landmarks are placed along a curve (or a surface) between two landmarks in a way that best describes the curvatures of the outline. In a second step, an algorithmic approach optimizes the approximation of the outline (*Zelditch, Swiderski & Sheets, 2012*).

In order to omit missing values in subsequent analyses, the data set was limited to whorl stage number 5.5. From a methodological point of view, it is more practical to rotate



Figure 2 Biostratigraphic distribution of fossil material and synoptic lithostratigraphic sections of the outcrops in the Muller Canyon and Fossil Hill area. Stratigraphic section of Muller Canyon adapted from *Bischof & Lehmann (2020)*, Fig. 2. Gray areas in stratigraphic column: Calcareous siltstone; white areas: lenticular limestone, box width refers to weathering profile.

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the shells by 90° compared to conventional illustrations (cf. *Stridsberg*, *1990*) into a lying position. Since ammonoid conchs are spiral-shaped, each whorl is cut in two parts when preparing the cross sections (see Fig. 6). The half whorls on the left side of the protoconch have odd numbers ("odd whorls"; here 0.5–5.5) and those on the right side have even numbers ("even whorls"; here 1.0–5.0"). Homologous landmarks were set in accordance to the axial plane.

Calculation of procrustes shape

All geometric morphometric analyses were carried out using the R software v 3.6.3. (*R Core Team, 2020*) packages *Morpho* v2.8 (*Schlager, 2017*), geomorph v3.3.1. (*Adams et al., 2020*) and *RRPP* v0.6.0 (*Collyer & Adams, 2018*; *Collyer & Adams, 2020*). Plots were drawn with the R package ggplot2 (*Wickham, 2016*). Using the *Morpho::procSym* function, the 2D

| Tuble 1 Morphological comparison of the species in focus, 1 of biostratigraphic distribution see 112, 2. | | | | | | | | |
|--|----|------------------------------------|---|--|--------------|-----------|------------------------|------------------|
| Species | Ν | Total number of volutions | Venter and conch outline | Sculpture | Dmax [mm] | U/D | W/D | Figure herein |
| Beyrichitinae <i>Spath</i> , 1934 | | | | | | | | |
| Billingsites cordeyi Mon- net & Bucher, 2005 | 6 | 6—6.5 | Slightly angular ventral shoulder | Falcoid, prorsiradi- ate ribs, sometimes branched | 34.3 | min: 0.17 | min: 0.28 max: 0.35 | 3A-D |
| | | | Very weak devel- oped keel | Nodes at branch- ing points | | max: 0.24 | | |
| Dixieceras lawsoni (Smith, 1914) | 10 | 6—7 | Stout, discoidal outline | Falcoid, prorsiradi- ate ribs, sometimes branched | 57.7 | min: 0.19 | min: 0.23 max: 0.44 | 3I–L |
| | | | Rounded ventral shoulders | Umbilical thicken- ing of whorls | | max: 0.25 | | |
| Frechites nevadanus (Mojsisovics, 1888) | 6 | 5.5—6 | Subrectangular outline | Strong, falcoid, prorsiradiate ribs, sometimes branched | 28.9 | min: 0.29 | min: 0.39 max: 0.46 | 3М-Р |
| | | | Clearly developed keel | Adults: Pronounced tubercles at lower flank | | max: 0.37 | | |
| Frechites occidentalis (Smith, 1914) | 7 | 6—7 | Angular ventral shoulder | Strong, slightly prorsiradiate ribs, some rare tubercles | 42.6 | min: 0.24 | min: 0.38 max: 0.43 | 4I–L |
| | | | Sometimes very weak developed keel | Towards maturity ribbing fades | | max: 0.27 | | |
| Gymnotoceras blakei (Gabb, 1864) | 5 | 5.5—6 | Discoidal outline | Falcoid, prorsira- diate, unbranched ribs | 37.8 | min: 0.15 | min: 0.30 max: 0.38 | 4A–D |
| | | | Rounded ventral shoulders, weak keel | Towards matu- rity fading ribs and megastriae | | max: 0.28 | | |
| Gymnotoceras mimetus Monnet & Bucher, 2005 | 9 | 6—6.5 | Discoidal to sub- rectangular outline | Megastriae and weak falcoid, prorsiradiate ribs, slightly swelling towards umbilicus | 43.0 | min: 0.14 | min: 0.29 max: 0.40 | 5K–N |
| | | | Rounded ventral shoulders, no keel | | | max: 0.22 | | |
| <i>Gymnotoceras rotelli-</i> <i>formis (Meek, 1877)</i> | 6 | 6 | Stout discoidal outline, very weak keel | Regular, slightly prorsiradiate ribs | 34.3 | min: 0.17 | min: 0.32 max: 0.38 | 4M-P |
| | | | Rounded ventral shoulders | Towards matu- rity ribbing slightly fades | | max: 0.26 | | |

 Table 1
 Morphological comparison of the species in focus. For biostratigraphic distribution see Fig. 2.

(continued on next page)

Table 1 (continued)

| Species | Ν | Total number of volutions | Venter and conch outline | Sculpture | Dmax [mm] | U/D | W/D | Figure herein |
|--|---|------------------------------------|--|--|--------------|-----------|------------------------|------------------|
| Gymnotoceras weitschati Monnet & Bucher, 2005 | 3 | 6 | Compressed, dis- coidal outline | Megastriae and weak falcoid, prorsiradiate ribs, slightly swelling towards umbilicus | 28.4 | min: 0.17 | min: 0.29 max: 0.33 | 5A–E |
| | | | Perfectly rounded shoulders, no keel | | | max: 0.20 | | |
| Parafrechites dunni (Smith, 1914) | 5 | 5.5—6.5 | Stout discoidal outline, sometimes keel | Regular but weak, slightly prorsiradi- ate ribs | 35.2 | min: 0.18 | min: 0.31 max: 0.42 | 4E–H |
| | | | Rounded to sub- angular ventral shoulders | Towards matu- rity ribbing slightly fades | | max: 0.20 | | |
| Parafrechites meeki (Mo- jsisovics, 1888) | 5 | 5.5—6 | Subrectangular outline | Strong and regular, falcoid, prorsiradi- ate ribs, sometimes branched | 32.1 | min: 0.22 | min: 0.34 max: 0.41 | 50–R |
| | | | Strong keel, sub- angular shoulders | | | max: 0.27 | | |
| Paraceratitinae <i>Silber-</i> <i>ling</i> , 1962 | | | | | | | | |
| Brackites vogdesi (Smith, 1904) | 4 | 6—7 | Subrectangular outline, slightly rounded shoulders | Regular, falcoid, branched, prorsir- adiate ribs | 29.6 | min: 0.28 | min: 0.35 max: 0.37 | 3E–H |
| | | | Tubercles at branc- ing point | | max: 0.37 | | | |
| Marcouxites spinifer (Smith, 1914) | 6 | 5.5—6 | Subrectangular outline, angular shoulder | Strong and regular, falcoid, prorsiradi- ate ribs | 25.8 | min: 0.26 | min: 0.38 max: 0.42 | 5F–J |
| | | | Clearly developed keel | Tubercles and spines at branching point | | max: 0.36 | | |

Notes.

N, Number of specimens; U, maximum umbilical diameter; W, maximum whorl width; D, maximum diameter of conch.

Measurement values and ratios based on material herein. More detailed information on the studied species was published by *Silberling & Nichols (1982)* and *Monnet & Bucher (2005)*.

landmark coordinates were subjected to a full generalized Procrustes alignment (GPA). The semilandmarks were slid minimizing Procrustes distance. The full Procrustes fit standardizes size, orientation and position, leaving only the Procrustes shape coordinates (*Bookstein, 1991*, chap. 7.1, p. 258–270; *Hammer & Harper, 2005*; *Zelditch, Swiderski & Sheets, 2012*). Since the "odd" and "even" whorls cannot be made congruent by any of these operations (i.e., alignment, translation, rotation), all "even" whorls were manually mirrored before the GPA.

The individual whorls were regarded as different structures of the ammonoid conch. Therefore, the GPA was performed separately for every whorl. The *procSym*



Figure 3 Ceratitid ammonoids from the Anisian (Middle Triassic) Fossil Hill Member of NW Nevada, USA. (A–D) *Billingsites cordeyi* (*Monnet & Bucher, 2005*), (A, B) GSUB C11082, (C, D) GSUB C11517; (E– H) Brackites vogdesi (*Smith, 1904*), (E, F) GSUB C11649, (G, H) GSUB C11646; (I–L) Dixieceras lawsoni (*Smith, 1914*), (I, J) GSUB C13801, (K, L) GSUB C13805; (M–P) Frechites nevadanus (*Mojsisovics, 1888*), (M, N) GSUB C12377, (O, P) GSUB C12382.

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function performs Procrustes superimposition including sliding of semi-landmarks on curves and accounts for the symmetry of the object. Subsequently, the R function *geomorph::combine.subsets* was used to normalize the configurations of all whorl stages to unit centroid size or with a customized weighting (see "Developmental morphospaces"). The centroid size (CS) is regarded as a proxy for the size of the whorls and equals the



Figure 4 Ceratitid ammonoids from the Anisian (Middle Triassic) Fossil Hill Member of NW Nevada, USA. (A–D) *Gymnotoceras blakei* (*Gabb*, *1864*), (A, B) GSUB C12243, (C, D) GSUB C12264; (E–H) Parafrechites dunni (*Smith*, *1914*), (E, F) GSUB C9946 (G, H) GSUB C12906; (I–L) *Frechites occidentalis* (*Smith*, *1914*), (I, J) GSUB C8998, (K, L) GSUB C13251; (M–P) *Gymnotoceras rotelliformis* (*Meek*, *1877*), (M, N) GSUB C11594, (O, P) GSUB C11702.

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Figure 5 Ceratitid ammonoids from the Anisian (Middle Triassic) Fossil Hill Member of NW Nevada, USA. (A–E) Gymnotoceras weitschati (Monnet & Bucher, 2005), (A, B) GSUB C11111, (C–E) GSUB C11158; (F–J) Marcouxites spinifer (Smith, 1914), (F, G) GSUB C10050, (H–J) GSUB C10137; (K–N) Gymnotoceras mimetus (Monnet & Bucher, 2005), (K, L) GSUB C15005, (M, N) GSUB C13811; (O–R) Parafrechites meeki (Mojsisovics, 1888), (O, P) GSUB C12534, (Q, R) GSUB C12618. Full-size DOI: 10.7717/peerj.10931/fig-5

square root of the summed squared distances of each landmark from the centroid of the landmark configuration before the GPA (*Zelditch, Swiderski & Sheets, 2012*). The function *geomorph::combine.subsets* was originally introduced to combine different parts of a body (e.g., heads and tails; *Collyer, Davis & Adams, 2020*).

To visualize the multivariate data in two-dimensional morphospaces, we ran a Principal Component Analysis (PCA) on the aligned Procrustes shape coordinates using the R



Figure 6 Digitized sketch of high-precision cross-section of an ammonoid specimen meeting the initial chamber (protoconch) with position of landmarks on last two half whorls. Filled crosses: fixed landmarks; empty crosses: sliding landmarks; black numbers: numbers. Definition of fixed landmarks: (1) venter of preceding whorl; (2) venter of whorl; (3) and (4) ventral shoulder or point of highest curvature; (5) and (6) maximum width; (7) and (8) Umbilical seam.

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function *stats::prcomp*. Thereby we used two different types of visualization: Ontogenetic trajectory spaces and developmental morphospaces.

Ontogenetic trajectory spaces

It is well-known that ammonoids have a very characteristic but also complex ontogenetic development (e.g., *Klug, 2001*). To visualize the ontogenetic development of ammonoids, there are different types of morphospaces. Ontogenetic trajectory spaces (originally called ontogenetic morphospaces), as defined by *Bischof & Lehmann (2020)*, p.2), illustrate the differences in total ontogenetic development of individuals. They show the data in an artificial state of combined morphologies of different ontogenetic stages. To calculate an ontogenetic trajectory space, all Procrustes shapes (i.e., whorls) of an individual are re-assembled before running the Principal Component Analysis. This means that, in an ontogenetic trajectory space, the ontogenetic trajectory of every individual is reduced to a single data point. Ontogenetic trajectory spaces are a tool to examine if the ontogenetic pathways of individuals differ, but they do not show how the trajectories vary. To test whether ontogenetic trajectories statistically differ between species, a multivariate analysis of variance (MANOVA) using the R function *stats::manova* was applied.

Developmental morphospaces

Developmental morphospaces as defined by *Eble* (2003, p. 40) are morphospaces that directly contain developmental information. In terms of this study, this means that every individual dot in the morphospace reflects a specific ontogenetic stage (i.e., half whorl) of an individual. By connecting all points of an individual, its ontogenetic trajectory can be obtained. In contrast to ontogenetic trajectory spaces, developmental morphospaces show how individual whorls differ from each other.

General Procrustes Analysis (GPA) removes all information about size from a given set of data leaving only the pure shape coordinates. However, as can be seen in Fig. 6, size differences between different whorl stages are tremendous. If normalized to unit centroid size (i.e., non-weighted morphospace), the earliest whorls of ammonoids therefore get enormously enlarged and the last whorls scaled down. In general, deviations (measurement uncertainties as well as actual morphological variation) are increased for the initial whorls and reduced for older whorl stages. Therefore, a second morphospace with weighted Procrustes shape coordinates was calculated. Thereby, the logarithmic centroid size (log₁₀CS) of all configurations of a whorl stage were normalized to the proportional centroid size of the respective stage to the sum of all whorl stages (log₁₀ CS_{whorl *i*/ \sum log₁₀ CS_{whorls}). The principal components of the PCA on the weighted shapes were called wPC (weighted principal components).}

If the relative \log_{10} CS is used to normalize the centroid size of the configurations, this approach is extremely similar to a Relative Warp Principal Component Analysis (RW-PCA; size-shape space) after *Mitteroecker et al.* (2004). To calculate a RW-PCA the shape matrix of a configuration is augmented by an additional column containing information about the \log_{10} CS of the configurations. Whereas the R function *geomorph::combine.subsets* scales every configuration accordingly, the size information in the RW-shape matrices are stored in the additional variable. The resulting RW size-shape space can be analyzed with an ordinary PCA. Typically, RW size-shapes spaces are strongly dominated by the \log_{10} CS and PC1 therefore often accounts for more than 90% of the variation. If proportional \log_{10} CS-values (\log_{10} CS_{configuration i}/ $\sum \log_{10}$ CS_{configurations}) are used, the analysis is less dominated by size, but the eigenvalues are very similar to the ones of the weighted PCA (wPCA). For simplicity the R function was used here.

Because weighting does not change the shapes itself, weighted and non-weighted developmental morphospaces look very similar. The main difference is the placement of the individual configurations within the morphospace. Whether weighted or non-weighted shape coordinates should be used, depends on what the analysis is intended to show. A summary of the three different approaches how to combine landmark configurations can be found in Table 2.

To model the shapes at the maximum and minimum PC-values, the R function *GeometricMorphometricsMix::reversePCA* (*Fruciano, 2019*) was used. The function is designed to recalculate artificial Procrustes shape variables from the extreme PC-values in a morphospace. The thin-plate spline deformation grids were calculated using the R function *geomorph::plotRefToTarget*.

Trajectory analysis

In morphometric studies, ontogenetic trajectories represent a series of measurement values of different ontogenetic stages of an individual or a group, called longitudinal data (*Klingenberg, 1998*). To quantify the differences of the ontogenetic trajectories of the individual species, the R function *RRPP::trajectory.analysis* with 999 iterations was used. The function calculates a linear model with at least one categorical interaction variable (here: Shape ~ Species * WhorlStage) and assesses differences in path distance (magnitude differences, length of trajectories), trajectory shape and the angle between the individual trajectories (trajectory correlation) (*Collyer & Adams, 2013*).

| Shape variables | Effect on shape data set | Advantages | Disadvantages |
|------------------------------|--|--|---|
| Non-weighted (wPCA) | • Minimizes shape difference be- tween all configurations of all o.s. | • No perturbation of original dataset | • Overestimation of deviations of earliest/small earliest/smallest o.s. and underestimation of lat- est/largest o.s. |
| Weighted (wPCA) | • Adds allometric/size informa- tion to analysis | • More complicated computa- tion than other methods | • Domination of size (less than in RW-PCA) |
| | • Maximizes shape differences between o.s. | • Deviations of earliest/smallest o.s. less overestimated and of lat- est/large o.s. less underestimated | • Partial loss of objectivity |
| | • Minimizes shape differences within an o.s. | | • May suppress potential true variation in earliest o.s. |
| | | | • Slight perturbation of original data set (creation of arbitrary co-variances) |
| Shape-size space (RW-PCA) | • Adds allometric/size informa- tion to analysis | • Easy to compute | • Extreme domination of size |
| | •Every configuration is scaled individually according to their centroid size | • Intuitive | • Strong perturbation of original data set, cannot be used for most subsequent analyses |

Table 2 Comparison of most important characteristics of weighted and non-weighted approaches to combine landmark configurations.

Notes.

o.s., ontogenetic stages (i.e., whorls).

If weighted shape-coordinates were used, the artificial size-shape relationship could overlay true differences between the trajectories. Therefore, only non-weighted shape coordinates were analyzed in the trajectory analysis.

RESULTS

Ontogenetic trajectory spaces

Ontogenetic trajectory spaces are means to visualize whether or not the ontogenetic development of two or more individuals differs. The first three components (PCs) of the Principal Component Analysis (PCA) on the shape coordinates with combined ontogenetic stages of an individual account for 58.3% (PC₁ = 38.3%, PC₂ = 13.4%, PC₃ = 6.6%) of the total variation. Especially considering that there are a total of 352 primary components (*x* and *y* coordinates of 176 landmarks), but only 72 specimens, this result can be regarded as satisfactory. The convex hulls of the ontogenetic trajectory space of most species reveal a large overlap (Fig. 7). Nevertheless, ontogenetic trajectory spaces differ significantly between species (MANOVA: Pillai's trace = 9.8962; ~ *F* (1, 11) = 1.6437, *p* < 0.001).

Since PC1 accounts for 38.3% of the total variation, the most important characteristic is the position of the individuals on the *x*-axis. In fact, there are certain species that primarily have negative PC1 values (*B. vogdesi*, *F. nevadanus*, *F. occidentalis*, *M. spinifer*, *P. meeki*) and some that are more restricted to positive PC1 values (*G. blakei*, *G. mimetus*, *G. rotelliformis*, *G. weitschati*, *P. dunni*). *B. cordeyi* and *D. lawsoni*, both cover a wider range of different PC1 values, but are generally restricted to negative PC2 and positive PC3 values.



Figure 7 Ontogenetic morphospace of all species analyzed. (A) Principal Component 1 and 2; (B) Principal component 1 and 3.

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Non-weighted developmental morphospace occupation

The first three components of the PCA on the non-weighted shape space account for 93.8% (PC₁ = 78.5%, PC₂ = 11.5%, PC₃ = 3.1%) of the total variation. The PCA plot (Fig. 8) of PC1 and PC2 shows that the whorls of early ontogenetic stages cover the lower left quadrant of the morphospace (negative PC1 and PC2 values), which characterizes extremely depressed, broad whorls with a flat venter (Fig. 9A). The center of the morphospace (PC1 equals 0 and PC2 is positive) is occupied by intermediate growth stages (juveniles), which have a more quadratic outline with an only slightly triangular venter (Fig. 9B). The lower right quadrant (high PC1 and low PC2 values) is associated with the latest ontogenetic stages (adults). Towards maturity, the whorls increase mainly in height and have a clearly triangular venter and sometimes a keel (Figs. 9C and 10). Overall, there are two extreme adult shapes: Type (A) describes rather depressed, stout conchs with only a slight overlap with the preceding whorl and are associated with much shorter ontogenetic trajectories. Type (B) describes compressed conchs with a clearly triangular venter and a higher degree of overlap and are associated with longer ontogenetic trajectories. For the assignment of the species to the two types see Table 3.

The developmental morphospace of beyrichitine and paraceratitine ammonoids comprises three basic shape stages, which are not separated by sharp borders (Fig. 9): (1) Earliest whorls: broad and very flat; (2) Juveniles: more rounded and depressed; (3) Adults: mostly high and compressed whorls. Since type A species stop their development at more rounded and depressed whorls, their adult whorls resemble the juvenile stages of type B (Fig. 10).

Ontogenetic trajectories in the non-weighted developmental morphospace

The ontogenetic trajectories of species in the non-weighted developmental morphospace share many similarities: They all have similar directions of propagation and a slight parabolic shape (Fig. 8). The variation detected by the trajectory analysis (R function *RRPP::trajectory.analysis*) revealed significant differences in trajectory length (path distance), trajectory shape and trajectory slope between most species (File S1, summarized results below).



Figure 8 Developmental morphospace with PCA of Procrustes shape variables. Point size refers to whorl stage. Deformation grids of the mean shape to the modeled shapes of the extreme values for PC1 and PC2.

Full-size DOI: 10.7717/peerj.10931/fig-8



Figure 9 Mean shapes of whorl stages 0.5, 3.0 and 5.5. Fixed landmarks are marked with a cross. Full-size 🖾 DOI: 10.7717/peerj.10931/fig-9

Members of the type A ontogeny have smaller magnitudes of shape change and different trajectory shapes than members from type B ontogeny (path distances A: 0.2005–0.0225; B: 0.2276–0.2778). For the assignment of the species to the two types see Table 3. Only pairwise differences of the path distance and the trajectory shape between type A and type B species are statistically significant (magnitude of shape change: 17/66 possible pairs; trajectory shape: 16/66 possible pairs).

Most species have statistically significant pairwise differences in trajectory slope (57/66 possible pairs). Species of all pairs with non-significant pairwise *p*-values are in the same ontogenetic group (i.e., both belong to type A or B). The trajectories of the nine pairs with non-significant differences in slope, have non-significant magnitudes and shapes also.



Figure 10 TPS spline of mean shape of whorl 5.5 of all species in this study (grey) plotted against the mean shape of whorl 5.5 of the respective species (black). Fixed landmarks are marked with a cross. Full-size DOI: 10.7717/peerj.10931/fig-10

 Table 3
 Summary and explanation on the three different ontogenetic types. Heterochronic terms as defined by *McNamara* (2012).

| Туре | Species | Heterochrony | Adult whorl shapes | | |
|------|------------------|----------------|--|--|--|
| A1 | F. nevadanus | | Depressed, stout conches, only slight overlap with preceding whorl | | |
| | F. occidentalis | | | | |
| | M. spinifer | Paedomorph | | | |
| | B. vogdesi | | | | |
| | P. meeki | | | | |
| В | B. cordeyi | | | | |
| | D. lawsoni | | | | |
| | G. blakei | Doromorph | Compressed conches, more | | |
| | G. mimetus | (Acceleration) | pronounced venter, more overlap with preceding whorl | | |
| | G. rotelliformis | () | | | |
| | G. weitschati | | | | |
| | P. dunni | | | | |

However, none of the species that share a common slope have overlapping biostratigraphic ranges.

Weighted developmental morphospace occupation

The first three components of the PCA on the weighted shape space account for 94.8% (wPC₁ = 85.8%, wPC₂ = 7%, wPC₃ = 3.0%) of the total variation. In comparison to the regular PCA, the wPCA is—by definition—more strongly controlled by the centroid size of the configurations, which is mainly expressed by the domination of PC1.

Similar to the regular PCA plot (Fig. 8), the wPCA morphospace (Fig. 11) can be divided into three main parts: (1) The extremely depressed earliest whorls cover the lower left quadrant (low PC1 and PC2 values); (2) the center of the plot (PC1 equals 0, PC2 positive) is occupied by the more depressed whorls of juveniles and (3) adult whorls are associated with positive PC1 values. In contrast to the PCA, the wPCA reveals a more distinct separation of the type A and type B groups of adult whorls (see Table 3). Representatives of the more depressed type B clearly occupy the lower right quadrant (positive PC1 and





Full-size DOI: 10.7717/peerj.10931/fig-11

negative PC2 values). This division into type A and B can also be seen in the mean shapes of the whorl 5.5 of the respective species (Fig. 10).

DISCUSSION

Members of the family Ceratitidae show high intraspecific variation and strongly overlapping morphospaces (Table 1, Figs. 3, 4 and 5). The ornamentation, which is often regarded as essential for the description of Mesozoic ammonoid groups (*Klug et al.*, 2015; Klug et al., 2015b), is not a unique characteristic among the family Ceratitidae. A better feature to delineate the ceratitids studied here appears to be the shape of the whorl section. The latter, however, cannot be quantified adequately by traditional morphometric methods (Neige, 1999). Accordingly, the utility of conventional taxonomic and morphological methods is limited in this regard. Here, we utilize landmarks and semi-landmarks on ontogenetic cross-sections. Since previous geometric morphometric studies on mollusks all focus either on conch shape or on single (isolated) ontogenetic stages (landmarks: e.g., Johnston, Tabachnick & Bookstein, 1991; Neige & Dommergues, 1995; Reyment & Kennedy, 1998; Stone, 1998; Neige, 1999; Reyment, 2003; Van Bocxlaer & Schultheiß, 2010; Knauss & Yacobucci, 2014; Fourier analysis: e.g., Courville & Crônier, 2005; Simon, Korn & Koenemann, 2010; Simon, Korn & Koenemann, 2011; Korn & Klug, 2012; Klein & Korn, 2014) they cannot be regarded as ontogenetic studies. This study, investigates the use of geometric morphometric methods (GMM) with respect to their usefulness in ontogenetic developmental studies and taxonomic descriptions.

Ontogenetic patterns in Ceratitidae

The ontogenetic trajectories of the studied species comprise the biphasic development from strongly depressed to weakly depressed to compressed whorl profiles (Figs. 8–11). It is commonly accepted that sudden changes in ontogenetic allometry often mark the onset of sexual maturity (i.e., *Kullmann & Scheuch*, 1970; *Klug*, 2001; *Klug et al.*, 2015*a*; *Klug et al.*, 2015*b*).

The studied species can be divided into two main ontogenetic groups: Type (A) Truncated trajectories that are associated with depressed adult whorls; type (B) longer, complete trajectories that lead to a compressed adult whorl shape. The process of lengthening and shortening of the trajectories (i.e., related to changes in rate and timing of the development) account for the ontogenetic differentiation of the species in focus. This contrasts a previous traditional morphometric analysis by *Bischof & Lehmann (2020)* of ptychitids, which revealed that the spherocone-cadicone morphospace is much more distinct. The highly ontogenetically differentiated genus *Ptychites* directly differed through characteristic ontogenetic trajectories.

While precise temporal growth rates of ammonoids are unknown (Lécuyer & Bucher, 2006; Knauss & Yacobucci, 2014), a basic assumption herein was that the individual species have similar coiling rates (i.e., the individual species develop the same number of whorls in the course of their life). Modified rate/timing of shape change from any ancestor to any descendent within an evolutionary framework is called heterochrony (Zelditch, Swiderski & Sheets, 2012, p. 317). Between type A and type B species, interspecific variation of the species in focus arises from an acceleration, a special case of peramorphosis; (for discussion of this term, see Alberch et al., 1979; McNamara, 2012) that allows type B species to occupy an extended portion of the morphospace characterized by more compressed whorls. Therefore, the studied ceratitids do not primarily differ in shape, but rather in the timing of the development of individual shapes. Heterochrony as a mechanism in macroevolution is known to be a key driving factor in phenotypic diversification (e.g., Gould, 1977; Alberch et al., 1979; McKinney & McNamara, 1991; Gerber, Neige & Eble, 2007; Gerber, 2011; Korn et al., 2013; Knauss & Yacobucci, 2014). The quantification of patterns of morphologic disparity and the relationship between size and shape (i.e., heterochrony) will be the subject of future studies.

Anisian ammonoid diversity

It is widely agreed that ammonoid diversity reached its maximum during the Triassic period (*House, 1993*; *Brayard et al., 2009*; *Whiteside & Ward, 2011*). Thereby, the late Anisian ammonoid diversity peak was dominated by members of the family Ceratitidae (*Brayard et al., 2009*; supporting material Fig. S2). However, there are a growing number of studies critically questioning diversity peaks by arguing that—to some extent—the high diversity might be artificially inflated by taxonomic over-splitting (*Forey et al., 2004; De Baets, Klug & Monnet, 2013; Knauss & Yacobucci, 2014*).

The results obtained here do justice to the general opinion that ontogenetic trajectories can be a powerful tool to describe (e.g., *Korn & Klug, 2007*) and discriminate ammonoid species (e.g., *Rieber, 1962; Bischof & Lehmann, 2020*): The newly introduced methods

succeeded in statistically discriminating the ontogenetic pathways of the pre-defined ceratitid species. Based on this analysis, the high diversity of the Anisian ammonoid assemblages of Nevada appears not to be artificially inflated and the alpha taxonomy is regarded to be adequate. However, the high morphological resemblances of the investigated species cannot be denied. Therefore, this study supports the main idea of *McGowan (2004)*, *McGowan (2005)* and *Brosse et al. (2013)* that taxonomic diversity and morphological disparity need not necessarily be closely linked.

It is important to be aware of the fact that GMM carry no direct biological information. They help to understand if and how configurations differ, but not what the underlying mechanisms for their morphological development are. In the complex discoidal morphospace landmark-based approaches have proven to be useful to evaluate a priori defined taxonomic groups. Nevertheless, geometric morphometric methods cannot be considered as being a phylogenetic or taxonomic tool per se. But they certainly represent an improvement and valuable supplement to traditional methods.

Why it is worth the effort

There is no doubt that preparation and analysis of ontogenetic cross sections involves a lot of work (*Korn, 2012*). However, geometric morphometric methods (GMMs) open the door to a new world of objectified, statistically quantifiable descriptions. For example, in the case of the fauna described herein, conventional descriptions and traditional morphometric methods did not succeed to differentiate species adequately. Landmarks and semi-landmarks, however, make it possible to statistically quantify shape variations of entire morphologies (*Neige, 1999*) and allow the analysis of shape and size separately (*Hammer & Harper, 2005*).

The high resolution of the ontogenetic trajectories of the herein studied material was achieved owing to the accretionary planispiral growth of ammonoids with conservation of previous growth stages (*Korn, 2012*), which adds an intuitive, relative time-component to the ontogenetic analysis. Even though it is likely that small-scale ontogenetic changes are overlooked at a measurement density of one measurement per 180 degrees, it can be assumed that no major developmental steps were skipped (*Tajika & Klug, 2020*). Leaving out complete ontogenetic stages would most likely prevent the recognition of ontogenetic processes such as heterochrony. If, for example, only the earliest and latest stages of the ceratitid development were analyzed, representatives of type A and type B would differ fundamentally. The accretionary growth of many ammonoid conchs therefore not only adds an individual time component to the analysis, but more importantly ensures that no major developmental steps have been overlooked. This reinforces the general opinion that ontogenetic trajectories of ammonoids are a powerful tool to study evolutionary processes.

CONCLUSIONS

The Anisian ammonoid diversity peak was dominated by the family Ceratitidae (*Brayard et al.*, 2009; supporting material Fig. S2). However, the investigated ceratitid species show high intraspecific variation and sometimes completely overlapping morphospaces. Using conventional methods, ceratitids are often difficult to distinguish. It was therefore assumed

that the high Anisian diversity in Nevada might be artificially inflated by taxonomic over-splitting.

Using a landmark-based geometric morphometric approach, this study succeeded in differentiating the ontogenetic growth of the pre-defined taxonomic entities in the fossil material from the late Anisian Fossil Hill Member in Nevada, USA. Based on the findings of this study, the high Anisian ammonoid diversity in western North America appears not to be unreasonably inflated. In this context, this study furthermore supports the hypothesis that taxonomic diversity and morphologic disparity of Triassic ammonoids were decoupled (*Brosse et al., 2013; McGowan, 2004; McGowan, 2005*). The largest interspecific differences of ceratitids are the result of alterations of the ontogenetic trajectories that are likely linked to heterochronic processes (i.e., differences in timing of ontogenetic changes). This means that the individual species of this group are not solely defined by the morphology they attain at a certain growth stage, but rather by the sum and timing of all of their ontogenetic stages. The statistical quantification of the relationship between size and shape (i.e., heterochrony) will be the subject of future studies. These processes make an ad hoc distinction of the different species particularly challenging.

For a reliable traditional taxonomic identification of the species herein, it is necessary to have several individuals (*Silberling*, 1962) with different ages of the same species from the same stratum. It has furthermore proven to be essential to analyze morphological variation of ceratitids not only between species but also across different ontogenetic stages. Therefore, the significance of ontogenetic studies on ammonoids with regard to taxonomic implications cannot be dismissed. The geometric morphometric methods introduced herein represent a big leap towards more quantitative and objective taxonomic descriptions of ammonoids.

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Competing Interests

The authors declare there are no competing interests.

Author Contributions

- Eva Alexandra Bischof conceived and designed the experiments, performed the experiments, analyzed the data, prepared figures and/or tables, authored or reviewed drafts of the paper, preparation of fossil specimens, and approved the final draft.
- Nils Schlüter conceived and designed the experiments, performed the experiments, analyzed the data, authored or reviewed drafts of the paper, and approved the final draft.
- Dieter Korn conceived and designed the experiments, authored or reviewed drafts of the paper, preparation of fossil specimens, and approved the final draft.
- Jens Lehmann conceived and designed the experiments, authored or reviewed drafts of the paper, and approved the final draft.

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Data Availability

The following information was supplied regarding data availability: Data and code are available in the Supplemental Files.

Supplemental Information

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